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Highlights

- > Filtration technique while maintaining hydrochemical conditions was developed.
- > Applicability of the filtration technique for analyzing colloids was demonstrated.
- > Complex formation of colloids with rare earth elements was suggested.



Graphical abstract

- ² Size and composition analyses of colloids in deep
- ³ granitic groundwater using
- ⁴ microfiltration/ultrafiltration while maintaining in
- ⁵ situ hydrochemical conditions
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12

13 Abstract

In studies of colloids in deep groundwater, a serious problem exists because the properties of the 14 colloids are influenced by changes in the chemistry of groundwater upon exposure to the 15 atmosphere and pressure release during sampling. To collect and analyze colloids in deep 16 groundwater without significantly changing the groundwater 17 chemistry. а 18 microfiltration/ultrafiltration technique, while maintaining in situ hydrochemical conditions (mainly hydraulic pressure and anaerobic condition), was developed. The size distributions and 19 compositions of colloids in granitic groundwater obtained at a depth of 300 m at the Mizunami 20 21 Underground Research Laboratory (MIU) in Japan were studied using the microfiltration/ultrafiltration technique. The shapes and compositions of the fractionated colloids 22 collected on the membrane filter were determined by scanning electron microscopy, 23 energy-dispersive X-ray spectroscopy, and attenuated total reflection infrared spectroscopy. 24 Additionally, partitioning of rare earth elements (REEs) based on colloid size was compared with 25 REE concentrations in the filtered and unfiltered groundwater samples using inductively coupled 26 plasma mass spectrometry. Different types of colloids consisting of inorganic and organic 27 substances were observed and complex formation between the colloids and REEs was suggested. 28 29 Furthermore, characteristic partitioning of REEs depending on colloids size was observed in complex natural environments without chemical disturbance. The current findings are useful for 30 31 understanding migration of radionuclides.

32

33 Keywords

34 Ultrafiltration, Microfiltration, Groundwater, Colloids, Rare earth elements

36 **1. Introduction**

Understanding the behavior of elements in underground environments is important for 37 geological disposal of high-level radioactive waste (HLW), carbon dioxide capture and storage, 38 and groundwater pollution. In particular, the migration of radionuclides is of great concern in the 39 geological disposal of HLW. The behavior of elements in underground environments is mainly 40 controlled by groundwater flow, precipitation, diffusion, and interaction with rocks (e.g., 41 adsorption and filtration). Additionally, it has been clearly shown that the migration velocity of 42 elements either increases or decreases in the presence of colloids (particles and macromolecules 43 44 in the size range from 1 to 1000 nm) [1-4]. Therefore, studying the physicochemical properties (e.g., concentration, size, shape, and chemical composition) of colloids is important for 45 understanding the behavior of elements in underground environments. 46

Colloids in deep groundwater have been studied in several papers. Degueldre et al. [5] 47 discussed the properties of colloids from various geological formations, ranging from crystalline 48 to sedimentary, from organic-rich to organic-poor systems, and from subsurface to very deep 49 groundwater. The authors summarized the influence of geochemical conditions, such as pH, 50 redox potential, and concentration of primary ions, on colloid properties. Geckeis et al. [6] 51 52 reported the size distribution of smectitic colloids and the interaction between humic colloids in groundwater and trace metal ions, such as U, Th, and rare earth elements (REEs), using 53 flow-field flow fractionation (FI-FFF) combined with online inductively coupled plasma mass 54 55 spectrometry (ICP-MS). However, there are difficulties in collecting and analyzing colloids in terms of their size, composition, and interaction with trace elements in groundwater because of 56 the low concentrations of colloids in groundwater [7, 8]. To address this issue, enrichment and 57 fractionation techniques based on colloid size have been developed. Saito et al. [9] reported a 58

colloid enrichment technique, first using ultrafiltration, and then using the focusing technique 59 with a large injection loop and the slot flow technique inherent to FI-FFF. However, the most 60 serious problem is that the physicochemical properties of colloids are likely to change owing to 61 variations in the physicochemical parameters of groundwater, such as pH, redox potential, and 62 hydraulic pressure, and exposure to the atmosphere during sampling [10]. For example, CO_2 63 64 degassing of groundwater resulting from a decrease in the hydraulic pressure may cause an increase in the pH. Precipitation of metal (hvdr)oxides can occur upon oxidation of groundwater. 65 These changes influence the properties of the colloids in groundwater. Thus, a method for 66 67 collecting and analyzing colloids while maintaining in situ hydrochemical conditions (mainly hydraulic pressure and anaerobic condition) without chemical disturbance is required. Moreover, 68 understanding the interaction between colloids and REEs in natural environments is important; 69 REEs can be regarded as analogues of trivalent actinides [11]. The hydrogeochemistry of REEs 70 in groundwater has been studied in terms of their usefulness as tracers for various geochemical 71 processes such as water-rock interactions and hydrothermal activity. REEs are known to have 72 high affinities for colloids owing to their large ionic potential (ionic charge/ionic radius) [12–21]. 73 However, the interaction between colloids and REEs is not well understood. Interactions 74 75 evaluated on a laboratory scale do not necessarily correspond to that measured in in situ experiments possibly because of chemical disturbance. For example, preferential adsorption of 76 heavy REEs (HREEs) onto iron (hydr)oxide colloids has been determined from laboratory 77 78 experiments [22], whereas preferential adsorption of light REEs (LREEs) onto colloids has been reported in in situ experiments [23–25]. 79

The objectives of this study are to develop a microfiltration/ultrafiltration technique while maintaining in situ hydraulic pressure and anaerobic condition and to understand the partitioning

of REEs in the presence of colloids in complex natural environment without chemical 82 disturbance. Using the technique developed herein, we sampled colloids and filtered 83 groundwater in granite at a depth of 300 m that corresponds to the depth of geological disposal 84 sites for HLW according to the Designated Radioactive Waste Final Disposal Act [26]. Moreover, 85 to understand the influence of hydrochemical conditions (mainly hydraulic pressure and 86 87 anaerobic condition) on colloid properties, the groundwater at a depth of 200 m was exposed to the atmosphere for different periods of time and fractionated by microfiltration/ultrafiltration. 88 This study provides useful information on the size distribution and composition of colloids, and 89 90 the interaction between colloids and REEs in complex natural environments.

91

92 **2.** Development of the microfiltration/ultrafiltration apparatus

The microfiltration/ultrafiltration apparatus developed in this study consisted of two filter 93 holders: an inlet pressure regulator and an outlet pressure regulator (Fig. 1). The filter holders 94 were made of stainless used steel (SUS) and designed to maintain anaerobic conditions under 95 high hydraulic pressures (10 MPa). To prevent contamination from SUS, a part of the apparatus 96 that had contact with groundwater was passivated. The concentration of Fe in ultrapure water 97 98 flowing through the apparatus was less than 0.01 μ g/L and that of REEs was less than 0.1 ng/L. The results confirmed that colloid characterization was not influenced by contamination from 99 SUS. Microfiltration/ultrafiltration membranes with different pore sizes were packed into the 100 101 filter holders. For example, when membranes with pore sizes of 0.45 and 0.01 µm were set into the pre- and post-filter holders, respectively, colloids in the size range of 0.01–0.45 µm were 102 103 collected. The inlet and outlet pressure regulator valves connected to the filter holders were used to adjust the differential pressure to prevent excess resistance pressure (~0.4 MPa) of the 104

microfiltration/ultrafiltration membrane filter in the filter holder. Thus, though the pressure of 105 groundwater is higher than the resistance pressure the membrane 106 of filter. microfiltration/ultrafiltration can be performed without failure of the membrane filter. The inlet 107 pressure regulator, outlet pressure regulator, and two filter holders were connected using 108 quick-connects (Swagelok, Solon, OH, USA) with double-end shut-off stems to prevent 109 110 exposure to the atmosphere. Though the filter holders are separated from the inlet and outlet pressure regulators after microfiltration/ultrafiltration, hydraulic pressure and anaerobic 111 condition in the filter holder can be maintained. Following discharge of residual groundwater in 112 113 the filter holder using inert gas (e.g., Ar), the membrane filters were collected from the filter holder in an anaerobic chamber purged with an inert gas. The relatively small package size (180 114 $mm \times 500 mm \times 120 mm$; 3.2 kg) of the microfiltration/ultrafiltration apparatus affords easy 115 transport and assembly in underground facilities. Moreover. the 116 present microfiltration/ultrafiltration technique can be conducted under artesian pressure as the driving 117 force without pumping that may cause artificial alteration. 118

119

120 **3. Materials and methods**

121 *3.1. Sampling point*

Groundwater samples were collected from boreholes 07MI07 and 09MI20. 07MI07 and 09MI20 boreholes are both horizontal boreholes with the lengths of 55 and 102 m, respectively, located on the -200 and -300 m access/research galleries of MIU, respectively. The geological site at MIU consists of Cretaceous granitic rocks (Toki granite; ~60–70 Ma) and Miocene sedimentary rocks (Mizunami Group; ~15–20 Ma). The Mizunami Group unconformably overlies the Toki granite at a depth of ~160 m at the MIU site. The 07MI07 and 09MI20

boreholes are designed for investigations of any hydrochemical changes related to facility 128 construction. The 07MI07 and 09MI20 boreholes are divided into six sections by impermeable 129 packers. The sections are numbered from 1 to 6 according to the distance from the base of the 130 07MI07 and 09MI20 boreholes. Groundwater samples were collected from each section through 131 nylon tubing without exposure to the atmosphere. The outlet of the nylon tubing was directly 132 133 connected to the microfiltration/ultrafiltration apparatus developed in this study using quick-connects with a double-end shut-off stem, i.e., a valve was used to simplify connection and 134 disconnection of the microfiltration/ultrafiltration apparatus from the nylon tubing. 135

136

137 *3.2. Air exposure experiment*

Microfiltration of groundwater samples collected from section 4 of borehole 07MI07 at a 138 depth of 200 m was performed under different exposure times to the atmosphere (0, 180, and 139 43,200 min) following pressure release. The sample exposed to the atmosphere for 0 min was 140 filtered using the apparatus mentioned above under artesian pressure (~1 MPa). A microfiltration 141 membrane with a pore size of 0.2 µm (H020A047A, Advantec Toyo Kaisha, Ltd., Tokyo, Japan) 142 was used. The samples exposed to the atmosphere for 180 and 43,200 min were filtered using a 143 144 filter holder (SAKE-142, Advantec Toyo Kaisha, Ltd., Tokyo, Japan) in the laboratory under gas pressure (~0.3 MPa) after exposure to atmosphere and pressure release in a Teflon bottle. A 145 microfiltration membrane with a pore size of 0.2 µm (H020A142C, Advantec Toyo Kaisha, Ltd., 146 147 Tokyo, Japan) was used.

148

149 3.3. Microfiltration/ultrafiltration under maintained in situ conditions

150 Microfiltration/ultrafiltration of groundwater samples collected from sections 1–6 of borehole

09MI20 at a depth of 300 m was performed using the apparatus mentioned above under artesian 151 pressure. To remove large particles, membrane filters with a pore size of 0.45 µm (A045A047A, 152 Advantec Toyo Kaisha, Ltd., Tokyo, Japan) were fitted into pre-filter holders. A microfiltration 153 membrane with a pore size of 0.2 µm (H020A047A, Advantec Tovo Kaisha, Ltd., Tokyo, Japan) 154 or an ultrafiltration membrane with a molecular weight cut-off (MWCO) of 10 kDa (13622, 155 156 Merck, Darmstadt, Germany) was used for the post-filter. The filtered volume was 1 L for both types of filtration. The duration of the micro- and ultrafiltration was ~ 0.5 and 10 h, respectively. 157 The in situ groundwater pressure as displayed on the inlet pressure gauge remained mostly 158 159 constant and was comparable with the in situ pressure (~ 2 MPa) during filtration. The differential pressure between the inlet and outlet pressures was adjusted to ~ 0.3 MPa at the start of filtration, 160 and remained mostly constant during filtration. The dissolved oxygen (DO) was measured by 161 colorimetric method. The DO concentration in the unfiltered groundwater was less than 0.02 162 mg/L, however, rapidly increased upon exposure to the atmosphere. The DO concentration of 163 groundwater filtered using the microfiltration/ultrafiltration apparatus was less than 0.02 mg/L 164 even after sample collection and differed slightly from that of unfiltered groundwater. This 165 indicates that the hydraulic pressure and anaerobic condition of groundwater are maintained 166 167 during the filtration process.

168

169 *3.4. Characterization of colloids*

170 *3.4.1. ICP–MS measurements*

The concentrations of Fe and REE in the filtered and unfiltered groundwater were determined by ICP–MS (ELAN DRC-II, PerkinElmer Inc., Waltham, MA, USA). The REE concentrations were determined after preconcentration using chelate resin (Chelate Disc Empore, Sumitomo 3M

Ltd., Tokyo, Japan). The calibration curves were constructed from 0.01, 0.05, 0.1, 0.5, 1, 5, and 10 μg/L REE solutions prepared from 10 mg/L REE standard solution (SPEX CertiPrep Ltd. London, UK). Sample preparation and ICP–MS measurements were conducted in a clean room (Class 10,000). The detection limit values of Fe and REEs were 0.01 μg/L and 0.1 ng/L, respectively.

179

180 *3.4.2. Scanning electron microscopy–energy-dispersive X-ray spectroscopy*

After filtration, the 10 kDa MWCO membrane filters were dried in an anaerobic chamber 181 under Ar atmosphere. The dry membrane filters were coated with Os. The filters were imaged 182 using a field-emission scanning electron microscope (FE-SEM, Quanta 200, FEI Co., Hillsboro, 183 OR, USA) equipped with an energy-dispersive X-ray analyzer (EDX, INCA Energy version 4.05, 184 Oxford Instruments plc., Abingdon, UK) under high vacuum conditions. Sample preparation was 185 performed in a clean room (Class 10,000). The accelerating voltage was 15 kV. The 186 magnification was in the range of $100-40,000\times$. The equivalent circle diameters of the colloids 187 were calculated by image processing of the obtained backscattered electron images. 188

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190 *3.4.3. Fourier transform infrared spectroscopy*

After filtration, the 10 kDa MWCO membrane filters were dried in an anaerobic chamber under Ar atmosphere. Each filter was examined by attenuated total reflection infrared spectroscopy (ATR-FTIR). The ATR-IR spectra (650-4000 cm⁻¹) were recorded on a Fourier transform infrared spectrophotometer (Spectrum Spotlight400, PerkinElmer Inc.).

195

196 **4. Results and discussion**

197 *4.1. Air exposure experiment*

To understand the influence of the hydrochemical conditions (mainly hydraulic pressure and anaerobic condition) on the colloid properties, filtrations were performed under different exposure times to the atmosphere (0,180, and 43,200 min) following pressure release.

Table 1 shows the concentration of Fe in the filtered groundwater for different exposure times 201 202 to the atmosphere. The Fe concentration in the filtered groundwater sample at 180 min was lower than that at 0 min. This indicates that Fe colloids larger than 0.2 um formed by oxidation within 203 180 min after sample collection. The Fe concentrations at 180 and 43.200 min were not markedly 204 205 different, suggesting that the formation of Fe colloids may be a rapid process reaching equilibrium quickly. The results of the Fe concentration indicate that exposure to the atmosphere 206 during filtration impedes on the concentration determination of redox-sensitive elements in the 207 filtered groundwater sample. For instance, Fe colloid formation as a result of oxidation may lead 208 to an overestimation of the colloid size. 209

Fig. 2 shows the relative REE concentrations in the filtered groundwater samples for different 210 exposure times to the atmosphere (180 and 43,200 min) after pressure release. Each 211 concentration was normalized to filtered groundwater at 0 min. The data at 180 and 0 min were 212 not markedly different, suggesting that REEs remained in the $<0.2 \mu m$ fraction despite formation 213 of Fe colloids. In contrast to the data at 180 min, the REE concentrations in the filtered 214 groundwater sample at 43,200 min were less than those at 0 min, especially for LREEs (La, Ce, 215 216 Pr, and Nd). This indicates that REE adsorption onto the Fe colloids occurred within 180-43,200 min; however, Fe colloids were formed within 180 min. This discrepancy may be due to the 217 218 degassing of carbon dioxide from groundwater after pressure release. Carbonate ions can form stable complexes with REEs in groundwater, especially with heavier REEs [27] under both 219

pressurized and anaerobic conditions. Most HREEs–carbonate complexes (<0.2 um) were stable 220 after 43,200 min. On the other hand, a part of LREE-carbonate complexes (<0.2 µm) might have 221 dissociated within 180–43,200 min upon degassing of carbon dioxide owing to exposure to the 222 atmosphere and pressure release because the stability of LREE-carbonate complexes are weaker 223 than that of HREEs-carbonate complexes. Then, the free ion of LREEs may be adsorbed onto Fe 224 colloids (>0.2 µm). However, this hypothesis was not verified because the concentration of 225 inorganic carbon and pH were not measured. Based on these results, the oxidation of 226 groundwater influenced the size distribution of the REEs though the REEs (except for Ce and 227 228 Eu) were insensitive to the redox conditions.

229

230 4.2. Size distribution of colloids and REE partitioning

The REE concentrations in the unfiltered groundwater of the 09MI20 borehole are shown in 231 Fig. 3. The groundwater samples collected from all sections contained 0-25 ng/L REEs. The 232 rock-water interaction is the main source of REEs because granitic rocks contain relatively large 233 amounts of REE [28]. The difference in the REE concentrations between the sections suggests 234 spatial variation of REEs. Mizuno et al. [29] reported that the salinity of groundwater changes 235 236 according to variations in the flow conditions of groundwater owing to the shaft construction at MIU. Therefore, the spatial variation of REEs may result from the shaft construction as well as 237 the salinity. Moreover, a short sampling time is required to understand the size distribution of 238 239 colloids and REEs in groundwater during construction of the facilities.

Fig. 4 shows the relative REE concentrations in the filtered groundwater samples collected from the six sections of 09MI20 borehole. The groundwater samples were filtered using membranes with different pore sizes (10 kDa MWCO and 0.2 μ m). Each concentration was

normalized to the unfiltered groundwater. All samples were collected on July 21, 2011 and July 243 26, 2011. For samples collected from section 2, the REE concentrations in the filtered 244 groundwater samples were higher than those of the unfiltered groundwater. This was believed to 245 be due to an analytical error during ICP-MS measurements because the REE concentrations of 246 samples from section 2 were the lowest among samples collected from all sections of 09MI20 247 borehole and comparable with the detection limit of the ICP-MS instrument. Thus, results 248 relating to samples from section 2 are not discussed. The REE concentrations in the filtered 249 groundwater decreased with decreasing membrane pore sizes (except for sample collected from 250 251 section 2). The results indicate REE association with colloids present with size ranges of 10 kDa-0.2 µm and 0.2-0.45 µm. About 50% of the REEs present in all sections (except for section 252 2) associated with colloids with sizes larger than 10 kDa. Preferential association of LREEs with 253 colloids with sizes ranging from 10 kDa to 0.2 µm and from 0.2 to 0.45 µm was observed in 254 samples collected from sections 5 and 6, whereas that of LREEs with colloids with sizes ranging 255 from 0.2 to 0.45 µm was observed in samples collected from section 1. Preferential association 256 of LREEs with colloids, as observed in groundwater, has also been observed in river water [30]. 257

258

4.3. Size distribution and composition of colloids on the ultrafiltration membrane

Fig. 5 shows representative SEM images of the inorganic colloids in the groundwater samples collected from all six sections of 09MI20 borehole. The equivalent circle diameters of the inorganic colloids were estimated by image processing of backscattered electron images. For all sections, the equivalent circle diameters of most of the inorganic colloids were in the range of $0.05-0.45 \mu m$. These results were consistent with the size distribution of the colloids estimated from the size distribution of REEs, as mentioned in section 4.2. In this study, it was difficult to measure the equivalent circle diameters of colloids smaller than $0.05 \ \mu\text{m}$. High-resolution analysis techniques, such as transmission electron microscopy and atomic force microscopy, are required for such measurements.

Fig. 6 (a) shows representative SEM images of the inorganic colloids in the groundwater 269 samples collected from all six sections of the 09MI20 borehole and Fig. 6 (b) shows EDX 270 spectra of the inorganic colloids shown in Fig. 6 (a). The spectra show the presence of C, O, and 271 S that are from the polysulfone membrane filter. Inorganic particles containing Fe believed to be 272 iron sulfide and iron (hydr)oxide were found in samples collected from sections 1 and 3. 273 274 Inorganic particles containing Al, Si, alkaline earths, and alkaline earth metals were found in samples collected from sections 2, 5, and 6. Inorganic particles containing Al and Mg were found 275 in samples collected from section 4. These particles are believed to be fragments of either granite 276 or clay minerals. Because Fe (hydr)oxide and clay minerals have high affinities for REEs 277 [14–18], inorganic colloids seem to form complexes with the REEs. 278

Vilks et al. [31] reported that organic colloids with a range of sizes were present together with 279 inorganic colloids. Thus, the effect of organic colloids on partitioning of REEs was considered. 280 Identification of organic matter by SEM was difficult because the membrane filter was made of 281 282 an organic substance. To confirm the presence of organic matter, IR analysis was performed. The FTIR spectra of all six sections and the Suwannee River humic acid standard (2S101H, 283 International Humic Substances Society) are shown in Fig. 7. All FTIR spectra were 284 285 characterized by a number of absorption bands, exhibiting varying relative intensities. The strong broad band at 2400–3700 cm⁻¹ is generally attributed to O–H stretching of hydroxyl groups 286 involved in hydrogen links, carboxyl groups, and phenol groups. The band at 3292 cm⁻¹, which 287 was weak in the humic acid spectrum, was attributed to the N-H stretch absorption band of 288

amines and amides [32,33]. There were also a few absorption bands at 2923 and 2853 cm⁻¹. 289 These bands were assigned to C-H stretching of methyl (CH₃) and methylene (CH₂) groups of 290 aliphatic chains [34–36]. The bands at 1730 and 1701 cm^{-1} are generally attributed to C=O 291 stretching vibrations, mainly of carboxyl groups and, to a lesser extent, of ketones and aldehydes 292 [34-36]. The band at 1644 cm⁻¹, which was absent from the humic acid spectrum, was ascribed 293 to C=C stretching of aromatic rings [37] and C=O stretching vibration in amides (amide I band) 294 [38]. The band at 1605 cm^{-1} , which was absent from the spectra in all groundwater samples, was 295 ascribed to stretching of the C=C group conjugated with another C=C group (aromatic ring), 296 C=O and/or COO⁻ in unsaturated ketones, carboxylic acids, and amides [39]. The band at 1545 297 cm⁻¹, which was absent from the humic acid spectrum, was ascribed to NH bending in secondary 298 amides (amide II band) [38]. All samples exhibit broad bands in the region 1000–1300 cm⁻¹, 299 which are probably because of the C–O stretching vibration in alcohols, ethers, phenols, 300 carboxylic acids, and esters [38, 40]. For the groundwater samples collected from all sections of 301 the borehole, it was difficult to detect absorption bands in the region of $650-1400 \text{ cm}^{-1}$ because 302 of the strong absorption bands of the blank. 303

The FTIR spectra of organic matter present in samples collected from all sections showed 304 similar shapes and intensities, suggesting that all sections of the 09MI20 borehole were similar 305 structures. Some differences were observed between the FTIR spectra of organic matter and 306 humic acid. Although the organic matter showed some absorption bands relating to nitrogen, 307 such as absorption bands at 1545, 1644, and 3292 cm^{-1} , these absorption bands were weak or 308 different from those of humic acid. These differences indicated that the nitrogen content of the 309 organic matter in the groundwater samples collected from all sections of the borehole was greater 310 than that of humic acid. Previous studies reported that most of the REEs in groundwater bound to 311

humic substances by forming stable complexes [41]. Organic matter present in 09MI20 borehole
may associate with REEs as well as inorganic colloids.

Vilks et al. [31] reported that inorganic colloids consisting of clay minerals, ferrioxide, and 314 carbonate minerals were present in a wide size range with organic matter in groundwater. Saito et 315 al. [9] reported that Al-, Mg-, and Fe-bearing colloids associated with organic matter at a depth 316 317 of 300 m at the MIU site. Additionally, in this study, we reveal that organic matter in groundwater associated with inorganic colloids and clarified the structures of organic matter. 318 Although the results obtained from this study are qualitatively similar to those of previous 319 320 studies [9, 31], our results may give more useful information owing to the newly developed sampling technique that can be operated under in situ hydrochemical conditions. 321

322

323 5. Conclusions

In this study, we developed a microfiltration/ultrafiltration technique that can maintain in situ physicochemical parameters and afford collection of colloids without chemical disturbance. The microfiltration/ultrafiltration technique can solve some of the uncertainties regarding colloid properties such as size distribution and interaction with REEs.

Using the microfiltration/ultrafiltration technique developed in this study, we obtained useful information on the presence of inorganic (e.g., Fe, Al, Mg, and Si) and organic colloids with a wide size range, and REEs characteristic partitioning in granite groundwater at a depth of 300 m. The latter depth corresponds to that of geological disposal sites for high-level radioactive waste. The inorganic colloids are primary minerals (i.e. fragments of granites) or secondary minerals such as clay minerals resulting from dissolution and precipitation of the host rock. Moreover, the findings indicated the co-existence of inorganic and organic colloids, and colloids bound metal ions. REEs characteristic partitioning, such as preferential association of LREEs with colloids with sizes ranging from 10 kDa to 0.2 μ m and from 0.2 to 0.45 μ m in a specific section of borehole, may be due to the influence of the wide variety of colloids. In future studies, we will investigate the effect of these colloids on radioactive behaviors using combined microfiltration/ultrafiltration technique and bulk analyses.

340

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346 **References**

- A.B. Kersting, D.W. Efurd, D.L. Finnegan, D.J. Rokop, D.K. Smith, J.L. Thompson,
 Migration of plutonium in ground water at the Nevada Test Site, Nature 397 (1999) 56–59.
- W.R. Penrose, W.L. Polzer, E.H. Essington, D.M. Nelson, K.A. Orlandini, Mobility of
 plutonium and americium through a shallow aquifer in a Semiarid region, Environ. Sci.
 Technol. 24 (1990) 228–234.
- [3] G.J. Moridis, Q. Hu, Y.S. Wu, G.S. Bodvarsson, Preliminary 3-D site-scale studies of
 radioactive colloid transport in the unsaturated zone at Yucca Mountain, Nevada, J. Contam.
 Hydrol. 60 (2003) 251–286.
- 355 [4] A.P. Novikov, S.N. Kalmykov, S. Utsunomiya, R.C. Ewing, F. Horreard, A. Merkulov, S.B.
- 356 Clark, V.V. Tkachev, B.F. Myasoedov, Colloid transport of plutonium in the far-field of the
- 357 Mayak Production Association, Russia, Science 314 (2006) 638–641.

- C. Degueldre, I. Triay, J.I. Kim, P. Vilks, M. Laaksoharju, N. Miekeley, Groundwater
 colloid properties: a global approach, Appl. Geochem. 15 (2000) 1043–1051.
- [6] H. Geckeis, T.N. Manh, M. Bouby, J.I. Kim, Aquatic colloids relevant to radionuclide
 migration: characterization by size fractionation and ICP-mass spectrometric detection,
 Colloids Surf. A: Physiochem. Eng. Aspects 217 (2003) 101–108.
- J. Gaillardet, J. Viers, B. Dupré, Trace Elements in River Waters, in: J.I. Drever (Ed.),
 Treatise on Geochemistry, Elsevier, Amsterdam (2003) 225–272.
- 365 [8] M. Plaschke, J. Romer, J.I. Kim, Characterization of Gorleben groundwater colloids by
 atomic force microscopy, Environ. Sci. Technol. 36 (2002) 4483–4488.
- 367 [9] T. Saito, Y. Suzuki, T. Mizuno, Size and elemental analyses of nano colloids in deep granitic
 368 groundwater: Implications for transport of trace elements, Colloids Surf. A: Physiochem.
 369 Eng. Aspects 435 (2013) 48–55.
- [10] C. Degueldre, M. Bolek, Modelling colloid association with plutonium: The effect of pH
 and redox potential, Appl. Geochem. 24 (2009) 310–318.
- [11] N. Chapman, J Smellie, Introduction and summary of the workshop, Chem. Geol. 55(3-4)
 (1986) 167–173.
- J.W. Tang, K.H. Johannesson, Speciation of rare earth elements in natural terrestrial waters:
 assessing the role of dissolved organic matter from the modeling approach, Geochim.
 Cosmochim. Acta 67 (13) (2003) 2321–2339.
- [13] Y. Yamamoto, Y. Takahashi, H. Shimizu, Systematic change in relative stabilities of
 REE-humic complexes at various metal loading levels, Geochem. J. 44 (2010) 39–63.
- [14] F.J. Millero, Stability constants for the formation of rare earth-inorganic complexes as a
- function of ionic strength, Geochim. Cosmochim. Acta 56(8) (1992) 3123–3132.

- [15] M. Bau, Scavenging of dissolved yttrium and rare earths by precipitating iron
 oxyhydroxide: experimental evidence for Ce oxidation, Y-Ho fractionation, and lanthanide
 tetrad effect, Geochim. Cosmochim. Acta 63 (1) (1999) 67–77.
- [16] A. Ohta, I. Kawabe, REE(III) adsorption onto Mn dioxide (δ-MnO2) and Fe oxyhydroxide:
- 385 Ce(III) oxidation by δ -MnO2, Geochim. Cosmochim. Acta 65(5) (2001) 695–703.
- [17] F. Coppin, G. Berger, A. Bauer, S. Castet, M. Loubet, Sorption of lanthanides on smectite
 and kaolinite, Chem. Geol. 182 (2002) 57–68.
- [18] Y. Takahashi, A. Tada, H. Shimizu, Distribution pattern of rare earth ions between water and
 montmorillonite and its relation to the sorbed species of the ions, Anal. Sci. 20(9) (2004)
 1301–1306.
- [19] M. Leybourne, W. Goodfellow, D. Boyle, G. Hall, Rapid development of negative Ce
 anomalies in surface waters and contrasting REE patterns in groundwaters associated with
 Zn–Pb massive sulphide deposits, Appl. Geochem. 15(6) (2000) 695–723.
- [20] B. Nelson, S. Wood, J. Osiensky, Partitioning of REE between solution and particulate
 matter in natural waters: a filtration study, J. Solid State Chem. 171(1-2) (2003) 51–56.
- [21] P. Rönnback, M. Åström, J. Gustafsson, Comparison of the behaviour of rare earth elements
 in surface waters, overburden groundwaters and bedrock groundwaters in two granitoidic
- settings, Eastern Sweden, Appl. Geochem. 23(7) (2008) 1862–1880.
- [22] L. Congqiang, W Jiahong, Y Wenhui, Controls of interactions between iron hydroxide
 colloid and water on REE fractionations in surface waters: Experimental study on
 pH-controlling mechanism, Science in China (Series D) 45 (2002) 449–458.
- 402 [23] H. De Baar, M. Bacon, P. Brewer, Rare-earth distributions with a positive Ce anomaly in the
 403 Western North Atlantic Ocean, Nature 301 (1983) 324–327.

- 404 [24] E. Sholkovitz, Chemical evolution of rare earth elements: fractionation between colloidal
 405 and solution phases of filtered river water, Earth Planet. Sci. Lett., 114 (1992) 77–84.
- 406 [25]H. Elderfield, R. Upstill-Goddard, E. Sholkovitz, The rare earth elements in rivers, estuaries,
- 407 and coastal seas and their significance to the composition of ocean waters, Geochim.
 408 Cosmochim Acta, 54 (1990) 971–991
- 409 [26] Minister of Economy, Trade and Industry, Designated Radioactive Waste Final Disposal Act
 410 (Act No. 117 of 2000) (2000) [in Japanese].
- 411 [27] J. Tang, K. Johannesson, Controls on the geochemistry of rare earth elements along a
 412 groundwater flow path in the Carrizo Sand aquifer, Texas, USA, Chem. Geol., 225(1-2)
 413 (2006) 156–171.
- 414 [28] Y. Takahashi, H. Yoshida, N. Sato, K. Hama, Y. Yusa, H. Shimizu, W- and M-type tetrad
 415 effects in REE patterns for water-rock systems in the Tono uranium deposit, central Japan,
 416 Chem. Geol. 184 (2002) 311–335.
- [29] T. Mizuno, D. Aosai, S Shingu, H. Hagiwara, Y. Yamamoto, A. Fukuda, Hydrochemical
 changes associated with construction of Mizunami Underground Research Laboratory,
 Transactions of the Atomic Energy Society of Japan 12 (2013) 89-102.
- 420
- [30] J. Ingri, A. Widerlund, M. Land, O. Gustafsson, P. Andersson, B. Ohlander, Temporal
 variations in the fractionation of the rare earth elements in a boreal river; the role of
 colloidal particles, Chem. Geol. 166(1–2) (2000) 23–45.
- P. Vilks, J. Cramer, D. Bachinski, D. Doern, and H. Miller, Studies of colloids and
 suspended particles, Cigar Lake Uranium Deposit, Saskatchewan, Canada, Appl. Geochem.
 8 (1993) 605–616.

428	[32] H.C. Kim, M.J. Yu, Characterization of natural organic matter in conventional water
429	treatment processes for selection of treatment processes focused on DBPs control, Water
430	Res. 39 (2005) 4779–4789.
431	[33] J.P. Simjouw, E.C. Minor, K. Mopper, Isolation and characterization of estuarine dissolved
432	organic matter: comparison of ultrafiltration and C-18 solid-phase extraction techniques.
433	Mar. Chem. 96 (2005) 219–235.
434	[34] R.M.B.O. Duarte, A.C. Duarte, Application of non-ionic solid sorbents (XAD resins) for the
435	isolation and fractionation of water-soluble organic compounds from atmospheric aerosols,
436	J. Atmos. Chem. 51 (2005) 79–93.
437	[35] V.I. Esteves, M. Otero, A.C. Duarte, Comparative characterization of humic substances
438	from the open ocean, estuarine water and fresh water, Org. Geochem. 40 (2009) 942–950.
439	[36] P.S.M. Santos, M. Otero, R.M.B.O. Duarte, A.C. Duarte, Spectroscopic characterization of
440	dissolved organic matter isolated from rainwater, Chemosphere 74 (2009) 1053–1061.
441	[37] S. Nagao, T. Iwatsuki, K. Hama, Characteristics of dissolved humic substances isolated
442	from groundwater in Tono area, Gifu Prefecture, Japan, J. Nucl. Fuel Cycle Environ. 15
443	(2008) 2 (in Japanese).
444	[38] R.M. Silverstein, F.X. Webster, in: Spectrometric Identification of Organic Compounds,
445	sixth ed., John Wiley & Sons, Inc., Canada. (1997).
446	[39] I. Simkovic, P. Dlapa, S.H. Doerr, J. Mataix-Solera, V. Sasinkova, Thermal destruction of
447	soil water repellency and associated changes to soil organic matter as observed by FTIR
448	spectroscopy, Catena 74 (2008) 205–211.
449	[40] J. Tarchitzky, O. Lerner, U. Shani, G. Arye, A. Lowengart-Aycicegi, A. Brener, Y. Chen,

- 450 Water distribution pattern in treated wastewater irrigated soils: hydrophobicity effect, Eur. J.
 451 Soil Sci. 58 (2007) 573–588.
- 452 [41] J.P.L. Dearlove, G. Longworth, M. Ivanovich, J.I. Kim, B. Delakowitz, P. Zeh, A study of
- 453 groundwater-colloids and their geochemical interactions with natural radionuclides in
- 454 Gorleben aquifer systems, Radiochim. Acta 52/53 (1991) 83–89.

456 Figure captions

- 457 Fig. 1. Schematic illustration of the microfiltration/ultrafiltration apparatus operating
 458 under pressurized/anaerobic conditions.
- 459 Fig. 2. Relative REE concentrations in the filtered groundwater samples at different
 460 exposure times to the atmosphere (180 and 43,200 min). Each concentration was
 461 normalized to filtered groundwater sample at 0 min.
- 462 Fig. 3. Concentrations of REEs in the groundwater samples collected from all six
 463 sections of the 09MI20 borehole.
- 464 Fig. 4. Relative REE concentrations in the filtered groundwater samples collected from
 465 all six sections of the 09MI20 borehole. The groundwater samples were filtered
 466 using membranes with different pore sizes (10 kDa MWCO and 0.2 μm). Each
 467 concentration was normalized to the unfiltered groundwater sample.
- 468 Fig. 5. Representative SEM images of inorganic colloids in the groundwater samples
 469 collected from all six sections of the 09MI20 borehole. The equivalent circle
 470 diameters of several colloids are shown.
- 471 Fig. 6. Representative SEM images and EDX spectra of inorganic colloids in the
 472 groundwater samples collected from all six sections of the 09MI20 borehole: (a)
- 473 SEM images and (b) EDX spectra of the inorganic colloids shown in Fig. 6 (a).
- 474 Fig. 7. FTIR spectra of the reference humic acid and organic colloids in the
 475 groundwater samples collected from all six sections of the 09MI20 borehole.



Fig. 1. Schematic illustration of the microfiltration/ultrafiltration apparatus

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Fig. 2. Relative REE concentrations in the filtered groundwater samples at different exposure times to the atmosphere (180 and 43,200 min). Each concentration was normalized to filtered groundwater sample at 0 min.



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Fig. 4. Relative REE concentrations in the filtered groundwater samples collected from all six sections of the 09MI20 borehole. The groundwater samples were filtered using membranes with different pore sizes (10 kDa MWCO and 0.2 μm). Each concentration was normalized to the unfiltered groundwater sample.



Fig. 5. Representative SEM images of inorganic colloids in the groundwater samples collected from all six sections of the 09MI20 borehole. The equivalent circle diameters of several colloids are shown.



Fig. 6(a). Representative SEM images and EDX spectra of inorganic colloids in the groundwater samples collected from all six sections of the 09MI20 borehole: (a) SEM images and (b) EDX spectra of the inorganic colloids shown in Fig. 6 (a).



Fig. 6(b). Representative SEM images and EDX spectra of inorganic colloids in the groundwater samples collected from all six sections of the 09MI20 borehole: (a) SEM images and (b) EDX spectra of the inorganic colloids shown in Fig. 6 (a).



Fig. 7. FTIR spectra of the reference humic acid and organic colloids in the groundwater samples collected from all six sections of the 09MI20 borehole. FTIR spectra of the reference humic acid and organic colloids in the groundwater samples collected from all six sections of the 09MI20 borehole.

Average concer	ntrations of tot	tal Fe in filtered	l groundwater.			
			Exposure time			
	0 (min)		180 (min)		43200 (min)	
	Avg.	S.D ^a	Avg.	S.D ^a	Avg.	S.D ^a
Fe (µg/L)	5.5	0.52	3.2	0.42	2.7	0.28
	1. 1.0					

Table 1Average concentrations of total Fe in filtered groundwater.

^aThe standard deviation of three samples.