

Development of Thermodynamic Databases for Geochemical Calculations

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1999

*Development of Thermodynamic Databases for Geochemical Calculations*Randolph C Arthur¹⁾, Hiroshi Sasamoto²⁾, Masahiro Shibata²⁾Mikazu Yui²⁾, Atsushi Neyama³⁾*Abstract*

Two thermodynamic databases for geochemical calculations supporting research and development on geological disposal concepts for high level radioactive waste are described in this report. One, SPRONS.JNC, is compatible with thermodynamic relations comprising the SUPCRT model and software, which permits calculation of the standard molal and partial molal thermodynamic properties of minerals, gases, aqueous species and reactions from 1 to 5000 bars and 0 to 1000°C. This database includes standard molal Gibbs free energies and enthalpies of formation, standard molal entropies and volumes, and Maier-Kelly heat capacity coefficients at the reference pressure (1 bar) and temperature (25°C) for 195 minerals and 16 gases. It also includes standard partial molal Gibbs free energies and enthalpies of formation, standard partial molal entropies, and Helgeson, Kirkham and Flowers (HKF) equation-of-state coefficients at the reference pressure and temperature for 1147 inorganic and organic aqueous ions and complexes. SPRONS.JNC extends similar databases described elsewhere by incorporating new and revised data published in the peer-reviewed literature since 1991.

The other database, PHREEQE.JNC, is compatible with the PHREEQE series of geochemical modeling codes. It includes equilibrium constants at 25°C and 1 bar for mineral-dissolution, gas-solubility, aqueous-association and oxidation-reduction reactions. Reaction enthalpies, or coefficients in an empirical $\log K(T)$ function, are also included in this database, which permits calculation of equilibrium constants between 0 and 100°C at 1 bar.

All equilibrium constants, reaction enthalpies, and $\log K(T)$ coefficients in PHREEQE.JNC are calculated using SUPCRT and SPRONS.JNC, which ensures that these two databases are mutually consistent. They are also internally consistent insofar as all the data are compatible with basic thermodynamic definitions and functional relations in the SUPCRT model, and because primary experimental and field observations that constrain these data are consistently evaluated

within this modeling framework. The accuracy of the data in SPRONS.JNC is evaluated in the present study and elsewhere by comparison of calculated equilibrium constants with their experimental counterparts at pressures and temperatures that span much of the subcritical and supercritical regions of H₂O stability. Additional experimental investigation of mineral solubilities and aqueous reactions, particularly between 0 and 100°C, are needed to further assess, and refine if necessary, the reliability of these databases. Field studies on phase equilibria in near-surface geological environments may be useful for this purpose because associated reaction times are greater than can be accommodated experimentally. The effects on mineral-solution equilibria of metastability and solid solution, and differences in the crystallinity and state of order/disorder in minerals, must be determined, however, before reliable thermodynamic properties can be retrieved from field investigations.

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地球化学計算のための熱力学データベースの開発

(研究報告)

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要 旨

本報告書では、高レベル放射性廃棄物地層処分の研究開発をサポートする地球化学計算のための2種類の熱力学データベースについて報告する。一つは、SPRONS.JNCであり、この熱力学データベースは、SUPCRT モデル・ソフトウェアの中で考慮されている熱力学的な関係式を基にしており、SUPCRT ソフトウェアを用いることで、広範囲にわたる圧力・温度条件（圧力：1～5000 bars，温度：0～1000℃）下での鉱物、ガスおよび水溶液化学種の標準モル熱力学特性・部分モル熱力学特性を計算することができる。SPRONS.JNC は、195 種類の鉱物・16 種類のガスについて、25℃，1 bar の条件下での生成標準モルギブス自由エネルギー、生成標準モルエンタルピー、標準モルエントロピーおよび体積に関するデータを含むと共に、Maier-Kelly 式による熱容量の温度依存式を適用した場合の各係数についても整備している。また SPRONS.JNC は、1147 種類におよぶ無機・有機イオンや錯体について、25℃，1 bar の条件下での生成標準モルギブス自由エネルギー、生成標準モルエンタルピー、標準部分モルエントロピーおよび Helgeson-Kirkham-Flowers (HKF) 式による温度依存に係わる係数も整備している。SPRONS.JNC は、1991 年以降に公開された新たなデータや改訂されたデータを取り入れたものであり、他の同様な熱力学データベースの拡張版である。

もう一つのデータベースは、PHREEQE.JNC であり、これは、地球化学計算コードである「PHREEQE」で用いることができる熱力学データベースである。PHREEQE.JNC は、鉱物の溶解反応、ガスの溶解反応、水溶液化学種を含む反応、酸化還元反応について、25℃，1 bar の条件下での平衡定数を整備している。また PHREEQE.JNC では、0～100℃，1 bar の条件下での平衡定数を算出できる様に反応エンタルピーあるいは平衡定数 ($\log K$) の温度依存に係わる各係数についても整備している。

PHREEQE.JNC に含まれる平衡定数、反応エンタルピーおよび平衡定数の温度依存に係わる係数の全ては、SUPCRT ソフトウェアおよび SPRONS.JNC を用いて計算されているので、PHREEQE.JNC と SPRONS.JNC は、同一の熱力学データベースである。またこれらの熱力学データベースに含まれる全てのデータは、基本的な熱力学の定義や SUPCRT モデルで用い

られている数式に適合し、さらに、基礎的な実験データやフィールドでの観察結果をもとに整合性がとられているという点では、これらの熱力学データベースは、内部整合性がとれたデータベースであると言える。SPRONS.JNCに含まれるデータの精度に関しては、本報告書の中でも確認されており、また他の研究者らによる、水の臨界領域付近まで温度・圧力条件をパラメータとした実験で求められた平衡定数と計算値との比較等からも確認されている。今後は、特に0～100℃の条件下における鉱物の溶解度や溶液化学種の反応に関する実験的研究が必要であり、また必要に応じて、これらデータベースの信頼性を向上させていく必要がある。但し、これらのデータを実験的に研究するには、比較的長い期間を要するため、浅地中環境下での相平衡に関するフィールド調査が有効であるかもしれない。しかしながら、信頼性のある熱力学的特性がフィールド調査からもたらされる前に、準安定、固溶体、結晶度による違い、鉱物における秩序－無秩序（転移）が鉱物－溶液間の平衡に与える影響について把握しておく必要がある。

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1 Introduction

The Japanese repository concept for permanent disposal of high-level nuclear wastes is based on an integrated system of natural and engineered barriers to radionuclide migration. A crystalline or sedimentary host rock (into which the wastes will be emplaced at depths of 1000 m, or 500 m, respectively) will function as the natural barrier. Engineered barriers will include a bentonite buffer, steel canister and vitrified-glass waste form. The natural and engineered barriers will act in concert, such that the wastes are isolated from the geosphere for long periods of time (e.g., 1000 years), and subsequent releases of radionuclides to the biosphere are below levels that would pose an unacceptable risk to the public's health.

The Japan Nuclear Cycle Development Institute (JNC) is carrying out an applied R&D program to evaluate the feasibility of this concept¹. The program is broad in scope, and includes field studies of various geological environments (Tono, Kamaishi and Mobarra sites), laboratory investigations of chemical and transport phenomena (ENTRY and QUALITY facilities), and modeling-based assessments of repository performance (PNC, 1992). A number of these studies deal specifically with geochemical and chemical engineering issues, including:

- the stability of engineered barriers in the disposal environment over time scales of thousands to millions of years,
- the solubility, aqueous-speciation and sorption behavior of radioelements and non-radioactive elements, and
- general and site-specific geochemical processes controlling groundwater evolution.

These issues can be resolved with the aid of thermodynamic models, which are routinely used by JNC and others to interpret the results of experimental and field studies (Sasamoto *et al.*, 1999a, b, c; Iwatsuki and Yoshida, 1999; Oda *et al.*, 1996; Sasaki *et al.*, 1995), to predict the long-term chemical evolution of natural and engineered-barrier systems (Oda *et al.*, 1999; Yui *et al.*, 1999; PNC, 1992, Yui *et al.*, 1992b), and to evaluate associated impacts on repository performance (Azuma *et al.*, 1999; PNC, 1992; Yui *et al.*, 1992a).

Thermodynamic models require basic thermodynamic data over an appropriate range of temperatures and pressures. It is essential that the basic data are reliable, because, as stated by Silva *et al.* (1995), "... the quality of the models cannot be better than the quality of the data they are based on". Although the validity of

¹ JNC was established in 1998, and is responsible for many of the repository R&D functions formerly assigned to the Power Reactor and Nuclear Fuel Development Corporation (PNC).

this statement is beyond dispute, it is useful to clarify what the terms "reliable" and "quality" mean in the context of thermodynamic databases. Here we adopt a pragmatic definition of both terms that is consistent with concepts underlying recent advances in the development and management of thermodynamic databases (Helgeson *et al.*, 1978; Berman, 1988, Holland and Powell, 1990; Engi, 1992; Grenthe *et al.*, 1992; Silva *et al.*, 1995; Gottschalk, 1997). Thus, a database is considered reliable (or to be of high quality) if the following conditions are met, or are approximated as closely as possible:

- *internal consistency* - internal consistency is established among all the data, and is maintained as the database is updated and revised (the concept of internal consistency is discussed further in Section 3),
- *experimental basis* - selected values of thermodynamic parameters are based on evaluations of all available types of relevant experimental (and field) studies, not just on subsets of preferred experimental technique, or over arbitrary ranges of temperature and pressure,
- *documentation* - all procedures in the evaluations are documented such that selected data are *traceable* to original experimental sources, and that these sources are *scientifically defensible* and *reproducible*, and
- *verification of accuracy* - the accuracy of data retrieved from the evaluations is verified by comparison of calculated thermodynamic quantities with their experimental counterparts.

Unfortunately, most published compilations of thermodynamic data are not internally consistent, have not been evaluated on the basis of all (or even most) available types of experimental and field studies, and are poorly documented with respect to data sources and reasons for data selection or rejection (Engi, 1992; Nordstrom and Munoz, 1985; - important exceptions are summarized in Section 3, however). For this reason, PNC¹ commissioned a two-year project beginning in 1996 to develop a thermodynamic database that adheres as closely as possible to the conditions stated above. The project was limited in scope to data for minerals, gases and aqueous species characteristic of geologic systems, and some engineered barriers (e.g., bentonite "clays" and corrosion products of the steel canister). The project was also intended to complement, and to be consistent with, similar efforts carried out by PNC to obtain reliable thermodynamic data for actinide and fission-product elements (Yui *et al.*, 1992a).

The results of the project are summarized in this report, where two new databases, SPRONS.JNC and PHREEQE.JNC, are described, and where the basic philosophy and associated procedures used to generate them are documented. The report is organized as follows. General comments addressing the objectives of this study in relation to JNC's overall R&D program are summarized in Section 2. A survey of existing, reliable thermodynamic databases appropriate for geologic systems is summarized in Section 3. The

results of the survey are used to define a general modeling-based strategy for database development and management. The model, SUPCRT² (Johnson *et al.*, 1991; Johnson *et al.*, 1992) is consistent with SPRONS.JNC, and is briefly described in Section 4. The SPRONS.JNC and PHREEQE.JNC databases are described in Section 5, and listings of these databases are provided in Appendices A and B, respectively. The accuracy of these databases is assessed in Section 6 by comparison of experimental and calculated equilibrium constants for selected heterogeneous reactions (mineral and gas solubilities, Appendix C) and aqueous-speciation reactions (Appendix D).

² SUPCRT is an abbreviation of "supercritical"; SPRONS stands for Sequential-access (ASCII) file of PROPERTIES of Natural Substances

2 Objectives

As noted by Johnson *et al.* (1991), thermodynamic databases are never completed, only abandoned. There is always a need to uncover and correct any errors in the data (not all of which are merely typographical), and to update and revise the database as the results of new experimental and field studies become available. The present study is therefore considered to be an initial effort to implement a general strategy for managing thermodynamic data within a complex and evolving R&D program. SPRONS.JNC and PHREEQE.JNC are the first products of this strategy, but it is reasonable to expect that these databases, and others, will be refined and improved by JNC in the future.

With this in mind, the objective of the study is to develop a thermodynamic database that is both *reliable* and *relevant* for conditions considered in the Japanese repository concept. The range of temperatures (T), pressures (P) and fluid- and solid-phase compositions embraced by this concept are bounded by:

- temperatures between 0 and 100°C, but also higher temperatures that may be considered in alternative scenarios of repository evolution (e.g., the so-called volcanism scenario),
- a total (fluid) pressure between 1 bar and hydrostatic pressures at the depth of the repository ($\approx$ 50 to 100 bars), and
- minerals, gases and aqueous species characteristic of geologic environments and engineered barriers (essentially all components of the barriers that are not radioactive).

As noted in the introduction, however, the study is also based on the premise that a reliable thermodynamic database must be compatible with all relevant types of experimental studies (e.g., calorimetry, phase equilibrium, solubility, electrochemical, etc.) over as broad a range of temperatures and pressures as possible. For this reason, the first two conditions noted above are expanded in scope to consideration of P - T - ρ (where ρ refers to density) regions of stability of fluid phases of H₂O that include both subcritical and supercritical domains. This includes high-temperature ($\approx$ 200 - 1000°C) and high-pressure ($\approx$ 1 - 50 kb) conditions considered in hydrothermal and phase-equilibrium experiments.

Two databases are developed in this study to best accommodate this objective. One is adopted as a *primary* database. It is a *type S* database in the sense defined by Engi (1992), and therefore includes standard molal and partial molal properties of minerals, gases and aqueous species at the reference temperature and pressure, and empirical coefficients enabling calculation of corresponding thermodynamic properties at other temperatures and pressures. These data are obtained from a number of literature sources (Section 5), where procedures to retrieve the data from the results of experimental or field studies (see Section

4.3), or by various estimation techniques (see Section 4.4), are documented. The accuracy of these data is assessed in the present study by comparing calculated equilibrium constants with corresponding values determined experimentally, or by inferences drawn from observations of mineral stability relations in natural systems (Section 6). These comparisons are often at conditions of temperature and pressure that lie considerably outside the range of conditions nominally considered in the Japanese repository concept. Thus the primary database, and SUPCRT, are used to

- assess the accuracy of basic thermodynamic data over a broad range of experimental temperatures and pressures, and
- calculate thermodynamic properties of relevant reactions at temperatures and pressures considered in the repository concept.

These dual uses of the primary database address the requirements of reliability and relevance, respectively, which underlie the main objective of this project.

SPRONS.JNC is the primary database developed in this study. For minerals and gases, it includes standard molal Gibbs free energies and enthalpies of formation, and standard molal entropies and volumes, at a reference pressure of 1 bar and reference temperature of 25°C. Empirical coefficients enabling calculation of standard molal heat capacities and associated thermodynamic quantities at other temperatures and pressures are also documented. For aqueous species, standard partial molal Gibbs free energies and enthalpies of formation, and standard partial molal entropies at the reference temperature and pressure are reported, as are equation-of-state parameters for calculation of standard partial molal volumes and heat capacities, and associated standard partial molal properties as functions of temperature and pressure. The thermodynamic properties of reactions, including equilibrium constants and reaction enthalpies, are calculated using SUPCRT and SPRONS.JNC over a range of temperatures and pressures that encompass the subcritical and supercritical regions of H₂O stability (Section 4).

The other database is considered an *applications* database, or *type K* database according to Engi's classification (Engi, 1992). Thermodynamic data are in the form of equilibrium constants (*K*) for hydrolysis and aqueous-association reactions, and the temperature dependence of the equilibrium constant is calculated using either the van't Hoff equation (requiring the reaction enthalpy at the reference temperature and pressure; e.g., Langmuir, 1997) or an empirical equation representing $\log K(T)$. The application database directly supports geochemical models that are used in safety assessments based on JNC's reference scenario. Repository temperatures in this scenario are determined only by the geothermal gradient and a relatively short-lived thermal excursion due to radiogenic heating. A temperature range of 0 to 100°C fully brackets expected repository temperatures in this scenario. The corresponding pressures of interest

will, as noted earlier, be in the range of 1 bar to that determined by the hydrostatic pressure at the depth of the repository (50 - 100 bars). Because such low pressures have negligible effects on most heterogeneous and homogeneous reactions, however, a pressure of 1 bar is considered adequate for these conditions.

PHREEQE.JNC is the applications database developed in this project. It supports the PHREEQE series of geochemical modeling codes (Parkhurst *et al.*, 1980), and therefore consists of equilibrium constants and reaction enthalpies at the reference pressure and temperature, or empirical coefficients that can be used in a polynomial equation to calculate equilibrium constants between 0 and 100°C at 1 bar.

All thermodynamic data in PHREEQE.JNC are calculated using SUPCRT and SPRONS.JNC. Consistency between the primary and applications databases is therefore ensured. This level of consistency among these type S and type K databases can be extended to other modeling software, which is advantageous because several geochemical modeling codes, including EQ3/6 (Wolery, 1992), the Geochemist's Workbench (GWB; Bethke, 1996) and several of the PHREEQE series of codes (PHREEQE, PHREEQE-60 and PHREEQC), will be used by JNC in specialized research areas. Consistency among thermodynamic data in the primary database and other databases used by JNC (i.e., data for minerals, gases and aqueous species of radioelements) are assessed by comparison of data that are common among the databases to determine if they are compatible (i.e., values are identical within stated uncertainty limits). Results are discussed in Section 5.2.6.

3 Status of Thermodynamic Databases

Four databases have been developed over the past twenty years that fulfill to varying degrees the conditions noted in the introduction that we have adopted as standards for a reliable database. Some remarks on these conditions are summarized below. This is followed by a brief description of the four databases. Supplementary comments on other thermodynamic databases are also included for completeness. We then outline the general strategy used to develop SPRONS.JNC and PHREEQE.JNC.

3.1 Database Requirements

3.1.1 Internal consistency

Internal consistency ensures that there are no sources of ambiguity in the database. Such ambiguities are revealed when mathematical manipulation of the data in various ways results in two (or more) different values for a given thermodynamic property. The two values are mutually incompatible, and therefore internally inconsistent with respect to the data used to calculate them. If, on the other hand, a database is internally consistent, and all such discrepancies are therefore resolved, then data that are in conflict with experimental observations can be attributed unequivocally to errors in the data, or to errors in the experimental results. Internal consistency is thus a conditional requirement, which must be met before the accuracy of a database can be unambiguously assessed.

Engi (1992) suggests that the internal consistency of a database is best evaluated in terms of a level of increasingly stringent conditions:

- all the data are compatible with basic thermodynamic definitions, and basic functional relations used to retrieve thermodynamic data from experimental results,
- a single set of reference values (e.g., reference temperature and pressure) and constants (e.g., gas constant, atomic weights, etc.) is adhered to,
- the interdependence of the data is minimized by simultaneous evaluation of experimental results for multiple reactions,
- parameter values are constrained by all relevant experimental (and field) data, except those that are rejected (or uncertainties “relaxed”) on the basis of experimental procedures employed.

Thermodynamic databases are then classified as:

- *formally* consistent if only the first two conditions are satisfied,

- *partially* consistent if the third condition is also satisfied, and
- *fully* consistent if all four conditions are satisfied.

According to this definition, the databases described in Sections 3.2 and 3.3 are *at least* partially consistent because parameter values are determined by simultaneous evaluations of experimental data for many reactions (see Section 4.3.4.3). The databases are not fully consistent in the strictest sense, however, because statistical techniques to identify discordant experimental data, and reasons for ultimately rejecting such data, are inadequately documented. The need for brevity in scientific journals, where all these databases are described, accounts for this general lack of full and complete documentation.

The definition of internal consistency given above is useful because it includes the concept of “levels-of-attainment”. A fully consistent database is thus an ideal standard, which may be extremely difficult to achieve in practice, and to maintain as the database is inevitably updated and revised. It is important to emphasize here that the level of internal consistency does not necessarily correlate in any meaningful way with the accuracy of the data (Section 3.5). Thus, thermodynamic data in an uncritical data compilation may still be accurate.

3.1.2 Experimental basis

We adopt the premise of Berman (1988) that “... *it is only possible to evaluate the validity of thermodynamic properties by comparison with all available data, because subsets of preferred experimental data are never sufficiently precise to define fully all thermodynamic properties*”. This view is based in part on the work of Helgeson *et al.* (1978), who realized that enthalpies of formation had not generally been determined calorimetrically with the accuracy needed to reproduce the results of high temperature and high pressure phase-equilibrium experiments [see also Berman and Brown (1985)]. Internally consistent databases developed since 1978 (Sections 3.2 and 3.3) are therefore constrained by both phase-equilibrium data and a limited amount of reference, or “anchor”, calorimetric values. Berman’s premise is further supported by more recent investigations, which show that thermodynamic properties retrieved from solubility and mineral-stability experiments at relatively low-temperatures and pressures are in some cases inconsistent with corresponding data retrieved from both calorimetric and phase-equilibrium studies (Sverjensky *et al.*, 1991; Pokrovskii and Helgeson, 1995; 1997b). This suggests that reliable thermodynamic data for minerals, gases and aqueous species must be properly constrained by all available types of experimental (and field) data spanning as broad a range of temperatures and pressures as possible.

Nordstrom *et al.* (1990) apparently contravene these views, however, with the reasonable argument that because reversible solubilities for many minerals have

never been convincingly demonstrated at low temperatures and pressures, equilibrium constants for corresponding reactions should not be included in a thermodynamic database appropriate for such conditions. We agree with this, but believe such considerations should be exercised when the data are applied to a particular problem, not when the database itself is being developed. There is always some level of interdependence among thermodynamic data, and valuable constraints on the thermodynamic properties of minerals that exhibit reversible equilibrium at low temperatures can therefore be extracted from the results of high-temperature and high-pressure experiments involving minerals that do not exhibit such behavior.

3.1.3 Documentation

Quality assurance is an important component of any applied R&D program, and particularly that being carried out by JNC for the permanent disposal of high-level nuclear wastes. A quality-assured thermodynamic database is sufficiently well documented that results of any calculation using the data can be traced back to original data sources. This requirement is addressed in the present study by documenting appropriate references, compiling a library of the referenced studies, and documenting assessments of the accuracy of selected values.

In addition to quality-assurance issues, it is also important that data analysis procedures used to retrieve thermodynamic parameters from experimental and field studies are adequately documented. This is because independent retrievals may lead to recommended values of thermodynamic properties that are in conflict with accepted values in the database. To resolve these conflicts, it is essential to understand the thermodynamic framework and assumptions supporting the procedures used to retrieve the database values, as well as the statistical/mathematical basis for selecting, and rejecting, experimental observations. Also, it is necessary that the database should be updated and revised as results of new experimental or field studies become available. This also requires familiarity with the details of the data analysis procedures supporting the database.

3.1.4 Verification of accuracy

Verifying the accuracy of thermodynamic data is necessary to build confidence in inferences drawn from the results of geochemical calculations that are based on the data. An essential step in assessing the accuracy of a database is to compare calculated values of thermodynamic parameters (e.g., standard Gibbs free energies of reaction, or corresponding equilibrium constants) with their experimental counterparts over as broad a range of temperatures and pressures as possible. Such comparisons are effectively documented in the form of diagrams, in which the reported uncertainty of each experimental value is also portrayed. A more rigorous test of the accuracy of the data is to compare calculated results

with constraints derived from geologic observations, particularly if these constraints refer to conditions of temperature, pressure and time that have not been simulated experimentally. Activity diagrams are especially useful in this regard because calculated equilibrium relations among minerals, gases and aqueous solutions in various chemical subsystems can be simultaneously correlated with the observed compositions of natural waters. Inconsistencies revealed by such correlations are common at relatively low temperatures and pressures, however, and result from the effects of metastability and from uncertainties in the state of crystallinity and degree of order/disorder in the minerals considered (e.g., Helgeson *et al.*, 1978).

3.2 Internally Consistent Databases for Minerals, Gases and Aqueous Species

Databases that partially fulfill the requirements discussed in Section 3.1, and which also include data for minerals, gases and aqueous species, have evolved more-or-less continuously since 1974 through the combined efforts of researchers affiliated with the Laboratory of Theoretical Geochemistry at the University of California, Berkeley. For convenience, the results of this collective effort are referred to here as the SUPCRT model, and its associated databases.

3.2.1 SUPCRT

An internally consistent thermodynamic database developed primarily on the basis of calorimetric and phase-equilibrium data was first developed by Helgeson *et al.* (1978). It includes standard thermodynamic properties for 83 minerals in the system $\text{Na}_2\text{O}-\text{K}_2\text{O}-\text{CaO}-\text{MgO}-\text{FeO}-\text{Fe}_2\text{O}_3-\text{Al}_2\text{O}_3-\text{SiO}_2-\text{TiO}_2-\text{H}_2\text{O}-\text{CO}_2$. Several of these minerals exhibit considerable non-stoichiometry and complications due to temperature-dependent disordering and phase transitions. This study also pioneered the use of equations of state for aqueous species [Na^+ , K^+ , $\text{O}_2(\text{aq})$ and $\text{SiO}_2(\text{aq})$; Helgeson and Kirkham (1974 a,b; 1976), Walther and Helgeson (1977) and Delany and Helgeson (1978)] to retrieve Gibbs free energies of formation of minerals from the results of solubility and mineral-stability experiments. The accuracy of the data is verified for more than 130 reactions in the form of diagrams, where experimental solution compositions, or phase-equilibrium reversal temperatures (see Section 4.3.1.1), are compared with calculated equilibrium constants, or univariant/divariant equilibrium curves.

Johnson *et al.* (1991) utilized subsequent improvements to theoretical models giving more accurate representation of the thermodynamic and dielectric properties of H_2O at subcritical and supercritical temperatures and pressures (see Section 4) to add data for 294 aqueous species, and an additional 96 minerals³,

³ Many of the data for minerals are from Table 9 of Helgeson *et al.* (1978), which are based solely on calorimetric data from several different sources. The reliability of these data was not assessed by Helgeson *et al.* (1978), nor by Johnson *et al.* (1991).

to the original Helgeson *et al.* (1978) database. The revised database is referred to as SPRONS92.DAT, and is included in the SUPCRT software package described by Johnson *et al.* (1991).

Oelkers *et al.* (1995) discuss results of calculations using SUPCRT and a modified version of SPRONS92.DAT, which includes revisions to thermodynamic data for minerals and aqueous species in the system $\text{Al}_2\text{O}_3\text{-H}_2\text{O-NaCl}$ (Pokrovskii and Helgeson, 1995), and unpublished revisions to data for minerals and aqueous species in the system $\text{Na}_2\text{O-Al}_2\text{O}_3\text{-SiO}_2\text{-HCl}$. Standard Gibbs free energies of formation are the only data documented by these authors, however, (the complete database will be described in a future report; E. H. Oelkers, *pers. comm.*).

Revisions and additions of new data to SPRONS92.DAT are also available from the Lawrence Livermore National Laboratory (LLNL), Livermore, California. The revised database (SPRONS96) can be obtained from LLNL at <ftp://s122.es.llnl.gov>. Databases supporting EQ3/6 and the Geochemist's Workbench, which incorporate data in SPRONS96, can also be obtained at this site.

3.3 Internally Consistent Databases for Minerals

Three databases that partially fulfill the requirements discussed in Section 3.1 have been developed for minerals. All these databases also consider the thermodynamic properties of H_2O and simple $\text{H}_2\text{O-CO}_2$ mixtures. None of the databases include data for other aqueous solutes, however, and they are therefore less comprehensive in this regard than the databases discussed in Section 3.2.

3.3.1 Berman (1988)

Berman (1988) derived an internally consistent database constrained primarily by the results of phase-equilibrium studies and a limited number of calorimetric reference values for 67 minerals in the system $\text{Na}_2\text{O-K}_2\text{O-CaO-MgO-FeO-Fe}_2\text{O}_3\text{-Al}_2\text{O}_3\text{-SiO}_2\text{-TiO}_2\text{-H}_2\text{O-CO}_2$. The minerals considered are those for which phase-equilibrium data were available, and for which complications due to solid-solution behavior are minimal. Retrieved values of standard molal Gibbs free energies and enthalpies of formation, and standard molal heat capacities and volumes are reported. Coefficients in an empirical heat capacity equation given by Berman and Brown (1985) are provided, as are coefficients in an empirical equation for calculation of standard molal volumes over an appropriate range of pressures and temperatures. Berman (1988) also lists empirical equations and coefficients to calculate the effects of temperature-dependent phase transitions on the standard molal heat capacity and associated thermodynamic properties of akermanite, alpha cristobalite, hematite, magnetite, alpha quartz and low tridymite, and similar effects of disordering on the standard molal heat capacity and associated thermodynamic properties of dolomite, gehlenite and K-feldspar.

Phase equilibrium data [approximately 1200 experimental results involving 180 equilibria] are used to retrieve values of the thermodynamic parameters using a mathematical programming technique (Berman *et al.*, 1986). The results of calculations using the retrieved data are compared with experimental phase-equilibrium results (44 figures). Good agreement overall is obtained. Less than 1% of the experimental data are discordant with calculated values.

3.3.2 Holland and Powell (1990)

Holland and Powell (1990) derived an internally consistent database on the basis of phase-equilibrium studies and a limited number of calorimetric reference values for 123 minerals in the system $\text{Na}_2\text{O}-\text{K}_2\text{O}-\text{CaO}-\text{MgO}-\text{MnO}-\text{FeO}-\text{Fe}_2\text{O}_3-\text{Al}_2\text{O}_3-\text{SiO}_2-\text{TiO}_2-\text{H}_2\text{O}-\text{CO}_2$. Values of standard molal enthalpies of formation, and standard molal entropies and volumes retrieved from experimental phase-equilibrium studies utilizing calorimetric enthalpies of formation of 23 reference minerals are reported. Coefficients in an empirical heat capacity equation are provided, as are coefficients of thermal expansion and compressibility. Retrieval of standard molal enthalpies of formation is carried out using a least-squares regression technique. Results are documented in a tabular format for several reactions. Good agreement is achieved overall with respect to both the experimental database, and data retrieved by Berman (1988).

3.3.3 Gottschalk (1997)

Gottschalk (1997) derived an internally consistent database constrained primarily by the results of phase-equilibrium studies and a limited number of calorimetric reference values for 94 minerals in the system $\text{Na}_2\text{O}-\text{K}_2\text{O}-\text{CaO}-\text{MgO}-\text{FeO}-\text{Fe}_2\text{O}_3-\text{Al}_2\text{O}_3-\text{SiO}_2-\text{TiO}_2-\text{H}_2\text{O}-\text{CO}_2$. Retrieved values of standard molal enthalpies of formation, and standard molal entropies are reported. Standard molal volumes and coefficients of thermal expansion and compressibilities from various sources are used in the retrievals. Coefficients in an empirical heat capacity equation from Haas and Fisher (1976) are provided. Gottschalk also lists coefficients in an empirical heat capacity equation from Berman and Brown (1985) to extrapolate values beyond the upper temperature range constrained by the calorimetric data. Phase equilibrium data (approximately 5300 experimental results involving 253 equilibria) are used to retrieve values of the thermodynamic parameters using an iterative regression technique. An advantage of this approach is that all available experimental results are considered in the regression procedure, whether they are consistent with other data in the set or conflicting. The results of calculations using the retrieved data are compared with experimental phase-equilibrium results [all figures and references to experimental data are available at <http://www.gfz-potsdam.de/> - navigate to "major projects", then "thermodynamische daten für gesteinsbildende minerale (thermodynamic data for rock-forming minerals)"]. Good agreement overall is obtained, reproducing 92% of the available phase-equilibrium results reported in the literature.

3.4 Other Internally Consistent Databases, and Data Compilations

Other internally consistent databases for minerals that are derived on the basis of phase-equilibrium and calorimetric data are described by Robinson *et al.* (1983), Halbach and Chatterjee (1984), Holland and Powell (1985), Berman *et al.* (1986), Saxena *et al.* (1993) and Berman and Aranovich (1996). These databases pertain to minerals in relatively small chemical subsystems, however, and are therefore of limited use for general geochemical calculations. Although they are not considered here for this reason, these studies are important additional sources of reliable thermodynamic data, which can in some cases be merged with other internally consistent databases with little, if any, conflict (Saxena, 1997).

Robie *et al.* (1978) and Hemingway *et al.* (1982) describe databases for a large number of minerals, and a limited number of gases and aqueous species. These databases are internally consistent, but reported values are evaluated exclusively on the basis of calorimetric investigations carried out by the U.S. Geological Survey. These calorimetric databases have been superseded by the databases described in Sections 3.2 and 3.3 for reasons discussed in Section 3.1.2. The calorimetric data reported by Robie *et al.* (1978) and Hemingway *et al.* (1982) nevertheless provide important constraints on evaluations of phase equilibrium data, and are used to varying degrees in all the databases described in Sections 3.2 and 3.3.

Numerous databases, including most that support geochemical modeling codes such as PHREEQE and EQ3/6, are compilations of log K values obtained from many different sources. These data have not been evaluated simultaneously, nor are they necessarily compatible with all available calorimetric, phase-equilibrium and other experimental constraints. These data compilations are *at best* only formally consistent, and are not considered further here.

There are numerous compilations of thermodynamic data for geochemical calculations that are also, at best, only formally consistent. Nordstrom and Munoz (1985), for example, cite 206 literature sources of thermodynamic data for minerals, gases and aqueous species that are appropriate for geologic systems, and 40 sources on the thermodynamic properties of H₂O.

3.5 Comparison of Reliable Thermodynamic Data for Minerals

Arthur *et al.* (1997) compared standard molal volumes and entropies, and standard molal Gibbs free energies and enthalpies of formation, of minerals that are common to the Helgeson *et al.* (1978), Berman (1988), Holland and Powell (1990), Johnson *et al.* (1991) and Oelkers *et al.* (1995) databases at 25°C and 1 bar [the Gottschalk (1997) database had not been published at the time this comparison was carried out]. Comparisons among any two of these databases are of interest because discrepancies in reported values for a given mineral

indicate that at least one of the values must be in error. The comparisons do not provide a basis for deciding which value, if either, is actually correct, however.

The results are summarized in Figs. 3.5_1 - 3.5_4 where, for convenience, the abbreviations 78HEL, 88BER, 90HOL, 91JOH and 95OEL refer to Helgeson *et al.* (1978), Berman (1988) Holland and Powell (1990), Johnson *et al.* (1991), and Oelkers *et al.* (1995), respectively. Mineral names are abbreviated in the figures as:

- *elements* - gph (graphite), iron (elemental Fe),
- *oxides* - bcri (beta cristobalite), coe (coesite), cor (corundum), hem (hematite), ilm (ilmenite), lime [CaO(*c*)], mt (magnetite), mang (manganosite), per (periclase), q (quartz), bq (beta quartz), ru (rutile), sp (spinel),
- *hydroxides* - br (brucite), dia (diaspore),
- *carbonates* - arag (aragonite), cc (calcite), dol (dolomite), fdol (Fe-dolomite) mag (magnesite), rhc (rhodocrosite), sid (siderite), and
- *silicates* - acm (acmite), ak (akermanite), ab (albite), abh (albite-high), abl (albite-low), alm (almandine), and (andalusite), andr (andradite), ann (annite), an (anorthite), ant (antigorite), anth (anthophyllite), cap (Ca-Al pyroxene), cel (celadonite), chr (chrysotile), clin (clinocllore), cz (clinozoisite), crd (cordierite), cumm (cummingtonite), di (diopside), ed (edenite), ep (epidote), fa (fayalite), fo (forsterite), fs (ferrosilite), geh (gehlenite), gl (glaucophane), gr (grossular), grun (grunerite), hed (hedenbergite), jd (jadeite), kao (kaolinite), ksp (K-feldspar), kals (kalsilite), ky (kyanite), law (lawsonite), ma (margarite), me (meionite), merw (merwinite), mont (monticellite), mu (muscovite), ne (nepheline), pa (paragonite), parg (pargasite), phl (phlogopite), pre (prehnite), pswo (pseudo-wollastonite), py (pyrope), pyhl (pyrophyllite), san (sanidine), sill (sillimanite), sph (sphen), ta (talc), tr (tremolite) and wo (wollastonite).

A notation in the figures, such as 90HOL - 78HEL, signifies that the indicated parameter in one database (i.e., from Helgeson *et al.*, 1978) is subtracted from that of another (i.e., Holland and Powell, 1990).

Differences among standard molal volumes retrieved by Holland and Powell (1990) and those determined by Helgeson *et al.* (1978), Johnson *et al.* (1991) and Berman (1988) are shown in Fig. 3.5_1. As can be seen, there is little discrepancy among these data. This is because molar volume data are calculated on the basis of X-ray cell refinements, which are routinely determined with small standard errors, in the range of 0.1-0.2%. Molar volumes in the 91JOH database for celadonite, cummingtonite, glaucophane, grunerite, merwinite, and nepheline differ significantly compared with corresponding values in the 90HOL and 88BER databases, however.

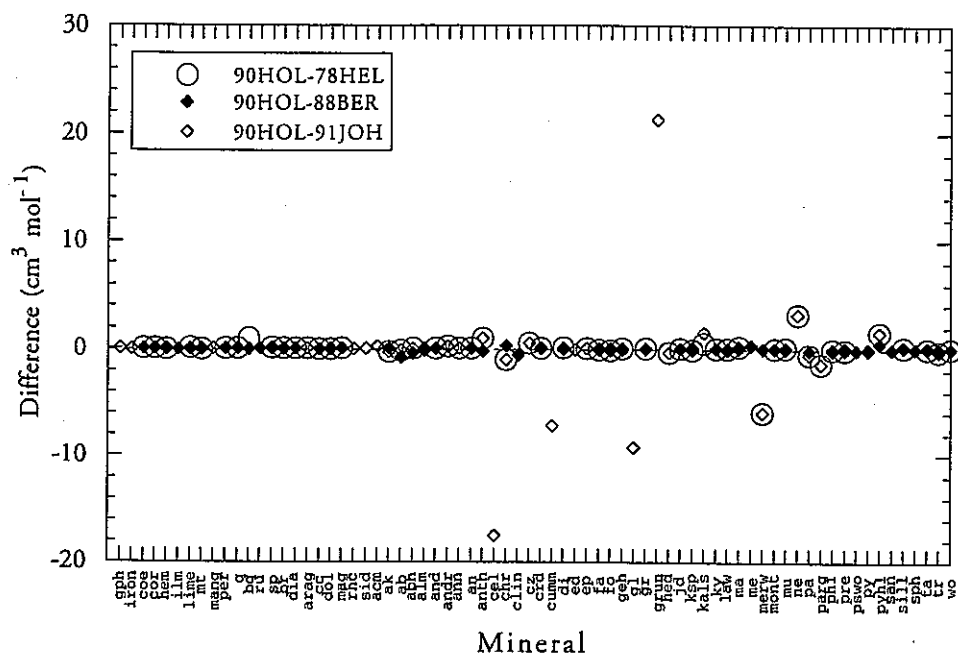


Figure 3.5_1. Differences in standard volumes of minerals between the 90HOL and 78HEL, 88BER and 91JOH databases.

A plot similar to Fig. 3.5_1 is shown for standard entropies in Fig. 3.5_2, where it can be seen that the entropies determined by Holland and Powell (1990) and Berman (1988) agree in most cases, but discrepancies exist for albite, clinocllore, cordierite, gehlenite and phlogopite. Berman (1988) assigns entropy values for

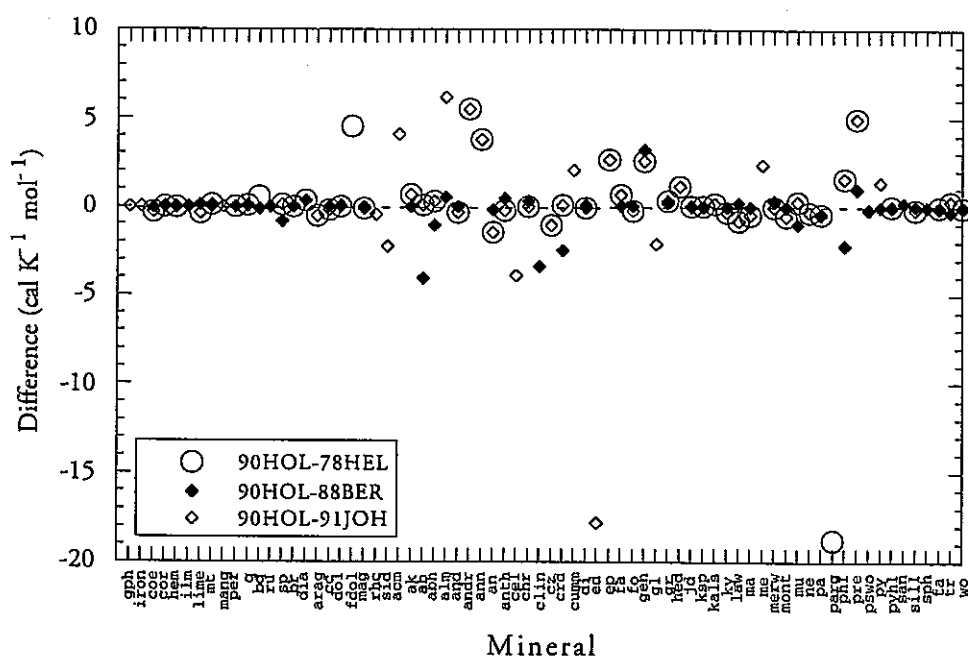


Figure 3.5_2. Differences in standard entropies of minerals between the 90HOL and 78HEL, 88BER and 91JOH databases.

“high” albite to albite, whereas Holland and Powell (1990), Helgeson *et al.* (1978) and Johnson *et al.* (1991), attribute values for “low” albite to this phase. Significant discrepancies (e.g., greater than ± 2 cal K⁻¹ mol⁻¹) exist between the 90HOL and 78HEL and 91JOH databases for ferrous dolomite, siderite, acmite (aegeirine), almandine, andradite, annite, celadonite, edenite, epidote, gehlenite, glaucophane, nepheline, pargasite and prehnite.

A plot similar to Fig. 3.5_1 is shown for standard enthalpies in Fig. 3.5_3. The 90HOL and 88BER data agree reasonably well, but data for albite, high albite, clinochlore, cordierite, and meionite differ significantly (e.g., greater than ± 2 kcal mol⁻¹). The discrepancy for albite may be more apparent than real for reasons noted above. A number of such discrepancies also exist between the 90HOL and 78HEL and 91JOH databases, i.e., for corundum, spinel, akermanite, high albite, andalusite, andradite, anorthite, anthophyllite, clinozoisite, cordierite, grossular, jadeite, kalsilite, kyanite, lawsonite, margarite, merwinite, monticellite, muscovite, nepheline, paragonite, phlogopite, pyrophyllite, sillimanite and tremolite. Many of these minerals are aluminous phases, and the discrepancies could therefore reflect a systematic error in the 78HEL and 91JOH databases (Section 4.3.4.3).

Differences among standard Gibbs free energies of formation retrieved by Berman (1988) and those of Helgeson *et al.* (1978) and Johnson *et al.* (1991) are shown in Fig. 3.5_4 [the 90HOL database does not include free energy values]. Data from Oelkers *et al.* (1995), which, as noted earlier, include several revisions to the 78HEL database, are also shown in the figure for comparison. Significant

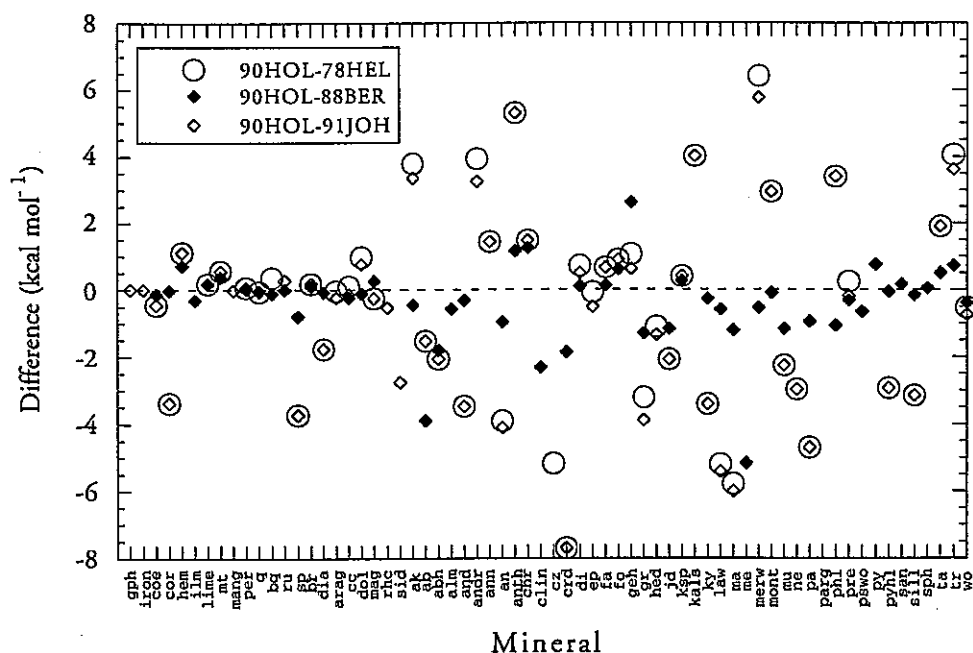


Figure 3.5_3. Differences in standard enthalpies of formation of minerals between the 90HOL and 78HEL, 88BER and 91JOH databases.

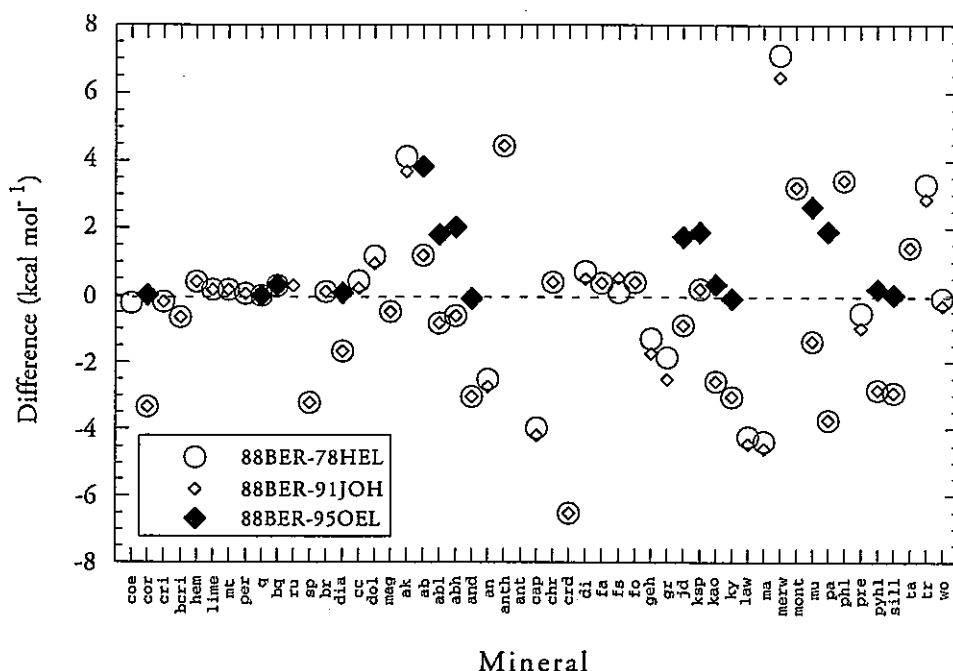


Figure 3.5_4. Differences in standard Gibbs free energies of formation of minerals between the 88BER and 78HEL, 91JOH and 95OEL databases.

discrepancies (e.g., exceeding ± 2 kcal mol⁻¹) are apparent between the 88BER and 78HEL and 91JOH databases for corundum, spinel, akermanite, andalusite, anorthite, anthophyllite, Ca-Al pyroxene, cordierite, grossular, kaolinite, kyanite, lawsonite, margarite, merwinite, monticellite, paragonite, phlogopite, pyrophyllite, sillimanite and tremolite. Revisions incorporated into the 95OEL database, however, agree substantially better with the 88BER data. This improvement includes revised data for corundum, diaspore, andalusite, kaolinite, kyanite, paragonite, pyrophyllite and sillimanite. Moreover, major discrepancies ($> \pm 2$ kcal mol⁻¹) remain only for albite and muscovite. The discrepancy involving albite is probably due to differences between the 88BER and other databases in assigning the properties of either low, or high, albite to albite.

Figures 3.5_1 - 3.5_4 reveal generally good agreement among all the databases with regard to standard molal volumes and entropies, but rather poor agreement with regard to standard enthalpies and Gibbs free energies of formation. This may seem surprising because all the databases are internally consistent and are constrained by similar, and in some cases identical, calorimetric and phase-equilibrium data. Discrepant values could result for a number of reasons, including differences among the databases in adopted calorimetric reference data, and uncertainties that are inherent in extrapolating data that are retrieved from high-temperature, high-pressure experiments to the much lower reference pressure and temperature (Engi, 1992). The 88BER and 90HOL databases generally agree better with each other than they do with either of the 78HEL and 91JOH databases. It is encouraging, however, that

recent revisions incorporated by Oelkers *et al.* (1995) bring a number of formerly discrepant values in the 78HEL and 91JOH databases into significantly better agreement with those of 88BER and 90HOL. The evolution of these databases therefore appears to be on a convergent path. Incorporation of constraints derived from solubility and mineral-stability experiments at relatively low temperatures and pressures may be needed to resolve remaining discrepancies.

3.6 Reliable Databases for Geochemical Calculations

3.6.1 Recommended databases

The databases described in Sections 3.2 and 3.3 are considered the *most reliable* available. Each is internally consistent, and selected values of thermodynamic properties are based on evaluations of at least calorimetric and phase-equilibrium data. Moreover, the accuracy of the selected values is documented in the form of diagrams comparing calculated and experimental values of equilibrium constants, or calculated and experimentally observed pressure-temperature coordinates of univariant or divariant equilibria. Future revisions and expansions of the databases described by Berman (1988), Holland and Powell (1990), Johnson *et al.* (1991) and Gottschalk (1997) are expected.

Of these preferred databases, the SUPCRT model and associated databases, such as that described partially by Oelkers *et al.* (1995), are the *most relevant* to JNC's R&D program. This is because the database includes the thermodynamic properties of minerals, gases and aqueous species, all of which are needed to evaluate mineral-water equilibria at the relatively low temperatures and pressures of primary interest. The databases published by Berman (1988), Holland and Powell (1990) and Gottschalk (1997) are less useful, because they do not include data for aqueous species, and consider only the thermodynamic properties of pure fluid phases of H₂O, and H₂O-CO₂ mixtures.

3.6.2 Databases not recommended

The other databases noted in Section 3.4 are not recommended because:

- they are derived primarily or entirely from calorimetric data (see Section 3.1.2), and/or
- they are uncritical compilations, for which the internal consistency among reported values is highly questionable.

The latter compilations also suffer in many cases from inadequate documentation, and little or no evidence is given that the accuracy of the data has been verified by direct comparison with all available experimental results.

3.6.3 Development strategy

The review of thermodynamic databases summarized above is the foundation of a strategy to develop reliable thermodynamic databases appropriate for JNC's geochemical calculations (Arthur *et al.*, 1997). This strategy includes the following steps:

1. adopt the thermodynamic model and database (SPRONS92.DAT) comprising the SUPCRT software package (Johnson *et al.*, 1991),
2. update SPRONS92.DAT with revised and new data for minerals, gases and aqueous species reported in the literature since 1991,
3. verify the accuracy of the revised database by comparison of calculated and experimental equilibrium constants for important heterogeneous reactions (mineral and gas solubilities), and important aqueous-speciation reactions,
4. revise equilibrium constants and reaction enthalpies at 25°C and 1 bar in the PHREEQE databases, using SUPCRT and SPRONS.JNC, and
5. calculate the values of coefficients in a polynomial expression used by PHREEQE to account for the temperature dependence of equilibrium constants by regression of corresponding values of equilibrium constants calculated using SUPCRT and SPRONS.JNC.

Implementation of this strategy is the subject of the remainder of this report. Section 4 provides background information on the thermodynamic framework of the SUPCRT model. Steps 2, 4 and 5 are discussed in Section 5, and documented in Appendices A and B. The results of Step 3 are discussed in Section 6, with reference to figures in Appendices C and D.

4 Thermodynamic Framework

The SUPCRT model includes:

- an equation of state for H₂O (Levelt-Sengers *et al.*, 1983; Haar *et al.*, 1984),
- equations enabling calculation of the dielectric properties of the solvent as functions of temperature and pressure (Johnson and Norton, 1991),
- the revised HKF (Helgeson-Kirkham-Flowers; Helgeson *et al.*, 1981) equations of state for aqueous species, and equations to calculate their apparent standard partial molal thermodynamic properties as functions of temperature and pressure (Tanger and Helgeson, 1988), and
- equations of state for minerals and gases, and equations to calculate their apparent standard molal thermodynamic properties as functions of temperature and pressure (Helgeson *et al.*, 1978).

The thermodynamic framework represented by these equations as a whole, and their temperature, pressure and solvent-density ($\rho_{\text{H}_2\text{O}}$) regions of validity, has evolved more or less continuously since 1974 (Helgeson and Kirkham, 1974a,b; Helgeson and Kirkham, 1976; Helgeson *et al.*, 1978; Helgeson *et al.*, 1981; Shock and Helgeson, 1988; Tanger and Helgeson, 1988; Johnson *et al.*, 1991 (see also Johnson *et al.*, 1992); Shock *et al.*, 1992; Sassani and Shock, 1992; Shock and Koretsky, 1995; Haas *et al.*, 1995; Shock *et al.*, 1997a,b; Sverjensky *et al.*, 1997). The SUPCRT model permits calculation of the apparent standard molal thermodynamic properties of minerals and gases from 1 to 10000 bars and from 0°C to a maximum value consistent with the Maier-Kelly (Maier and Kelly, 1932) heat-capacity coefficients for the mineral or gas of interest (see below). For charged aqueous species, calculation of the apparent standard partial molal Gibbs free energy of formation is restricted to $\rho_{\text{H}_2\text{O}} \geq 0.35 \text{ g-cm}^3$, and to $T \leq 350^\circ\text{C}$, or $T \geq 400^\circ\text{C}$ at $P < 500$ bars. Calculation of the standard partial molal enthalpy of formation of charged species and their standard partial molal entropy, heat capacity and volume is limited to $\rho_{\text{H}_2\text{O}} \geq 0.35 \text{ g-cm}^3$ and $T \leq 350^\circ\text{C}$ at $P < 1000$ bars.

Thermodynamic relations in the SUPCRT model are briefly summarized in this section with the objective of identifying the primary data needed to calculate the thermodynamic properties of minerals, gases, aqueous species and reactions as functions of P and T (see Sections 4.2.1.1 and 4.2.2.5). These data comprise the primary database (SPRONS.JNC) described in Section 5. In-depth descriptions of the SUPCRT model are included in the references noted above.

4.1 Standard-State and Other Conventions

The standard state for thermodynamic components that are stoichiometrically

equivalent to minerals calls for unit activity of the pure component at all pressures and temperatures. For liquids, including $\text{H}_2\text{O}(l)$, the standard state corresponds to unit activity of the pure solvent at any pressure and temperature. The standard state for gases is unit fugacity of the hypothetical ideal gas at 1 bar and any temperature. For aqueous species other than H_2O , the standard state is unit activity in a hypothetical ideal one molal solution referenced to infinite dilution at any pressure and temperature.

Standard molal Gibbs free energies and enthalpies of formation of minerals, gases and aqueous species are calculated in terms of their *apparent* standard molal counterparts⁴. The apparent standard molal Gibbs free energy and enthalpy of formation ($\Delta G_{P,T}^\circ$ and $\Delta H_{P,T}^\circ$, respectively) are defined at a given pressure and temperature by:

$$\Delta G_{P,T}^\circ \equiv \Delta G_f^\circ + (G_{P,T}^\circ - G_{P_r,T_r}^\circ) \quad (4.1_1)$$

and

$$\Delta H_{P,T}^\circ \equiv \Delta H_f^\circ + (H_{P,T}^\circ - H_{P_r,T_r}^\circ), \quad (4.1_2)$$

where ΔG_f° and ΔH_f° represent the standard molal Gibbs free energy and enthalpy of formation from the elements in their stable phase at the reference pressure ($P_r = 1$ bar) and temperature ($T_r = 25^\circ\text{C}$). The parenthetical terms in these equations refer to differences in the absolute standard Gibbs free energy (G°) and absolute standard enthalpy (H°) caused by changes in pressure ($P - P_r$) and temperature ($T - T_r$).

The thermodynamic properties of aqueous ionic species refer to two half-cell reactions: one involving formation of the ion from its element and the other, by convention, formation of H^+ from $\text{H}_2(g)$. The reactions are always written such that the mass of electrons is balanced when the two half-cell reactions are combined. The number of electrons produced in the combined reaction is thus equal to the number consumed, and it is therefore unnecessary to specify their thermodynamic properties explicitly (e.g., Garrels and Christ, 1965). The standard molal properties of ions derived in this manner are referred to as *conventional* properties⁵. The conventional standard molal properties of aqueous ionic species are related to their *absolute* counterparts by:

4- Apparent standard molal quantities are defined as a matter of convenience to avoid the necessity of accounting for changes in the stable phase of the elements with changes in temperature and pressure (e.g., Anderson and Crerar, 1993). It is unnecessary to account for these changes because the thermodynamic properties of the elements always cancel in any balanced chemical reaction.

5- Conventional standard molal properties of aqueous solutes are equivalent to their conventional standard *partial* molal analogs (Helgeson *et al.*, 1981). These two terms are used interchangeably in the following discussion.

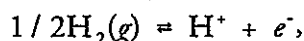
$$\Delta \Xi_{j,P,T}^{\circ} \equiv \Delta \Xi_{j,P,T}^{\circ,abs} - Z_j \Delta \Xi_{H^+,P,T}^{\circ,abs} \quad (4.1_3)$$

and

$$\Xi_{j,P,T}^{\circ} \equiv \Xi_{j,P,T}^{\circ,abs} - Z_j \Xi_{H^+,P,T}^{\circ,abs} \quad (4.1_4)$$

where $\Delta \Xi_{j,P,T}^{\circ}$ and $\Delta \Xi_{j,P,T}^{\circ,abs}$ denote the conventional and absolute standard molal Gibbs free energy or enthalpy of formation of the j -th ionic species, $\Xi_{j,P,T}^{\circ}$ and $\Xi_{j,P,T}^{\circ,abs}$ stand for the conventional and absolute standard molal entropy, heat capacity or volume of the j -th ion, and Z_j represents the ion's charge. The quantity $\Delta \Xi_{H^+,P,T}^{\circ,abs}$ stands for the absolute standard molal Gibbs free energy or enthalpy of H^+ formation, and $\Xi_{H^+,P,T}^{\circ,abs}$ represents the hydrogen ion's absolute standard molal entropy, heat capacity or volume. All conventional standard molal properties of the hydrogen ion are thus equal to zero at any temperature and pressure.

Although the thermodynamic properties of the aqueous electron, e^- , are not considered explicitly in SUPCRT, it is convenient to define these properties in a conventional sense that maintains consistency with the basic thermodynamic framework in this model. This enables the results of calculations using SUPCRT to be directly incorporated into databases that support certain geochemical modeling codes, such as PHREEQE, in which oxidation-reduction reactions are written in terms of e^- . Although the thermodynamic properties of the aqueous electron are in fact unknown (Hostetler, 1984), its conventional thermodynamic properties can nevertheless be defined by assuming that the reaction for formation of the hydrogen ion,



also refers to "formation" of e^- (e.g., Drever, 1982). Thus,

$$\Delta \Xi_{e^-,P,T}^{\circ,abs} = \Delta \Xi_{H^+,P,T}^{\circ,abs} \quad (4.1_5)$$

$$\Xi_{e^-,P,T}^{\circ,abs} = \Xi_{H^+,P,T}^{\circ,abs} \quad (4.1_6)$$

and, as follows from Eqns. (4.1_3) and (4.1_4),

$$\Delta \Xi_{e^-,P,T}^{\circ} = \Xi_{e^-,P,T}^{\circ} = 0. \quad (4.1_7)$$

Thus, as for H^+ , all *conventional* standard molal properties of e^- are equal to zero at any temperature and pressure. An additional assumption, often made in electrochemical studies, that

$$\Delta \Xi_{H^+,P,T}^{\circ,abs} = \left(\Xi_{H^+,P,T}^{\circ,abs} + \Xi_{e^-,P,T}^{\circ,abs} - 1/2 \Xi_{H_2(g),P,T}^{\circ,abs} \right) = 0$$

(i.e., the so-called "hydrogen-ion convention"; Denbigh, 1971; Puigdomenech *et al.*, 1997) is not adopted here, because it results in only a slight simplification of notation, and is otherwise unnecessary (Anderson and Crerar, 1993).

Conventional standard molal properties of an aqueous electrolyte are related to the corresponding conventional properties of its constituent ions by additivity constraints (Shock *et al.*, 1992):

$$\Delta \Xi_{k,P,T}^{\circ} = \sum_j \nu_{j,k} \Delta \Xi_{j,P,T}^{\circ} \quad (4.1_8)$$

and

$$\Xi_{k,P,T}^{\circ} = \sum_j \nu_{j,k} \Xi_{j,P,T}^{\circ} \quad (4.1_9)$$

where subscripts k and j refer to the electrolyte and ion, respectively, and $\nu_{j,k}$ stands for the number of moles of the j -th ion in one mole of the k -th electrolyte. Equations (4.1_3), (4.1_4), (4.1_8) and (4.1_9) imply that the standard molal thermodynamic properties of aqueous anions are equivalent to those of the corresponding fully associated acid electrolyte, and that the standard molal thermodynamic properties of an aqueous electrolyte are equal to the stoichiometric sum of the corresponding properties of the cation and anion.

The units of thermodynamic quantities considered in the SUPCRT model are temperature in Kelvins (°K), pressure in bars, energy in calories (cal), mass in moles (mol) and volume in cubic centimeters (cm³). We retain these units in favor of SI standards to avoid confusion.

4.2 Thermodynamic Relations

The parenthetical terms in Eqns. (4.1_1) and (4.1_2) are evaluated using equations of state for pure minerals and gases, and equations of state for aqueous species that account for changes in both the thermodynamic and dielectric properties of fluid phases of H₂O as functions of temperature and pressure. Thermodynamic functions relate equation-of-state parameters characterizing the standard molal volume to corresponding values of the isobaric heat capacity, entropy, and Gibbs free energy and enthalpy of formation. The functions contain adjustable coefficients determined either by evaluation of experimental data (Section 4.3) or by estimation techniques based on empirical correlations among known values of standard molal or partial molal quantities (Section 4.4).

4.2.1 Minerals and Gases

The equation of state for minerals adopted in the SUPCRT model is given by (Helgeson *et al.*, 1978):

$$V_{P,T}^{\circ} = V_{P_r,T_r}^{\circ} + \sum_{i=1}^{\phi_{PT}} \Delta V_{t,i}^{\circ} \quad (4.2.1_1)$$

where V° refers to the standard molal volume of the mineral and $\Delta V_{t,i}^{\circ}$ stands for the change in the standard molal volume caused by the i -th of ϕ_{PT} phase transitions that occur between P_r, T_r and P, T . Although the effects of variations in temperature and pressure on the standard molal volume are not considered in this equation-of-state, this simplifying approximation is appropriate for most minerals over a broad range of pressures and temperatures. This is because the effects of thermal expansion and compression cause opposing changes in V° on the order of 3% or less at temperatures less than 800°C at 1 bar, and pressures less than 20 kb at 25°C (Helgeson *et al.*, 1978; Berman *et al.*, 1986). Evaluation of mineral compressibilities and expansivities may be important, however, for low-entropy (e.g., solid-solid) reactions and reactions involving highly contrasting volumetric properties (Berman, 1988; Holland and Powell, 1990; Gottschalk, 1997). The equation of state for minerals is also used for gases, in which case $\phi_{PT} = 0$, and $V_{P,T}^{\circ} = V_{P_r,T_r}^{\circ}$.

The apparent standard molal Gibbs free energy and enthalpy of formation of a mineral or gas is calculated using Eqns. (4.1_1), (4.1_2) and (4.2.1_1), and the following equations from Helgeson *et al.* (1978) for the parenthetical terms in Eqn. (4.1_1) and (4.1_2):

$$\begin{aligned} G_{P,T}^{\circ} - G_{P_r,T_r}^{\circ} = & -S_{P_r,T_r}^{\circ}(T - T_r) + \sum_{i=1}^{1+\phi_r} \int_{T_i}^{T_{i+1}} C_{P_{ti}}^{\circ} dT - T \sum_{i=1}^{1+\phi_r} \int_{T_i}^{T_{i+1}} C_{P_{ti}}^{\circ} d \ln T \\ & + \int_{P_r}^P V_{P_r,T}^{\circ} dP - \sum_{i=1}^{\phi_p} \int_{P_{ti}}^P \Delta V_{t,i}^{\circ} dP - \sum_{i=1}^{\phi_r} \frac{\Delta H_{t,i}^{\circ}}{T_{i+1}} (T - T_{i+1}), \end{aligned} \quad (4.2.1_2)$$

and

$$H_{P,T}^{\circ} - H_{P_r,T_r}^{\circ} = \sum_{i=1}^{1+\phi_r} \int_{T_i}^{T_{i+1}} C_{P_{ti}}^{\circ} dT + \int_{P_r}^P V_{P_r,T}^{\circ} dP - \sum_{i=1}^{\phi_p} \int_{P_{ti}}^P \Delta V_{t,i}^{\circ} dP + \sum_{i=1}^{\phi_r} \Delta H_{t,i}^{\circ}. \quad (4.2.1_3)$$

An expression for the standard molal entropy consistent with Eqns. (4.2.1_2) and (4.2.1_3) is given by:

$$S_{P,T}^{\circ} = S_{P_r,T_r}^{\circ} + \sum_{i=1}^{1+\phi_r} \int_{T_i}^{T_{i+1}} C_{P_{ti}}^{\circ} d \ln T + \sum_{i=1}^{\phi_r} \frac{\Delta H_{t,i}^{\circ}}{T_{i+1}} \quad (4.2.1_4)$$

In Eqns. (4.2.1_2) - (4.2.1_4), ϕ_T refers to the number of phase transitions from P_r, T_r to P_r, T (where $T_l = T_r$); ϕ_P stands for the number of phase transitions from P_r, T to P, T ; $P_{i,i,T}$ represents the pressure of the i -th phase transition at T , and $\Delta H_{i,i}^\circ$ denotes the change in the standard molal enthalpy of the i -th phase transition. The isobaric heat capacity, $C_{P,i}^\circ$, is expressed in terms of the Maier-Kelly power function (Maier and Kelly, 1932):

$$C_{P,i}^\circ = a_i + b_i T + c_i T^{-2}, \quad (4.2.1_5)$$

where a_i , b_i , and c_i stand for regression coefficients valid over the temperature range T_i to T_{i+1} .

Substitution of Eqn. (4.2.1_5) into Eqns. (4.2.1_2), (4.2.1_3) and (4.2.1_4) leads to (Helgeson *et al.*, 1978; Johnson *et al.*, 1991):

$$\begin{aligned} G_{P,T}^\circ - G_{P_r,T_r}^\circ = & -S_{P_r,T_r}^\circ(T - T_r) + \sum_{i=1}^{1+\phi_T} a_i \left\{ T_{i+1} - T_i - T_{i+1} \ln \left(\frac{T_{i+1}}{T_i} \right) \right\} \\ & + \sum_{i=1}^{1+\phi_T} \left\{ \frac{(-c_i - b_i T_{i+1} T_i^2)(T_{i+1} - T_i)^2}{2 T_{i+1} T_i^2} \right\} + V_{P_r,T}^\circ(P - P_r) \\ & - \sum_{i=1}^{\phi_P} \int_{P_{i,i,T}}^P \Delta V_{i,i}^\circ dP - \sum_{i=1}^{\phi_T} \frac{\Delta H_{i,i}^\circ}{T_{i,i}}(T - T_{i,i}); \end{aligned} \quad (4.2.1_6)$$

$$\begin{aligned} H_{P,T}^\circ - H_{P_r,T_r}^\circ = & \sum_{i=1}^{1+\phi_T} \left\{ a_i(T_{i+1} - T_i) + \frac{b_i}{2}(T_{i+1}^2 - T_i^2) - c_i \left(\frac{1}{T_{i+1}} - \frac{1}{T_i} \right) \right\} \\ & + V_{P_r,T}^\circ(P - P_r) - \sum_{i=1}^{\phi_P} \int_{P_{i,i,T}}^P \Delta V_{i,i}^\circ dP + \sum_{i=1}^{\phi_T} \Delta H_{i,i}^\circ; \end{aligned} \quad (4.2.1_7)$$

and,

$$S_{P,T}^\circ = S_{P_r,T_r}^\circ + \sum_{i=1}^{1+\phi_T} \left\{ a_i \ln \left(\frac{T_{i+1}}{T_i} \right) + b_i(T_{i+1} - T_i) - \frac{c_i}{2} \left(\frac{1}{T_{i+1}^2} - \frac{1}{T_i^2} \right) \right\} + \sum_{i=1}^{\phi_T} \frac{\Delta H_{i,i}^\circ}{T_{i,i}} \quad (4.2.1_8)$$

Equations (4.1_1), (4.1_2) and (4.2.1_6) - (4.2.1_8) are the fundamental equations in SUPCRT that are used to calculate thermodynamic properties of minerals and gases over a wide range of pressures and temperatures.

4.2.1.1 Primary thermodynamic data for minerals and gases. Inspection of Eqns. (4.1_1), (4.1_2) and (4.2.1_6) - (4.2.1_8) indicates that for gases and minerals that do not undergo phase transitions, the following data are needed to calculate their apparent standard molal thermodynamic properties at a given temperature and pressure: ΔG_f° , ΔH_f° , S_{P_r, T_r}° , V_{P_r, T_r}° , a_i , b_i , c_i and T_{max} (where T_{max} refers to the upper temperature limit of validity of the Maier-Kelly coefficients). These data are also required for minerals that experience phase transitions, in addition to the enthalpy and volume change at each transition temperature, and Maier-Kelly coefficients for each transitional phase and the upper temperature limits of their validity. In practice, the only first-order phase transitions⁶ that have been considered in this framework are those for which the Clapeyron slope is positive (Helgeson *et al.*, 1978). Lambda transitions have also been evaluated to account for the effects of substitutional order/disorder in albite, dolomite, K-feldspar and epidote (Helgeson *et al.*, 1978).

4.2.2 Aqueous species: Summary of the revised HKF model

The structure of the revised Helgeson, Kirkham and Flowers (HKF) model is represented by the total differential of the partial molal Gibbs free energy of the j -th ion (or neutral species) in a hypothetical ideal one molal solution corresponding to the standard state at any temperature and pressure:

$$d\bar{G}_j^\circ = -\bar{S}_j^\circ dT + \bar{V}_j^\circ dP.$$

Changes in $\bar{G}_j^\circ$ caused by variations in temperature and pressure are evaluated by integrating this equation from reference conditions to the temperature and pressure of interest:

$$\int_{P_r, T_r}^{P, T} d\bar{G}_j^\circ = - \int_{T_r}^T \bar{S}_{j, P_r}^\circ dT + \int_{P_r}^P \bar{V}_{j, T}^\circ dP.$$

The HKF model provides a framework for evaluating this equation using semi-empirical equations-of-state for $\bar{V}_j^\circ$, and associated equations for the isobaric standard partial molal heat capacity, $\bar{C}_{P_j}^\circ$ (which are needed to evaluate the first term on the right hand side of this equation). The equations-of-state are semi-empirical in the sense that they account for both solvation (non-empirical) and non-solvation (empirical) contributions to the standard partial molal volume and heat capacity. The equation of state for aqueous ions and neutral species at any

⁶ First-order phase transitions are characterized by discontinuities in the temperature dependence of the entropy, enthalpy, volume, heat capacity, expansivity and compressibility of a mineral. In contrast, lambda transitions result from continuous infinitesimal changes in the internal structure of the mineral, or its distribution of atoms, with variations in pressure and/or temperature. The dependence on temperature of the heat capacity, expansivity and compressibility of a mineral undergoing a lambda transition differ significantly before and after the transition, in contrast to the behavior of first-order transitions (Helgeson *et al.*, 1978).

temperature and pressure⁷ is given by (Tanger and Helgeson, 1988):

$$\bar{V}^{\circ} = \Delta \bar{V}_j^{\circ} + \Delta \bar{V}_n^{\circ} = \hat{C} \left\{ a_1 + \frac{a_2}{\Psi + P} + \frac{a_3}{T - \Theta} + \frac{a_4}{(\Psi + P)(T - \Theta)} - \omega Q + \left(\frac{1}{\varepsilon} - 1 \right) \left(\frac{\partial \omega}{\partial P} \right)_T \right\}, \quad (4.2.2_1)$$

where $\bar{V}^{\circ}$ stands for the standard partial molal volume of the j -th ion. The solvation contribution ($\Delta \bar{V}_j^{\circ}$) is represented by terms containing the a_i equation-of-state coefficients, which are pressure- and temperature-independent and unique to the j -th species. The non-solvation contribution ($\Delta \bar{V}_n^{\circ}$) is represented by remaining terms, which contain the conventional Born coefficient, ω (see below). The parameter $\hat{C}$ is a conversion factor equal to 41.84 bar cm³ cal⁻¹.

Comments on the solvation and non-solvation contributions to this equation of state, and associated equations for the isobaric heat capacity, are summarized in Sections 4.2.2.1 - 4.2.2.3. The comments address how various parameters that appear in expressions for the standard partial molal thermodynamic properties of aqueous solutes (Section 4.2.2.4) are related to the solute's partial molal volume and heat capacity.

4.2.2.1 Solvation contribution to $\bar{V}_{j,r,T}^{\circ}$. The solvation contribution to the standard partial molal volume arises from the change in free energy involved in moving an ion of radius r_e and charge, Z , from a vacuum and placing it in H₂O(*l*) of dielectric constant ε (Born, 1920):

$$\Delta \bar{G}_j^{\circ} = \frac{N_a (Z e)^2}{2 r_e} \left(\frac{1}{\varepsilon} - 1 \right), \quad (4.2.2_2)$$

where $\Delta \bar{G}_j^{\circ}$ denotes the standard partial molal Gibbs free energy of ion solvation, or *Born function* for the j -th ion, N_a stands for Avogadro's number (6.02252 x 10²³ mol⁻¹), and e represents the absolute electronic charge (4.80298 x 10⁻¹⁰ cm^{3/2} g^{1/2} s⁻¹). This equation, based only on theoretical consideration of coulombic forces, has been shown to predict enthalpies of hydration that are usually within an order of magnitude of values determined experimentally (e.g., Shock and Helgeson, 1988).

⁷ To simplify notation in the following, all properties are taken to apply to the j -th ion at a given temperature and pressure.

The *effective* ionic radius of the j -th aqueous ion⁸, r_e , is related to its corresponding crystallographic ionic radius (r_x) by

$$\begin{aligned} r_e &= r_x + |Z|g \text{ (anions), or} \\ r_e &= r_x + |Z|(0.94 + g) \text{ (cations),} \end{aligned} \quad (4.2.2_3)$$

where the “ g ” function is introduced in the revised HKF model to account for the effects on r_e of variations in T and P . Tanger and Helgeson (1988) derive an empirical function for g :

$$g = a_g(1 - \bar{p})^{b_g} - f,$$

in which $\bar{p}$ is a dimensionless variable (equal to $\rho_{H_2O}/1 \text{ g-cm}^3$), a_g and b_g denote regression coefficients (Shock *et al.*, 1992) and f stands for a pressure- and temperature-dependent difference function. Values of f are given for H_2O by Shock *et al.* (1992). Values of g have been retrieved from regressions of experimental data over a wide range of temperatures and pressures on the standard partial molal volumes and heat capacities of $NaCl(aq)$ (Tanger and Helgeson, 1988; Shock *et al.*, 1992). The results indicate $g = 0$ when $T < 150^\circ\text{C}$, or when $P > 2000$ bars.

The Born function [Eqn. (4.2.2_2)] refers to the properties of individual ions, which cannot be determined experimentally (i.e., only the thermodynamic properties of electrolytes containing the ion of interest are experimentally accessible). For this reason, a *conventional* electrostatic Born “coefficient” consistent with Eqn. (4.1_4) is defined:

$$\omega_j \equiv \omega_j^{abs} - Z_j \omega_H^{abs} = \frac{N_a(Z_j e)^2}{2r_{aj}} - Z_j \left(\frac{N_a(Z_H e)^2}{2r_{aH}} \right), \quad (4.2.2_4)$$

and the corresponding *conventional* Born function is then given by:

$$\Delta \bar{G}_{s,j}^\circ = \omega_j \left(\frac{1}{\epsilon} - 1 \right). \quad (4.2.2_5)$$

The solvation contribution to the partial molal volume is derived from the conventional Born function, i.e., (Tanger and Helgeson, 1988; Shock and Helgeson, 1988):

⁸ The effective radius of neutral aqueous species and polyatomic ionic species cannot be assessed unambiguously because it is not possible in these cases to compute r_x . Alternative approaches described by Shock *et al.* (1989) and Shock and Helgeson (1988), respectively, are used to estimate effective radii for these species.

$$\Delta \bar{V}_j^\circ = \left(\frac{\partial \Delta \bar{G}_j^\circ}{\partial P} \right)_T = -\omega Q + \left(\frac{1}{\varepsilon} - 1 \right) \left(\frac{\partial \omega}{\partial P} \right)_T, \quad (4.2.2_6)$$

where Q refers to a *solvent Born function* defined by (Helgeson and Kirkham, 1974a; Helgeson *et al.*, 1981)

$$Q \equiv \frac{1}{\varepsilon^2} \left(\frac{\partial \varepsilon}{\partial P} \right)_T. \quad (4.2.2_7)$$

The magnitude of the solvation contribution is thus a function of both the conventional Born coefficient, ω [and therefore of the charge and effective electrostatic radius of the j -th species, Eqn. (4.2.2_4)], and of the dielectric constant of the solvent and its associated pressure derivatives (Section 4.2.2.3).

4.2.2.2 Non-solvation contribution to $\bar{V}_{p,T}^\circ$. The revised HKF model assumes that any discrepancies in comparisons of experimental values of $\bar{V}^\circ$ with $\Delta \bar{V}_j^\circ$, which is calculated using the conventional Born function [Eqn. (4.2.2_5)] and the first identity in Eqn. (4.2.2_6), can be attributed to non-solvation contributions, and that the discrepancies can be accounted for with equations of the form (e.g., Shock and Helgeson, 1988):

$$\Delta \bar{V}_n^\circ = a_1 + \frac{a_2}{\Psi + P} + \frac{a_3}{T - \Theta} + \frac{a_4}{(\Psi + P)(T - \Theta)}. \quad (4.2.2_8)$$

where Ψ and Θ denote temperature- and pressure-independent constants equal to 2600 bars and 228°K, respectively and $a_{1..4}$ denote empirical pressure- and temperature-independent coefficients unique to the j -th aqueous solute. The non-solvation contributions are due to other interactions involved in the hydration process, primarily the energetics associated with disrupting the water structure near the hydrated ion. The values of the coefficients $a_{1..4}$ are obtained by regression of experimental data on the partial molal volumes of electrolytes containing the ion of interest (Section 4.3), or by estimation techniques based on correlations among standard partial molal thermodynamic properties (Section 4.4). Regression strategies are described by Shock and Helgeson (1988), Shock *et al.* (1989; 1997a) and Sverjensky *et al.*, (1997).

4.2.2.3 Heat capacities of aqueous solutes. Conventional standard partial molal heat capacities (at constant pressure) of aqueous ions are considered in terms of solvation and non-solvation contributions, in a manner analogous to that discussed above for conventional standard partial molal volumes, i.e.,

$$\bar{C}_p^\circ = \Delta \bar{C}_{p,s}^\circ + \Delta \bar{C}_{p,n}^\circ. \quad (4.2.2_9)$$

The solvation contribution to $\bar{C}_p^\circ$ is related to the conventional Born function [Eqn. (4.2.2_5)] by (Tanger and Helgeson, 1988; Shock and Helgeson, 1988):

$$\Delta \bar{S}_i^\circ = - \left(\frac{\partial \Delta \bar{G}_i^\circ}{\partial T} \right)_P, \quad (4.2.2_{10})$$

and

$$\begin{aligned} \Delta \bar{C}_{P,i}^\circ &= T \left(\frac{\partial \Delta \bar{S}_i^\circ}{\partial T} \right)_P \\ &= \omega TX + 2TY \left(\frac{\partial \omega}{\partial T} \right)_P - T \left(\frac{1}{\epsilon} - 1 \right) \left(\frac{\partial^2 \omega}{\partial T^2} \right)_P, \end{aligned} \quad (4.2.2_{11})$$

where the parameters X and Y denote additional solvent Born functions given by:

$$Y = \frac{1}{\epsilon^2} \left(\frac{\partial \epsilon}{\partial T} \right)_P, \text{ and } X = \frac{1}{\epsilon^2} \left(\frac{\partial^2 \epsilon}{\partial T^2} \right)_P - 2\epsilon Y^2.$$

Discrepancies in comparisons of experimental values of $\bar{C}_p^\circ$ and $\Delta \bar{C}_{P,i}^\circ$ are similarly attributed to non-solvation contributions, and are accounted for when $P = P_r$ with equations of the form (e.g., Shock and Helgeson, 1988):

$$\Delta \bar{C}_{P,i}^\circ = c_1 + \frac{c_2}{(T - \Theta)^2} \quad (4.2.2_{12})$$

where c_1 and c_2 refer to empirical pressure- and temperature-independent coefficients unique to the j -th ion. Regression of experimental data using strategies such as described by Shock *et al.* (1989) permits values for these coefficients to be retrieved. They are estimated if experimental data are unavailable (Section 4.4).

A relation describing the dependence of $\bar{C}_p^\circ$ on pressure is derived in terms of the partial molal volume, consistent with the identity:

$$\left(\frac{\partial \bar{C}_p^\circ}{\partial P} \right) = -T \left(\frac{\partial^2 \bar{V}^\circ}{\partial T^2} \right), \quad (4.2.2_{13})$$

or,

$$\bar{C}_{P,P,T}^\circ = \bar{C}_{P,P_r,T}^\circ + \int_{P_r}^P -T \left(\frac{\partial^2 \bar{V}^\circ}{\partial T^2} \right) dP. \quad (4.2.2_{14})$$

Integration yields (Tanger and Helgeson, 1988):

$$\begin{aligned}\bar{C}_{P,T}^{\circ} = c_1 + \frac{c_2}{(T-\Theta)^2} - \left[\frac{2T}{(T-\Theta)^3} \right] & \left[a_3(P-P_r) + a_4 \ln \left(\frac{\Psi+P}{\Psi+P_r} \right) \right] \\ & + \omega TX + 2TY \left(\frac{\partial \omega}{\partial T} \right)_P - T \left(\frac{1}{\varepsilon} - 1 \right) \left(\frac{\partial^2 \omega}{\partial T^2} \right)_P.\end{aligned}\quad (4.2.2_{15})$$

The pressure and temperature partial derivatives of the conventional Born coefficients in this equation [see also Eqn. (4.2.2_6)] are given by (Tanger and Helgeson, 1988):

$$\left(\frac{\partial \omega}{\partial P} \right)_T = -\eta \left\{ \frac{|Z|^3}{r_c^2} - \frac{Z}{(3.082+g)^2} \right\} \left(\frac{\partial g}{\partial P} \right)_T, \quad (4.2.2_{16})$$

$$\left(\frac{\partial \omega}{\partial T} \right)_P = -\eta \left\{ \frac{|Z|^3}{r_c^2} - \frac{Z}{(3.082+g)^2} \right\} \left(\frac{\partial g}{\partial T} \right)_P, \quad (4.2.2_{17})$$

and

$$\left(\frac{\partial^2 \omega}{\partial T^2} \right)_P = 2\eta \left\{ \frac{Z^4}{r_c^3} - \frac{Z}{(3.082+g)^3} \right\} \left(\frac{\partial g}{\partial T} \right)_P^2 - \eta \left\{ \frac{|Z|^3}{r_c^2} - \frac{Z}{(3.082+g)^2} \right\} \left(\frac{\partial^2 g}{\partial T^2} \right)_P \quad (4.2.2_{18})$$

where the partial derivatives of the g function for H_2O are given by Shock *et al.* (1992), and η represents a constant equal to $1.66027 \times 10^5 \text{ \AA cal mol}^{-1}$.

4.2.2.4 Apparent standard partial molal thermodynamic properties of aqueous ions and neutral species. The apparent standard partial molal Gibbs free energy and enthalpy of formation of aqueous species are calculated using Eqns. (4.1_1) and (4.1_2), equations analogous to Eqns. (4.2.1_2) and (4.2.1_3) for aqueous species (Tanger and Helgeson, 1988; Shock and Helgeson, 1988), and Eqns. (4.2.2_1) and (4.2.2_11):

$$\begin{aligned}\bar{G}_{P,T}^{\circ} - \bar{G}_{P_r,T_r}^{\circ} = & -\bar{S}_{P_r,T_r}^{\circ}(T-T_r) - c_1 \left[T \ln \left(\frac{T}{T_r} \right) - T + T_r \right] + a_1(P-P_r) + a_2 \ln \left(\frac{\Psi+P}{\Psi+P_r} \right) \\ & - c_2 \left\{ \left[\left(\frac{1}{T-\Theta} \right) - \left(\frac{1}{T_r-\Theta} \right) \right] \left[\frac{\Theta-T}{\Theta} \right] - \frac{T}{\Theta^2} \ln \left[\frac{T_r(T-\Theta)}{T(T_r-\Theta)} \right] \right\} \\ & + \left(\frac{1}{T-\Theta} \right) \left[a_3(P-P_r) + a_4 \ln \left(\frac{\Psi+P}{\Psi+P_r} \right) \right] + \omega \left(\frac{1}{\varepsilon} - 1 \right) - \omega_{P_r,T_r} \left(\frac{1}{\varepsilon_{P_r,T_r}} - 1 \right) \\ & + \omega_{P_r,T_r} Y_{P_r,T_r}(T-T_r),\end{aligned}\quad (4.2.2_{20})$$

and,

$$\begin{aligned} \bar{H}_{P,T}^{\circ} - \bar{H}_{P_r,T_r}^{\circ} = & c_1(T - T_r) - c_2 \left[\left(\frac{1}{T - \Theta} \right) - \left(\frac{1}{T_r - \Theta} \right) \right] + a_1(P - P_r) + a_2 \ln \left(\frac{\Psi + P}{\Psi + P_r} \right) \\ & + \left[\frac{2T - \Theta}{(T - \Theta)^2} \right] \left[a_3(P - P_r) + a_4 \ln \left(\frac{\Psi + P}{\Psi + P_r} \right) \right] + \omega \left(\frac{1}{\varepsilon} - 1 \right) + \omega T Y \\ & - T \left(\frac{1}{\varepsilon} - 1 \right) \left(\frac{\partial \omega}{\partial T} \right)_P - \omega_{P_r,T_r} \left(\frac{1}{\varepsilon_{P_r,T_r}} - 1 \right) - \omega_{P_r,T_r} T_r Y_{P_r,T_r}. \end{aligned} \quad (4.2.2_{21})$$

The corresponding expression for $\bar{S}_{j,P,T}^{\circ}$ is given by:

$$\begin{aligned} \bar{S}_{P,T}^{\circ} = \bar{S}_{P_r,T_r}^{\circ} + & c_1 \ln \left(\frac{T}{T_r} \right) - \frac{c_2}{\Theta} \left\{ \left[\frac{1}{T - \Theta} \right] - \left[\frac{1}{T_r - \Theta} \right] + \frac{1}{\Theta} \ln \left[\frac{T_r(T - \Theta)}{T(T_r - \Theta)} \right] \right\} \\ & + \left(\frac{1}{T - \Theta} \right)^2 \left[a_3(P - P_r) + a_4 \ln \left(\frac{\Psi + P}{\Psi + P_r} \right) \right] + \omega Y - \left(\frac{1}{\varepsilon} - 1 \right) \left(\frac{\partial \omega}{\partial T} \right)_P - \omega_{P_r,T_r} Y_{P_r,T_r} \end{aligned} \quad (4.2.2_{22})$$

Equations (4.1_1), (4.1_2) and (4.2.2_20) - (4.2.2_22) represent the HKF model as it is incorporated into SUPCRT.

4.2.2.5 Primary thermodynamic data for aqueous species. Inspection of Eqns. (4.1_1), (4.1_2) and (4.2.2_20) - (4.2.2_22) indicates that the following data are needed for the j -th aqueous solute to calculate its corresponding apparent standard molal thermodynamic properties at a given temperature and pressure: $\Delta \bar{G}_f^{\circ}$, $\Delta \bar{H}_f^{\circ}$, $\bar{S}_{P_r,T_r}^{\circ}$, $a_{1...4}$, $c_{1...2}$, ω_{P_r,T_r} and Z . Species that have been considered within this framework include aqueous ions and electrolytes (Shock and Helgeson, 1988; Tanger and Helgeson, 1988, Shock *et al.*, 1997a) inorganic neutral species (Shock *et al.*, 1989), organic species (Shock and Helgeson, 1990), and inorganic and organic metal complexes (Sverjensky *et al.*, 1997).

4.2.3 Reactions among minerals, gases and aqueous species

The change in the apparent standard molal Gibbs free energy or enthalpy of a reaction is given by

$$\Delta \bar{\Xi}_{r,P,T}^{\circ} = \sum_{i=1}^I \nu_i \Delta \bar{\Xi}_{i,P,T}^{\circ} + \sum_{j=1}^J \nu_j \Delta \bar{\Xi}_{j,P,T}^{\circ} + \nu_{H_2O} \Delta \bar{\Xi}_{H_2O,P,T}^{\circ} \quad (4.2.3_1)$$

where $\Delta \bar{\Xi}$ refers to ΔG or ΔH , ν_i denotes stoichiometric reaction coefficients of minerals (and/or gases), ν_j stands for those of aqueous species, and ν_{H_2O} refers to that of fluid H_2O under subcritical (liquid or gaseous) or supercritical conditions. The corresponding equilibrium constant ($K_{r,P,T}$) is given by:

$$\ln K_{r,R,T} = \frac{-\Delta G_{r,R,T}}{RT} \quad (4.2.3_2)$$

where R stands for the gas constant. The equilibrium constant can be calculated at any P and T using Eqns. (4.2.3_2), (4.2.3_1), (4.1_1) and (4.1_2), and

- Eqn. (4.2.1_6) for minerals and gases,
- Eqn. (4.2.2_20) for aqueous species, and Eqn. (4.2.2_4) for the conventional Born coefficient and its associated pressure and temperature derivatives [Eqns. (4.2.2_16) - (4.2.2_18)] plus associated values of the g function given by Tanger and Helgeson (1988) and Shock et al. (1992), and
- equations summarized by Johnson and Norton (1991) and Johnson *et al.* (1991) for calculation of the thermodynamic properties of H_2O as a function of temperature and pressure.

All these equations are incorporated into the SUPCRT software package (Johnson *et al.*, 1991), which, with appropriate thermodynamic data (Sections 4.2.1.1 and 4.2.2.5), permits calculation of equilibrium constants and other reaction properties over the range of temperatures and pressures noted in the introductory paragraph of Section 4.

4.3 Retrieval of Thermodynamic Properties from Experimental Data

Experimental data are evaluated using thermodynamic relations summarized in Section 4.2 to retrieve values of the thermodynamic parameters identified in Sections 4.2.1.1 and 4.2.2.5. The evaluation approach is summarized below for several types of experimental methods that have been used to derive much of the data in SPRONS.JNC. This is followed by a brief description of statistical/mathematical techniques to analyze experimental results, and to retrieve an internally consistent set of thermodynamic properties. Alternative methods described in Section 4.4 are used to estimate the data identified in Sections 4.2.1.1 and 4.2.2.5 if appropriate experimental results are unavailable.

4.3.1 Reaction data

Reaction equilibria among minerals and a fluid phase are studied using a variety of experimental techniques. Two alternative statements of Eqn. (4.2.3_1) are appropriate for such experiments, depending on whether the fluid is pure H_2O , or an aqueous solution, i.e.,:

$$\Delta \Xi_{r,R,T}^\circ = \sum_{i=1}^I v_i \Delta \Xi_{i,R,T}^\circ + v_{H_2O} \Delta \Xi_{H_2O,R,T}^\circ, \text{ or} \quad (4.3.1_1)$$

$$\Delta \Xi_{r,R,T}^\circ = \sum_{i=1}^I v_i \Delta \Xi_{i,R,T}^\circ + \sum_{j=1}^J v_j \Delta \Xi_{j,R,T}^\circ + v_{H_2O} \Delta \Xi_{H_2O,R,T}^\circ, \quad (4.3.1_2)$$

respectively. *Phase-equilibrium* experiments are evaluated using Eqn. (4.3.1_1). Although experiments evaluated using Eqn. (4.3.1_2) are also phase-equilibrium experiments in a strict sense, we refer to them as *solubility* experiments to emphasize the difference between Eqns. (4.3.1_2) and (4.3.1_1).

4.3.1.1 Phase-equilibrium experiments. These experiments characterize reversals in the relative stabilities of reactant and product mineral assemblages (Gordon, 1973; Carmichael, 1977; Newton 1987; Navrotsky, 1987; Wood, 1987; Kerrick, 1987; Lofgren, 1987). The reversals are determined by holding an initial reactant assemblage at fixed P - T conditions for a specified period of time. The pressure and temperature are then rapidly reduced and the reaction products, if any, identified. A new pressure and/or temperature are then selected, and the experiment repeated using a fresh set of reactants. Ideally, a number of such experiments will closely approximate equilibrium conditions, which are bounded by the experimentally determined set of P - T coordinates, or "half brackets", over a range of temperatures and pressures. The exact equilibrium temperature and pressure are rarely determined in phase-equilibrium studies, however.

Each phase-equilibrium experiment defines a specific P - T coordinate at which $\Delta G_{r,P,T}^\circ > 0$ (reactants stable), or $\Delta G_{r,P,T}^\circ < 0$ (products stable). This constitutes an inequality constraint on the standard molal Gibbs free energy of reaction, i.e.,:

$$\Delta G_{r,P,T}^\circ = \sum_{i=1}^I \nu_i \Delta G_{i,P,T}^\circ + \nu_{H_2O} \Delta G_{H_2O,P,T}^\circ \begin{pmatrix} > \\ < \end{pmatrix} 0, \quad (4.3.1.1_1)$$

which can be expanded such that all thermodynamic variables are written in terms of fitting parameters and standard molal properties. For example, simplifying the discussion in Section 4.2.1 by assuming no first-order or lambda transitions, Eqn. (4.3.1.1_1) can be written as:

$$\Delta G_{r,P,T}^\circ \begin{pmatrix} > \\ < \end{pmatrix} \sum_i \nu_i \left\{ \Delta G_{fi}^\circ - S_{i,P,T_r}^\circ (T - T_r) + a_i \left[T - T_r - T \ln \left(\frac{T}{T_r} \right) \right] + \frac{(-c_i - b_i T T_r^2)(T - T_r)^2}{2 T T_r^2} + V_{r,P,T}^\circ (P - P_r) \right\} + \nu_{H_2O} \Delta G_{H_2O,P,T}^\circ \quad (4.3.1.1_2)$$

To evaluate this equation, one or more sets of phase-equilibrium data are plotted on a P - T *diagram* (e.g., Fig. 4.3.1.1_1). Equilibrium conditions, where $\Delta G_{r,P,T}^\circ = 0$, are then estimated by calculating the best-fit curve through all the P - T half brackets, such that the curve, representing univariant equilibrium, passes between half-brackets where reactants are stable and those where products are stable. Regression techniques applied to the experimental dataset with respect to Eqn. (4.3.1.1_2) and the condition $\Delta G_{r,P,T}^\circ = 0$ are then used to retrieve values of

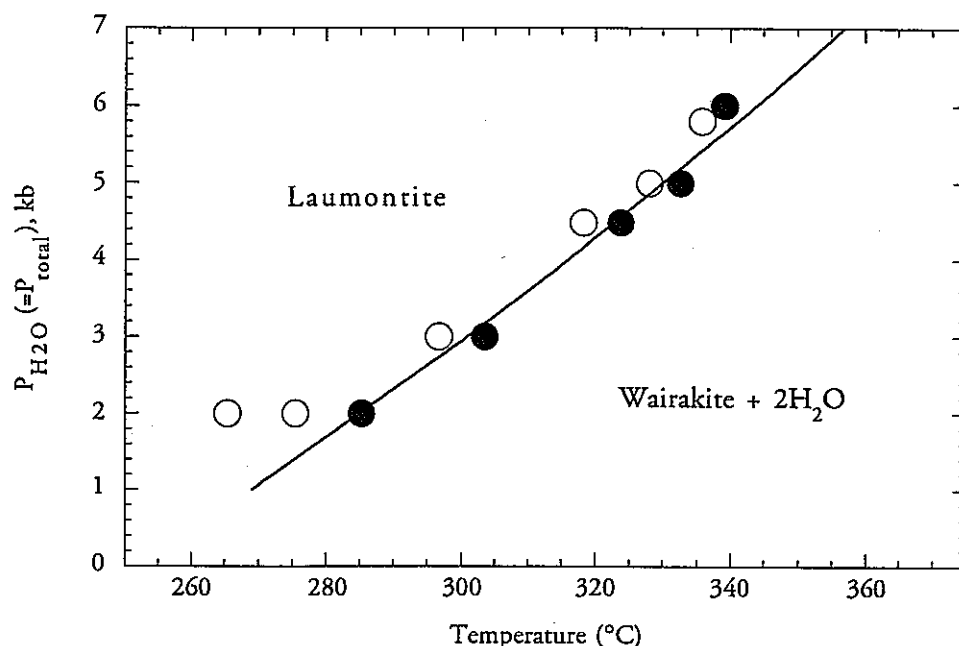


Figure 4.3.1_1. P - T diagram showing calculated and experimental equilibrium conditions for the reaction $\text{laumontite} = \text{wairakite} + 2 \text{H}_2\text{O}$. The curve is calculated using SUPCRT and SPRONS.JNC. Experimental data are from Liou (1971). Half brackets where the zeolite laumontite is the stable phase are indicated by open symbols; half brackets where the zeolite wairakite is stable are indicated by filled symbols (see Helgeson *et al.*, 1978, Fig. 91b). The calculated curve is consistent with all the experimental data, except the point indicating that wairakite is stable at 340°C and 6 kb.

unknown properties (e.g., ΔG_{fi}°). In practice, additional constraints on a few of the parameters in this equation (e.g., the a_i Maier-Kelly coefficients) are usually required to adequately constrain the regression calculations. These independent constraints are obtained using other experimental techniques, such as calorimetry (Section 4.3.2.1).

Alternative approaches to that described above have been used to evaluate phase-equilibrium data. Berman (1988) evaluates inequalities in equations similar to Eqn. (4.3.1.1_2) explicitly using a linear programming technique (Gordon, 1973; Berman *et al.*, 1986). The other internally consistent databases described in Section 3.3 are derived using alternative regression-based methods unique to each database. Differences in the regression methods deal mainly with the problem of locating true equilibrium temperatures and pressures between the experimentally determined P - T half brackets (Engi, 1992). The strengths and limitations of regression and linear programming techniques are discussed by Berman *et al.* (1986) and Kolassa (1990).

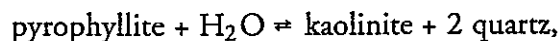
Thermodynamic data for 83 minerals in SPRONS.JNC were originally retrieved by Helgeson *et al.* (1978) using the approach described above. The data added to this database by Johnson *et al.* (1991) to generate SPRONS92.DAT (Section 3.2) are from Table 9 of Helgeson *et al.* (1978). The added data are for selected elements, oxides, halides, sulfates, carbonates and sulfides, and

are retrieved solely from calorimetric investigations. Although the reliability of these calorimetric data is untested on the basis of experimental phase-equilibrium constraints, they appear to be more consistent with phase relations observed in geologic systems (Helgeson *et al.*, 1978) than values adopted in previous studies (Helgeson, 1969).

4.3.1.2 Solubility experiments. These experiments use a variety of techniques⁹ to determine the equilibrium constants of reactions involving one or more minerals (and/or gases) and an aqueous solution. Attainment of equilibrium is thus the primary goal of these experiments, in contrast to phase-equilibrium studies, where it is only necessary to approximate equilibrium conditions in terms of bounding temperatures and pressures. This requires verification that the reaction is reversible, which can be demonstrated by approaching equilibrium from conditions of both undersaturation and supersaturation of the aqueous solution's composition with respect to the reaction of interest. Non-ideal behavior of the aqueous phase must also be accounted for to accurately calculate the activities of aqueous reactants and products that partially or completely determine the value of the equilibrium constant.

To retrieve thermodynamic properties from the results of solubility studies, the experimental equilibrium constants must first be converted into Gibbs free energies of reaction using Eqn. (4.2.3_2), and then cast into terms containing fitting parameters and standard molal properties using appropriate equations described in Sections 4.2.1 and 4.2.2.4. Unknown properties (e.g., ΔG_f°) are obtained by regression of the derived equations using the experimentally determined equilibrium constants as constraints over an appropriate range of temperatures and pressures. Additional independently determined parameters (e.g., Maier-Kelly coefficients) may be required to adequately constrain the regression calculations.

In practice, the results of solubility experiments have been used primarily to test and refine the accuracy of thermodynamic properties retrieved from phase-equilibrium and calorimetric data. An example of this use of solubility data is illustrated by comparison of Figs. 4.3.1.2_1 and 4.3.1.2_2 (see also Figs. 48a and 47a, respectively, in Helgeson *et al.*, 1978). Fig. 4.3.1.2_1 shows experimental phase-equilibrium data for the reaction:



where the corresponding univariant equilibrium curve calculated using SUPCRT and SPRONS.JNC is also indicated for comparison. As can be seen, the calculated curve is poorly constrained by the limited amount of experimental data. The reliability of most of these data (Thompson, 1970) may also be

⁹ References cited in Section 6 include detailed descriptions of these techniques. The book edited by Ulmer and Barnes (1987) describes associated experimental equipment.

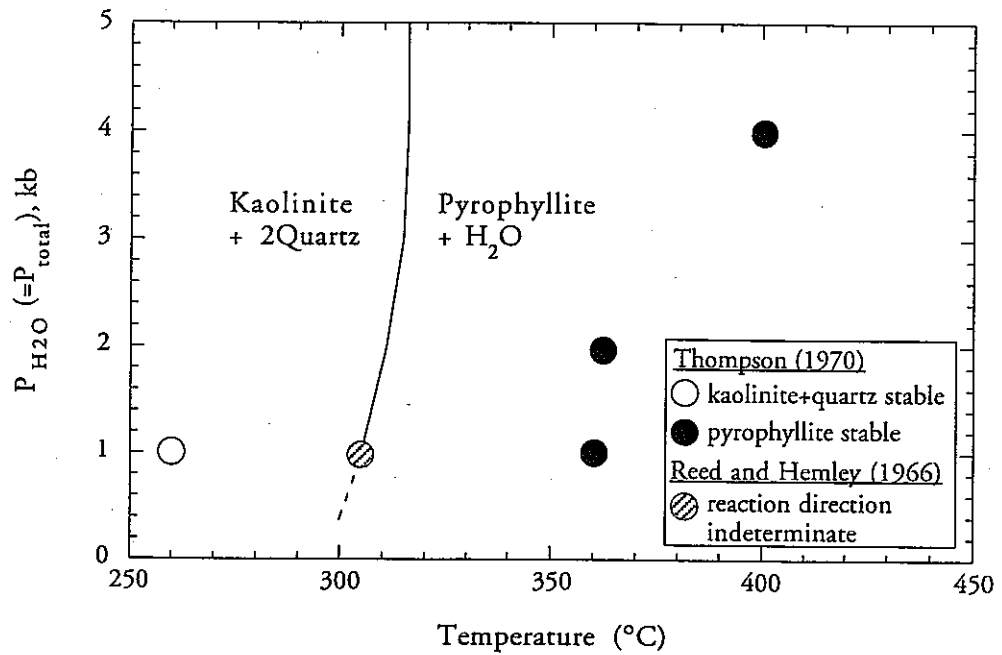


Figure 4.3.1.2_1. Calculated univariant equilibrium curve and experimental observations (symbols) of phase equilibria among kaolinite, quartz, pyrophyllite and H₂O.

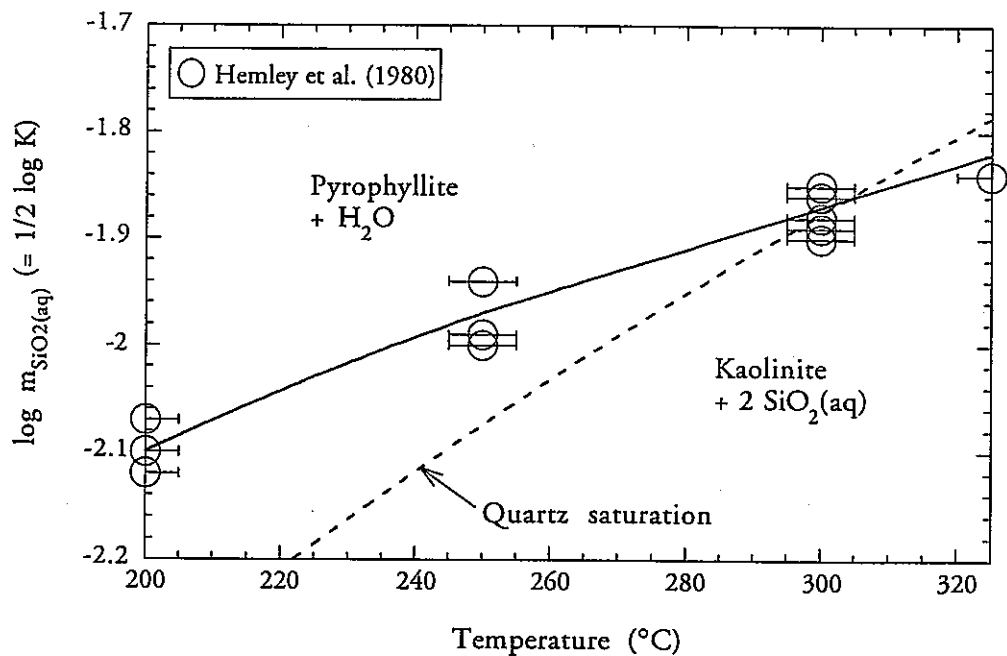
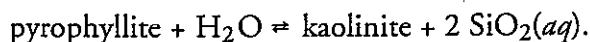


Figure 4.3.1.2_2. Calculated (solid curve) and experimental (symbols) equilibrium constants for the reaction pyrophyllite + H₂O $\rightleftharpoons$ kaolinite + 2 SiO₂(aq). The size of the symbols is greater than the analytical uncertainty assigned by Hemley *et al.* (1980) to SiO₂(aq) measurements ($\pm 5\%$ of the reported value). The uncertainty in temperature measurements is $\pm 5^\circ\text{C}$ (Hemley *et al.*, 1980). Note that $m_{SiO_2(aq)} \approx a_{SiO_2(aq)}$ for these experimental conditions.

questionable (Section 4.3.4.2).

Fig. 4.3.1.2_2 shows experimental equilibrium constants at 1 kb, and their calculated counterparts (using SUPCRT and SPRONS.JNC), for the similar, solubility reaction:



As can be seen, the calculated curve is in good agreement with the available experimental data, within stated uncertainty limits. Moreover, this curve intersects the dashed curve representing quartz saturation at approximately 303°C. This partially confirms results from the phase equilibrium experiments (Fig. 4.3.1.2_1), which indicate that pyrophyllite, kaolinite and quartz should equilibrate at 1 kb when $T \approx 305^\circ\text{C}$.

Similar uses of solubility data to test and refine thermodynamic properties retrieved from phase-equilibrium and calorimetric experiments are described by Helgeson *et al.* (1978) for various minerals in the systems: MgO-SiO₂-H₂O, CaO-MgO-SiO₂-CO₂-H₂O, Al₂O₃-SiO₂-H₂O, K₂O-Al₂O₃-SiO₂-HCl-H₂O, Na₂O-Al₂O₃-SiO₂-HCl-H₂O, Na₂O-K₂O-Al₂O₃-SiO₂-H₂O, CaO-FeO-Fe₂O₃-SiO₂-H₂O, and CaO-FeO-Fe₂O₃-Al₂O₃-SiO₂-CO₂-H₂O. The thermodynamic properties of these minerals (forsterite, talc, chrysotile, antigorite, wollastonite, anthophyllite, enstatite, quartz, sepiolite, calcite, dolomite, diopside, tremolite, gibbsite, boehmite, diaspore, corundum, kaolinite, pyrophyllite, andalusite, albite, low albite, high albite, K-feldspar, high sanidine, maximum microcline, analcime, paragonite, muscovite, anorthite, kalsilite, nepheline, epidote, andradite, hedenbergite, magnetite and hematite) are thus constrained by three different types of experimental data. Most of these data pertain to relatively high temperatures and pressures, however, which are far removed from conditions of primary interest in the Japanese repository concept (see Section 6). It is similarly noted that Johnson *et al.* (1991) revised Gibbs free energies and enthalpies of formation of several Ca-bearing minerals in the Helgeson *et al.* (1978) database to be consistent with updated values of Gibbs free energies of Ca²⁺ and CO₃²⁻ formation retrieved from experimental determinations of calcite and aragonite solubility reported by Plummer and Busenberg (1982).

As noted in Section 3.2, all the data for minerals from Helgeson *et al.* (1978), and the revisions incorporated by Johnson *et al.* (1991), are included in SPRONS.JNC. More recent uses of solubility data to test and refine thermodynamic properties retrieved from phase-equilibrium and calorimetric experiments are reported by Sverjensky *et al.* (1991) and Pokrovskii and Helgeson (1995; 1997b). The results of these studies are also incorporated in SPRONS.JNC (Section 5.2.2).

4.3.2 Mineral and gas data

The heat capacities and molar volumes of minerals and gases are independent of all other thermodynamic properties because they can be determined directly using calorimetric methods (heat capacity), or calculated from refinements of X-ray data on a mineral's unit-cell dimensions, or from the ideal gas law (molar volumes). Calorimetric experimental techniques and apparatus are summarized by McCullough and Scott (1968) and Robie (1987). Nordstrom and Munoz (1985) describe the procedure to calculate molar volumes from unit-cell data. The unit-cell dimensions (Bloss, 1971) of many rock-forming minerals are described by Deer et al. (1962-1963).

To evaluate these experimental data, various empirical expressions are assumed for the heat capacity as a function of temperature (i.e., Eqn. 4.2.1_5; see also Anderson and Crerar, 1993), or molar volume as a function of pressure and temperature (e.g., Berman, 1988; Holland and Powell, 1990; Gottschalk, 1997). The fitting parameters in these expressions are then determined by regression of one or more experimental datasets that span as broad a range of temperatures or pressures as possible.

4.3.2.1 Standard molal heat capacities. The empirical $a_{1...3}$ parameters in SPRONS.JNC used with Eqn. (4.2.1_5) to calculate isobaric heat capacities as a function to temperature are:

- from calorimetric investigations (e.g., Kelly, 1960; Pankratz, 1964a,b; Walther and Helgeson, 1977; King *et al.*, 1975; Pankratz and King, 1965, Pankratz and King, 1970; Pankratz and Kelley, 1964; Pankratz, 1968a,b; King and Weller, 1961; Mah *et al.*, 1967; Mills, 1974),
- generated by Helgeson *et al.* (1978) by calculation or regression of heat capacity data reported by others (Stull and Prophet, 1971; Holm *et al.*, 1967; Perkins *et al.*, 1977; Kiseleva *et al.*, 1972, 1974; Holm and Kleppa, 1966; Thompson *et al.*, 1974; Krupka *et al.*, 1977a,b; Kelley *et al.*, 1953; Robie and Hemingway, 1973), or
- estimated using equations described by Helgeson *et al.* (1978) and calorimetric data from King (1955) and King and Weller (1961).

The coefficients are valid from 25°C to an upper temperature limit corresponding to the maximum temperature investigated experimentally, or to the temperature of phase transitions (see Section 5 and Appendix A). Equation (4.2.1_5) yields fits of experimental heat capacities in this temperature range to within $\pm 0.1 \text{ cal mol}^{-1} \text{ K}^{-1}$, or less (Helgeson *et al.*, 1978). Less accurate fits are obtained, however, in the case of minerals with high Debye temperatures (Denbigh, 1971; Berman and Brown, 1985). The effects of solid-state phase transitions, order/disorder, and vacancy defects on the heat capacity and other thermodynamic properties of minerals are evaluated by Helgeson *et al.* (1978),

Robie *et al.* (1978) and Berman and Brown (1985). These effects were accounted for in the original Helgeson *et al.* (1978) database, and are thus incorporated in SPRONS.JNC.

4.3.2.2 Standard molal volumes. The volumetric properties of minerals and gases in SPRONS.JNC are limited to standard molal volumes at the reference temperature and pressure. This is because the equation of state for minerals and gases adopted in SUPCRT (Section 4.2.1) does not consider the effects of variations in temperature and pressure on the standard molal volume. Associated data characterizing thermal expansivities and compressibilities of minerals are therefore unnecessary (Helgeson *et al.*, 1978). Standard molal volumes in SPRONS.JNC are:

- reported values (Robie and Waldbaum, 1968; Robie *et al.*, 1966; Fyfe *et al.*, 1958; Ernst, 1968; Chatterjee, 1970; Kunze, 1961, Helgeson and Kirkham, 1974a),
- computed by Helgeson *et al.* (1978) using published data on unit-cell dimensions (Donnay and Ondik, 1973; Burnham, 1965; Deer *et al.*, 1962-1963; Berry, 1974), or
- estimated using procedures described by Helgeson *et al.* (1978) and data from Sheppard and Gude (1970) and Steinfink (1958).

The reported values for minerals, and values computed by Helgeson *et al.*, (1978), are based on refinements of unit-cell data determined by X-ray diffraction. Standard volumes determined in this manner are generally accurate to within ± 0.05 to $\pm 0.25\%$ (Nordstrom and Munoz, 1985). It is not uncommon, however, for several volume determinations to disagree by more than the limits of reported uncertainties. This is often observed when the molar volumes of natural and synthetic samples of the same mineral are determined.

4.3.3 Electrolyte data

Equation-of-state parameters ($a_{1...4}$, c_1 and c_2 , see Section 4.2.2.4) for aqueous solutes are retrieved from experimental measurements of the *apparent* molal volumes and heat capacities of electrolytes over a range of temperatures and electrolyte concentrations. For a single electrolyte solution, apparent molal properties ($\phi_{\pm}$) are related to their partial molal analogs by:

$$\phi_{\pm} = \frac{1}{m_k} \int_0^{m_k} \Xi_k dm_k, \quad (4.3.3_1)$$

where,

$$\Xi_k = \left(\frac{\partial \Xi}{\partial m_k} \right)_{P,T}, \quad (4.3.3_2)$$

and m_k stands for the molality of the k -th electrolyte (Helgeson, 1981). The retrieval procedure therefore involves:

- extrapolation of the apparent molal properties to reference-state conditions of infinite dilution, at which point the corresponding standard partial molal properties are defined,
- calculation of non-solvation contributions to the standard partial molal volume, or heat capacity,
- regression of the non-solvation volumes or heat capacities as a function of temperature and pressure to determine respective values of the empirical coefficients, $a_{1...4}$, or $c_{1...2}$, and
- calculation of corresponding equation-of-state parameters for the electrolyte's constituent ions.

An approach to carry out the first of these tasks is described in Section 4.3.3.1. Regression strategies that address the three latter tasks are discussed in Section 4.3.3.2. Calorimetric methods for experimental determination of apparent molal heat capacities are described by Robie (1987) and Chen *et al.* (1994). Experimental methods for determination of apparent molal volumes are described by Hepler *et al.* (1965), Ellis (1968) and Barta and Hepler (1986).

4.3.3.1 Partial molal volumes and heat capacities of aqueous electrolytes. An approach to extrapolate experimental apparent molal properties to infinite dilution is described by Helgeson *et al.* (1981) and Helgeson (1981). The approach is based on calculation of differences between experimental apparent molal volumes or heat capacities and values of these properties that are estimated using a version of the extended Debye-Hückel equation described by Helgeson *et al.* (1981). These so-called *difference functions* are defined by (Helgeson, 1981; Shock and Helgeson, 1988):

$$\rho_{V,k}^* \equiv \bar{V}_k^\circ + \frac{v_k b_{V,k} I}{4}, \text{ and} \quad (4.3.3.1_1)$$

$$\rho_{J,k}^* \equiv \bar{C}_{P,k}^\circ - \frac{v_k b_{J,k} I}{4} \quad (4.3.3.1_2)$$

where $\rho_{V,k}^*$ and $\rho_{J,k}^*$ denote difference functions for the partial molal volume and heat capacity of the k -th electrolyte, respectively, $b_{V,k}$ and $b_{J,k}$ are terms that can be calculated from temperature and pressure derivatives of the extended term of the Debye-Hückel equation (Helgeson, 1981), I represents the *effective* ionic strength,

$$I = 1/2 \sum_j Z_j^2 m_j, \quad (4.3.3.1_3)$$

where Z stands for the ion's charge and m represents its molality, and v_k is given by

$$v_k = \sum_j v_{j,k}, \quad (4.3.3.1_4)$$

where $v_{j,k}$ refers to the number of moles of the j -th ion in the electrolyte.

Helgeson (1981) notes that for completely dissociated electrolytes, the effective and *stoichiometric* ionic strength,

$$I \equiv 1/2 \sum_j v_{j,k} Z_j^2 m_k, \quad (4.3.3.1_5)$$

are equivalent, and that a plot of Eqn. (4.3.3.1_1) or (4.3.3.1_2) versus I at a given temperature and pressure should therefore be linear. Extrapolation of the linear curve in such plots to the intercept, $I = 0$, will then yield $\bar{V}_k^\circ$ or $\bar{C}_{p,k}^\circ$, respectively. This behavior is confirmed for a large number of electrolytes in relatively dilute solutions, but deviations from linear behavior are often observed at higher ionic strengths (Helgeson *et al.* 1981; Shock and Helgeson, 1988¹⁰). Representative results are shown in Figs. 4.3.3.1_1 and 4.3.3.1_2, for example. The deviations are caused by the effects of ion-association, which become more pronounced with increasing ionic strength and increasing temperature.

Shock and Helgeson (1988) and Shock *et al.* (1997a) employ this approach to determine standard partial molal volumes of 75 electrolytes, and standard partial molal heat capacities of 80 electrolytes. These data serve as primary constraints in the regression strategies described below to determine standard partial molal properties of the corresponding ions. Correlations among the retrieved ionic properties enable estimates to be made of standard partial molal properties of other ions and complexes, for which experimental data are unavailable (Section 4.4).

4.3.3.2. Equation-of-state parameters for aqueous ions. These parameters are obtained by regression of non-solvation contributions to the standard partial molal volumes and heat capacities of corresponding electrolytes. The non-solvation parameters are determined by rearranging Eqns. (4.2.2_1) and (4.2.2_9), i.e.,

$$\Delta \bar{V}_{n,k}^\circ = \bar{V}_k^\circ - \Delta \bar{V}_{s,k}^\circ, \text{ and}$$

$$\Delta \bar{C}_{p,n,k}^\circ = \bar{C}_{p,k}^\circ - \Delta \bar{C}_{p,s,k}^\circ,$$

and thus by subtracting the respective solvation contributions calculated using

¹⁰ The appendix to this reference contains 37 pages of supplemental data supporting this conclusion, and is available in microfiche form from the National Auxiliary Publication Service (document No. 04597), P.O. Box 3513, Grand Central Station, New York, NY 10163-3513.

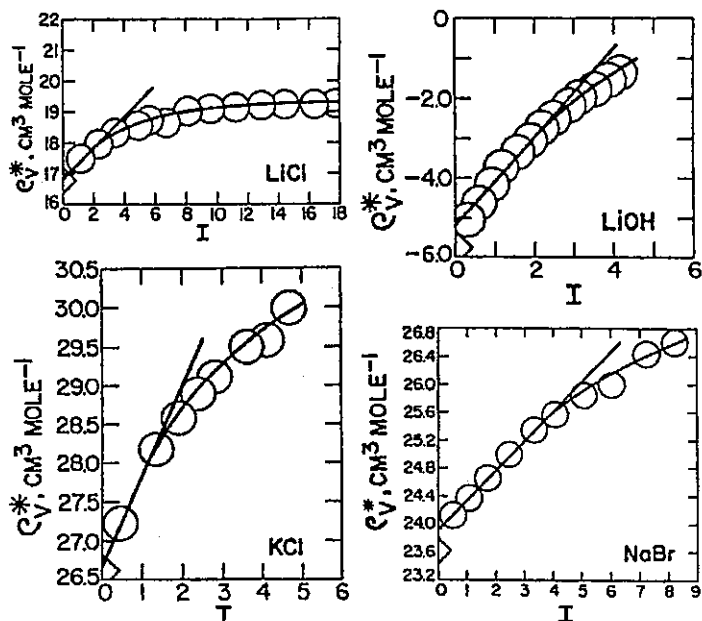


Figure 4.3.3.1_1. Calculated values (symbols) of the difference function [Eqn. (4.3.3.1_1)] for the apparent molal volumes of various electrolytes at 25°C and 1 bar (from Helgeson, 1981; Fig 7.20). Extrapolation to conditions of zero ionic strength, indicated by the linear portion of the curves, defines the standard partial molal volume. Similar plots by Helgeson (1981), Helgeson *et al.* (1981) and Shock and Helgeson (1988), document sources of experimental data.

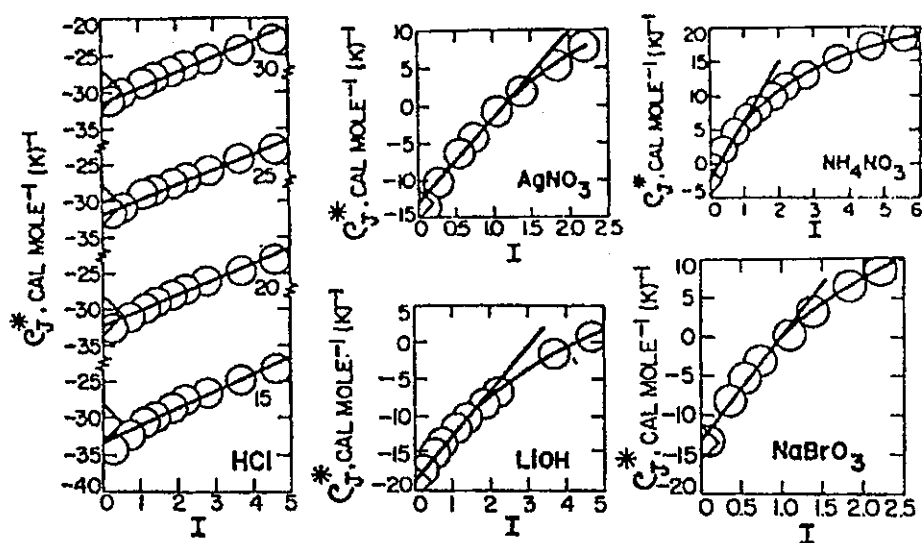


Figure 4.3.3.1_2. Calculated values (symbols) of the difference function [Eqn. (4.3.3.1_2)] for the apparent molal heat capacities of various electrolytes at 1 bar and temperatures between 15 and 30°C (from Helgeson, 1981; Fig 7.19). Extrapolation to conditions of zero ionic strength, indicated by the linear portion of the curves, defines the standard partial molal heat capacity. Similar plots by Helgeson (1981), Helgeson *et al.* (1981) and Shock and Helgeson (1988), document sources of experimental data.

Eqns. (4.2.2_6) or (4.2.2_11) from the experimentally constrained values of the standard partial molal volume or heat capacity (i.e., Section 4.3.3_1).

With the non-solvation parameters in hand, the regression strategy focusses on retrieving values for the empirical coefficients $a_{1...4}$, and $c_{1...2}$. One approach that has been used for this purpose (e.g., Shock and Helgeson, 1988; Shock *et al.*, 1997a) to retrieve all the equation-of-state parameters for aqueous solutes in SPRONS.JNC (Section 5.2.1) combines terms in Eqn. (4.2.2_8),

$$\sigma = a_1 + \frac{a_2}{\Psi + P}, \text{ and} \quad (4.3.3.2_1)$$

$$\xi = a_3 + \frac{a_4}{\Psi + P}, \quad (4.3.3.2_2)$$

such that the non-solvation volume is represented by a linear equation with respect to temperature given by,

$$\Delta \bar{V}_n^\circ = \sigma + \frac{\xi}{T - \Theta}. \quad (4.3.3.2_3)$$

The analogous equation for the non-solvation heat capacity is Eqn. (4.2.2_12):

$$\Delta \bar{C}_{Pn}^\circ = c_1 + \frac{c_2}{(T - \Theta)^2},$$

which is linear with respect to $(T - \Theta)^2$. Plots of $\Delta \bar{V}_{n,k}^\circ$ versus $1/(T - \Theta)$, or $\Delta \bar{C}_{Pn,k}^\circ$ versus $1/(T - \Theta)^2$, can thus be fit to experimentally constrained data on non-solvation volumes or heat capacities, respectively, to determine σ , or c_1 , from the intercepts of the resultant linear curves (at $T = 25^\circ\text{C}$), and ξ , or c_2 , from their slopes. Plots of this kind are illustrated in Figs 4.3.3.2_1. Corresponding data for the ionic constituents of the electrolyte are determined using Eqn. (4.1_9).

Shock and Helgeson (1988) and Shock *et al.* (1989) demonstrated that equation-of-state parameters for aqueous species obtained by regression of experimental standard partial molal volumes and heat capacities in the manner described above correlate linearly with the standard partial molal volume or heat capacity of the species at 25°C and 1 bar. The derived correlations for ionic species are given by (Sverjensky *et al.*, 1997):

$$\bullet \quad a_1 = (0.013684) \left\{ \bar{V}_{P_h, T_r}^\circ + (41.84) \omega_{P_h, T_r} Q_{P_h, T_r} \right\} + 0.1765 \quad (4.3.3.2_4)$$

$$\bullet \quad a_2 = (33.423) \left\{ \bar{V}_{P_r, T_r}^\circ + (41.84) \omega_{P_h, T_r} Q_{P_h, T_r} \right\} - 347.23 \quad (4.3.3.2_5)$$

$$\bullet \quad a_3 = -(0.1435) \left\{ \bar{V}_{P_h, T_r}^\circ + (41.84) \omega_{P_h, T_r} Q_{P_h, T_r} \right\} + 7.0274 \quad (4.3.3.2_6)$$

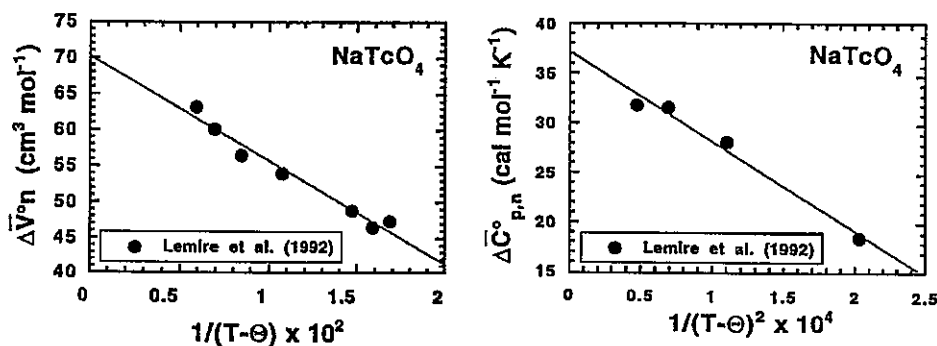


Figure 4.3.3.2_1. Non-solvation contributions to the standard partial molal volume and heat capacity of NaTcO₄ plotted versus temperature functions given by Eqns. (4.3.3.2_3) and (4.2.2_12), respectively (from Shock *et al.*, 1997a).

$$\bullet \quad a_4 = -(138.17) \left\{ \bar{V}^{\circ}_{P_r, T_r} + (41.84) \omega_{P_r, T_r} Q_{P_r, T_r} \right\} - 26355 \quad (4.3.3.2_7)$$

$$\bullet \quad c_1 = (0.6087) \bar{C}^{\circ}_{P_r, T_r} + \omega_{P_r, T_r} T_r X_{P_r, T_r} + 5.85 \quad (4.3.3.2_8)$$

$$\bullet \quad c_2 = (2037) \bar{C}^{\circ}_{P_r, T_r} - 30460 \quad (4.3.3.2_9)$$

where Q_{P_r, T_r} and X_{P_r, T_r} (see Sections 4.2.2.1 and Section 4.2.2.3, respectively) are equal to 5.903×10^{-7} (bar⁻¹) and -3.09×10^{-7} (K⁻²), respectively, at $T_r = 25^{\circ}\text{C}$ and $P_r = 1$ bar, and ω_{P_r, T_r} is given by Eqn. (4.2.2_4). These correlations can be substituted into Eqns. (4.1_1), (4.1_2) and (4.2.2_20) - (4.2.2_22) to derive regression equations for the standard partial molal Gibbs free energy and enthalpy of formation, and the standard partial molal entropy, in which the regression parameters are only $\bar{V}^{\circ}_{P_r, T_r}$, $\bar{C}^{\circ}_{P_r, T_r}$, and ω_{P_r, T_r} (Sverjensky *et al.*, 1997).

4.3.4 Statistical/Mathematical analyses

Retrieval of thermodynamic properties from experimental data is carried out using statistical/ mathematical techniques on three levels:

- analysis of properties,
- analysis of reactions, and
- simultaneous evaluation of multiple reactions

The techniques applicable to each of these levels are briefly described below.

4.3.4.1 Properties. Regression techniques to evaluate standard molal heat capacities and volumes seek to minimize the *residual* function defined by:

$$R = \sum_{i=1}^n W_i (Q_i - q_i)^2, \quad (4.3.4.1_1)$$

where R stands for the residual among the set of $i \dots n$ determinations, q_i refers to parameter values used as estimators of the measured data, Q_i , and W_i denotes a weighting factor, which is usually taken as being inversely proportional to the measurement precision [i.e., the more precise (smaller in precision magnitude) data are accorded more weight in the regression] (e.g., Engi, 1992). For example, Q_i could represent measured values over a range of temperatures, q_i would then denote corresponding heat capacities calculated using Eqn. (4.2.1_5), and W_i would refer to the inverse of the precision of each measurement. The problem then becomes one of finding values for the temperature-independent linear fit parameters (a_i , b_i , c_i) in the Maier-Kelly equation that lead to a minimization of the residual function.

The next step in the analysis entails comparison of individual measurements to corresponding values of q_i , which, to continue with the example in the preceding paragraph, would be calculated using the Maier-Kelly equation and the optimized values for a_i , b_i , and c_i at the measurement temperature. Error plots are useful devices for visualizing such comparisons (Robinson *et al.*, 1983). The error is defined by:

$$\text{Error} = W_i(Q_i - q_i),$$

and a plot of values over the temperature range of the measurements provides a visual picture of the overall quality of the data set. An error plot is shown, for example, in Fig. 4.3.4.1_1 for a set of calorimetric measurements of anorthite's heat capacity reported by Krupka *et al.* (1979). Ideally, the errors should be centered about the zero axis, and should not exceed the precision of the measurement technique (two standard deviations are reported by Krupka *et al.*, 1979). As can be seen, all the data meet these requirements, and all are therefore considered consistent.

In similar experiments, some of the data may exceed two standard deviations from the weighted average among all the data in the set, however, and these data are then considered to be *discordant*. Rejection of discordant data is not automatic, and two approaches are taken for dealing with such data. First, the data may be rejected if an examination of experimental technique points to possible causes for the discrepancies. Potential sources of experimental error include inadequate compositional analyses of starting materials, lack of information concerning the ordering/disordering state of the reactants, and other errors or ambiguities, which can only be detected by thorough assessments of the experimental technique.

If errors in the experimental procedures cannot be identified, then the discordant data must be retained in the evaluation because there is no scientific basis for rejecting these data (i.e., rejection criteria are not based solely on statistical evaluations). In such cases, the influence of discordant data on

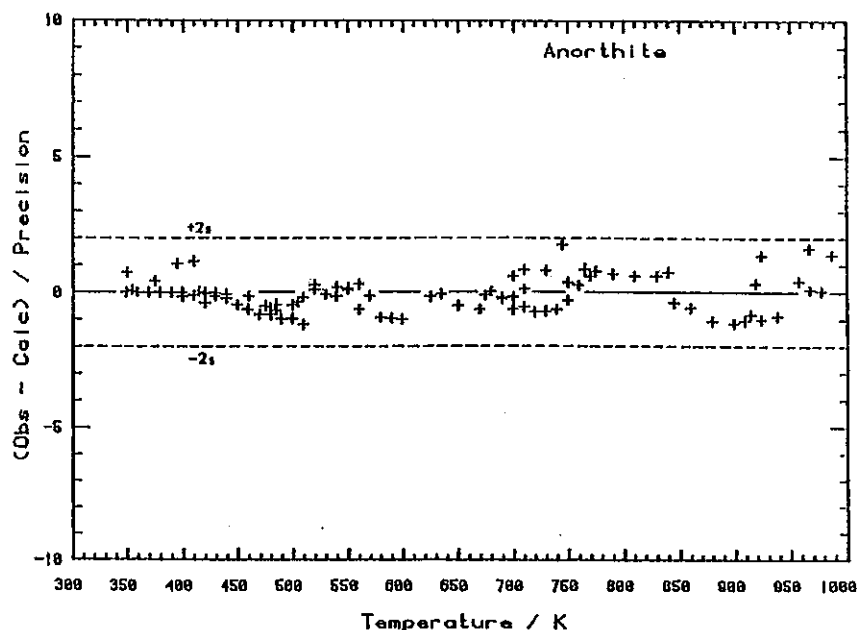


Figure 4.3.4.1_1. Error plot as a function of temperature (from Robinson *et al.*, 1983) for differential scanning calorimetry measurements of the heat capacity of anorthite reported by Krupka *et al.* (1979).

regression parameters is minimized through adjustment of weighting parameters. Such adjustments may involve increasing the magnitude of the measurement precision (i.e., acknowledging greater uncertainty in these data) thus decreasing the weight accorded to the discrepant data and their influence on the residual function. Regression techniques for evaluation of thermodynamic data are described by Haas (1974), Haas and Fisher (1976), Robinson *et al.* (1983), and Nordstrom and Munoz (1985).

4.3.4.2 Reactions. Standard molal thermodynamic properties retrieved from experiments involving reactions are not independent because they are determined by the individual thermodynamic properties of all reactant and product phases and aqueous species. Reaction-based data that constrain the internally consistent databases described in Sections 3.2 and 3.3 are predominantly constrained by phase-equilibrium results.

Two statistical/mathematical approaches have been developed for evaluation of thermodynamic properties from phase-equilibrium studies (Berman *et al.*, 1986):

- mathematical (or linear) programming (MAP), and
- regression analysis.

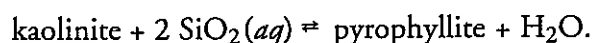
The MAP approach translates each P - T half-bracket into an inequality relation [Eqn. (4.3.1.1_1)] and a corresponding inequality constraint equation given by the following *objective* function (Engi, 1992):

$$O_b = \sum_{i=1}^n W'_i (Q'_i - q'_i)^2, \quad (4.3.4.2_1)$$

where the primed symbols are similar to those discussed above for the residual function evaluated in regression analyses, but differ in the sense that parameter values may be determined using both equality and inequality constraints. In this sense, MAP differs from regression methods because the objective function is capable of operating on individual half-brackets, rather than on pairs of half-brackets and associated assumptions about the true location of the equilibrium point, and more generally about the probability distribution of the equilibrium position between half-brackets. The merits and deficiencies of both approaches are considered by Berman *et al.* (1986), Holland and Powell (1990) and Kolassa (1990). Both approaches have been used successfully to evaluate thermodynamic data from phase-equilibrium studies using a minimum of additional calorimetric constraints.

The regression-based methods for data selection described earlier for individual properties are similar in principle to selection methods based on regression applied to individual reactions. Selection criteria derived from MAP evaluations are also basically determined by comparison of the weighted average of a set of thermodynamic properties with observed values. Discordant data are readily identified in such comparisons, and may be rejected if warranted by a subsequent review of experimental procedures.

Robinson *et al.* (1983), for example, used regression methods to evaluate solubility data from Hemley *et al.* (1980) and phase-equilibrium data reported by Thompson (1970) for the reaction (see Section 4.3.1.2):



The Gibbs free energy change for this reaction over a broad temperature range is shown in Fig. 4.3.4.2_1, where it can be seen that the Hemley *et al.* (1980) data are closely approximated by the weighted average of the full data set. The Thompson (1970) data, however, are consistently displaced upwards with respect to the weighted average.

An error plot for these data (Fig. 4.3.4.2_2) reveals that the solubility data all lie within one standard deviation of the weighted average (one standard deviation is the error of precision assigned to these data by Hemley *et al.*, 1980). In contrast, the Thompson (1970) data exceed two standard deviations of the weighted average [two standard deviations is the precision reported by Thompson (1970)]. These results suggest that the Thompson (1970) data are, at best, discordant. A review of the experimental technique used by Thompson (1970) suggests that the reason for the discordance could be due to the finely ground kaolinite used in the experiments and/or to the short duration of the

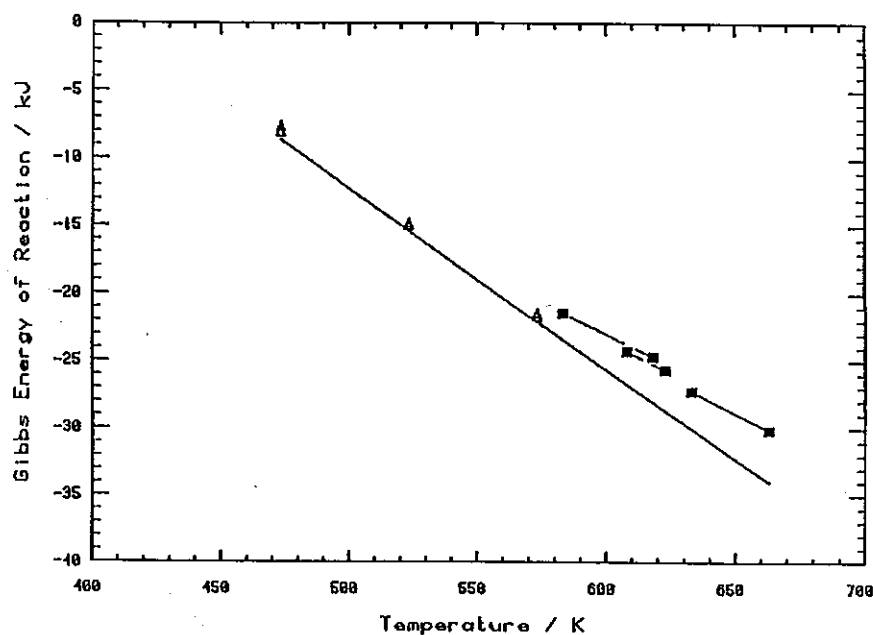


Figure 4.3.4.2_1. Standard Gibbs free energies of reaction as a function of temperature for the reaction: kaolinite + 2 SiO₂(aq) = pyrophyllite + H₂O (from Robinson *et al.*, 1983). Open symbols represent calculations based on measurements of dissolved silica concentrations reported by Hemley *et al.* (1980). Solid symbols connected by lines denote values computed from the phase-equilibrium experiments of Thompson (1970).

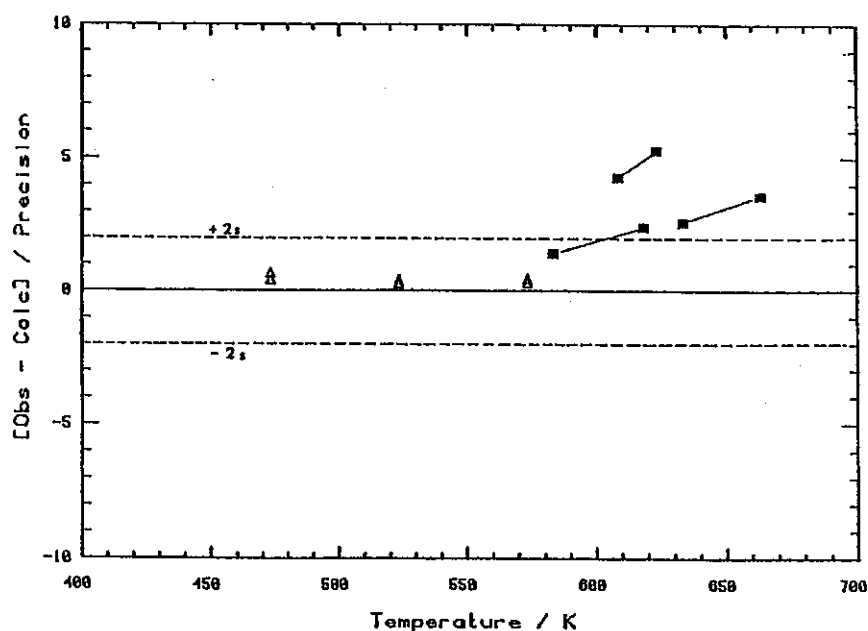


Figure 4.3.4.2_2. Error plot as a function of temperature for the reaction: kaolinite + 2 SiO₂(aq) = pyrophyllite + H₂O (from Robinson *et al.*, 1983). Open symbols represent the experimental results of Hemley *et al.* (1980). Closed symbols connected by lines represent *P-T* half brackets determined by Thompson (1970). The dashed lines denote two times the experimental precision reported by Hemley *et al.* (1980), or two times the width of the Gibbs free energy of reaction "bracket" corresponding to the *P-T* half brackets.

experiments (Robinson *et al.*, 1983). The Thompson (1970) data were finally rejected by Robinson *et al.* (1983) for these reasons, not simply because the data are discordant based on the statistical evaluation of the full data set.

This example again illustrates that discordant data can be readily identified using statistical methods, but that the cause of the disagreement may not always be as straight-forward as its detection. Thus, thermodynamic relations and statistical/mathematical evaluations are useful for detecting consistent and discordant data, but technical competence and experience in laboratory investigations is indispensable for final data selection/rejection decisions.

4.3.4.3 Simultaneous evaluation of multiple reactions. Matrix analytical methods are used in simultaneous evaluations of multiple reactions to set up a *general optimization* problem similar to that described in Section 4.3.4.2. These methods also ensure internal consistency, and have been used to construct all the internally consistent databases discussed in Sections 3.2 and 3.3.

Matrix methods are based on a constraint matrix, which consists of an ordered array of constraints written in matrix format. Constraint matrices developed for MAP analyses include two sub-matrices, which are defined separately using equality and inequality constraints. A single matrix consisting of equality constraints is set up for regression-based methods.

An instructive example of the approach is given by Powell and Holland (1985), who adopt a regression method based on calculation of enthalpies of reaction from phase-equilibrium data (Section 4.3.1.1) to construct the following matrix:

$$\mathbf{b} = \mathbf{R}\mathbf{h}, \quad (4.3.4.3_1)$$

where $\mathbf{R}$ denotes the matrix of stoichiometric coefficients among a set of reactions, $\mathbf{h}$ stands for the vector of enthalpies of formation for each of the minerals considered (i.e., these are the primary values being calculated in this example), and $\mathbf{b}$ refers to the vector of reaction enthalpies, which are calculated using an appropriate statement of Eqn. (4.2.3_1). For this particular application, $\mathbf{b}$ represents a vector of "centers" of reaction enthalpies corresponding to pairs of P - T half brackets determined in a set of phase-equilibrium experiments.

The objective of the calculations in this regression approach is to minimize values of the following quantity:

$$\|\mathbf{b} - \mathbf{R}\mathbf{h}\|^2 \quad (4.3.4.3_2)$$

which, for the example considered by Powell and Holland (1985) involving high albite (*ab*), paragonite (*pa*), andalusite (*and*), kyanite (*ky*), jadeite (*jd*), corun-

dum (*cor*), quartz (*q*) and water (H_2O) is given by:

h^T = enthalpies of formation of

$$R = \begin{bmatrix} ab & pa & and & ky & jd & cor & q & H_2O \\ 1 & 0 & 0 & 0 & -1 & 0 & -1 & 0 \\ 0 & -1 & 0 & 1 & 1 & 0 & 0 & 1 \\ 1 & -1 & 0 & 0 & 0 & 1 & 0 & 1 \\ 1 & -1 & 1 & 0 & 0 & 0 & -1 & 1 \\ 1 & -1 & 0 & 1 & 0 & 0 & -1 & 1 \end{bmatrix}$$

$$b = \begin{bmatrix} \Delta H_{r,P,T}^\circ (jd+q = ab) \\ \Delta H_{r,P,T}^\circ (pa = jd+ky+H_2O) \\ \Delta H_{r,P,T}^\circ (pa = ab+cor+H_2O) \\ \Delta H_{r,P,T}^\circ (pa+q=ab+and+H_2O) \\ \Delta H_{r,P,T}^\circ (pa+q = ab+ky+H_2O) \end{bmatrix}$$

Although the matrix can be solved in principle using conventional methods, i.e.,

$$h = (R^T R)^{-1} R^T b,$$

the problem considered in this example is not amenable to solution because there are five equations (fixing the enthalpies of reaction) and eight unknowns (enthalpies of formation of *ab*, *pa*, *and*, *ky*, *jd*, *cor*, *q*, and H_2O). To overcome this problem, Powell and Holland (1985) expand R by six rows of the form:

$$[0, 0, \dots, 0, 1, 0, \dots, 0, 0]$$

where the value, 1, is centered on the position of an "anchor" mineral, for which a corresponding row in b is added containing a calorimetrically determined reference value for the enthalpy of formation of the anchor phase.

The anchor minerals in this example are assumed to be *and*, *ky*, *jd*, *cor*, *q* and H_2O , for which calorimetric values for the enthalpy of formation at 1 bar and 25°C from Robie *et al.* (1978) are used to calculate enthalpies of formation for *ab* and *pa* (Powell and Holland, 1985). These authors also demonstrate how weighting factors can be incorporated into the matrix formulation of the residual function in regression evaluations.

Matrix analytical methods are useful because they ensure that derived databases are internally consistent (Section 3.1.1). There are two important limitations in this approach, however. First, errors in any one anchor datum will be propagated throughout the database. The database derived by Helgeson *et al.* (1978), for example, appears to suffer from this problem (Nordstrom and Munoz, 1985). Standard molal Gibbs free energies and enthalpies of formation for aluminous minerals in this database contain a systematic error of 1.55 kcal mol⁻¹, which was traced by Hemingway *et al.* (1982) to the use of an incorrect value adopted for

the reaction kaolinite + H₂O = 2 gibbsite + 2 SiO₂(aq). The incorrect value was subsequently used by Helgeson *et al.* (1978) to calculate the enthalpy of formation of kaolinite. Kaolinite was adopted as a reference mineral, and the error was therefore propagated throughout all aluminous phases in this database. Re-evaluation of the Al₂O₃-H₂O-NaCl system (Pokrovskii and Helgeson, 1995) may have partially corrected this problem (Section 3.5), but associated effects on the internal consistency of the database are uncertain as a result (Section 6.4.1).

The second weakness in the use of matrix analytical methods for simultaneous evaluation of thermodynamic data is simply the fact that the matrices are difficult to evaluate for complex, multi-component systems. Thus it is not a simple matter to undertake periodic updates and revisions of internally consistent thermodynamic databases.

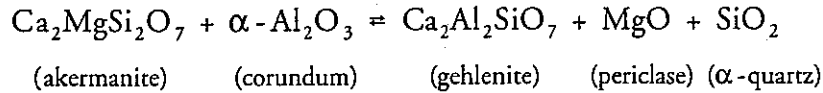
4.4 Estimation of Thermodynamic Properties and Equation-of-State-Parameters

The standard molal volume, entropy, heat capacity and equation-of-state parameters of several minerals in SPRONS.JNC are estimated by Helgeson *et al.* (1978) using techniques that are briefly described in Section 4.4.1. Similarly, the standard partial molal volume, entropy, heat capacity and equation-of-state parameters of many aqueous ions are estimated by Shock and Helgeson (1988), Shock *et al.* (1997a), Sverjensky *et al.* (1997) and others, based on correlations that are deduced empirically from relations observed among experimentally determined standard partial molal properties of other ionic solutes. The correlations that have been used to estimate thermodynamic properties in SPRONS.JNC are described in Section 4.4.2.

4.4.1 Minerals

Helgeson *et al.* (1978) use the concept of structurally identical, or similar, classes of minerals as the basis for estimation of standard molal entropies, volumes and heat capacities. The structural classes considered include ortho and ring silicates, chain and band silicates, framework silicates, sheet silicates and carbonates. Estimation techniques for standard molal volumes and entropies are described below. This is followed by a brief description of the technique used to estimate standard molal heat capacities.

Reasonably accurate estimates of the standard molal entropy or volume of minerals can be obtained by first writing a reversible reaction between the mineral for which the standard entropy or volume is to be estimated, and other minerals in the same or similar structural class whose entropies or volumes are known, using oxide formula units to balance the reaction if necessary. If, for example the standard entropy, or volume, of the mineral akermanite is unknown, then an appropriate reaction for estimation of these quantities is:



for which,

$$\Delta S_r^\circ = S_{\text{gehlenite}}^\circ + S_{\text{periclase}}^\circ + S_{\alpha\text{-quartz}}^\circ - S_{\text{corundum}}^\circ - S_{\text{akermanite}}^\circ, \quad (4.4.1_1)$$

and,

$$\Delta V_r^\circ = V_{\text{gehlenite}}^\circ + V_{\text{periclase}}^\circ + V_{\alpha\text{-quartz}}^\circ - V_{\text{corundum}}^\circ - V_{\text{akermanite}}^\circ. \quad (4.4.1_2)$$

Helgeson *et al.* (1978) rearrange equations such as Eqn. (4.4.1_1) and (4.4.1_2) such that unknown parameters, including the unknown entropy and volume of reaction, are grouped together on the left-hand sides of these equations, i.e.,:

$$S_{\text{akermanite}}^\circ + \Delta S_r^\circ = S_{\text{gehlenite}}^\circ + S_{\text{periclase}}^\circ + S_{\alpha\text{-quartz}}^\circ - S_{\text{corundum}}^\circ, \quad (4.4.1_3)$$

and

$$V_{\text{akermanite}}^\circ + \Delta V_r^\circ = V_{\text{gehlenite}}^\circ + V_{\text{periclase}}^\circ + V_{\alpha\text{-quartz}}^\circ - V_{\text{corundum}}^\circ. \quad (4.4.1_4)$$

Then, designating the left-hand sides of such equations generally as S_i° and V_i° , respectively, the standard molal entropy of the i -th mineral (e.g., akermanite) can be estimated using:

$$S_{i, P_r, T_r}^\circ = \frac{S_{i, P_r, T_r}^\circ (V_{s, P_r, T_r}^\circ + V_{i, P_r, T_r}^\circ)}{2V_{s, P_r, T_r}^\circ}, \quad (4.4.1_5)$$

assuming the standard molal volume of the mineral is known. Conversely, this equation can be re-arranged to estimate the standard molal volume of the mineral if the standard molal entropy is known.

Equation (4.4.1_5) is used to estimate the standard molal entropy of several non-ferrous silicate minerals in SPRONS.JNC (see Table 3 of Helgeson *et al.*, 1978). It is also appropriate for non-ferrous carbonates, chlorides, sulfates and oxides, but inappropriate for ferrous silicates, due to crystal field stabilization effects imposed on silicate structures by the ferrous ion. As a first approximation, Helgeson *et al.* (1978) use a modified version of Eqn. (4.4.1_5) for such minerals given by:

$$S_{i, P_r, T_r}^\circ = \frac{S_{s, P_r, T_r}^\circ (V_{s, P_r, T_r}^\circ + V_{i, P_r, T_r}^\circ)}{2V_{s, P_r, T_r}^\circ} - 2\nu_{Fe, i}, \quad (4.4.1_6)$$

where $\nu_{Fe, i}$ stands for the number of moles of ferrous iron per mole of the mineral. Equation (4.4.1_6) is used to estimate the standard molal entropy of

several ferrous silicate minerals in SPRONS.JNC (see Table 3 of Helgeson *et al.*, 1978).

Similarly, the contribution of structural H₂O to the standard molal entropies of hydrous silicates is estimated by assigning a value of 9.6 cal mol⁻¹ K⁻¹ at 25°C and 1 bar to the oxide formula unit ($S_{H_2O}^\circ$) representing H₂O that is tightly bound in crystallographic lattices (Helgeson *et al.*, 1978). Loosely bound H₂O in the structural channels of zeolites is assigned a value of 8 cal mol⁻¹ K⁻¹.

Helgeson *et al.* (1978) have shown that the estimated entropy values of ferrous and non-ferrous silicates at 25°C and 1 bar generally agree to within ±1% of values determined calorimetrically. The effects of errors of this magnitude on calculated entropies of reactions at high temperatures are negligible. The effects of such errors on calculated Gibbs free energies of formation at high temperatures may exceed several hundred cal mol⁻¹ K⁻¹, but the errors tend to cancel if all the estimated data are consistent with one another, and if more than one estimated value is involved in the calculation.

Heat capacities of minerals are estimated by Helgeson *et al.* (1978) on the basis of additivity constraints and the assumption that the heat capacity of reaction among oxide and silicate minerals of the same structural class is equal to zero at all temperatures. This approach is formulated as a *structural algorithm* given by:

$$C_{P_x}^\circ = \sum_i^x \nu_{i,r} C_{P_i}^\circ / \nu_{x,r} \quad (4.4.1_7)$$

where the properties to be estimated are represented by the subscript x , those that are known are distinguished by the subscript i , and $\nu_{x,r}$ and $\nu_{i,r}$ denote stoichiometric coefficients (positive for products and negative for reactants) of the unknown and known minerals, respectively, in a given reference reaction among structurally similar phases. Comparison of heat capacities estimated using this algorithm with corresponding data determined calorimetrically indicates that discrepancies are generally less than 2% of the calorimetric values.

Restating Eqn. (4.4.1_7) in terms of equation-of-state parameters for minerals (a_i , b_i , c_i) consistent with the Maier-Kelly equation [Eqn. (4.2.1_5)] allows estimates of these parameters to be made in accordance with the following equations (Helgeson *et al.*, 1978):

$$a_x = \sum_i^{x-1} \nu_{i,r} a_i / \nu_{x,r}, \quad (4.4.1_8)$$

$$b_x = \sum_i^{x-1} \nu_{i,r} b_i / \nu_{x,r}, \quad (4.4.1_9)$$

and

$$c_x = \sum_i^{x-1} v_{i,r} c_i / v_{x,r} \quad (4.4.1_{10})$$

Maier-Kelly coefficients for 66 minerals in SPRONS. JNC are estimated using Eqns. (4.4.1_8) - (4.4.1_10) (Helgeson *et al.*, 1978 - Table 7). Comparison of similar estimates with experimental counterparts indicates that the estimated coefficients may differ considerably with experimental values (Helgeson *et al.*, 1978), but that corresponding heat capacities are nevertheless closely approximated at both low and high temperatures. This is because discrepancies among the coefficients tend to cancel when they are used in Eqn. (4.2.1_5).

4.4.2 Aqueous solutes

As noted in Section 4.3.3, values of the standard partial molal entropy, heat capacity, volume and conventional Born coefficient of an aqueous solute are needed at the reference temperature and pressure in order to calculate the standard partial molal Gibbs free energy and enthalpy of formation of the solute at other temperatures and pressures. Although the appropriate data have been determined experimentally for many aqueous solutes, there are numerous ions and other aqueous species for which experimental data are presently lacking.

To address this situation, Shock and Helgeson (1988) proposed several empirical correlations among $\bar{S}^\circ$, $\bar{C}_p^\circ$, and $\bar{V}^\circ$. The correlations plot as straight lines on diagrams of $\bar{S}^\circ$ vs. $\bar{V}^\circ$, $\bar{S}^\circ$ vs. $\bar{C}_p^\circ$, or $\bar{V}^\circ$ vs. $\bar{C}_p^\circ$, where the slopes and intercepts of the lines are determined by groups of ions or ionic species of similar type and charge. Shock *et al.* (1997a) recently proposed revisions to the original correlations of Shock and Helgeson (1988), and have also proposed new correlations for additional types of ionic species. Additional correlations for metal-ligand complexes have also been recently generated by Sverjensky *et al.* (1997). Because the correlations developed by Shock and Helgeson (1988), Shock *et al.* (1997a) and Sverjensky *et al.* (1997) are the basis for estimates of partial molal properties and equation-of-state parameters for numerous aqueous solutes in SPRONS.JNC, the correlation equations and parameters, and supporting figures illustrating correlation behavior are summarized in Table 4.4.2_1 and Figs. 4.4.2_1 to 4.4.2_23. Figures 4.4.2_1 to 4.4.2_15 are from Shock *et al.* (1997). The remaining figures are from Sverjensky *et al.* (1997).

Uncertainties in standard partial molal properties estimated using the correlations in Table 4.4.2_1 are evaluated by Sverjensky *et al.* (1997). Uncertainties in $\bar{S}^\circ$ and $\bar{C}_p^\circ$ are estimated to be within ± 5 cal mol⁻¹ K⁻¹. Uncertainties in $\bar{V}^\circ$ are estimated to be approximately ± 5 cm³ mol⁻¹. The uncertainties in the standard partial molal entropy and heat capacity are more significant than uncertainties in the standard partial molal volume, and would, for example, generate uncertainties in calculated values of the partial molal Gibbs free energy of formation at 500°C of ± 2400 cal mol⁻¹ and ± 1100 cal mol⁻¹,

Table 4.4.2_1. Correlations among standard partial molal volumes (V), heat capacities (C_p) and entropies (S) of cationic and anionic aqueous species at 25°C and 1 bar¹.

Species	Correlation Eqn.	Parameters		Figure
Ions, Oxygen-Bearing Ionic Species and Non-Conventional Hydroxide Complexes ^a				
	$V = y S + z$	y	z	
monovalent, monatomic cations, except Li ⁺		1.29	-20.5	4.4.2_1
divalent transition cations		0.92	0.0	
alkaline earth cations, and Pb ²⁺		0.19	-14.8	
trivalent cations excluding rare earths		0.50	-4.0	
monovalent, monatomic halide anions		1.29	1.8	see ref. c
other monovalent anions		1.06	0.07	4.4.2_4
non-metal oxygen bearing species (#O = 2)		1.05	-3.3	4.4.2_5
non-metal oxygen bearing species (#O = 3)		0.87	3.0	
non-metal oxygen bearing species (#O = 4)		0.78	11.6	
metal oxygen bearing species (#O = 2)		0.25	11.7	4.4.2_6a
1st row transition elements (#O=4)		0.75	9.5	4.4.2_6c
1st row transition elements (#O=4)		0.48	-8.2	4.4.2_6d
higher-order transition elements (#O=4)		0.44	27.3	4.4.2_6b
tetravalent cations		0.1	-43.3	est.
monovalent, MO ⁻		0.45	-12	est.
divalent MOH ⁺ (transition metals)		0.45	-12	est.
divalent MOH ⁺ (alkaline earths)		0.16	4.9	est.
divalent MO(aq) (transition metals)		0.45	-12	est.
divalent MO(aq) (alkaline earths)		0.16	4.9	est.
divalent HMO ₂ ⁻ (transition metals)		0.54	-4.8	est.
divalent HMO ₂ ⁻ (alkaline earths)		0.25	11.7	est.
divalent MO ₂ ²⁻ (transition metals)		0.54	-4.8	est.
divalent MO ₂ ²⁻ (alkaline earths)		0.25	11.7	est.
trivalent MOH ²⁺ (1st row transition metals)		0.45	-12	est.
trivalent MOH ²⁺ (all others)		0.16	4.9	est.
trivalent MO ⁺ (1st row transition metals)		0.45	-12	est.
trivalent MO ⁺ (all others)		0.16	4.9	est.
trivalent HMO ₂ (1st row transition metals)		0.54	-4.8	est.
trivalent HMO ₂ (all others)		0.25	11.7	est.
tetravalent MOH ³⁺		0.16	4.9	est.
tetravalent MO ²⁺		0.16	4.9	est.
tetravalent HMO ₂ ⁺		0.25	11.7	est.
tetravalent MO ₂ (aq)	0.25	11.7	est.	
tetravalent HMO ₃ ⁻	0.35	19.3	est.	
	$C_p = u S + v$	u	v	
monovalent, monatomic cations, except Li ⁺		-0.88	22.2	4.4.2_2

Species	Correlation Eqn.	Parameters		Figure
other monovalent cations		0.064	14.2	4.4.2_2
divalent transition cations		0.53	5.5	
alkaline earth cations and Pb ²⁺		-0.19	-11.9	
trivalent cations excluding rare earths		0.91	41.7	
monovalent, monatomic halide anions		0.0	-28.2	see ref. c
other monovalent anions		0.62	-31.6	4.4.2_4
non-metal oxygen-bearing anions ($Z_j = -1$)		0.60	-33.8	4.4.2_7
non-metal oxygen-bearing anions ($Z_j = -2$)		0.60	-63.3	
non-metal oxygen-bearing anions ($Z_j = -3$)		0.60	-92.8	est.
oxyacids and acid oxyanions		1.65	-46.8	4.4.2_8
metal oxygen bearing species (unprotonated)		-2.06	-34.5	4.4.2_9
metal oxygen-bearing species (protonated)		-2.06	95.5	
tetravalent cations		0.31	31.1	est.
monovalent MOH(aq)		-1.14	9	est.
monovalent MO ⁻		-1.14	15.5	est.
divalent MOH ⁺		-1.14	9.0	4.4.2_10
divalent MO(aq)		-1.14	-15.5	
divalent HMO ₂ ⁻		-2.28	-24.	4.4.2_11
divalent MO ₂ ²⁻		-2.28	-106.2	
trivalent MOH ²⁺		-1.14	-37.2	est.
trivalent MO ⁺		-1.14	-60.8	est.
trivalent HMO ₂ (aq)		-2.28	-24	est.
trivalent MO ₂ ⁻		-2.06	-34.5	est.
tetravalent MOH ³⁺		-1.14	-37.2	est.
tetravalent MO ²⁺		-1.14	-60.8	est.
tetravalent HMO ₂ ⁺		-2.28	-24	est.
tetravalent MO ₂ (aq)		-2.06	-34.5	est.
tetravalent HMO ₃ ⁻		-2.28	-24	est.
	$C_p = w V + x$	w	x	
monovalent, monatomic cations, except Li ⁺		-0.68	8.3	4.4.2_3
divalent transition cations		0.58	5.5	
alkaline earth cations, and Pb ²⁺		-1.0	-26.7	
trivalent cations excluding rare earths		1.82	49.0	
monovalent, monatomic halide anions		0.0	-28.2	est.
non-metal oxygen-bearing anions ($Z_j = -1$)		0.55	-32.	4.4.2_12
non-metal oxygen-bearing anions ($Z_j = -2$)		0.55	-70.	
non-metal oxygen-bearing anions ($Z_j = -3$)		0.55	-108.	est.
oxyacids and acid oxyanions		1.94	-65.8	4.4.2_13
metal oxygen bearing species (#O = 2)		-0.12	7.3	4.4.2_14
metal oxygen bearing species (#O = 4)		0.40	45.8	

Species	Correlation Eqn.	Parameters		Figure
tetravalent cations		3.1	165.3	est.
	<i>S</i> correlations ²	e	f	
monovalent MOH(<i>aq</i>)		1.32	-6	est.
monovalent MO ⁻		1.42	-11	est.
divalent MOH ⁺		1.32	24.5	4.4.2_15
divalent MO(<i>aq</i>)		1.42	20.5	
divalent HMO ₂ ⁻		1.52	15.5	
divalent MO ₂ ²⁻		1.62	11.	est.
trivalent MOH ²⁺		1.32	37	est.
trivalent MO ⁺		1.42	83	est.
trivalent HMO ₂ (<i>aq</i>)		1.52	123	est.
trivalent MO ₂ ⁻		1.62	118	est.
tetravalent MOH ³⁺		1.32	74	est.
tetravalent MO ₂ ⁺		1.42	108	est.
tetravalent HMO ₂ ⁺		1.52	140	est.
tetravalent MO ₂ (<i>aq</i>)		1.62	135	est.
tetravalent HMO ₃ ⁻		1.72	130	est.
Metal-Ligand Complexes ⁶				
	<i>S</i> correlations ³			
neutral Cl complexes, <i>y</i> = 1 or 2		footnote #3		4.4.2_16
monovalent cations and ligands, <i>y</i> = 1 or 2		footnote #7		4.4.2_20
divalent cations, monovalent ligands, <i>y</i> = 1		footnote #8		4.4.2_21
sulfate complexes		footnote #9		4.4.2_22
carbonate complexes		footnote #10		4.4.2_23
	<i>C_P</i> correlations ⁴	d		
monovalent cations, <i>y</i> = Cl ⁻		18.0		4.4.2_17
divalent cations, <i>y</i> = Cl ⁻		63.3		
divalent cations, <i>y</i> = CH ₃ COO ⁻		93.8		
	<i>V</i> correlations ^{5,6}			
complexes containing monovalent ligands		footnote #5		4.4.2_18
metal sulfate complexes		footnote #6		4.4.2_19

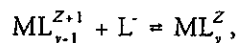
^a - Shock *et al.*, 1997a; ^b - Sverjensky *et al.*, 1997; ^c - Shock and Helgeson, 1988.

- 1- #O refers to the number of oxygen atoms in the complex; "est." stands for estimated correlation parameters; M denotes a cation; *Z_j* represents the charge of the *j*-th hydrogen-free anion.
- 2- The following correlation equation is used to relate the standard partial molal entropy of non conventional hydroxide complexes (Section 5.2.5) to the standard partial molal entropy of the corresponding cation:

$$\bar{S}_{\text{complex}}^{\circ} = e \bar{S}_{M^{Z+}}^{\circ} + f,$$

where $\bar{S}_{\text{complex}}^{\circ}$ and $\bar{S}_{M^{Z+}}^{\circ}$ denote the entropy of the complex and cation, respectively.

- 3- Following the notation of Sverjensky *et al.* (1997) for stepwise association reactions involving monovalent ligands,



where M refers to a metal cation, y stands for the number of ligands (L) and Z denotes the charge on the y -th complex, the entropy correlation equation is given by:

$$\Delta \bar{S}_{r,y}^\circ = \alpha_{Z,L} \bar{S}_{ML_{y-1}^{Z+1}}^\circ + \beta_{Z,L},$$

where,

$$\Delta \bar{S}_{r,y}^\circ = \bar{S}_{ML_y^Z}^\circ - \bar{S}_{ML_{y-1}^{Z+1}}^\circ - \bar{S}_L^\circ,$$

and, for neutral complexes with $L = Cl^-$,

$$\alpha_{Z=0,Cl} = 0.357, \text{ and } \beta_{Z=0,Cl} = -4.085.$$

- 4- For the stepwise association reaction (footnote #3), the heat-capacity correlation equation for mono- and divalent cation complexes with Cl^- and CH_3COO^- (acetate) ligands is given by:

$$\Delta \bar{C}_{p,y=1 \text{ or } 2}^\circ = 1.25 \bar{C}_{p,M^{Z+1}}^\circ + d,$$

where $\Delta \bar{C}_{p,y=1 \text{ or } 2}^\circ$ refers to the standard partial molal heat capacity of the association reaction, $\bar{C}_{p,M^{Z+1}}^\circ$ stands for the standard partial molal heat capacity of the metal cation, and Z refers to the charge on the complex.

- 5- For the stepwise association reaction (footnote #3), the volume correlation equation for complexes containing monovalent ligands is given by:

$$\Delta \bar{V}_{r,y}^\circ = 0.11419 \bar{V}_{M^{Z+1}}^\circ + 8.9432,$$

where $\Delta \bar{V}_{r,y}^\circ$ represents the standard partial molal volume of the association reaction, $\bar{V}_{M^{Z+1}}^\circ$ refers to the standard partial molal volume of the metal cation, and Z refers to the charge on the complex.

- 6- For the stepwise association reaction (footnote #3), the volume correlation equation for complexes containing divalent ligands is given by:

$$\Delta \bar{V}_{r,y=1}^\circ = \gamma_Z \bar{V}_{M^{2+}}^\circ + \delta_Z,$$

where $\Delta \bar{V}_{r,y}^\circ$ represents the standard partial molal volume of the association reaction, $\bar{V}_{M^{2+}}^\circ$ refers to the standard partial molal volume of the metal cation, and Z refers to the charge on the complex. For metal-sulfate complexes:

$$\gamma_Z = -0.25Z - 0.3 \text{ and } \delta_Z = -2.88Z + 3.32$$

where Z denotes the charge on the complex.

- 7- For the stepwise association reaction (footnote #3), the parameters in analogous correlation equations for the standard partial molal entropies of association of complexes formed from

monovalent metal cations and monovalent ligands are:

$$\alpha_{Z=0,L} = -0.000479\bar{S}_L^\circ + 0.3185, \text{ and } \beta_{Z=0,L} = -0.05757\bar{S}_L^\circ - 3.4494,$$

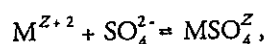
where $\bar{S}_L^\circ$ denotes the standard partial molal entropy of ligands, which are taken by Sverjensky *et al.* (1997) from Shock and Helgeson (1988).

- 8- Parameters in correlation equations for the standard partial molal entropies of association of complexes formed from divalent cations and monovalent ligands (footnote #3) are:

$$\alpha_{Z=1,L} = 0.015762\bar{S}_L^\circ + 0.03874, \text{ and } \beta_{Z=1,L} = 0.10344\bar{S}_L^\circ + 7.5807$$

where $\bar{S}_L^\circ$ denotes the standard partial molal entropy of ligands, which are taken by Sverjensky *et al.* (1997) from Shock and Helgeson (1988).

- 9- Parameters for standard partial molal entropies of association that are consistent with the reaction:



where M^{Z+2} refers to mono-, di- and trivalent cations, and the corresponding correlation equation;

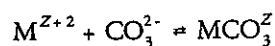
$$\Delta\bar{S}_{r,y=1}^\circ = \alpha_{Z,SO_4^{2-}}\bar{S}_{M^{Z+2}}^\circ + \beta_{Z,SO_4^{2-}},$$

where $\bar{S}_{M^{Z+2}}^\circ$ refers to the standard partial molal of the cation (taken from Shock and Helgeson, 1988), are given by:

$$\alpha_{Z,SO_4^{2-}} = -0.055Z + 0.055, \text{ and } \beta_{Z,SO_4^{2-}} = 13.84Z + 18.16,$$

where Z refers to the charge on the complex.

- 10- Parameters in correlation equations analogous to that in footnote #9 for the standard partial molal entropy of association of complexes formed from mono- and divalent cations and carbonate ligands, consistent with the reaction:



are given by:

$$\alpha_{Z,CO_3^{2-}} = -1.617Z + 0.213, \text{ and } \beta_{Z,CO_3^{2-}} = 66.82Z + 28.67,$$

where Z refers to the charge on the complex.

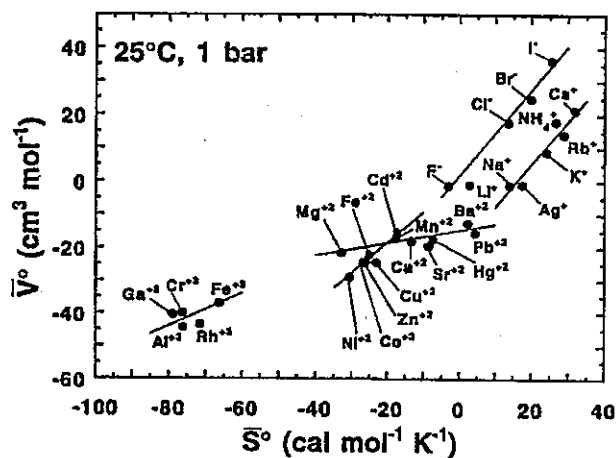


Figure 4.4.2_1. Correlation of standard partial molal volumes with entropies for monatomic ions (and NH_4^+) at 25°C and 1 bar. Symbols represent values retrieved by Shock *et al.* (1997a) and Shock and Helgeson (1988) from experimental data using the approach described in Section 4.3.3.

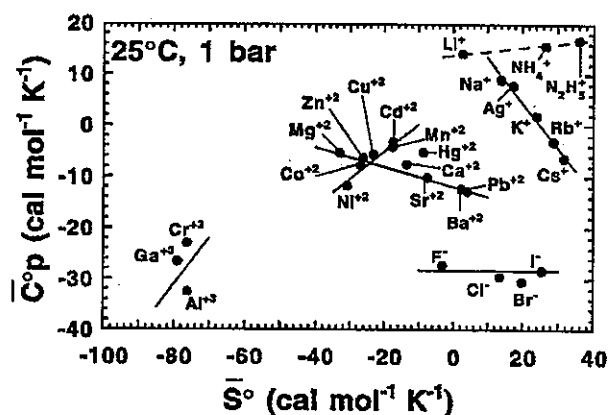


Figure 4.4.2_2. Correlation of standard partial molal heat capacities with entropies for monatomic ions (and NH_4^+ and N_2H_5^+) at 25°C and 1 bar. Symbols represent values retrieved by Shock *et al.* (1997a) and Shock and Helgeson (1988) from experimental results following the approach described in Section 4.3.3.

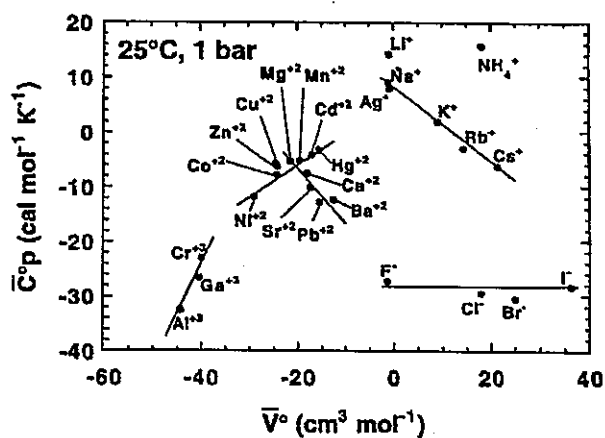
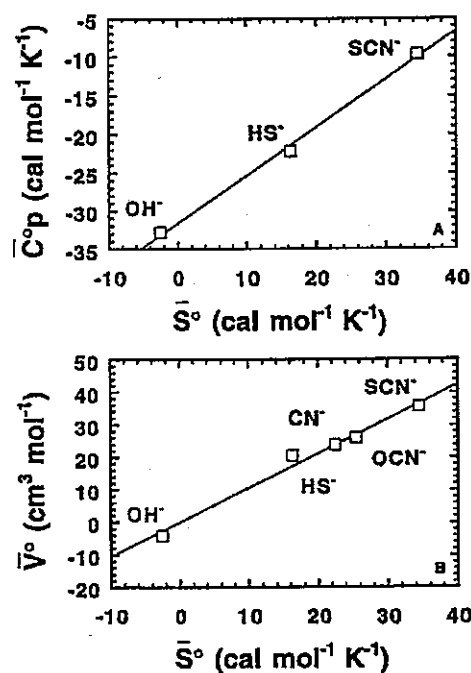


Figure 4.4.2_3. Correlation of standard partial molal heat capacities with volumes for monatomic ions (and NH_4^+) at 25°C and 1 bar. Symbols represent values retrieved by Shock *et al.* (1997a) and Shock and Helgeson (1988) from experimental data using the approach described in Section 4.3.3.



Figures 4.4.2_4 (above) and 4.4.2_5 (three figures to the right). Correlations of standard partial molal volumes with entropies at 25°C, 1 bar for monovalent anions (Fig.4.4.2_4) and non-metal oxygen-bearing species (Fig.4.4.2_5).

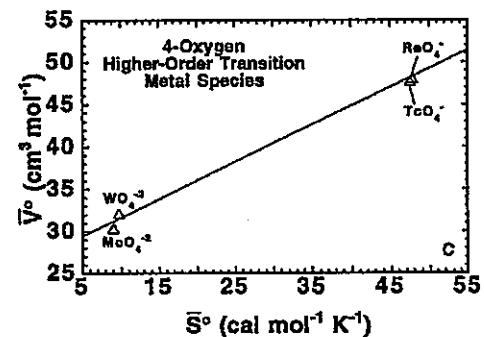
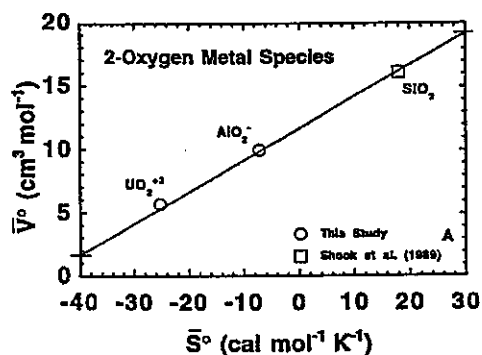
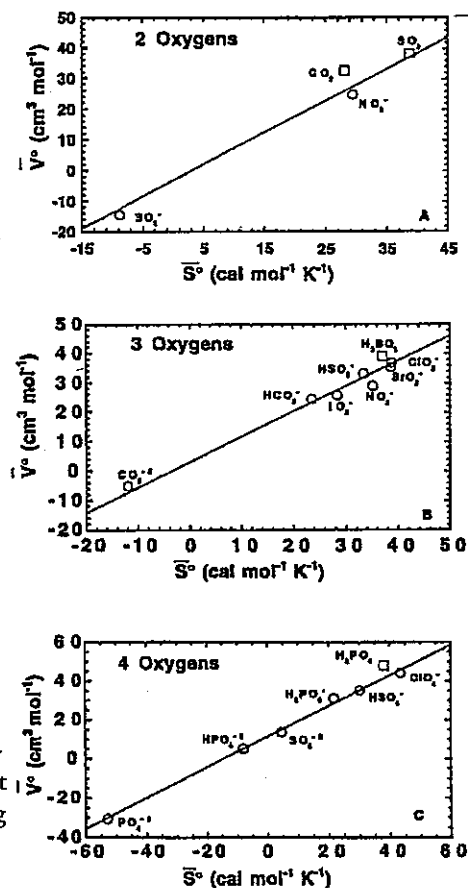
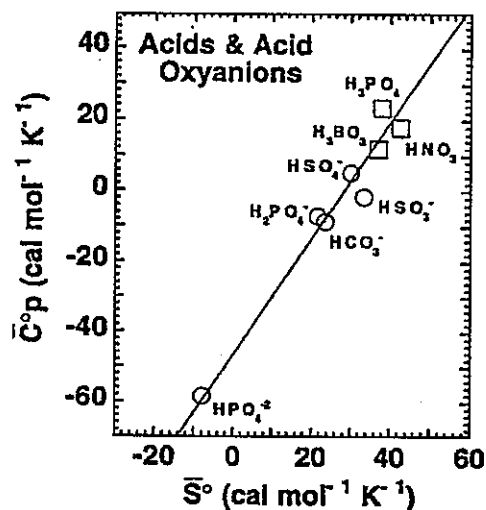
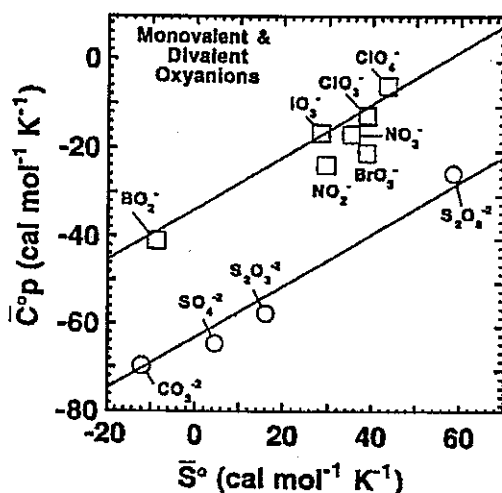
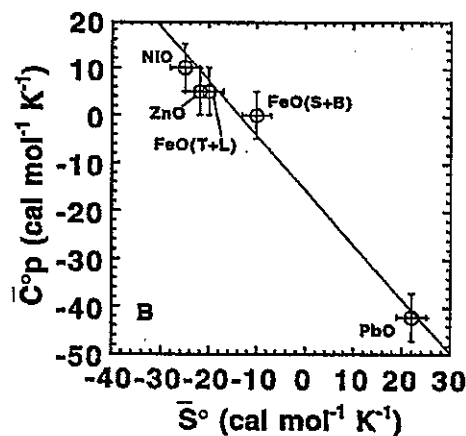
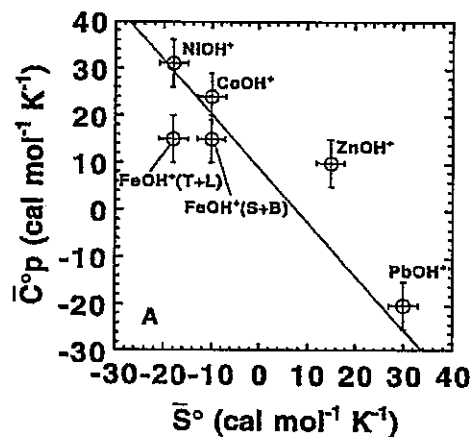
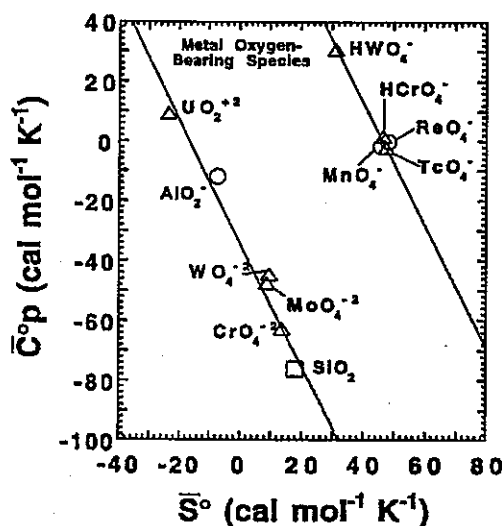


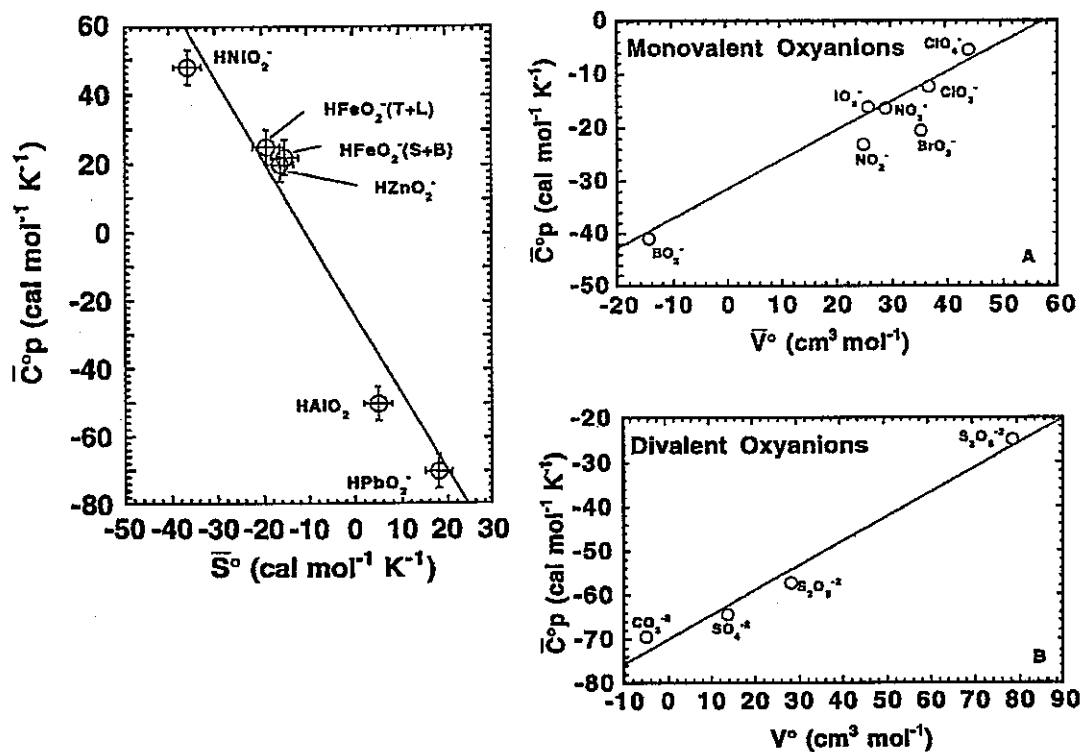
Figure 4.4.2_6 (a-d). Correlations of standard partial molal volumes with entropies at 25°C and 1 bar for metal oxygen-bearing species.



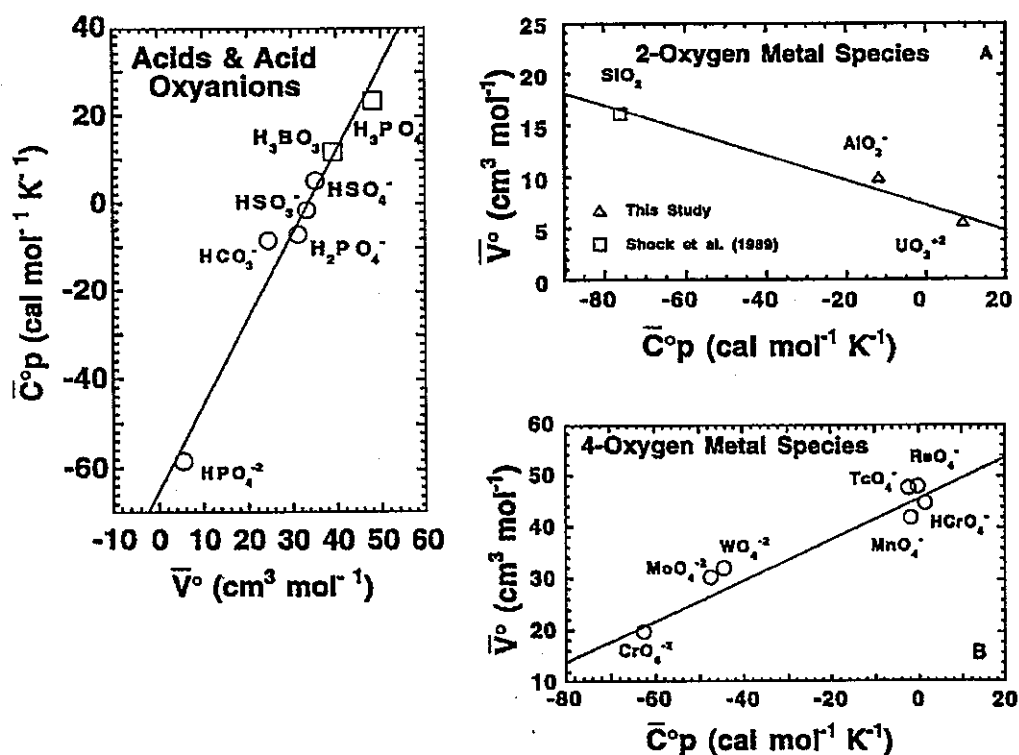
Figures 4.4.2_7 (left) and 4.4.2_8 (right). Correlations of standard partial molal heat capacities with entropies at 25°C and 1 bar for nonmetal oxygen bearing species [Fig. 4.4.2_7; monovalent anions (squares), divalent anions (circles)], and non-metal acid and acid oxyanion species [Fig. 4.4.2_8; circles from Shock *et al.* (1997a), squares from Shock and Helgeson (1988)]



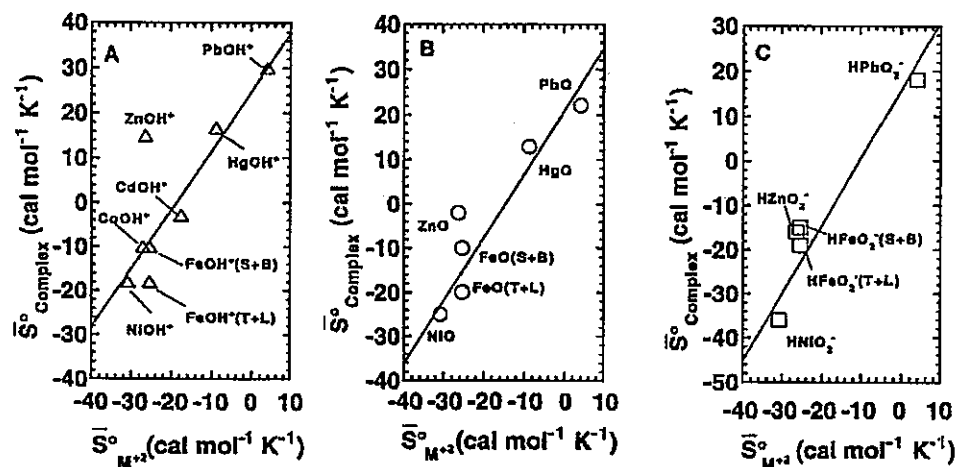
Figures 4.4.2_9 (above) and 4.4.2_10 (two figures to the right). Correlations of standard partial molal heat capacities with entropies for metal oxygen-bearing species (Fig. 4.4.2_9) and non-conventional hydroxide complexes of divalent cations (Fig. 4.4.2_10 - see Section 5.2.5).



Figures 4.4.2_11 (above) and 4.4.2_12 (two figures to the right). Correlations among $\bar{C}_p^\circ$ and $\bar{V}^\circ$ for non-conventional third hydroxide complexes of divalent metals (Fig. 4.4.2_11, see Section 5.2.5) and monovalent and divalent non-metal oxygen-bearing species (Fig. 4.4.2_12).



Figures 4.4.2_13 (above) and 4.4.2_14 (two figures to the right). Correlations among $\bar{C}_p^\circ$ and $\bar{V}^\circ$ for non-metal oxygen-bearing species that contain hydrogen (Fig. 4.4.2_13) and metal oxygen-bearing species (Fig. 4.4.2_14).



Figures 4.4.2_15. Correlations of standard partial molal entropies of non-conventional hydroxide complexes (Section 5.2.5) with the entropies of corresponding divalent metal cations at 25°C and 1 bar. Symbols refer to values regressed from experimental results by Shock *et al.* (1997a), and lines are consistent with correlation equations and associated parameters in Table 4.4.2_1.

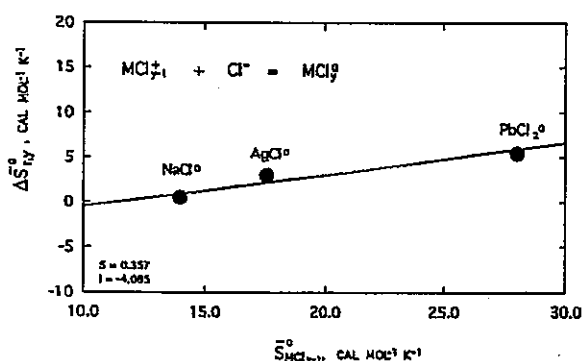


Figure 4.4.2_16. Correlation of the stepwise standard partial molal entropy of association of chloride complexes with the standard partial molal entropy of the cation at 25°C and 1 bar (see Table 4.4.2_1, footnote #3). Symbols refer to calculated values from Sverjensky *et al.* (1997) based on thermodynamic data that are incorporated in SPRONS.JNC.

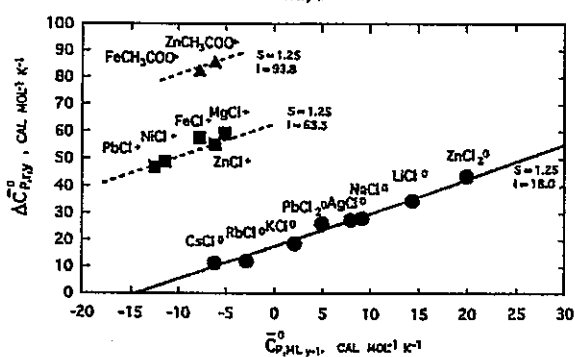


Figure 4.4.2_17. Correlation of the stepwise standard partial molal heat capacity of association with the standard partial molal heat capacity of the cation at 25°C and 1 bar (see Table 4.4.2_1, footnote #4). Symbols refer to calculated values from Sverjensky *et al.* (1997) based on thermodynamic data that are incorporated in SPRONS.JNC.

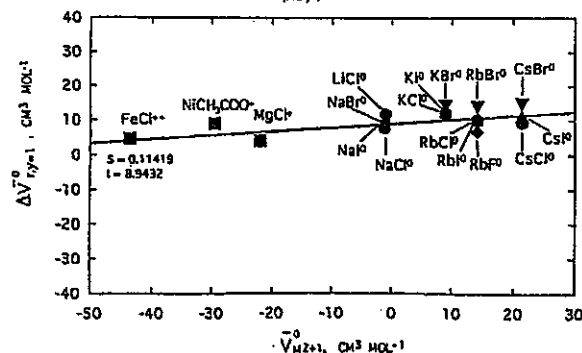


Figure 4.4.2_18. Correlation of the stepwise standard partial molal volume of association with the standard partial molal volume of the cation at 25°C and 1 bar (see Table 4.4.2_1, footnote #5). Symbols refer to calculated values from Sverjensky *et al.* (1997) based on thermodynamic data that are incorporated in SPRONS.JNC.

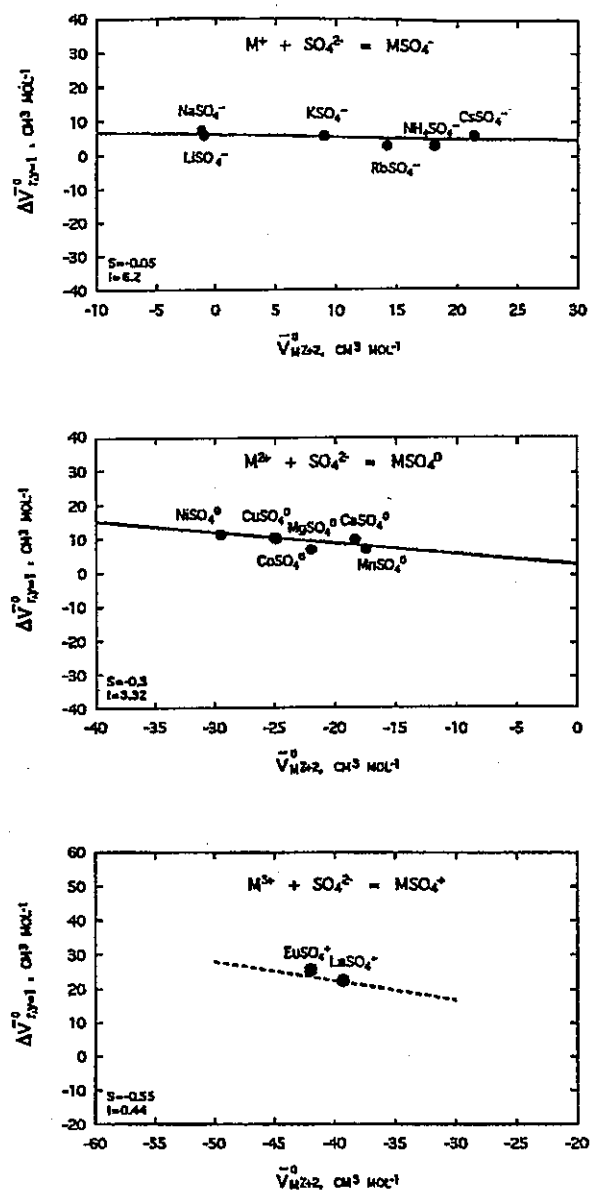


Figure 4.4.2_19. Correlation of the standard partial molal volumes of association of sulfate ion pairs involving mono-, di- and trivalent cations with the standard partial molal volumes of the cations at 25°C and 1 bar (see Table 4.2.2_1, footnote #6). Symbols refer to calculated values from Sverjensky *et al.* (1997) based on thermodynamic data that are incorporated in SPRONS.JNC.

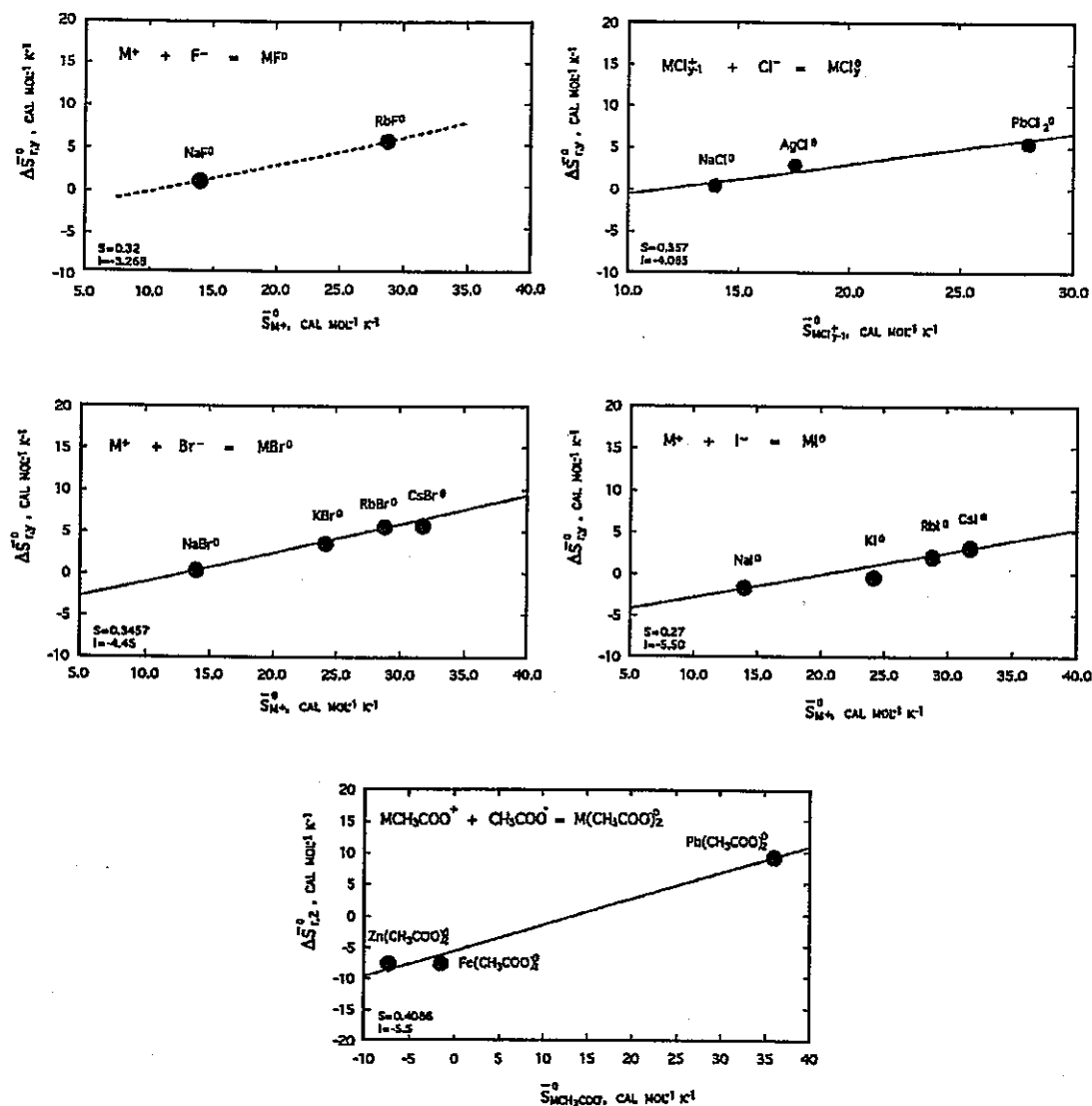
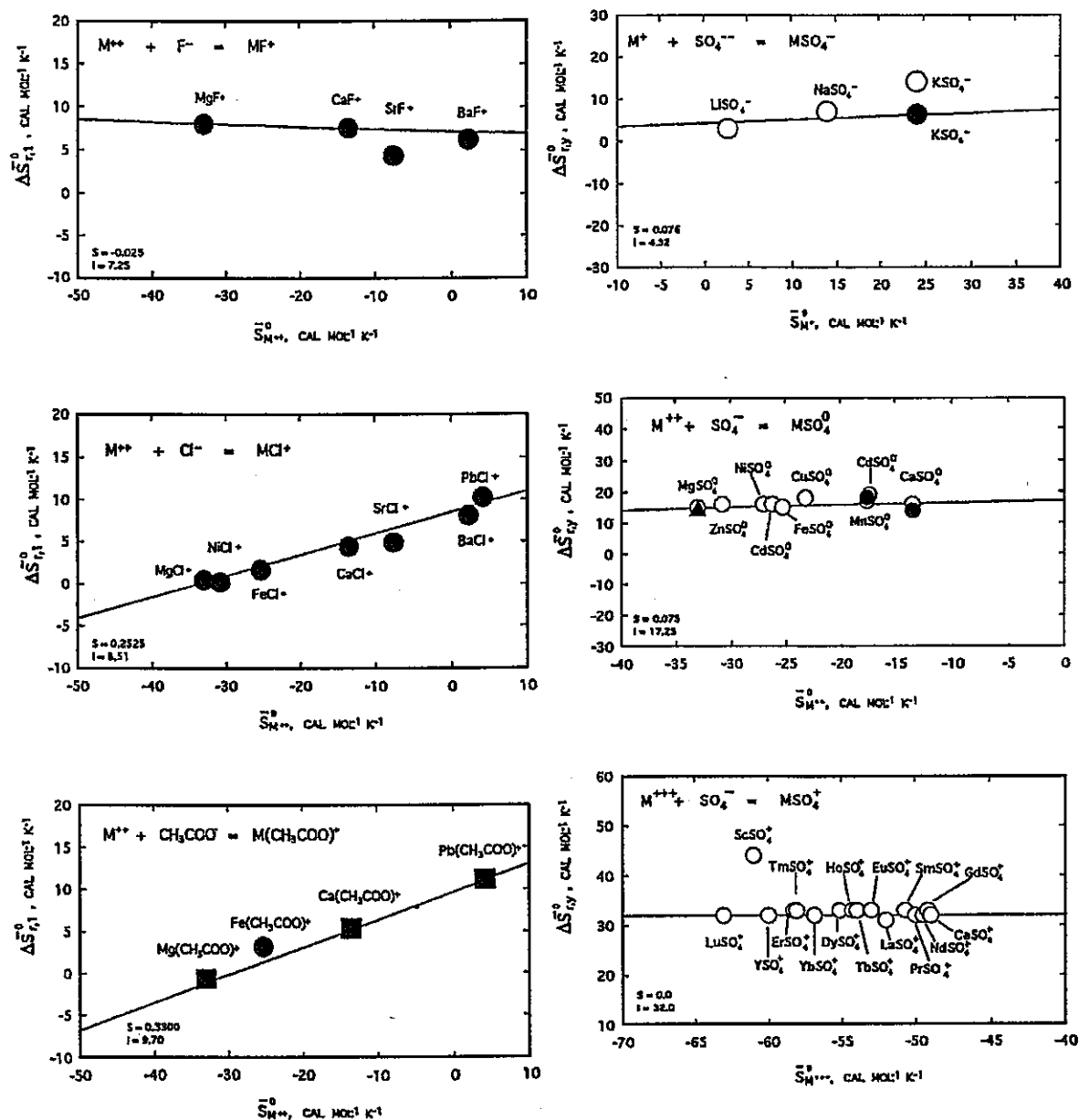


Figure 4.4.2_20. Correlation of the standard partial molal entropy of association for addition of a monovalent ligand to a cationic species with the standard partial molal entropy of the cationic species at 25°C and 1 bar (see Table 4.2.2_1, footnote #7). Symbols refer to calculated values from Sverjensky *et al.* (1997) based on thermodynamic data that are incorporated in SPRONS.JNC.



Figures 4.4.2_21 (three figures on the left) and 4.4.2_22 (three figures on the right). Correlation of the standard partial molal entropy of association for addition of a monovalent ligand to a divalent cation (Fig. 4.4.2_21), and for sulfate complexes of mono-, di- and trivalent cations (Fig. 4.4.2_22), with the standard partial molal entropies of the cations at 25°C and 1 bar. [see Table 4.2.2_1, footnotes #8 (Fig. 4.4.2_21) and #9 (Fig. 4.4.2_22)]. Symbols refer to calculated values from Sverjensky *et al.* (1997) based on thermodynamic data that are incorporated in SPRONS.JNC.

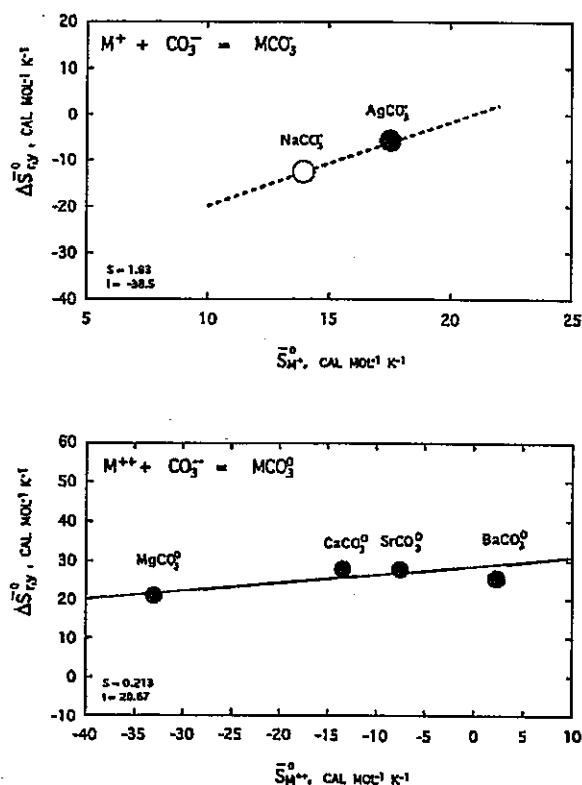


Figure 4.4.2_23. Correlation of the standard partial molal entropies of association of carbonate complexes involving mono- and divalent cations with the standard partial molal entropy of the cations at 25°C and 1 bar (see Table 4.2.2_1, footnote #10). Symbols refer to calculated values from Sverjensky *et al.* (1997) based on thermodynamic data that are incorporated in SPRONS.JNC.

respectively. Combined uncertainties in estimated equation-of-state parameters may lead to uncertainties in ΔG° as high as $\pm 700 \text{ cal mol}^{-1}$ at 500°C and 1 kb, or $\pm 1500 \text{ cal mol}^{-1}$ at 1000 °C and 5 kb (Shock and Helgeson, 1988).

5 Description of Thermodynamic Databases

The thermodynamic databases SPRONS.JNC and PHREEQE.JNC are described in this section. Section 5.1 provides an overview of the approach used to generate these databases. SPRONS.JNC is described in Section 5.2. PHREEQE.JNC is described in Section 5.3. The databases are listed in Appendices A (SPRONS.JNC) and B (PHREEQE.JNC).

5.1 Overview

The procedure to develop SPRONS.JNC and PHREEQE.JNC is illustrated schematically in Fig. 5.1_1. SPRONS92.DAT (Johnson *et al.*, 1991) is the basic source of thermodynamic data in SPRONS.JNC, which also includes new and revised data published in the peer-reviewed scientific literature since 1991. PHREEQE.JNC is derived from SPRONS.JNC using SUPCRT to calculate thermodynamic properties of relevant reactions. Remarks concerning these calculations, and associated software, are summarized in the following paragraphs. SUPCRT is described by Johnson *et al.* (1991), and is available from the Laboratory of Theoretical Geochemistry, University of California, Berkeley.

MPRONS is a pre-processor used to modify the SPRONS database. It is an interactive program that asks the user for additions, deletions or modifications to the database, executes the requested changes, and generates a new database that incorporates the new or revised data. In practice, any suitable text editor can also be used for this purpose, in which case it is essential that formatting and

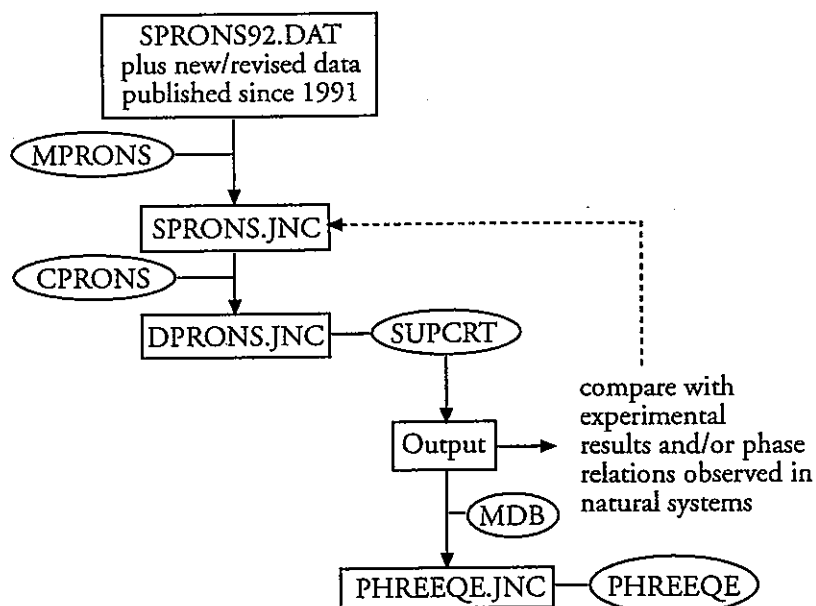


Figure 5.1_1. Schematic diagram of the procedure used to develop SPRONS.JNC and PHREEQE.JNC. Boxes refer to data or databases, ovals refer to software (see text).

unit conventions are strictly adhered to (Tables 5.1_1 to 5.1_3), and that the total numbers of minerals, gases, aqueous species and comment lines in the database are accurately updated (these totals are appended to the end of the database, as can be seen in Appendix A, for example).

CPRONS is an interactive pre-processor that converts the sequential access file SPRONS.JNC into its direct access equivalent, DPRONS.JNC. DPRONS.JNC increases the run-time efficiency of calculations carried out by SUPCRT.

SUPCRT consists of the following subprogram modules:

- *SUP* - an executive program that reads user-generated input files and the database file DPRONS.JNC, passes these data to the *REAC* module and interfaces calculated reaction properties with the reporting module *REP*,
- *REAC* - calculates the apparent standard molal Gibbs free energy and enthalpy, and the associated standard molal entropy, heat capacity and volume, of reactions among minerals, gases and aqueous species as a function of temperature and pressure,
- *H₂O* - calculates the electrostatic and standard molal thermodynamic properties of H₂O as a function of independent state conditions specified by the user, and
- *REP* - writes an output file, which lists all relevant data used in the calculations extracted from DPRONS.JNC, together with calculated standard molal thermodynamic properties of reactions (a "PLT" file, which contains calculated values in a format suitable for two-dimensional plotting, may also be written at the user's discretion).

Two input files are required for calculations using SUPCRT. One, the CON (conditions) file, specifies the range of independent state conditions to be considered in the calculations, including either:

- the liquid-phase side of the H₂O vaporization boundary, which requires user specification of the temperature, or pressure, range, or
- single-phase regions of H₂O stability, which requires user specification of temperatures and pressures, temperatures and H₂O densities, or specification of *T* and log *K*, or *P* and log *K*.

The temperature, pressure and solvent-density regions of validity of the SUPCRT model are summarized in the introductory paragraph of Section 4.

The other user-specified input file required by SUPCRT is the RXN (reactions) file. This file specifies reactions among minerals, gases and aqueous species whose standard molal thermodynamic properties are to be calculated by SUPCRT over

Line	Var	Var	Var	Var	Format
1	name	scform			1x, a20, a30
2	abbrev	ecform			1x, a15, 5x, a40
3	ref	date			1x, a5, 15x, a9
4	ΔG°_f	ΔH°_f	$S^\circ_{Pr, Tr}$	$V^\circ_{Pr, Tr}$	4x, 2(2x, f12.1), 2(2x, f8.3)
5	$a(10^0)$	$b(10^3)$	$c(10^{-5})$		4x, 3(2x, f12.6)
6	T_{max} (for a, b, c)				8x, f7.2

Table 5.1_1. Generic species data block in SPRONS.JNC for gases, and minerals that do not undergo phase transitions. Abbreviations: Var. (variable), Format (FORTRAN77 format), abbrev. (abbreviation), scform (structural chemical formula), ecform (elemental chemical formula), ref (literature reference - see Appendix A) and date (date of last revision). Units: ΔG°_f and ΔH°_f (cal mol⁻¹), $S^\circ_{Pr, Tr}$ (cal mol⁻¹ K⁻¹), $V^\circ_{Pr, Tr}$ (cm³ mol⁻¹), a (cal mol⁻¹ K⁻¹), b (cal mol⁻¹ K⁻²), c (cal K mol⁻¹), and T_{max} (K) [see Section 4.2.1].

Line	Var	Var	Var	Var	Format
1	name	scform			2x, a20, a30
2	abbrev	ecform			2x, a15, 5x, a40
3	ref	date			2x, a5, 15x, a9
4	ΔG°_f	ΔH°_f	$S^\circ_{Pr, Tr}$	$V^\circ_{Pr, Tr}$	4x, 2(2x, f12.1), 2(2x, f8.3)
5	$a_i(10^0)$	$b_i(10^3)$	$c_i(10^{-5})$	$T_{Ti, Pr}$	4x, 3(2x, f12.6), 2x, f7.2
	ΔH°_{Ti}	ΔV°_{Ti}	$(dP/dT)_{Ti}$		2x, f8.1, 2(2x, f10.3)
6	$a_n(10^0)$	$b_n(10^3)$	$c_n(10^{-5})$		4x, 3(2x, f12.6)
7	$T_{max, n}$ (for a_n, b_n, c_n)				8x, f7.2

Table 5.1_2. Generic species data block in SPRONS.JNC for minerals that undergo phase transitions. Abbreviations: as in Table 5.1_1. Units: as in Table 5.1_1, and $T_{Ti, Pr}$ (K), ΔH°_{Ti} (cal mol⁻¹), ΔV°_{Ti} (cm³ mol⁻¹), $(dP/dT)_{Ti}$ (bar K⁻¹), $a_{i, n}$ (cal mol⁻¹ K⁻¹), $b_{i, n}$ (cal mol⁻¹ K⁻²), $c_{i, n}$ (cal K mol⁻¹), and $T_{max, n}$ (K) [see Section 4.2.1].

Line	Var	Var	Var	Var	Format
1	name	scform			2x, a20, a30
2	abbrev	ecform			6x, a15, 5x, a40
3	ref	date			6x, a5, 15x, a9
4	ΔG°_f	ΔH°_f	$S^\circ_{Pr, Tr}$		4x, 2(2x, f10.0), 4x, f8.3
5	$a_1(10^1)$	$a_2(10^{-2})$	$a_3(10^0)$	$a_4(10^{-4})$	4x, 4(2x, f8.4, 2x)
6	$c_1(10^0)$	$c_2(10^{-4})$	$\omega_{Pr, Tr}(10^{-5})$	charge	4x, 3(2x, f8.4, 2x), 9x, f3.0

Table 5.1_3. Generic species data block in SPRONS.JNC for aqueous species. Abbreviations: as in Table 5.1_1. Units: as in Table 5.1_1, and a_1 (cal mol⁻¹ bar⁻¹), a_2 (cal mol⁻¹), a_3 (cal K mol⁻¹ bar⁻¹), a_4 (cal K mol⁻¹), c_1 (cal mol⁻¹ K⁻¹), c_2 (cal K mol⁻¹), $\omega_{Pr, Tr}$ (cal mol⁻¹) and charge (dimensionless) [see Section 4.2.2].

the range of conditions specified in the CON file. The reactions are specified by the user in terms of the appropriate reaction stoichiometry and corresponding names of mineral, gas and aqueous reactants and products.

The output from SUPCRT is processed by MDB (Modify DataBase; Arthur *et al.*, 1997) to generate revised versions of PHREEQE.JNC. The SUPCRT output file tabulates the standard molal thermodynamic properties of reactions at the reference temperature and pressure, and also over a range of T - P conditions specified in the CON file. The temperature dependence of the equilibrium constant for corresponding reactions in PHREEQE is calculated using either the van't Hoff equation or a polynomial expression constrained by empirical temperature-independent coefficients. In the former case, MDB extracts values of $\log K$ and the reaction enthalpy at 25°C from the SUPCRT output file, and incorporates these values into datablocks for the corresponding reaction in PHREEQE.JNC. In the latter case, MDB calculates values of the empirical coefficients by regression of $\log K$ *vs.* T values calculated by SUPCRT, and then writes these coefficients, together with $\log K$ at 25°C, into appropriate datablocks in PHREEQE.JNC. For each such modification, MDB appends a record to PHREEQE.JNC specifying data that were added or revised, as well as the date and identity of the person making the revisions.

The accuracy of the data in SPRONS.JNC is assessed by comparing calculated equilibrium constants for selected reactions with their experimental counterparts. A number of such assessments are documented in Section 6, where it can be seen that in some cases discrepancies exist between calculated and experimental results. Unfortunately, determining the reason for such discrepancies is not straightforward, because they may be generated by errors in the thermodynamic properties of one or more reactants or products, and/or by inappropriate experimental technique or mis-interpretation of experimental results. We expect that future revisions to SPRONS.JNC, as noted in Section 2, will partially or completely eliminate all such discrepancies (as suggested by the dashed line in Fig. 5.1_1), but this may require additional experimental work by JNC and others. Similarly, we note that despite efforts to avoid completely any errors of transcription from original data sources, some errors of this type may exist in SPRONS.JNC. Thus, if discrepancies are encountered between calculated and experimental results, it is recommended that the relevant data should first be checked to confirm that they have been correctly entered into the database.

5.2 SPRONS.JNC

5.2.1 Description

SPRONS.JNC includes standard molal Gibbs free energies and enthalpies of formation, and standard molal entropies and volumes at the reference pressure (1 bar) and temperature (25°C) for 143 minerals that do not undergo phase

transitions, 33 minerals that undergo one phase transition, 17 minerals that undergo two phase transitions, 2 minerals that undergo three phase transitions, and 16 gases. Also included are Maier-Kelly heat capacity coefficients (Eqn. 4.2.1_5) that permit calculation of corresponding apparent standard molal Gibbs free energies and enthalpies of formation, and standard molal entropies at pressures from 1 to 10000 bars and temperatures from 0°C to a maximum value consistent with the Maier-Kelly coefficients. SPRONS.JNC also includes standard molal Gibbs free energies and enthalpies of formation and standard molal entropies at the reference pressure and temperature, and HKF equation-of-state coefficients (a_1 ... a_4 , c_1 , c_2 , and ω ; Section 4.2.2) for 1147 aqueous species. Together with equations of state for fluid phases of H₂O and equations characterizing the dielectric properties of the solvent as functions of pressure and temperature (encoded in SUPCRT's subprogram module H_2O), apparent standard partial molal Gibbs free energies and enthalpies of formation, and standard partial molal entropies of these aqueous species can be calculated over a broad range of temperatures and pressures. The apparent standard molal thermodynamic properties of minerals, gases and aqueous species in SPRONS.JNC can be used with SUPCRT to calculate corresponding properties of reactions.

The data in SPRONS.JNC that are taken from SPRONS92.DAT are described in Section 5.2.2. Sources of new data, and of revisions to data in SPRONS92.DAT, are discussed in Sections 5.2.3 and 5.2.4.

5.2.2 Sources of unrevised data

Unless otherwise noted (Section 5.2.3 and 5.2.4), thermodynamic data in SPRONS.JNC are from references 1 - 19, Appendix A. The data for minerals are mostly from Helgeson *et al.* (1978), but additional data for Sn-bearing minerals (Jackson and Helgeson, 1985), and a small amount of calorimetric data (references 6, 7, 8, 15 and 16, Appendix A), are also included. Standard molal Gibbs free energies and enthalpies of formation are not available for several minerals, including:

- bromellite, chabazite, chloritoid, cummingtonite, dickite, ferropargasite, ferrotremolite, fluorophlogopite, fluortremolite, glaucophane, grunerite, leonhardite, richterite, staurolite, zincite, almandine, 14Å-amesite, edenite, epistilbite, ferroedenite, fluoredenite, heulandite, natrolite, Ca-, K- and Na-phillipsite, pyrope, spessartine, stilbite, aegerine, larnite, riebeckite, 7Å-cronstedtite, ferrogedrite, hastingsite, magnesiohastingsite, magnesio-riebeckite, and PD (proton-deficient)-oxyannite

(an entry in SPRONS.JNC of "999999" is the default null value assigned to the Gibbs free energy and enthalpy of formation when these data are unavailable). Calculated results involving these minerals therefore refer only to the parenthetical terms in Eqns. (4.1_1) and (4.1_2).

The standard Gibbs free energies of formation of calcite and aragonite retrieved by Helgeson *et al.* (1978) are corrected by Johnson *et al.* (1991) in order to be consistent with the experimental solubilities of these minerals determined by Plummer and Busenberg (1982). The Gibbs free energies and enthalpies of several other Ca-bearing minerals evaluated by Helgeson *et al.* (1978), including anorthite, clinozoisite, diopside, dolomite, epidote, laumontite, wairakite, wollastonite, lawsonite, and margarite, are adjusted to be consistent with the revised data for calcite and aragonite (see reference 19, Appendix A).

Thermodynamic data and equation-of-state coefficients for aqueous ions and ionic species are from Shock and Helgeson (1988). Similar data for inorganic neutral species are from Shock *et al.* (1989). Data for inorganic metal complexes attributed to unpublished work in SPRONS92.DAT (reference 5, Appendix A) have since been published, and in some cases revised (reference 22, Appendix A). These newer data supersede those from reference 5, and are incorporated in SPRONS.JNC (Section 5.2.3). Data for many of the 86 organic species that are included in SPRONS92.DAT (reference 4, Appendix A) are omitted in SPRONS.JNC because these species are not relevant to the Japanese repository concept, nor are they helpful in retrieving thermodynamic properties of other, more relevant, aqueous species from experimental data and field observations. Recent data for organic acids and metal-organic complexes are, however, included in SPRONS.JNC (see Section 5.2.3), because these species are known to be generated by diagenetic alteration of organic matter in sedimentary environments (e.g., Kharaka *et al.*, 1983; Barth and Bjorlykke, 1993). These species may therefore be relevant to the Japanese repository concept, and may also affect groundwater chemistry at field sites that are being investigated by JNC [e.g., the Mobarra site; Kamei (1997); Sasamoto *et al.* (1999c)]. Data for gases are from Wagman *et al.* (1982) and Kelley (1960).

It should perhaps be emphasized that data from the sources described above are themselves constrained by supplementary data from numerous other sources, as is generally true of all thermodynamic databases. The supplementary data for minerals and gases are documented in footnotes appended to Tables 2, 3, 7, 8 and 9 of Helgeson *et al.* (1978). These data include standard molal entropies of several elements from Wagman *et al.* (1968; 1969), and standard molal Gibbs free energies and enthalpies of formation, and standard molal entropies, derived from calorimetric studies.

Sources of supplementary data for aqueous ions and ionic species are referenced in footnotes to Tables 11 and 12 of Shock and Helgeson (1988). Standard partial molal volumes and heat capacities at the reference temperature and pressure are retrieved from experimental data on apparent molal properties of electrolytes (Section 4.3.3.1). Additional data are from Tanger and Helgeson (1988), who similarly retrieved these data from experimental measurements of apparent molal properties. Supplemental standard partial molal entropies, and

enthalpies and Gibbs free energies of formation at the reference temperature and pressure, are taken primarily from CODATA (1978) and Wagman *et al.* (1982), and also from Berg and Vanderzee (1978), Jackson and Helgeson (1985) and (for lanthanide cations) from Morss (1976).

Sources of supplementary data for inorganic neutral species are referenced in footnotes to Tables 4 and 7 of Shock *et al.* (1989). Supplemental standard partial molal volumes and heat capacities at the reference temperature and pressure are from Olofsson *et al.* (1985), Tiepel and Gubbins (1972), Allred and Woolley (1981), Barbero *et al.* (1983) and Pierotti (1965). Supplemental standard partial molal entropies, and enthalpies and Gibbs free energies of formation, are from Wagman *et al.* (1982) and Olofsson *et al.* (1985).

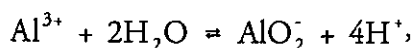
5.2.3 Sources of new and revised data

References 20 - 40 (Appendix A) document new and revised thermodynamic data included SPRONS.JNC. These references are briefly summarized below.

Ref. 20: Pokrovskii and Helgeson (1995) evaluate experimental solubilities reported in the literature for gibbsite, boehmite, diaspore and corundum to derive an internally consistent set of thermodynamic data for Al^{3+} , $\text{Al}(\text{OH})^{2+}$, $\text{Al}(\text{OH})_3(\text{aq})$ and $\text{Al}(\text{OH})_4^-$, and to retrieve standard molal Gibbs free energies of formation of boehmite and diaspore at the reference pressure and temperature. This is an important study because all available calorimetric, phase-equilibrium, solubility and potentiometric data pertinent to the $\text{Al}_2\text{O}_3\text{-H}_2\text{O-NaCl}$ system are considered over a wide range of pressures and temperatures, and because the retrieved data for this chemical subsystem are internally consistent. Standard molal thermodynamic properties and Maier-Kelley heat capacity coefficients for corundum, gibbsite, boehmite and diaspore from Table 2 of this reference are included in SPRONS.JNC. Standard partial molal thermodynamic properties and HKF equation-of-state coefficients for the aqueous aluminum species noted above, and for $\text{NaAlO}_2(\text{aq})$ and $\text{NaOH}(\text{aq})$, from Table 3 of this reference are also included in SPRONS.JNC. These internally consistent data are also consistent with data for Na^+ , Cl^- and OH^- from Shock and Helgeson (1988) and data for $\text{NaCl}(\text{aq})$ from Shock *et al.* (1989), and thus are consistent with the data for these species in SPRONS.JNC.

The data for Al^{3+} and $\text{Al}(\text{OH})_4^-$ (i.e., AlO_2^- , see Section 5.2.5) from Pokrovskii and Helgeson (1995) are accepted here rather than alternative data for these species recommended by Shock *et al.* (1997a). This is because the Pokrovskii and Helgeson (1995) data are consistent with the available solubility data for gibbsite and boehmite, and because they are internally consistent with corresponding data for the other aqueous aluminum species noted above. Shock *et al.* (1997a), on the other hand, assert that there is considerable disagreement among experimental solubilities reported for gibbsite and boehmite at low pH. To avoid this confusion, they use experimental data from Couturier *et al.* (1984)

for the hydrolysis reaction:



to retrieve the standard partial molal entropy of Al^{3+} . While we agree there is some scatter in the solubility data (Figs. C.11 and C.15, Appendix C), we do not believe this is sufficient reason to adopt this approach. The available solubility data provide important experimental constraints on the thermodynamic properties of Al^{3+} and $\text{Al}(\text{OH})_4^-$, and we feel it is better to account for these constraints, despite the scatter in the data, than to ignore them.

It is of interest to note that the thermodynamic properties determined by Pokrovskii and Helgeson (1995) and Shock *et al.* (1997a) for $\text{Al}(\text{OH})_4^-$ are virtually identical, but that the standard partial molal Gibbs free energy of formation of Al^{3+} reported by Shock *et al.* (1997a) is nearly 1 kcal mol⁻¹ more positive than that retrieved by Pokrovskii and Helgeson (1995). Similarly, the partial molal enthalpy of formation is nearly 2 kcal mol⁻¹ more positive, and the partial molal entropy is approximately 3 cal mol⁻¹ K⁻¹ more positive, than corresponding values retrieved by Pokrovskii and Helgeson (1995).

Ref. 21: Shock *et al.* (1997a) use correlations among experimentally determined standard partial molal volumes, heat capacities and entropies to estimate standard partial molal properties and HKF equation-of-state parameters for more than 300 inorganic cations and anions, polyatomic anions, oxyanions, acid oxyanions, neutral oxy-acid species, dissolved gases and hydroxide complexes of metal cations. In addition to these new data, revised data for 80 species from Shock and Helgeson (1988) are also provided. These revisions are motivated by improvements to previous correlations (Sassani and Shock, 1992; 1994). The revised correlations account for the coordination of aqueous ions, and associated effects on estimates of effective ionic radii and partial molal entropies (see also Puigdomenech *et al.*, 1997). The data summarized in Tables 4, 5 and 10 of this reference, and all the revisions to data from Shock and Helgeson (1988), are included in SPRONS. JNC.

Ref. 22: Sverjensky *et al.* (1997) regressed experimental ion-association data reported in the literature to retrieve standard partial molal properties and HKF equation-of-state coefficients for 57 aqueous metal complexes with halogen (Cl^- , F^- , Br^- , I^-), carbonate, sulfate, nitrate, acetate, and H_3SiO_4^- ligands. Correlations developed on the basis of experimental standard partial molal heat capacities, volumes and entropies of association, combined with similar data from Shock and Helgeson (1988) for aqueous cations and anionic ligands, are used to calculate provisional estimates of standard partial molal properties and equation-of-state coefficients for an additional 52 complexes, for which little or no experimental data are available. The data summarized in Tables 2, 3 and 12 of this reference are included in SPRONS. JNC.

Ref. 23: Pokrovskii and Helgeson (1997a) regressed experimental dissociation constants reported in the literature to retrieve standard partial molal properties and HKF equation-of-state parameters for $\text{KCl}(aq)$. The retrieval procedure used data from Shock and Helgeson (1988) for the corresponding properties of K^+ and Cl^- . The data for $\text{KCl}(aq)$ are therefore consistent with these data, and with corresponding data in SPRONS.JNC.

Ref. 24: Pokrovskii and Helgeson (1997b) regressed experimental data on the solubilities of gibbsite and corundum in aqueous KOH solutions to retrieve standard partial molal properties and HKF equation-of-state coefficients for $\text{KAlO}_2(aq)$ and $\text{KOH}(aq)$. These data are consistent with the thermodynamic properties of K^+ and OH^- in SPRONS.JNC retrieved by Shock and Helgeson (1988), and with the revised properties of AlO_2^- in SPRONS.JNC (Ref. 20) determined by Pokrovskii and Helgeson (1995).

Ref. 25: Haas *et al.* (1995) used experimental data from the literature and correlation algorithms to retrieve or estimate standard partial molal thermodynamic properties and HKF equation-of-state parameters for 246 inorganic aqueous lanthanide complexes with halogen (Cl^- , F^-), hydroxide, carbonate, sulfate, bicarbonate, nitrate and orthophosphate ligands. The retrieval and estimation procedures are constrained by data from Shock and Helgeson (1988), and reported values are therefore consistent with the data in SPRONS.JNC. The data are listed in the appendix of this reference, and were obtained from the authors (E. L. Shock, at <http://zonvark.wustl.edu/geopig/>).

Ref. 27: Shock and Koretsky (1995) regressed experimental equilibrium constants for association of metal-organic complexes involving monovalent organic acid ligands to retrieve, or estimate, standard partial molal thermodynamic properties and HKF equation-of-state coefficients for 270 metal-organic species. The organic ligands and associated acids that are represented in SPRONS.JNC include formate, acetate, propanoate, butanoate, pentanoate, glycolate, lactate, oxalate and succinate [the partial molal thermodynamic properties of these ligands and their corresponding acids are from Ref. 26 (Shock, 1995)]. Analogous data for 114 metal acetate complexes from Ref. 33 (Shock and Koretsky, 1993) are also included in SPRONS.JNC. All the data from Refs. 26, 27 and 33 are consistent with data for inorganic aqueous ions and ionic species in SPRONS.JNC from Shock and Helgeson (1988) and Shock *et al.* (1997a). The organic-species data are reported in various tables and appendices in Refs. 26, 27 and 33, and were obtained in the present study from the senior author (E. L. Shock, at <http://zonvark.wustl.edu/geopig/>).

The data for the organic acids and ligands noted above are included in SPRONS.JNC because these species are found commonly in the formation waters of sedimentary basins (e.g., Carothers and Kharaka, 1978; Kharaka *et al.*, 1993). These species may also affect mineral diagenesis, and the aqueous

speciation behavior of inorganic elements such as Ca, Al and Fe (Kharaka *et al.*, 1986; Lundegard and Kharaka, 1990; Harrison and Thyne, 1992).

Ref. 29: Saccocia and Seyfried (1993) reinterpreted published experimental phase-equilibrium data involving chamosite (synonymous according these authors with 14Å-daphnite) and retrieved values of the standard molal Gibbs free energy and enthalpy of formation for 14Å-daphnite that are consistent with the standard entropy, volume and Maier-Kelley heat capacity coefficients for this mineral evaluated by Helgeson *et al.* (1978). The recommended data are included in SPRONS.JNC.

Ref. 31: Sverjensky *et al.* (1991) reconciled inconsistencies in the thermodynamic properties of minerals in the systems K_2O - and $Na_2O-Al_2O_3-SiO_2-H_2O-HCl$ using an analysis procedure that integrates the results of phase-equilibrium and calorimetric experiments with the results of solubility experiments at high temperatures. Recommended revisions to values reported by Helgeson *et al.* (1978) for the Gibbs free energy of formation of muscovite, phlogopite, K-feldspar, sanidine and microcline are incorporated in SPRONS.JNC.

Standard partial molal thermodynamic properties and HKF equation-of-state coefficients for $HCl(aq)$ are also retrieved by Sverjensky *et al.* (1991). These data are not used in SPRONS.JNC, however, because newer data for this species (Tagirov *et al.*, 1997; see Ref. 40) are more consistent with all experimentally determined dissociation constants for $HCl(aq)$ over a wider range of temperatures, pressures and H_2O densities than were considered by Sverjensky *et al.* (1991). Calculated thermodynamic properties of this species nevertheless agree closely with values calculated using the data of Sverjensky *et al.* (1991).

Ref. 35: Oelkers *et al.* (1995) determined standard partial molal thermodynamic properties and HKF equation-of-state coefficients for two aluminum acetate complexes $[Al(CH_3COO)_2^+]$ and $[Al(CH_3COO)_2^+]$ using the data for Al^{3+} from Ref. 20 and the procedure described in Ref. 33. These data are accepted rather than data for these species from Ref. 33 because they are consistent with the data for Al^{3+} from Ref. 20.

Ref. 36: Shock *et al.* (1999) use experimental data from the literature and correlations to estimate partial molal thermodynamic properties and HKF equation-of-state coefficients for divalent, trivalent, quadravalent, pentavalent and hexavalent actinide ions. These data are included in SPRONS.JNC, and were obtained from the senior author (E. L. Shock, *pers. comm.*).

Ref. 37: Shock *et al.* (1997b) use the available experimental data from the literature to estimate HKF equation-of-state parameters for aqueous uranium cations and non-conventional hydroxide complexes (Section 5.2.5), and to retrieve standard molal thermodynamic properties and Maier-Kelley heat

capacity coefficients for uraninite that are consistent with thermodynamic properties reported by Grenthe *et al.* (1992) and experimental solubilities at high temperatures determined by Parks and Pohl (1988). The data summarized in Tables 1 and 2 of Shock *et al.* (1997b) are included in SPRONS.JNC.

Ref. 38: Standard partial molal thermodynamic properties and HKF equation-of-state coefficients for Pd^{2+} , PdCl^+ , $\text{PdCl}_2(\text{aq})$, PdCl_3^- , PdCl_4^{2-} , $\text{Pd}(\text{OH})^+$ and $\text{PdO}(\text{aq})$ are estimated by Sassani and Shock (1990) using the limited available experimental data on dissociation constants from Droll *et al.* (1957), Biryukov and Shlenskaya (1964) and Izatt *et al.* (1967). Standard molal thermodynamic properties and Maier-Kelley heat capacity coefficients for $\text{Pd}(\text{c})$ and $\text{PdO}(\text{c})$ are also reported. The data incorporated into SPRONS.JNC are from Tables 1 and 2 of this reference. These data are provisional estimates, which may be revised in whole or in part by Sassani and Shock (1999).

Ref. 39: Standard partial molal thermodynamic properties and HKF equation-of-state parameters for the ion-pairs $\text{NaB}(\text{OH})_4(\text{aq})$ and NaSO_4^- are retrieved by Pokrovski *et al.* (1995) from potentiometric measurements of the respective association constants between 50 and 200°C. The data from Table 5 of this reference are included in SPRONS.JNC, and are consistent with thermodynamic properties and HKF equation-of-state parameters in this database for Na^+ , OH^- and SO_4^{2-} from Shock and Helgeson (1988) and Shock *et al.* (1997a).

Ref. 40: Tagirov *et al.* (1997) retrieved standard partial molal thermodynamic properties and HKF equation-of-state coefficients for $\text{HCl}(\text{aq})$ based on association constants consistent with measured solubilities of $\text{AgCl}(\text{c})$ and $\text{Ag}(\text{c})$. The experimental data determined in this study are evaluated simultaneously with association constants retrieved from other solubility and electrical conductance measurements reported in the literature between 25 and 700°C, and between 1 bar and 5 kb. The retrieved data for $\text{HCl}(\text{aq})$ from Table 6 of Tagirov *et al.* (1997) are incorporated in SPRONS.JNC. Standard partial molal properties and HKF parameters for AgCl_2^- were also determined by Tagirov *et al.* (1997), and are included in SPRONS.JNC for completeness. The data for both $\text{HCl}(\text{aq})$ and AgCl_2^- are consistent with data in SPRONS.JNC because they were retrieved using thermodynamic properties and HKF equation-of-state coefficients for Cl^- from Shock and Helgeson (1988), which are unrevised by Shock *et al.* (1997a), and thermodynamic properties and Maier-Kelley coefficients for $\text{AlCl}(\text{c})$ (chlorargyrite) and $\text{Ag}(\text{c})$ from Helgeson *et al.* (1978).

5.2.4 Data for minerals not considered by Helgeson *et al.* (1978)

Thermodynamic data in SPRONS.JNC and PHREEQE.JNC support different models of hydrolysis reactions involving non-stoichiometric solid solutions. Helgeson *et al.* (1978) prefer a modeling approach that considers mixing of atoms among energetically equivalent sites in the crystalline lattice, as well as exchange of atoms between energetically distinct sites. A model consistent with

this approach is described by these authors and involves calculation of the activities of thermodynamic components of the solid solution from site-mixing approximations. Component activities are then correlated with the compositions of the minerals and used as descriptive variables in activity diagrams. SPRONS.JNC includes data for minerals that are stoichiometrically equivalent to the thermodynamic components of many solid solutions, including several that are important in relatively low-temperature groundwater systems. Aagaard and Helgeson (1983) and Giggenbach (1985) use similar data and the model described above to calculate equilibrium mineral-stability relations appropriate for such systems. Arthur *et al.* (1993) and Arthur (1994) use this approach to model equilibrium porewater compositions in engineered bentonite barriers.

The PHREEQE series of codes do not include solid-solution models in the general sense considered by Helgeson *et al.* (1978), although ion-exchange reactions can be simulated. Rather, solid-solution behavior is grossly approximated using a *surrogate* stoichiometric phase to represent the solid solution. The stoichiometry of the surrogate phase is defined such that it is within the range of compositions exhibited by the solid solution it represents, but the stoichiometry is not variable, as would otherwise be consistent with solid-solution behavior. The advantage of this approach is that the thermodynamic properties of the surrogate phase can be estimated using a variety of techniques, such as described by Tardy and Garrels (1974), Mattigod and Sposito (1978) and Chermak and Rimstidt (1989)¹¹.

For this reason, it is necessary to include thermodynamic data for such surrogate phases in SPRONS.JNC in order to calculate equilibrium constants and enthalpies using SUPCRT, which can then be incorporated into PHREEQE .JNC. Estimated thermodynamic properties for a limited number of such phases representing smectite (Na-, K-, Ca- and Mg- nontronites, beidellites and montmorillonites), saponite (greenalite and celadonite), chlorite (chamosite and 7Å-daphnite) and illite solid solutions are included in SPRONS.JNC.

These data are from Refs. 30, 32 and 34 (Appendix A). Most of the data are estimated by Wolery (1976), who used a re-calibrated and slightly modified version of the technique described by Tardy and Garrels (1974) to estimate Gibbs free energies of formation, and methods described by Helgeson *et al.* (1978) to estimate corresponding standard molal entropies and Maier-Kelly heat capacity coefficients. Gibbs free energies and enthalpies of formation of celadonite, greenalite, 7Å-chamosite and 7Å-daphnite are combined with

¹¹- In passing, it is of interest to note that a third modeling approach appropriate for solid-solution minerals is described by Bourcier (1985), and is incorporated in the EQ3/6 software package (Wolery *et al.*, 1990). This approach is similar to that described by Helgeson *et al.* (1978), but considers ideal mixing of end-member thermodynamic components of the solid-solution, rather than ideal mixing of atoms on energetically equivalent sites.

standard entropies, volumes and heat capacities retrieved by Helgeson *et al.* (1978). All the thermodynamic properties of illite and Na-, K-, Ca- and Mg-nonttronites, beidellites and montmorillonites are estimated (Wolery, 1976). It is reasonable to include these data in SPRONS.JNC because they are consistent with thermodynamic data for the other minerals retrieved by Helgeson *et al.* (1978)[T. Wolery, *pers. comm.*]. It is not possible to assess the accuracy of these data because experimental solubilities corresponding to the exact stoichiometry of the surrogate phases are not available. It is possible, however, to assess whether the data are reasonably consistent with mineral stability relations observed in natural systems. A few such assessments are described in Section 6.11 for stoichiometric surrogates of smectite and illite solid solutions.

Halloysite is metastable with respect to its isostructural and isochemical analog kaolinite, and may persist for significant periods of time in relatively low temperature geologic systems. The standard molal entropy and volume of halloysite, and associated Maier-Kelley heat capacity coefficients, are reported by Helgeson *et al.* (1978), and are included in SPRONS.JNC, but its Gibbs free energy and enthalpy of formation are not determined by these authors. Because halloysite may be important in partially controlling groundwater compositions at some of JNC's field sites (Kamaishi, Tono and Mobarra), we calculated provisional estimates of its Gibbs free energy and enthalpy of formation (Ref. 28). This was accomplished by calculating the difference between calorimetric Gibbs free energies of formation of kaolinite and halloysite retrieved by Robie *et al.* (1978), and adding this difference to the Gibbs free energy of formation of kaolinite determined by Helgeson *et al.* (1978). A similar procedure was used to estimate hallosyite's enthalpy of formation. These are considered to be provisional estimates, but it is cautioned that the validity of this estimation procedure may not be supported by recent experiments on the relative stabilities of kaolinite and another of its isochemical analogs, dickite (Zotov *et al.*, 1998).

5.2.5 Conventional and non-conventional aqueous species

The non-conventional form of hydroxide complexes and other aqueous species are adopted in SPRONS.JNC, but conventional forms are represented in PHREEQE.JNC. The thermodynamic properties of conventional species are obtained by adding H₂O to the stoichiometry of their non-conventional counterparts, assuming the thermodynamic properties of the corresponding reactions, including the logarithm of the equilibrium constant, are equal to zero at any temperature and pressure (Wagman *et al.*, 1982). Such reactions are summarized in Table 5.2.5_1 for species that could be important in geologic systems. It is essential that reactions in PHREEQE.JNC that include conventional hydroxide complexes and other conventional aqueous species are written in terms of their non-conventional counterparts before thermodynamic properties of the reaction are calculated using SUPCRT and SPRONS.JNC.

Table 5.2.5_1. Selected conventional and non-conventional aqueous species in PHREEQE.JNC and SPRONS.JNC, respectively, and reactions for converting non-conventional species into their conventional counterparts (log K for all such reactions equals zero at any temperature and pressure).

Conventional Species (PHREEQE.JNC)	Non-conventional Species (SPRONS.JNC)	Conversion Reaction
$\text{Al}(\text{OH})_2^+$	AlO^+	$\text{AlO}^+ + \text{H}_2\text{O} = \text{Al}(\text{OH})_2^+$
$\text{Al}(\text{OH})_3(\text{aq})$	$\text{HAlO}_2(\text{aq})$	$\text{HAlO}_2(\text{aq}) + \text{H}_2\text{O} = \text{Al}(\text{OH})_3(\text{aq})$
$\text{Al}(\text{OH})_4^-$	AlO_2^-	$\text{AlO}_2^- + 2\text{H}_2\text{O} = \text{Al}(\text{OH})_4^-$
$\text{KAl}(\text{OH})_4(\text{aq})$	$\text{KAlO}_2(\text{aq})$	$\text{KAlO}_2(\text{aq}) + 2\text{H}_2\text{O} = \text{KAl}(\text{OH})_4(\text{aq})$
$\text{NaAl}(\text{OH})_4(\text{aq})$	$\text{NaAlO}_2(\text{aq})$	$\text{NaAlO}_2(\text{aq}) + 2\text{H}_2\text{O} = \text{NaAl}(\text{OH})_4(\text{aq})$
$\text{Bi}(\text{OH})_2^+$	BiO^+	$\text{BiO}^+ + \text{H}_2\text{O} = \text{Bi}(\text{OH})_2^+$
$\text{Bi}(\text{OH})_3(\text{aq})$	$\text{HBiO}_2(\text{aq})$	$\text{HBiO}_2(\text{aq}) + \text{H}_2\text{O} = \text{Bi}(\text{OH})_3(\text{aq})$
$\text{Bi}(\text{OH})_4^-$	BiO_2^-	$\text{BiO}_2^- + 2\text{H}_2\text{O} = \text{Bi}(\text{OH})_4^-$
$\text{Cd}(\text{OH})_2(\text{aq})$	$\text{CdO}(\text{aq})$	$\text{CdO}(\text{aq}) + \text{H}_2\text{O} = \text{Cd}(\text{OH})_2(\text{aq})$
$\text{Cd}(\text{OH})_3^-$	HCdO_2^-	$\text{HCdO}_2^- + \text{H}_2\text{O} = \text{Cd}(\text{OH})_3^-$
$\text{Cd}(\text{OH})_4^{2-}$	CdO_2^{2-}	$\text{CdO}_2^{2-} + 2\text{H}_2\text{O} = \text{Cd}(\text{OH})_4^{2-}$
$\text{H}_2\text{CO}_3(\text{aq})$	$\text{CO}_2(\text{aq})$	$\text{CO}_2(\text{aq}) + \text{H}_2\text{O} = \text{H}_2\text{CO}_3(\text{aq})$
$\text{Ca}(\text{H}_3\text{SiO}_4)^+$	$\text{Ca}(\text{HSiO}_3)^+$	$\text{Ca}(\text{HSiO}_3)^+ + \text{H}_2\text{O} = \text{Ca}(\text{H}_3\text{SiO}_4)^+$
$\text{Fe}(\text{OH})_2(\text{aq})$	$\text{FeO}(\text{aq})$	$\text{FeO}(\text{aq}) + \text{H}_2\text{O} = \text{Fe}(\text{OH})_2(\text{aq})$
$\text{Fe}(\text{OH})_3^-$	HFeO_2^-	$\text{HFeO}_2^- + \text{H}_2\text{O} = \text{Fe}(\text{OH})_3^-$
$\text{Fe}(\text{OH})_2^+$	FeO^+	$\text{FeO}^+ + \text{H}_2\text{O} = \text{Fe}(\text{OH})_2^+$
$\text{Fe}(\text{OH})_3(\text{aq})$	$\text{HFeO}_2(\text{aq})$	$\text{HFeO}_2(\text{aq}) + \text{H}_2\text{O} = \text{Fe}(\text{OH})_3(\text{aq})$
$\text{Fe}(\text{OH})_4^-$	FeO_2^-	$\text{FeO}_2^- + 2\text{H}_2\text{O} = \text{Fe}(\text{OH})_4^-$
$\text{Mg}(\text{H}_3\text{SiO}_4)^+$	$\text{Mg}(\text{HSiO}_3)^+$	$\text{Mg}(\text{HSiO}_3)^+ + \text{H}_2\text{O} = \text{Mg}(\text{H}_3\text{SiO}_4)^+$
$\text{Mn}(\text{OH})_2(\text{aq})$	$\text{MnO}(\text{aq})$	$\text{MnO}(\text{aq}) + \text{H}_2\text{O} = \text{Mn}(\text{OH})_2(\text{aq})$
$\text{Mn}(\text{OH})_3^-$	HMnO_2^-	$\text{HMnO}_2^- + \text{H}_2\text{O} = \text{Mn}(\text{OH})_3^-$
$\text{Mn}(\text{OH})_4^{2-}$	MnO_2^{2-}	$\text{MnO}_2^{2-} + 2\text{H}_2\text{O} = \text{Mn}(\text{OH})_4^{2-}$
$\text{Na}(\text{H}_3\text{SiO}_4)(\text{aq})$	$\text{Na}(\text{HSiO}_3)(\text{aq})$	$\text{Na}(\text{HSiO}_3)(\text{aq}) + \text{H}_2\text{O} = \text{Na}(\text{H}_3\text{SiO}_4)(\text{aq})$
$\text{Ni}(\text{OH})_2(\text{aq})$	$\text{NiO}(\text{aq})$	$\text{NiO}(\text{aq}) + \text{H}_2\text{O} = \text{Ni}(\text{OH})_2(\text{aq})$
$\text{Ni}(\text{OH})_3^-$	HNiO_2^-	$\text{HNiO}_2^- + \text{H}_2\text{O} = \text{Ni}(\text{OH})_3^-$
$\text{Ni}(\text{OH})_4^{2-}$	NiO_2^{2-}	$\text{NiO}_2^{2-} + 2\text{H}_2\text{O} = \text{Ni}(\text{OH})_4^{2-}$
$\text{H}_4\text{SiO}_4(\text{aq})$	$\text{SiO}_2(\text{aq})$	$\text{SiO}_2(\text{aq}) + 2\text{H}_2\text{O} = \text{H}_4\text{SiO}_4(\text{aq})$
H_3SiO_4^-	HSiO_3^-	$\text{HSiO}_3^- + \text{H}_2\text{O} = \text{H}_3\text{SiO}_4^-$
$\text{Pb}(\text{OH})_2(\text{aq})$	$\text{PbO}(\text{aq})$	$\text{PbO}(\text{aq}) + \text{H}_2\text{O} = \text{Pb}(\text{OH})_2(\text{aq})$
$\text{Pb}(\text{OH})_3^-$	HPbO_2^-	$\text{HPbO}_2^- + \text{H}_2\text{O} = \text{Pb}(\text{OH})_3^-$
$\text{Sn}(\text{OH})_2(\text{aq})$	$\text{SnO}(\text{aq})$	$\text{SnO}(\text{aq}) + \text{H}_2\text{O} = \text{Sn}(\text{OH})_2(\text{aq})$
$\text{Sn}(\text{OH})_3^-$	HSnO_2^-	$\text{HSnO}_2^- + \text{H}_2\text{O} = \text{Sn}(\text{OH})_3^-$
$\text{U}(\text{OH})_2^{2+}$	UO_2^+	$\text{UO}_2^+ + \text{H}_2\text{O} = \text{U}(\text{OH})_2^{2+}$
$\text{U}(\text{OH})_3^+$	H_2UO_2^+	$\text{H}_2\text{UO}_2^+ + \text{H}_2\text{O} = \text{U}(\text{OH})_3^+$
$\text{U}(\text{OH})_4(\text{aq})$	$\text{UO}_2(\text{aq})$	$\text{UO}_2(\text{aq}) + 2\text{H}_2\text{O} = \text{U}(\text{OH})_4(\text{aq})$
$\text{U}(\text{OH})_5^-$	H_2UO_3^-	$\text{H}_2\text{UO}_3^- + 2\text{H}_2\text{O} = \text{U}(\text{OH})_5^-$
$\text{UO}_2(\text{OH})_2(\text{aq})$	$\text{UO}_3(\text{aq})$	$\text{UO}_3(\text{aq}) + \text{H}_2\text{O} = \text{UO}_2(\text{OH})_2(\text{aq})$

Conventional Species (PHREEQE.JNC)	Non-conventional Species (SPRONS.JNC)	Conversion Reaction
$\text{UO}_2(\text{OH})_3^-$	HUO_4^-	$\text{HUO}_4^- + \text{H}_2\text{O} = \text{UO}_2(\text{OH})_3^-$
$\text{UO}_2(\text{OH})_4^{2-}$	UO_4^{2-}	$\text{UO}_4^{2-} + 2\text{H}_2\text{O} = \text{UO}_2(\text{OH})_4^{2-}$
$\text{Zn}(\text{OH})_2(\text{aq})$	$\text{ZnO}(\text{aq})$	$\text{ZnO}(\text{aq}) + \text{H}_2\text{O} = \text{Zn}(\text{OH})_2(\text{aq})$
$\text{Zn}(\text{OH})_3^-$	HZnO_2^-	$\text{HZnO}_2^- + \text{H}_2\text{O} = \text{Zn}(\text{OH})_3^-$
$\text{Zn}(\text{OH})_4^{2-}$	ZnO_2^{2-}	$\text{ZnO}_2^{2-} + 2\text{H}_2\text{O} = \text{Zn}(\text{OH})_4^{2-}$

5.2.6 Consistency between JNC's core thermodynamic database and SPRONS.JNC

As noted in Section 1, JNC is developing in parallel to the present study an additional thermodynamic database appropriate for minerals, gases and aqueous species of radioactive elements. Although this database and SPRONS.JNC are developed independently, there are a number of minerals, gases and aqueous species that are common to both databases. These primary, or "core", thermodynamic data in JNC's database for radioelements are from the OECD/NEA Thermodynamic Database Project (Silva *et al.*, 1995), which contains primarily CODATA key values, and additional data recommended by Grenthe *et al.* (1992). In this section we compare these data with their counterparts in SPRONS.JNC to determine if the data that are common to both databases are self-consistent, i.e., if they are identical within stated uncertainty limits.

Although uncertainties are assigned to the core data [at the statistically defined 95% confidence interval (Silva *et al.*, 1995)], absolute uncertainties in the SPRONS.JNC data cannot be assessed unequivocally. This is because there are multiple sources of uncertainty and ambiguity in interpretive methods used to retrieve these data from experimental results, and additional limitations associated with simultaneous evaluation of thermodynamic properties for multiple reactions (Section 4.3.4.3; Helgeson *et al.*, 1978). For these reasons it is considered preferable not to assign absolute uncertainties to the SPRONS.JNC data, rather than to estimate uncertainties that may be both inaccurate and misleading. We therefore consider the data in these two databases to be conservatively self consistent if reported values in SPRONS.JNC lie within the range of uncertainties assigned to data in the core database, but we do not explicitly account for uncertainties in the SPRONS.JNC data.

Minerals that are common to both SPRONS.JNC and the JNC core database include Cu(c), C(c)(graphite), corundum, lime, periclase, and α -quartz. The standard molal entropy of lime in SPRONS.JNC ($9.5 \text{ cal mol}^{-1} \text{ K}^{-1}$) differs from the corresponding nominal value and associated uncertainty limits in the core database ($9.106 \pm 0.096 \text{ cal mol}^{-1} \text{ K}^{-1}$). The data for minerals that are common to these databases are otherwise self-consistent

Discrepancies between SPRONS.JNC and the core database in ΔG_f° , ΔH_f° and S° for aqueous species and gases are shown in Figs. 5.2.6_1 - 5.2.6_3, respectively. The ordinate in these figures refers to the respective value in SPRONS.JNC minus the corresponding nominal value in the core database. When this difference exceeds uncertainty limits assigned to the nominal value in the core database (indicated by the error bars in Figs. 5.2.6_1 - 5.2.6_3), we take this to mean that the respective data in SPRONS.JNC and the core database are discrepant, in the conservative sense noted above. As can be seen in Fig. 5.2.6_1, the discrepancies defined in this manner in Gibbs free energies of formation are less than 1 kcal mol⁻¹, except for H₃PO₄(aq), H₂PO₄⁻, HPO₄²⁻, PO₄³⁻ and HF₂⁻. Discrepancies in enthalpies of formation, as shown in Fig. 5.2.6_2, are also less than 1 kcal mol⁻¹, except for the phosphate species, HF₂⁻ and SiO₂(aq). Discrepancies in entropies (Fig. 5.2.6_3) are less than 1 cal mol⁻¹ K⁻¹, except for HF(aq) and HSO₄⁻.

The significance of these apparent discrepancies is difficult to assess, however. It is possible that even crude estimates of absolute uncertainties in SPRONS.JNC, although ill-advised for reasons noted above, would eliminate most of the discrepancies, but probably not all [e.g., ΔG_f° and ΔH_f° for the phosphate species and ΔH_f° for SiO₂(aq)]. It is also possible to incorporate the core data into SPRONS.JNC, but this would entail revisions to the thermodynamic properties of other minerals, gases and aqueous species that are directly or indirectly

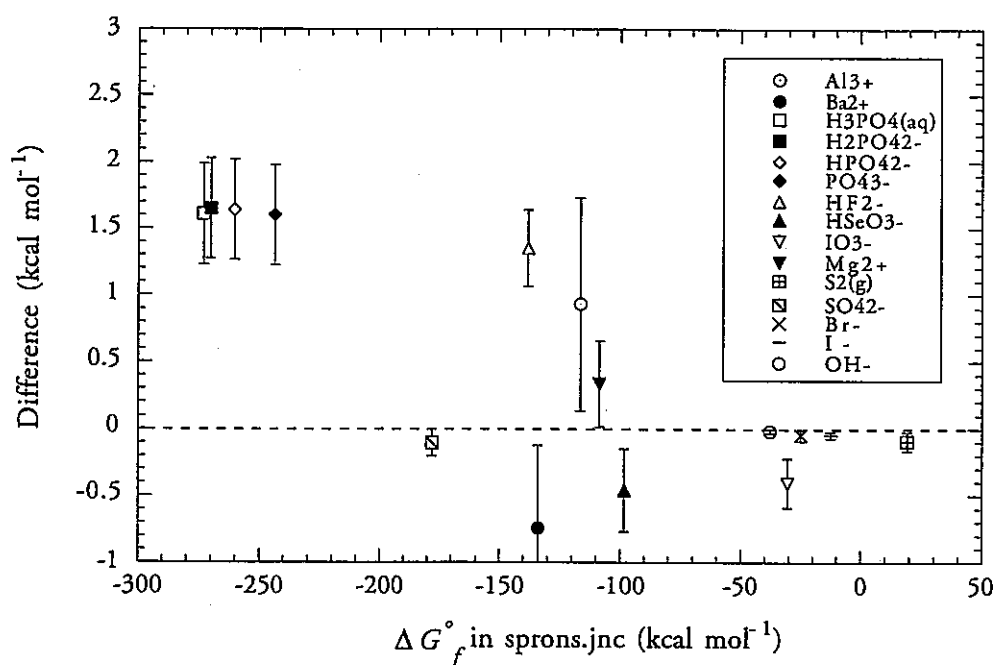


Figure 5.2.6_1. Differences between the standard Gibbs free energies of formation of aqueous species and S₂(g) in SPRONS.JNC and JNC's core thermodynamic database at 25°C and 1 bar versus ΔG_f° for each species or gas in SPRONS.JNC. Discrepancies in the data for these species are possible because calculated differences exceed uncertainty limits in the core database (indicated by error bars).

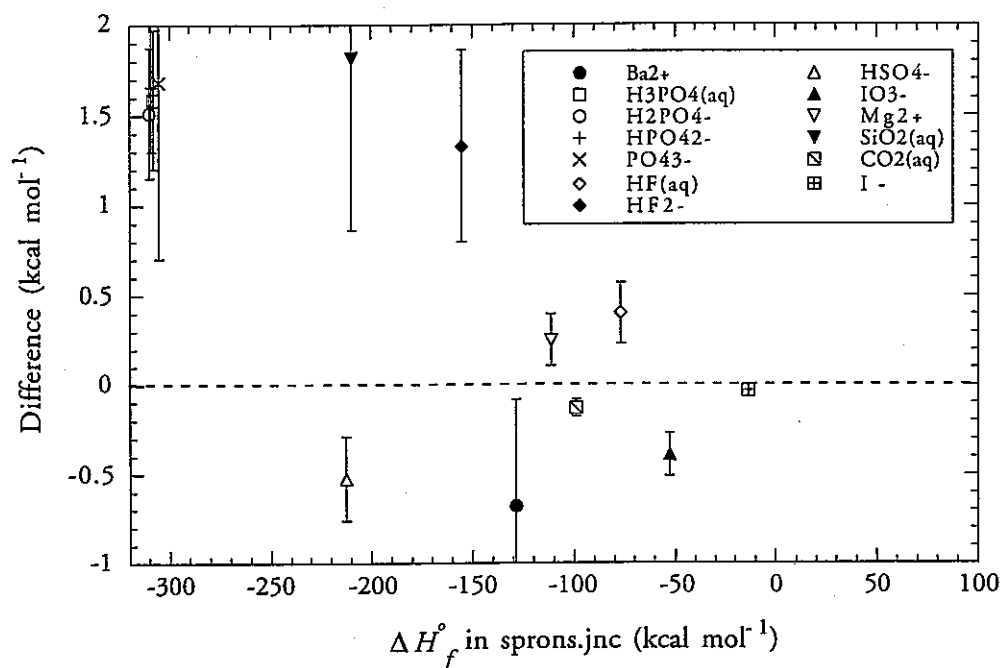


Figure 5.2.6_2. Differences between the standard partial molal enthalpies of formation of aqueous species in SPRONS.JNC and JNC's core thermodynamic database at 25°C and 1 bar versus ΔH_f° for each species in SPRONS.JNC. Discrepancies in the data for these solutes are possible because calculated differences exceed uncertainty limits in the core database (indicated by error bars).

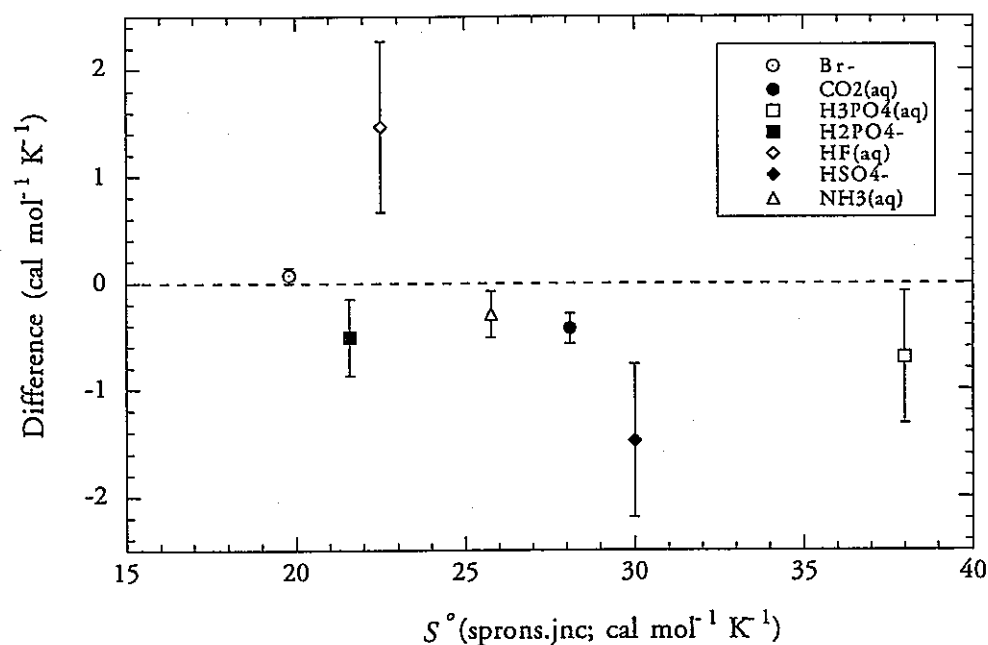


Figure 5.2.6_3. Differences between the standard partial molal entropies of aqueous species in SPRONS.JNC and JNC's core thermodynamic database at 25°C and 1 bar versus S° for each species in SPRONS.JNC. Discrepancies in the data for these aqueous ions and neutral species are possible because calculated differences exceed uncertainty limits in the core database (indicated by error bars).

constrained by the substituted data. This is certain to require a considerable amount of time that cannot be accommodated in the present study. Thus, although it is important to resolve the possible discrepancies noted above, such work must be deferred until future revisions to SPRONS.JNC are undertaken.

5.3 PHREEQE.JNC

5.3.1 Description

PHREEQE.JNC supports geochemical models of groundwater chemistry and evolution, laboratory investigations of water-rock interaction, and safety assessments of the Japanese repository concept consistent with JNC's reference scenario (Section 2). Relevant temperatures for these investigations are between 0 and 100°C. Corresponding pressures are in the range of 1 to 100 bars. Because such low pressures have negligible effects on most heterogeneous and homogeneous equilibria, however, a total pressure of 1 bar is considered adequate for these studies.

The PHREEQE.JNC database is listed in Appendix B. Format and unit conventions are specified in Tables 5.3.1_1 to 5.3.1_3. Thermodynamic data include equilibrium constants for mineral-dissolution and aqueous-association reactions at the reference pressure (1 bar) and reference temperature (25°C). As noted in Section 5.1, the temperature dependence of the equilibrium constant is calculated in the PHREEQE series of codes using either the van't Hoff equation, or empirical coefficients in an analytical expression representing $\log K(T)$. In the former case, MDB (Section 5.1) extracts values of $\log K$ and the reaction enthalpy at 25°C from the corresponding SUPCRT output file, and incorporates these values into datablocks for the reaction in PHREEQE.JNC. In the latter case, MDB calculates values of the empirical coefficients by regression of $\log K$ vs. T values calculated by SUPCRT, and then writes these coefficients, together with $\log K$ at 25°C, into appropriate datablocks in PHREEQE.JNC. Both approaches require the reaction enthalpy at the reference pressure and temperature, and these data are therefore included in PHREEQE.JNC.

All equilibrium constants and reaction enthalpies in PHREEQE.JNC are calculated using SUPCRT and SPRONS.JNC following the procedure described in Section 5.1 and illustrated in Fig. 5.1_1. This ensures that thermodynamic data in PHREEQE.JNC are in all cases consistent with the data in SPRONS.JNC. The sources of these data are documented in Sections 5.2.2 and 5.2.3.

PHREEQE.JNC includes data for minerals, gases and aqueous species that are relevant to geologic systems, and thus which contain the *geoelements* O, H, Na, Ca, Mg, K, P, F, Cl, Fe, S, C, N, Si, Al, B, Mn and I. Revisions to this database are planned, and may include data for minerals, gases and aqueous species of *radioelements*, and similar data for other elements (e.g., lanthanides and organic

Line	Var	Var	Var	Format
1	tname	nelt	tgfw	a8, 2x, i2, 3x, 3 f10.0

Table 5.3.1_1. Generic data block for elements in PHREEQE.JNC. Abbreviations: Var. (variable), tname (alphanumeric name of the element), nelt (index number of the element: number must be between 4 and 99), tgfw (gram formula weight).

Line	Var	Var	Var	Var	Var	Var	Var	Var	Var	Var	Format
1	I										i4
2	sname	nsp	kflag	gflag	zsp	thsp	dha	adhsp1	adhsp2	alksp	a8, 2x, i3, 2i1, 6f10.3
3	lktosp	dhsp	asp1	asp2	asp3	asp4	asp5				2f10.3, 5e12.5
4	lsp(I)	csp(I)									6(i3, f 7.3)

Table 5.3.1_2. Generic data block for aqueous association reactions in PHREEQE.JNC. Abbreviations: Var. (variable), I (index number assigned to the aqueous species formed in the reaction; number must be between 4 and 99 for master species, or between 100 and 1500 for other species), sname (alphanumeric species name), nsp (number of master species in the reaction), kflag [= 0 - van't Hoff equation is used to calculate $\log K(T)$; =1 analytical expression is used to calculate $\log K(T)$, see below], gflag (activity coefficient flag, see Parkhurst *et al.* (1980)), zsp (species charge), thsp [sum of operational valences of redox species involved in the reaction, see Parkhurst *et al.* (1980)], dha (the extended Debye-Hückel a°_i term), adhsp1 [the a_i term used in the WATEQ Debye-Hückel equation, see Parkhurst *et al.* (1980)], adhsp2 [the b_i term used in the WATEQ Debye-Hückel equation, see Parkhurst *et al.* (1980)], alksp (species alkalinity), lktosp (logarithm of the association constant), dhsp [enthalpy of association (kcal mol⁻¹)], asp1-5 [temperature independent coefficients used in the analytical expression $\log K(T) = \text{asp1} + \text{asp2}T + \text{asp3}/T + \text{asp4} \log T + \text{asp5}/T^2$, for $T(^{\circ}\text{K})$], lsp(I) (index number of master species in the reaction), csp(I) (stoichiometric coefficient of master species I in the reaction (positive for reactants, negative for products of the association reaction)).

Line	Var	Var	Var	Var	Var	Var	Var	Format
1	mname	nmin0	thmin	lktom	dhmin	mflag	simin	a8, 2x, i2, 3x, 3 f10.2, 5x, i1, 9x, f10.3
2	lmin(I)	cmin(I)						5(i4, f11.3)
3	amin1	amin2	amin3	amin4	amin5			5e12.5

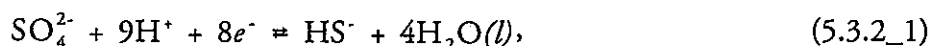
Table 5.3.1_3. Generic data block for mineral dissociation reactions in PHREEQE.JNC. Abbreviations: Var. (variable), mname (alphanumeric name of the mineral), nmin0 (number of aqueous species in the reaction), thmin [sum of the operational valences of redox species in the reaction, see Parkhurst *et al.* (1980)], lktom (logarithm of the equilibrium constant at 25°C), dhmin (reaction enthalpy, kcal mol⁻¹), mflag [= 0 - van't Hoff equation is used to calculate $\log K(T)$; =1 analytical expression is used to calculate $\log K(T)$, see below], simin [saturation index desired in the final solution (simin = 0 corresponds to equilibrium)], lmin(I) [index number of an aqueous species (I) in the reaction], cmin(I) [stoichiometric coefficient of an aqueous species (I) in the reaction], amin1-5 [temperature independent coefficients used in the analytical expression $\log K(T) = \text{amin1} + \text{amin2}T + \text{amin3}/T + \text{amin4} \log T + \text{amin5}/T^2$, for $T(^{\circ}\text{K})$].

aqueous species). To accommodate these future revisions, PHREEQE.JNC is dimensioned to include data for 1500 aqueous species. The first 99 index numbers are reserved for basis species. The first non-basis species (OH^-) is assigned the index number 100.

Mineral names in PHREEQE.JNC are in some cases shortened using the corresponding abbreviations in SPRONS.JNC (Appendix A). This is necessary because PHREEQE requires mineral names to be less than 9 characters in length. The abbreviations for some minerals are cryptic, including: CaAlpyrx (Ca-Al-pyroxene); Chm-7A (7Å chamosite); Cnc-7A (7Å clinocllore); Cnc-14A (14Å clinocllore); Czo (clinozoisite) a-Crs (α -cristobalite); b-Crs (β -cristobalite); K-fs (K-feldspar); Grn (greenalite); Hd (hedenbergite); Lmt (laumontite); Mng (manganosite); Max-mc (maximum microcline); Mnn (minnesotaite); Mtc (monticellite); Nesque (nesquehonite) and Hi-sa (high sanidine).

5.3.2 Redox reactions

Because the thermodynamic properties of the aqueous electron and H^+ are taken by convention to be equal to zero at any temperature and pressure (Section 4.1), equilibrium constants and reaction enthalpies for heterogeneous and homogeneous redox reactions in PHREEQE.JNC can be directly calculated using SUPCRT and SPRONS.JNC. For example, the conventional standard molal Gibbs free energy of the reaction:



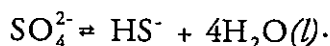
given by,

$$\Delta G_r^\circ = \left(\Delta \bar{G}_{f,\text{HS}^-}^\circ + 4\Delta \bar{G}_{f,\text{H}_2\text{O}}^\circ \right) - \left(\Delta \bar{G}_{f,\text{SO}_4^{2-}}^\circ + 9\Delta \bar{G}_{f,\text{H}^+}^\circ + 8\Delta \bar{G}_{f,e}^\circ \right),$$

is equivalent to

$$\Delta G_r^\circ = \left(\Delta \bar{G}_{f,\text{HS}^-}^\circ + 4\Delta \bar{G}_{f,\text{H}_2\text{O}}^\circ \right) - \Delta \bar{G}_{f,\text{SO}_4^{2-}}^\circ, \quad (5.3.2_2)$$

because the conventional Gibbs free energies of formation of H^+ and e^- are equal to zero. The quantities on the right-hand side of Eqn. (5.3.2_2) are evaluated using SUPCRT by specifying the following *fictive* reaction in the RXN file (see Section 5.1).



This "reaction" is chemically unrealistic because mass and charge are not conserved, but SUPCRT nevertheless calculates the standard Gibbs free energy change according to Eqn. (5.3.2_2). The result is therefore consistent with the

conventional thermodynamic properties of H^+ and e^- , and hence with the conventional thermodynamic properties of reaction (5.3.2_1). Equilibrium constants and enthalpies for all redox reactions in PHREEQE.JNC are calculated in this manner.

5.3.3 Comparison of thermodynamic data in PHREEQE.JNC and the H3 TDB

JNC's Heisei 3 (H3) thermodynamic database (TDB) is compared in this section with PHREEQE.JNC. The purpose of the comparison is to establish some continuity with regard to previous work carried out by JNC that includes thermodynamic calculations using the H3 database. That is, we assess the level of agreement among data that are common to both databases, and also identify minerals, gases and aqueous species that are included in one database, but not the other. The H3 database is described by Yui *et al.* (1992b). It includes data from several sources in a format that is compatible with PHREEQE software.

Equilibrium constants and enthalpies of aqueous association reactions in the H3 TDB and PHREEQE.JNC are shown in Table 5.3.3_1. The table includes all aqueous ions and complexes that are relevant to geologic systems, and also a limited number of radioelement ions and complexes that are common to both databases. As can be seen, many of the geologically relevant species in the H3 database are also present in PHREEQE.JNC. Equilibrium constants (25°C) are generally in close agreement. Reaction enthalpies (25°C) also compare favorably, although discrepancies exceeding a few kcal mol⁻¹ are not uncommon. Many aqueous species in the H3 database are not present in PHREEQE.JNC, however, and *vice versa*. Phosphate, carbonate, sulfate and nitrate complexes are generally better represented in the H3 TDB, whereas hydroxide and halogen complexes are more prevalent in PHREEQE.JNC.

Distinctly different approaches are associated with the H3 TDB and PHREEQE.JNC to account for the temperature dependence of the equilibrium constant of aqueous association reactions. As can be seen in Table 5.3.3_1, the H3 database requires equilibrium constants at temperatures other than 25°C to be estimated using the van't Hoff equation [i.e., flag (=kflag) = 0 (Table 5.3.1_2)]. Moreover, the association enthalpies of many complexes in this database are equal to zero, and equilibrium constants for the corresponding reactions at any temperature must therefore be equal to the value of the equilibrium constant at 25°C. In contrast, equilibrium constants between 0 and 100°C in PHREEQE.JNC are derived from calculations using SUPCRT and SPRONS.JNC (i.e., flag = 1; see Section 5.1). A comparison of log *K* values calculated in this manner at 25 and 60°C (Table 5.3.3_1) suggests that even relatively small increases in temperature can cause the equilibrium constant to increase, or decrease, by several orders of magnitude. This suggests that the van't Hoff equation may not be appropriate for estimating equilibrium constants at temperatures that are relevant to the Japanese repository concept, particularly if it is assumed that the reaction enthalpy equals zero. Moreover, the good agreement noted above among

Table 5.3.3_1. Equilibrium constants and reaction enthalpies for aqueous association reactions in the H3 TDB and PHREEQE.JNC.

Aqueous Species	H3 TDB			PHREEQE.JNC			
	flag*	ΔH_r° (kcal/mol)	Log K (25°C)	flag*	ΔH_r° (kcal/mol)	Log K (25°C)	Log K (60°C)
OH ⁻	0	13.345	-14.000	1	13.340	-13.995	
O ₂ (aq)	0	134.79	-86.08	1	133.734	-86.003	
H ₂ (aq)	0	-1.759	-3.15	1	-1.000	-3.105	
HCO ₃ ⁻	1	-3.561	10.33	1	-3.513	10.329	
H ₂ CO ₃ (aq)	1	-5.738	16.68	1	-5.832	16.673	
HSO ₄ ⁻	1	4.91	2.000	1	4.90	1.979	
H ₂ S(aq)	0	-65.179	40.681	1	-64.868	40.677	
HS ⁻	0	-40.14	33.652	1	-59.717	33.690	
H ₂ S ₂ O ₃ (aq)				1	-57.184	40.739	
HS ₂ O ₃ ⁻	0	0.	39.605	1	-60.684	40.153	38.615
HSO ₃ ⁻	0	0.	3.823	1	-0.587	3.583	4.35
H ₂ BO ₃ ⁻	0	3.224	-9.240	1	3.907	-9.249	
HPO ₄ ²⁻	0	-3.53	12.350	1	-3.515	12.322	
H ₂ PO ₄ ⁻	0	-4.52	19.55	1	-4.52	19.527	
HF(aq)	0	3.46	3.17	1	3.315	3.168	
HF ₂ ⁻	0	4.55	3.749	1	4.960	2.551	
H ₃ SiO ₄ ⁻	0	0.	-9.810	1	4.220	-9.585	-9.241
H ₂ SiO ₄ ²⁻	0	0.	-23.141				
HCl(aq)				1	9.648	-0.710	
HNO ₃ (aq)				1	4.019	-1.303	
H ₃ PO ₄ (aq)				1	-2.620	21.697	
HIO ₃ (aq)	0	168.0	-111.0	1	168.05	-110.518	
H ₂ Se(aq)	0	0.	85.1				
HSe ⁻	0	0.	81.2	1	-126.267	81.181	74.463
HSeO ₃ ⁻	0	0.	36.3	1	-48.097	36.304	33.416
H ₂ SeO ₃ (aq)	0	0.	38.9	1	-46.407	38.877	36.152
HSeO ₄ ⁻	0	0.	1.9	1	4.200	1.906	2.309
AlOH ²⁺	0	11.9	-4.99	1	11.902	-4.957	
Al(OH) ₂ ⁺	0	0.	-10.1	1	23.490	-10.594	-8.745
Al(OH) ₃ (aq)	0	0	-16.0	1	34.585	-16.433	-13.725
Al(OH) ₄ ⁻	0	44.06	-23.0	1	43.236	-22.883	
AlF ₂ ⁺	0	0.	7.01				
AlF ₂ ⁺	0	20.0	12.75				
AlF ₃ (aq)	0	2.5	17.02				
AlF ₄ ⁻	0	0.	19.72				
AlSO ₄ ⁺	0	2.15	3.02				
Al(SO ₄) ₂ ⁻	0	2.84	4.92				

Aqueous Species	H3 TDB			PHREEQE.JNC			
	flag*	ΔH_r° (kcal/mol)	Log K (25°C)	flag*	ΔH_r° (kcal/mol)	Log K (25°C)	Log K (60°C)
BaOH ⁺	0	15.095	-13.358	1	20.917	-13.500	
BaCl ⁺				1	3.110	-0.498	
BaF ⁺				1	2.140	-0.183	
BaCO ₃ (aq)				1	4.035	2.645	
BaHCO ₃ ⁺				1	1.999	11.311	
BiOH ₂ ⁺				1	4.157	-1.105	
Bi(OH) ₂ ⁺				1	18.657	-3.304	
Bi(OH) ₃ (aq)				1	30.974	-8.206	
Bi(OH) ₄ ⁻				1	45.674	-21.107	
BF(OH) ₃ ⁻	0	1.85	-0.40				
BF ₂ (OH) ₂ ⁻	0	1.635	7.628				
BF ₃ OH ⁻	0	-1.58	13.666				
BF ₄ ⁻	0	-1.795	20.274				
CaOH ⁺	0	14.535	-12.598	1	18.517	-12.833	
CaCl ⁺				1	1.126	-0.292	
CaCl ₂ (aq)				1	-1.394	-0.644	
CaF ⁺	0	3.798	0.940	1	1.350	0.682	
Ca(HCO ₃) ⁺	1	-0.869	11.435	1	-3.165	11.376	
CaCO ₃ (aq)	1	3.547	3.225	1	3.795	3.327	
CaSO ₄ (aq)	1	1.470	2.309	1	1.300	2.111	
CaPO ₄ ⁻	0	3.100	6.459				
CaHPO ₄ (aq)	0	-0.230	15.085				
CaH ₂ PO ₄ ⁺	0	-1.120	20.961				
CaH ₃ SiO ₄ ⁺				1	4.783	-8.575	
CH ₄ (aq)	0	-61.039	41.071	1	-64.575	38.194	
FeOH ⁺	0	13.2	-9.5	1	12.287	-9.315	
Fe(OH) ₂ (aq)	0	28.565	-20.57	1	27.417	-20.405	
Fe(OH) ₃ ⁻	0	30.3	-31.0	1	32.984	-29.207	
FeCl ⁺				1	0.723	-0.161	
Fe(Cl) ₂ (aq)				1	23.426	-8.172	
FeF ⁺				1	0.748	1.427	
FeSO ₄ (aq)	0	3.23	2.25				
Fe(HS) ₂ (aq)	0	-120.280	76.250				
Fe(HS) ₃ ⁻	0	-180.420	111.937				
FeHPO ₄ (aq)	0	-3.530	15.946				
FeH ₂ PO ₄ ⁺	0	-4.520	22.253				
Fe ³⁺	0	10.0	-13.032	1	10.200	-13.011	
FeOH ²⁺	0	20.4	-15.22	1	20.367	-15.216	

Aqueous Species	H3 TDB			PHREEQE.JNC			
	flag*	ΔH_r° (kcal/mol)	Log K (25°C)	flag*	ΔH_r° (kcal/mol)	Log K (25°C)	Log K (60°C)
Fe(OH) ₂ ⁺	0	10.0	-18.70	1	29.367	-18.661	
Fe(OH) ₃ (aq)	0	10.0	-26.63	1	38.384	-25.029	
Fe(OH) ₄ ⁻	0	10.0	-34.63	1	47.984	-34.631	
Fe ₂ (OH) ₂ ⁴⁺	0	33.5	-29.01				
Fe ₃ (OH) ₄ ⁵⁺	0	44.3	-45.4				
FeCl ₂ ⁺	0	15.6	-11.55	1	11.163	-11.531	
FeCl ₂ ⁺	0	10.0	-10.90				
FeCl ₃ (aq)	0	10.0	-11.90				
FeF ₂ ⁺	0	12.7	-6.80	1	15.059	-7.011	
FeF ₂ ⁺	0	14.7	-2.2				
FeF ₃ (aq)	0	15.4	0.97				
FeSO ₄ ⁺	0	13.91	-9.11				
Fe(SO ₄) ₂ ⁻	0	14.60	-7.61				
FeHPO ₄ ⁺	0	12.23	4.74				
FeH ₂ P ₂ ⁺	0	5.48	11.95				
IO ₃ ⁻	0	0.	-111.0	1	165.65	-111.324	-100.769
IO ₄ ⁻	0	251.0	-165.0	1	250.667	-165.044	
KOH(aq)				1	16.483	-14.895	
KCl(aq)				1	2.803	-1.751	
KBr(aq)				1	2.990	-1.737	
KI(aq)				1	2.041	-1.598	
KHSO ₄ (aq)				1	277.67	-179.909	
KSO ₄ ⁻	0	2.25	0.85	1	0.690	0.880	
KHPO ₄ ⁻	0	-3.530	12.636				
KAl(OH) ₄ (aq)				1	50.603	-24.224	
LiOH(aq)				1	13.369	-13.649	
LiCl(aq)				1	0.805	-1.511	
LiSO ₄ ⁻	0	0.	0.64				
MgOH ⁺	0	15.419	-11.794	1	26.729	-11.682	
MgCl ⁺				1	0.367	-0.135	
MgF ⁺	0	4.674	1.82	1	0.567	1.352	
Mg(HCO ₃) ⁺	1	-2.775	11.397	1	-2.998	11.365	
MgCO ₃ (aq)	1	2.535	2.981	1	2.182	2.979	
MgSO ₄ (aq)	0	1.4	2.25				
MgPO ₄ ⁻	0	3.100	6.589				
MgHPO ₄ (aq)	0	-0.2	15.216				
MgH ₂ PO ₄ ⁺	0	-1.120	21.066				
MgH ₃ SiO ₄ ⁺				1	3.759	-8.325	

Aqueous Species	H3 TDB			PHREEQE.JNC			
	flag*	ΔH°_r (kcal/mol)	Log K (25°C)	flag*	ΔH°_r (kcal/mol)	Log K (25°C)	Log K (60°C)
MnOH ⁺	0	14.40	-10.59	1	14.417	-10.59	
Mn(OH) ₂ (aq)				1	22.117	-22.201	
Mn(OH) ₃ ⁻	0	0.	-34.8	1	39.634	-34.800	-31.794
Mn(OH) ₄ ²⁻				1	56.234	-48.287	
MnCl ⁺	0	0.	0.607	1	4.553	-0.139	0.282
MnCl ₂ (aq)	0	0.	0.041				
MnCl ₃ ⁻	0	0.	-0.305				
MnF ⁺	0	0.	0.85	1	0.596	0.88	0.997
MnHCO ₃ ⁺	0	-3.604	11.60				
MnSO ₄ (aq)	0	2.17	2.26	1	3.55	1.913	
Mn(NO ₃) ₂ (aq)	0	-0.3	0.6				
Mn ³⁺	0	25.76	-25.507	1	22.2	-25.509	
MnO ₄ ²⁻	0	150.02	-118.44	1	169.567	-118.345	
NaOH(aq)				1	12.823	-14.795	
NaCl(aq)				1	1.206	-0.777	
NaF(aq)				1	1.723	-0.998	
NaBr(aq)				1	1.643	-1.357	
NaI(aq)				1	1.932	-1.540	
NaHCO ₃ (aq)	0	-3.604	10.080				
NaCO ₃ ⁻	0	8.911	1.268				
NaSO ₄ ⁻	0	1.12	0.70	1	-0.647	0.923	
NaHPO ₄ ⁻	0	-3.53	12.636				
NaAl(OH) ₄ (aq)				1	45.489	-23.627	
NaH ₂ BO ₃ (aq)				1	-64.557	32.581	
NaH ₃ SiO ₄ (aq)				1	1.631	-7.754	
NdOH ²⁺	0	0.	-8.0	1	19.317	-8.127	-6.608
Nd(OH) ₂ ⁺	0	0.	-16.9	1	27.817	-17.070	-14.293
Nd(OH) ₃ (aq)	0	0.	-26.5	1	55.134	-26.37	-22.284
Nd(OH) ₄ ⁻	0	0.	-37.1	1	71.434	-37.072	-32.068
NiOH ⁺	0	0.	-9.9	1	13.567	-10.803	-9.731
Ni(OH) ₂ (aq)	0	0.	-19.0	1	24.987	-20.706	-17.630
Ni(OH) ₃ ⁻	0	0.	-30.0	1	30.774	-31.003	-28.610
Ni(OH) ₄ ²⁻				1	47.734	-44.036	
Ni ₂ (OH) ₃ ⁺	0	0.	-10.7				
Ni ₄ (OH) ₄ ⁴⁺	0	0.	-27.7				
NiSO ₄ (aq)	0	0.	2.3				
NiCl ⁺				1	1.433	-0.996	
NiF ⁺				1	0.735	1.120	

Aqueous Species	H3 TDB			PHREEQE.JNC			
	flag*	ΔH_r° (kcal/mol)	Log K (25°C)	flag*	ΔH_r° (kcal/mol)	Log K (25°C)	Log K (60°C)
Np ³⁺	0	0.	3.03	1	7.000	3.445	4.304
NpO ₂ ⁺	0	35.73	-11.37	1	35.834	-9.145	
NpO ₂ ²⁺	0	0.	-32.29	1	63.934	-30.035	-25.859
NH ₃ (aq)	0	-174.58	109.83	1	-174.961	109.906	
NH ₄ ⁺	1	-187.055	119.070	1	-187.371	119.147	
NH ₄ SO ₄ ⁻	1	-187.055	120.19				
N ₂ (aq)	0	-312.13	207.08	1	-313.538	207.269	
NO ₂ ⁻	0	-43.76	28.57	1	-43.888	27.767	
PbOH ⁺	0	0.	-7.7	1	-0.803	-6.192	-6.283
Pb(OH) ₂ (aq)	0	0.	-17.1	1	23.397	-16.894	-15.152
Pb(OH) ₃ ⁻	0	0.	-27.3	1	32.214	-27.917	-25.548
Pb ₂ OH ³⁺	0	0.	-6.4				
Pb ₃ (OH) ₄ ²⁺	0	0.	-23.9				
Pb ₄ (OH) ₄ ⁴⁺	0	0.	-20.88				
Pb ₆ (OH) ₈ ⁴⁺	0	0.	-43.6				
PbCl ⁺	0	0.	1.61	1	1.083	1.437	1.578
PbCl ₂ (aq)	0	0.	2.44	1	1.946	2.003	2.24
PbCl ₃ ⁻	0	0.	2.01	1	1.879	1.688	1.941
PbCl ₄ ²⁻				1	0.303	1.43	
PbF ⁺				1	-0.661	2.06	
PbF ₂ (aq)				1	-2.703	3.420	
Pb(HS) ₂ (aq)				1	-134.805	82.089	
Pb(HS) ₃ ⁻				1	-196.525	117.078	
PbCO ₃ (aq)	0	0.	7.3				
Pb(CO ₃) ₂ ²⁻	0	0.	10.78				
PbSO ₄ (aq)	0	0.	2.59				
Pb(SO ₄) ₂ ²⁻	0	0.	3.5				
PdOH ⁺	0	0.	-2.12	1	-0.763	-1.091	-1.131
Pd(OH) ₂ (aq)	0	0.	-4.62	1	1.537	-2.190	-2.071
Pu ³⁺	0	0.	16.9	1	-13.300	17.006	16.399
PuO ₂ ⁺	0	46.23	-18.6	1	46.234	-17.941	-14.713
PuO ₂ ²⁺	0	0.	-34.9	1	68.334	-34.213	-29.603
S ₂ ⁻	0	-28.04	20.735				
S ₂ ²⁻				1	-104.534	57.645	
S ₃ ²⁻				1	-161.401	94.457	
S ₄ ²⁻				1	-217.968	131.049	
S ₅ ²⁻				1	-274.235	167.422	
S ₂ O ₃ ²⁻	0	0.	38.015	1	-61.784	38.467	36.701

Aqueous Species	H3 TDB			PHREEQE.JNC			
	flag*	ΔH_r° (kcal/mol)	Log K (25°C)	flag*	ΔH_r° (kcal/mol)	Log K (25°C)	Log K (60°C)
S ₂ O ₄ ²⁻				1	-18.567	10.549	
S ₂ O ₅ ²⁻				1	-2.150	2.348	
S ₂ O ₆ ²⁻				1	17.766	-8.418	
S ₂ O ₈ ²⁻				1	113.400	-65.501	
S ₃ O ₆ ²⁻				1	-36.701	25.902	
S ₄ O ₆ ²⁻				1	-106.168	76.128	
S ₅ O ₆ ²⁻				1	-150.435	97.474	
SO ₃ ²⁻	0	0.	-3.397	1	-2.817	-3.623	-3.105
SO ₂ (aq)				1	3.572	5.443	
Se ²⁻	0	0.	66.3				
SeO ₃ ²⁻	0	0.	29.0	1	-46.817	29.018	26.145
SnOH ⁺	0	0.	-3.89	1	6.617	-3.414	-2.927
Sn(OH) ₂ (aq)	0	0.	-7.90	1	10.317	-7.079	-6.33
Sn(OH) ₃ ⁻	0	0.	-17.39	1	16.634	-16.599	-15.397
SnCl ⁺	0	0.	1.05				
SnCl ₂ (aq)	0	0.	1.71				
SnCl ₃ ⁻	0	0.	1.69				
SnF ⁺	0	0.	4.08				
SnF ₂ (aq)	0	0.	6.68				
SnF ₃ ⁻	0	0.	9.46				
Sn ⁴⁺	0	0.	-5.38				
SnOH ³⁺	0	0.	-4.77				
Sn(OH) ₄ (aq)	0	0.	-4.53				
Sn(OH) ₂ ²⁺	0	0.	-5.56				
Sn(OH) ₃ ⁺	0	0.	-4.85				
SnF ³⁺	0	0.	-6.57				
SnF ₂ ²⁺	0	0.	-4.6				
SnF ₃ ⁺	0	0.	-6.31				
SnF ₄ (aq)	0	0.	4.34				
SnSO ₄ ²⁺	0	0.	-19.6				
Sn(SO ₄) ₂ (aq)	0	0.	-6.14				
SrOH ⁺	0	14.495	-13.178	1	19.987	-13.302	
SrCl ⁺	0	0.	-0.2	1	1.183	-0.248	-0.047
SrCl ₂ (aq)	0	0.	0.				
SrF ⁺				1	1.150	0.139	
SrHCO ₃ ⁺	1	2.486	11.513	1	2.489	11.509	
SrCO ₃ (aq)	1	5.217	2.805	1	4.435	2.865	
SrSO ₄ (aq)	0	1.6	2.55				

Aqueous Species	H3 TDB			PHREEQE.JNC			
	flag*	ΔH_r° (kcal/mol)	Log K (25°C)	flag*	ΔH_r° (kcal/mol)	Log K (25°C)	Log K (60°C)
SrNO ₃ ⁺	0	0.	0.8				
SrPO ₄ ⁻	0	0.	4.2				
SiF ₆ ²⁻	0	-16.26	30.18	1	-16.959	26.275	
U ³⁺	0	24.414	-8.615	1	24.400	-9.602	
UOH ³⁺	0	11.217	-0.540	1	11.217	-0.541	
U(OH) ₄ (aq)	0	0.	-4.479	1	18.234	-4.563	-3.185
UO ₂ ⁺	0	32.932	-7.507	1	32.934	-7.576	
UO ₂ OH(aq)				1	50.35	-25.738	
UO ₂ (OH) ₂ ⁻				1	72.650	-44.063	
UO ₂ ²⁺	0	34.4	-8.993	1	34.384	-9.049	
UO ₂ OH ⁺	0	44.723	-14.193	1	44.750	-14.267	
UO ₂ (OH) ₂ (aq)	0	0.	-20.994	1	46.650	-19.361	-16.557
UO ₂ (OH) ₃ ⁻	0	0.	-28.994	1	51.667	-28.295	-25.099
UO ₂ (OH) ₄ ²⁻	0	0.	-41.994	1	68.367	-42.075	-37.671
ZrOH ³⁺	0	0.	0.3	1	6.147	0.324	0.828
Zr(OH) ₂ ²⁺				1	14.347	-1.728	
Zr(OH) ₃ ⁺				1	21.664	-5.091	
Zr(OH) ₄ (aq)	0	0.	-9.7	1	23.164	-9.709	-7.904
Zr(OH) ₅ ⁻	0	0.	-16.0	1	21.980	-16.004	-13.699
Zr ₃ (OH) ₄ ⁸⁺	0	0.	-0.6				
Zr ₄ (OH) ₈ ⁸⁺	0	0.	6.0				
ZrSO ₄ ²⁺	0	0.	2.5				

* - kflag, see Table 5.3.1_2.

equilibrium constants at 25°C is probably not valid at other temperatures.

Thermodynamic data for a limited number of mineral dissolution reactions that have been considered by JNC in previous modeling studies of groundwater evolution are shown in Table 5.3.3_2. As can be seen, equilibrium constants at 25°C are generally in close agreement, but significant discrepancies (e.g., greater than about one order of magnitude) are evident for dolomite, kaolinite, K-feldspar, clinocllore, tremolite and pyrite dissolution. Reaction enthalpies (25°C) are also similar, but discrepancies exceeding a few kcal mol⁻¹ are apparent for dolomite, siderite, kaolinite and K-feldspar. Comments noted above on the use of the van't Hoff equation for estimating equilibrium constants at high temperatures also pertain to mineral dissolution reactions, as can be seen by comparing log K values at 25 and 60°C in the PHREEQE.JNC data base.

Mineral	H3 TDB			PHREEQE.JNC			
	flag*	ΔH_r° (kcal/mol)	Log K (25°C)	flag*	ΔH_r° (kcal/mol)	Log K (25°C)	Log K (60°C)
Calcite	1	-2.30	-8.48	1	-2.633	-8.480	
Aragonite	1	-2.59	-8.34	1	-2.654	-8.336	
Dolomite	0	-9.44	-17.09	1	-7.306	-18.144	
Siderite	0	-6.14	-10.57	1	-4.262	-10.521	
Strontianite	1	-0.40	-9.27	1	1.545	-10.643	
Gypsum	1	-0.03	-4.60				
Anhydrite	0	-4.30	-4.38	1	-4.44	-4.306	
Celestite	0	0.23	-6.58	1	-1.770	-5.677	
Barite	0	6.14	-9.98	1	6.20	-9.971	
Fluorite	0	4.71	-10.96	1	2.90	-10.037	
Amorph. Fe(OH) ₃	0	0.	4.89				
Quartz			-3.78	1	7.875	-3.999	
Gibbsite	0	-22.80	8.77	1	-24.566	7.756	
Chalcedony			-3.49	1	7.507	-3.728	
Kaolinite	0	49.15	-36.92	1	58.637	-38.957	
Phlogopite	0	0.	36.33	1	-73.975	37.267	31.355
K-feldspar	0	-12.47	1.78	1	-5.486	-0.448	
Albite**	0	0.	3.54	1	-12.393	2.764	1.568
Anorthite**	0	0.	26.37	1	-72.427	26.578	20.809
Clinocllore	0	0.	68.35	1	-146.359	67.239	55.772
Tremolite	0	0.	57.7	1	-97.131	61.237	53.171
Epidote	0	0.	45.43	1	-102.562	45.940	25.551
Muscovite	0	-59.38	14.60	1	-57.895	14.05	
Pyrite**	0	130.0	-86.0	1	130.684	-83.610	

* - mflag, see Table 5.3.1_3; ** - reaction stoichiometry in the H3 TDB

Table 5.3.3_2. Equilibrium constants and reaction enthalpies for selected mineral dissolution reactions in the H3 TDB and PHREEQE.JNC.

5.3.4 Excluded reactions

As noted in the preceding section, there are a number of aqueous species and minerals that are not represented in PHREEQE.JNC for which thermodynamic data are available in the H3 TDB (and numerous other databases). Several of these species and minerals could be important in experimental, field and modeling studies carried out by JNC. This raises the question whether such data should be included in PHREEQE.JNC.

Consistent with the objectives of this project (Section 2), we believe such data should not be included until an evaluation is made of their reliability at reference

conditions, and perhaps more importantly at other temperatures considered in the Japanese repository concept, and until it can be determined that they are internally consistent with respect to other data in PHREEQE.JNC and SPRONS.JNC. Due to time constraints in the present study, such assessments will have to be made by individual users as the need arises.

6 Evaluation of Database Reliability

The reliability of thermodynamic data in SPRONS.JNC (and, hence, PHREEQE.JNC) is evaluated in this section, where equilibrium constants for selected reactions calculated using SUPCRT and SPRONS.JNC are compared with their experimental counterparts. The comparisons are summarized in the form of diagrams showing calculated and experimental values of $\log K$ as a function of temperature at pressures corresponding to the liquid-phase side of the H_2O vaporization boundary (P_{sat}), or at higher pressures representing single-phase regions of H_2O stability. Diagrams for heterogeneous reactions (gas-solubility and mineral-solution equilibria) comprise Appendix C. Similar diagrams for homogeneous reactions (aqueous-speciation reactions) are included in Appendix D.

Equilibrium constants are chosen as the basis for this evaluation because they represent the most direct link between the thermodynamic properties of a reaction and experimental measurements (e.g., Nordstrom *et al.*, 1990). Some refinement of basic experimental data is usually necessary to determine values of $\log K$, however. For example, an appropriate chemical model of the aqueous system and total analytical solute concentrations may be needed to calculate the activities of species at reference conditions of infinite dilution, which partially or completely define the equilibrium constant. Time constraints in the present study do not permit independent retrievals of equilibrium constants from primary experimental results, however, and $\log K$ values reported in the literature must therefore be accepted at face value. For this reason, our objective is limited to identifying reactions for which calculated and experimental equilibrium constants agree favorably, as well as reactions where the agreement is less favorable, suggesting that future revisions to data in SPRONS.JNC and PHREEQE.JNC may be needed.

The thermodynamic properties of minerals in SPRONS.JNC are constrained primarily by the results of high-temperature and high-pressure phase-equilibrium experiments, and by a limited amount of calorimetric data (Section 3.2.1). The reliability of these data is evaluated by Helgeson *et al.* (1978), who compared experimental solution compositions, or phase-equilibrium reversal temperatures, for over 130 reactions with calculated mineral-stability relations, or univariant/divariant equilibrium curves. Comparisons of calculated and experimental equilibrium constants for mineral-solution equilibria and aqueous-speciation reactions are similarly documented in other sources of thermodynamic data incorporated in SPRONS.JNC (see Sections 5.2.2 and 5.2.3).

We supplement these earlier evaluations in the present study, focussing on heterogeneous and homogeneous equilibria at relatively low temperatures and pressures. Although the available experimental data referring to these conditions is still rather limited, the evaluations documented below are important because

they constitute independent checks on the reliability of data in SPRONS.JNC and PHREEQE.JNC at conditions that are most relevant to the Japanese repository concept. We also consider a few reactions involving common rock-forming minerals that are known to dissolve incongruently at low temperatures, because the dissolution and precipitation kinetics of these reactions is partially controlled by the equilibrium constant under near-to-equilibrium conditions. Comments on the reliability of thermodynamic data in SPRONS.JNC and PHREEQE.JNC are summarized in Sections 6.1 - 6.12. Laboratory and field experiments needed to improve these databases, and procedures to manage them, are summarized in Section 6.13.

6.1 Gas Solubilities and Acid Dissociation Constants

Plots of calculated and experimental equilibrium constants for gas solubility reactions are shown in the following figures (Appendix C):

- Fig. C.1; $\text{CO}_2(g)$,
- Fig. C.2; $\text{CH}_4(g)$,
- Fig. C.3; $\text{H}_2\text{S}(g)$,
- Fig. C.4; $\text{H}_2(g)$, and
- Fig. C.5; $\text{Ne}(g)$, $\text{Kr}(g)$, $\text{N}_2(g)$, $\text{Rn}(g)$, $\text{SO}_2(g)$, $\text{Ar}(g)$ and $\text{O}_2(g)$,

where sources of the experimentally determined equilibrium constants are also documented. As can be seen, calculated equilibrium constants represented by the curves in these figures are in close agreement with most of their experimental counterparts, represented by symbols (experimental uncertainties are indicated by the size of the symbols in Fig. C.5, and are estimated to be slightly larger than the size of the symbols in Figs. C.1-C.4 ($\approx \pm 0.1 \log K$ units)). Thermodynamic data and equation-of-state parameters for gases from Helgeson *et al.* (1978) and corresponding data for neutral aqueous species retrieved by Shock *et al.* (1989) are incorporated in SPRONS.JNC and were used to calculate all the curves in Figs C.1 - C.5. Shock *et al.* (1989) generate figures similar to Figs C.1 and C.4, and also document comparisons between calculated and experimental equilibrium constants for solubility reactions involving $\text{He}(g)$ and $\text{Xe}(g)$. The curves in Fig. C.5 for reactions involving $\text{Ne}(g)$, $\text{Kr}(g)$, $\text{SO}_2(g)$ and $\text{Ar}(g)$ are independent of the experimental data at temperatures greater than 25°C , because equation-of-state parameters for these gases are estimated using empirical correlation algorithms, as described by Shock *et al.* (1989). The agreement shown in Fig C.5 between calculated and experimental equilibrium constants at temperatures greater than 25°C is therefore noteworthy. Shock *et al.* (1989) also cite numerous references to experimental gas solubility data not shown in Figs. C.1 - C.5, which they assert are also closely consistent with the calculated curves in these figures. This wealth of experimental gas solubility data leaves little doubt that accurate gas solubilities can be calculated using SUPCRT and SPRONS.JNC between 0 and 100°C at 1 bar, and at considerably higher

temperatures and pressures.

Several of the gases noted above form weak acids when dissolved in aqueous solution. Plots of calculated and experimentally determined dissociation constants for these acids, and related ionic species, are shown in the following figures (Appendix D):

- Fig. D.1; H_2O ,
- Fig. D.2; $\text{CO}_2(aq)$ (i.e., $\text{H}_2\text{CO}_3(aq)$ - see Section 5.2.5),
- Fig. D.3; HCO_3^- ,
- Fig. D.4; $\text{H}_2\text{S}(aq)$,
- Fig. D.5; HSO_4^- ,
- Fig. D.6; $\text{HF}(aq)$,
- Fig. D.7; $\text{H}_3\text{BO}_3(aq)$,
- Fig. D.8; $\text{H}_3\text{PO}_4(aq)$,
- Fig. D.9; NH_4^+ , and
- Fig. D.10; CH_3COOH .

As can be seen, the calculated equilibrium constants are closely consistent with most of their experimental counterparts. Diagrams similar to Figs. D.2, D.4, D.6, D.7, D.8 and D.9 are generated by Shock *et al.* (1989), who also include diagrams for $\text{HNO}_3(aq)$ and $\text{SO}_2(aq)$ dissociation. These authors note that experimental equilibrium constants retrieved by Ellis and Giggensbach (1971) for $\text{H}_2\text{S}(aq)$ dissociation at temperatures greater than 100°C (Fig. D.4), and the equilibrium constants retrieved by Read (1975) for $\text{H}_2\text{CO}_3(aq)$ dissociation between 1 and 2 kb at temperatures less than 150°C (conditions not considered in Fig. D.2), are the only experimental values available that differ by more than $\pm 0.2 \log K$ units from values calculated using SUPCRT and thermodynamic data that are incorporated in SPRONS.JNC.

6.2 System $\text{SiO}_2\text{-H}_2\text{O}$

6.2.1 Quartz

The solubility of quartz has been experimentally investigated over a broad range of pressures and temperatures (Kennedy, 1950; Morey and Hesselgesser, 1951; Khitarov, 1956; Kitahara, 1960; van Lier *et al.*, 1960; Siever, 1962; Morey *et al.*, 1962; Weill and Fyfe, 1964; Anderson and Burnham, 1965, 1967; Crerar and Anderson, 1971; Hemley *et al.*, 1980; Walther and Orville, 1983; and Rimstidt, 1997). Shock *et al.* (1989) regressed experimental values of $\log K$ reported in many of these studies using the revised HKF model (Section 4.2.2) to retrieve equation-of-state coefficients and the standard partial molal volume, entropy and heat capacity of $\text{SiO}_2(aq)$ at 25°C and 1 bar. The conventional born coefficient for $\text{SiO}_2(aq)$ determined by Walther and Helgeson (1977) and thermodynamic properties of quartz from Helgeson *et al.* (1978) were adopted

in the regression procedure. The retrieved coefficients and properties of $\text{SiO}_2(aq)$ are used to generate the curves shown in Fig. C.6, where it can be seen that the curves accurately represent the available experimental data at temperatures between 0 and 900°C and pressures from 1 bar to 9 kb. The thermodynamic properties of $\text{SiO}_2(aq)$ retrieved by Shock *et al.* (1989) and the thermodynamic properties of quartz from Helgeson *et al.* (1978) are included in SPRONS.JNC.

Quartz solubilities calculated in the present study between 0 and 300°C at P_{sat} are shown in Fig C.7, which represents an expanded view of a portion of diagram A, Fig. C.6. As can be seen, the calculated equilibrium curve lies near all the experimental points, but it is apparent that the curve consistently under-predicts measured solubilities slightly over this temperature range. Error bars representing experimental uncertainties are probably smaller than the size of the plot symbols in this figure. These uncertainties are not documented in all the experimental studies, but we note that silica concentrations are routinely measured (e.g., using colorimetric methods; see Govett, 1961) with a precision of $\pm 5\%$, and that temperatures are probably accurate to at least $\pm 5^\circ\text{C}$ over this temperature range. The systematic deviation of the calculated curve from the experimental data in Fig C.7 is therefore probably real, even when experimental uncertainties are taken into account.

Figure C.7 incorporates recent experimental data from Rimstidt (1997) on quartz solubilities between 0 and 100°C. These experiments were designed to approach equilibrium from conditions of undersaturation in pure water at 21, 50, 74 and 96°C over equilibration periods ranging from 54 days (74°C) to 4917 days (21°C). Rimstidt (1997) combined his results with critically selected quartz solubility data from the literature to derive an empirical $\log K(T)$ function, which predicts that quartz solubility at 25°C is consistent with equilibrium $\text{SiO}_2(aq)$ concentrations of 10.8 ppm, rather than the widely accepted value of 6 ppm (Walther and Helgeson, 1977; Fournier and Potter, 1982). Using this function, Rimstidt (1997) calculates revised standard partial molal properties of $\text{SiO}_2(aq)$, which are consistent with the solubility of amorphous silica, and with silica concentrations in ancient groundwaters.

Quartz solubilities determined by Rimstidt (1997) between 0 and 100°C are compared in Fig. C.8 with solubilities calculated using SUPCRT and SPRONS.JNC. Also shown in the figure are experimental data from van Lier *et al.* (1960), which Rimstidt (1997) concludes are from equilibrated experiments, and data from Mackenzie and Gees (1971), Morey *et al.* (1962) and Fournier (1960), which Rimstidt (1997) suggests are from experiments that did not reach equilibrium. As can be seen, the calculated curve under-predicts the solubility data of Rimstidt (1997) and van Lier *et al.* (1960) by a factor less than 2 over this temperature range.

The observations summarized above with regard to Figs. C.7 and C.8 suggest that the thermodynamic properties of $\text{SiO}_2(\text{aq})$ in SPRONS.JNC may need to be revised. Although such revisions are probably minor, they are important because this would entail revising the properties of several other minerals in SPRONS.JNC, which were retrieved by Helgeson *et al.* (1978) using the standard Gibbs free energy of formation of $\text{SiO}_2(\text{aq})$ determined by Walther and Helgeson (1977).

6.2.2 Amorphous silica

The solubility of amorphous silica has been experimentally investigated over a range of pressures and temperatures (Hitchen, 1935; Okamoto *et al.*, 1957; Elmer and Nordberg, 1958; Kitahara, 1960; Morey *et al.*, 1964; Fournier and Rowe, 1977; Chen and Marshall, 1982). The thermodynamic properties of amorphous silica in SPRONS.JNC are from Helgeson *et al.* (1978), who adopted the standard Gibbs free energy at 25°C recommended by Walther and Helgeson (1977).

The solubilities of amorphous silica between 0 and 300°C calculated using SUPCRT and SPRONS.JNC are compared with experimental solubilities in Fig. C.9. The calculated values agree closely with their experimental counterparts over this temperature range. In contrast, solubilities determined in some of the older studies noted above do not agree with the calculated values. This is apparently because the earlier studies failed to allow sufficient time for the solid and solution phases to equilibrate (Morey *et al.*, 1964).

6.2.3 Chalcedony

The solubility of chalcedony was determined by Fournier and Rowe (1966) between 125 and 255°C at P_{sat} . Fournier (1973) later measured its solubility between 92 and 259°C. Fournier (1977) fit the combined data points from these two studies to an empirical $\log K(T)$ function. Walther and Helgeson (1977) digitized figures displaying the solubility data of Fournier and Rowe (1966) and Fournier (1973) and similarly fit these data to an empirical $\log K(T)$ function, which enabled the standard Gibbs free energy of formation of chalcedony to be estimated at 25°C and 1 bar, assuming chalcedony in fact consists solely of microcrystalline quartz (see Section 6.2.4, however). This value was adopted by Helgeson *et al.* (1978), and is incorporated in SPRONS.JNC.

Chalcedony solubilities between 0 and 100°C calculated using SUPCRT and SPRONS.JNC are compared in Fig. C.10 with values calculated using the temperature functions of Walther and Helgeson (1977) and Fournier (1977). The close agreement among solubilities calculated using SPRONS.JNC and the empirical function of Walther and Helgeson (1977) is expected because the same Gibbs free energy for chalcedony constrains both these calculations. Solubilities calculated using the $\log K(T)$ function of Fournier (1977) are greater

than those calculated using SPRONS.JNC at low temperatures, but converge at higher temperatures as the lower temperature limit of the experimentally calibrated range (92°C) is approached.

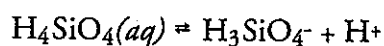
6.2.4 Moganite

Heaney and Post (1992) suggest that chalcedony is a mixture of microcrystalline quartz and moganite, a little known silica polymorph. Gislason *et al.* (1997) use the principle of detailed balancing and experimentally determined rates of moganite dissolution to recommend provisional estimates of its standard Gibbs free energy and enthalpy of formation, and standard entropy, at 25°C and 1 bar. The corresponding value of $\log K$ for moganite hydrolysis equals -3.14, which corresponds to a dissolved silica concentration of 44 mg kg⁻¹. We have not included these data in SPRONS.JNC, given the provisional nature of these results, but such data should be included in future revisions of this database when more definitive experimental data characterizing moganite's thermodynamic properties become available.

6.2.5 Aqueous speciation

As noted in Section 6.2.1, Shock *et al.* (1989) regressed experimental constants for the quartz solubility reaction using the revised HKF model to retrieve equation-of-state coefficients and the standard partial molal volume, entropy and heat capacity of SiO₂(aq) at 25°C and 1 bar. The retrieved coefficients and properties of SiO₂(aq) are used to generate the curves in Fig. C.6, where it can be seen that calculated and experimental quartz solubilities are in close agreement over a broad range of pressures and temperatures. The thermodynamic properties of SiO₂(aq) may need to be revised, however, to be consistent with recent experimental quartz solubilities determined by Rimstidt (1997) between 0 and 100°C.

Sverjensky *et al.* (1997) regressed experimental equilibrium constants between 0 and 350°C at *P_{sat}* reported by Harman (1928), Roller and Ervin (1940), Greenberg and Price (1957), Schwarz and Muller (1958), Ingri (1959), Lagerstrom (1959), van Lier *et al.* (1960), Volosov *et al.* (1972), Seward (1974) and Busey and Mesmer (1977) for the reaction



to retrieve thermodynamic properties and HKF equation-of-state coefficients for H₃SiO₄⁻, using thermodynamic data for H₄SiO₄(aq) from Shock *et al.* (1989). The results are shown in Fig. D.11, where the curve represents values calculated using SUPCRT and SPRONS.JNC and symbols refer to most of the experimental studies noted above. As can be seen, the calculated values agree with experimental counterparts to within about ±0.1 log *K* units at temperatures greater than 60°C, but the agreement is notably poorer at lower temperatures.

6.3 System $\text{Al}_2\text{O}_3\text{-H}_2\text{O}$

Mineral solubilities and the thermodynamic properties of minerals and aqueous species in this chemical subsystem have been studied intensively over the past 100 years. A thorough review of the experimental studies is summarized by Pokrovskii and Helgeson (1995). As noted in Section 5.2.3, thermodynamic data and equation-of-state coefficients in SPRONS.JNC for minerals and aqueous species comprising this subsystem are from this reference.

6.3.1 *Gibbsite*

Gibbsite solubilities in dilute acidic solutions, where Al^{3+} is the dominant Al species in solution, have been experimentally determined between 0 and 100°C by Frink and Peech (1962), Kittrick (1966a), Singh (1974), May *et al.* (1979), Bloom and Weaver (1982), Peryea and Kittrick (1988), Nagy and Lasaga (1992) and Palmer and Wesolowski (1992). The standard molal Gibbs free energy and enthalpy of formation, and standard entropy and volume, of gibbsite at the reference temperature and pressure in SPRONS.JNC are from Pokrovskii and Helgeson (1995), who adopted the calorimetric values recommended by Robie *et al.* (1978). The experimental solubilities of gibbsite at 25°C determined by Peryea and Kittrick (1988) and Palmer and Wesolowski (1992) were selected by Pokrovskii and Helgeson (1995) to calculate the standard partial molal Gibbs free energy of formation of Al^{3+} . This value and the corresponding standard partial molal enthalpy of formation from Cox *et al.* (1989) are used to calculate the standard partial molal entropy of Al^{3+} at 25°C and 1 bar. Calculated equilibrium constants using SUPCRT and SPRONS.JNC for the gibbsite solubility reaction between 0 and 100°C at 1 bar are compared with experimental values determined in the studies noted above in Fig. C.11, where it can be seen that the calculated curve is consistent with most of the experimental data. The experimentally determined equilibrium constants determined by Frink and Peech (1962) and Kittrick (1966a) were later questioned by Bloom and Weaver (1982) because the gibbsite used in these older experiments was covered with a reactive surface material suspected to be microcrystalline gibbsite.

Gibbsite solubilities in dilute alkaline solutions, where AlO_2^- [or, equivalently, $\text{Al}(\text{OH})_4^-$; see Section 5.2.5] is the dominant Al species in solution, have been experimentally determined between 6.4 and 130°C by Fricke and Jucaitis (1930), Tsirlina (1936), Fulda and Ginsberg (1951), Russell *et al.* (1955), Ikkatai and Okada (1962), Lyapunov *et al.* (1964), Kittrick (1966a), Apps (1970), Berecz and Szita (1970), Packter (1979), Apps and Neil (1990), Wesolowski (1992) and Verdes *et al.* (1992). Pokrovskii and Helgeson (1995) regressed experimental equilibrium constants retrieved from these studies to determine the standard partial molal Gibbs free energy of formation of AlO_2^- . The regression curve generated by Pokrovskii and Helgeson (1995) is compared in Fig. C.12 with experimental equilibrium constants. As can be seen, the calculated curve agrees

closely with most of the experimental values. The same agreement is expected using SUPCRT and SPRONS.JNC because thermodynamic data from Pokrovskii and Helgeson (1995) are adopted in the present study. This is confirmed in Fig. C.13, where equilibrium constants calculated using SUPCRT and SPRONS.JNC for the gibbsite solubility reaction in alkaline solutions are compared with experimental values determined by Apps and Neil (1990), Wesolowski (1992) and Verdes *et al.* (1992).

6.3.2 Boehmite

Experimental data characterizing the solubility of gibbsite at low temperatures in slightly acidic to moderately alkaline solutions, where AlOH^{2+} , $\text{Al}(\text{OH})_2^+$ and/or $\text{Al}(\text{OH})_3(\text{aq})$ are the dominant Al species, are difficult to evaluate because the solubilities are extremely low. At higher temperatures, gibbsite dehydrates to form boehmite (Russell *et al.*, 1955; Khodakovskiy *et al.*, 1980). Pokrovskii and Helgeson (1995) therefore evaluate the thermodynamic properties and HKF equation-of-state coefficients for AlOH^{2+} , $\text{Al}(\text{OH})_2^+$ and $\text{Al}(\text{OH})_3(\text{aq})$ from experimental data on boehmite solubilities at relatively high temperatures (≈ 50 to 300°C).

Experimental solubilities of boehmite [$\gamma\text{AlO}(\text{OH})$] in alkaline solutions (Russell *et al.*, 1955; Kuyunko *et al.*, 1983; Apps *et al.*, 1989; Verdes, 1990; Verdes *et al.*, 1992; Bourcier *et al.*, 1993; Castet *et al.*, 1993) are used by Pokrovskii and Helgeson (1995), together with the thermodynamic properties of AlO_2^- determined as noted above (and properties of $\text{NaAlO}_2(\text{aq})$ also determined by these authors), to calculate the standard Gibbs free energy of formation of boehmite at 25°C and 1 bar. Calculations using the thermodynamic data retrieved by Pokrovskii and Helgeson (1995) of equilibrium constants for the boehmite solubility reaction in alkaline solutions are compared with their experimental counterparts in Fig. C.14. As can be seen, the calculated values are closely consistent with the available experimental data between 50 and 300°C at P_{sat} , but there is a tendency for the calculations to slightly underpredict the experimentally determined values at temperatures less than about 160°C . Experimental uncertainties are not reported in all cases, but we estimate them conservatively to be about $\pm 0.1 \log K$ units and $\pm 5^\circ\text{C}$, which is roughly comparable to the size of the plot symbols in this figure.

A similar comparison of calculated and experimental equilibrium constants for the boehmite solubility reaction in acidic solutions (where Al^{3+} is the dominant Al species in solution) is shown in Fig. C.15, where it can be seen that the calculated curve agrees closely with all the experimental data, except equilibrium constants retrieved by Verdes (1990) at 70°C and Verdes *et al.* (1992) at 90°C . The otherwise close correspondence between calculated and experimental values is noteworthy because the thermodynamic properties of Al^{3+} retrieved by Pokrovskii and Helgeson (1995) are independent of the experimental data on boehmite solubilities. The size of the plot symbols in Fig. C.15 is probably

comparable to, or smaller than, experimental uncertainties, which are attributable primarily to the low concentrations of Al^{3+} in the experimental solutions, and to ambiguities in evaluating the activity coefficient of Al^{3+} in solutions with high ionic strength (Pokrovskii and Helgeson, 1995).

As noted above, thermodynamic properties and HKF equation-of-state coefficients for AlOH_2^+ , $\text{Al}(\text{OH})_2^+$ and $\text{Al}(\text{OH})_3(\text{aq})$ in SPRONS.JNC are based on experimental measurements of boehmite solubility in moderately acidic to moderately alkaline solutions (Kuyunko *et al.*, 1983; Verdes, 1990; Verdes *et al.*, 1992; Bourcier *et al.*, 1993; Castet *et al.*, 1993). Comparisons of calculated and experimental equilibrium constants for the boehmite solubility reaction in systems where AlOH_2^+ , $\text{Al}(\text{OH})_2^+$ and $\text{Al}(\text{OH})_3(\text{aq})$ are the dominant Al species in solution are depicted in Figs. C.16, C.17 and C.18, respectively, as functions of temperature at *Psat*. As can be seen in these figures, there is considerable scatter in the experimental data. The calculated curves in Figs. C.16 and C.17 are within the combined experimental and computational uncertainty reported by Bourcier *et al.* (1993). The thermodynamic properties and equation-of-state parameters for AlOH_2^+ and $\text{Al}(\text{OH})_2^+$ in SPRONS.JNC are therefore consistent with these experimental results, but not with the results of the other experimental measurements noted in the figures. The curve in Fig. C.18 was generated by Pokrovskii and Helgeson (1995) by regression of the experimental values of $\log K$ for the indicated reaction, and using supplemental data on corundum solubilities (see below). The curve falls within the scatter of the data, but there is considerable disagreement among the experimental values, and between most of these values and the calculated curve.

The experimental confusion regarding boehmite solubilities in moderately acidic to moderately alkaline solutions, as shown in Figs. C.16 - C.18, can only be resolved by additional experimentation. The thermodynamic properties and HKF coefficients for AlOH_2^+ , $\text{Al}(\text{OH})_2^+$ and $\text{Al}(\text{OH})_3(\text{aq})$ recommended by Pokrovskii and Helgeson (1995) are provisionally included in SPRONS.JNC pending experimental clarification of boehmite's solubility behavior.

6.3.3 Diaspore

Experimental data on the solubility of diaspore in NaOH solutions between 200 and 350°C at *Psat* (Druzhinina, 1955; Bernshtein and Matsenok, 1965; Wefers, 1967; Chang *et al.*, 1979; Verdes *et al.*, 1992) were regressed by Pokrovskii and Helgeson (1995) using their thermodynamic properties and HKF equation-of-state parameters for AlO_2^- and $\text{NaAlO}_2(\text{aq})$ to determine the standard Gibbs free energy of formation of diaspore at 25°C and 1 bar. Diaspore solubilities in dilute NaCl - NaOH solutions between 100 at 325°C and *Psat* are calculated by Pokrovskii and Helgeson (1995) and compared with experimental values determined by Apps *et al.* (1989), which were not considered in the regression procedure. The results (Fig. 9, Pokrovskii and Helgeson, 1995) indicate that the thermodynamic properties of diaspore and AlO_2^- are consistent with the

experimentally determined solubilities.

6.3.4 Corundum

Experimental data from Becker *et al.* (1983) on the solubility of corundum between 2.5 and 6 kb at 670°C were regressed by Pokrovskii and Helgeson (1995) using their retrieved thermodynamic properties and HKF equation-of-state coefficients for AlO_2^- , and thermodynamic properties for corundum at 25°C and 1 bar from Cox *et al.* (1989), to determine the partial molal thermodynamic properties of $\text{Al}(\text{OH})_3(\text{aq})$. The retrieved data are used to calculate corundum solubilities between 1 and 2.5 kb at 500°C, and the predicted values are compared with experimental values reported by Morey (1957) and Ragnarsdottir and Walther (1985). The calculated values of $\log K$ are within 0.3 log units or less of their experimental counterparts (see Fig. 23. Pokrovskii and Helgeson, 1995).

6.3.5 Aqueous speciation

As noted in Section 6.3.1, partial molal thermodynamic properties and HKF equation-of-state coefficients for Al^{3+} are determined by Pokrovskii and Helgeson (1995) by regression of experimental data on gibbsite solubilities. These properties are used by Pokrovskii and Helgeson (1995), together with potentiometric data on Al^{3+} hydrolysis and boehmite solubility data reported in the literature, to calculate corresponding thermodynamic properties and equation-of-state coefficients for $\text{Al}(\text{OH})_2^+$. Equilibrium constants for the hydrolysis reaction calculated using SUPCRT and SPRONS.JNC are compared with their experimental counterparts between 0 and 100°C at 1 bar in Fig. D.12, where it can be seen that the calculated constants agree to within 0.1 log K units (approximately equal to the size of plot symbols) with most of the experimental data. Additional experimental data for this reaction between 50 and 250°C reported by Nikolaeva and Tolpygina (1969), Verdes (1990) Castet *et al.* (1993) and Bourcier *et al.* (1993) are considered by Pokrovskii and Helgeson (1995), but are not shown in Fig. D.12. The experimental equilibrium constants determined by Verdes (1990), and the value reported by Bourcier *et al.* (1993) at 250°C, differ by more than 0.1 log K units from the values calculated using SUPCRT and SPRONS.JNC.

The reliability of thermodynamic data and equation-of-state coefficients for the species $\text{Al}(\text{OH})_2^+$ and $\text{Al}(\text{OH})_3(\text{aq})$ is difficult to assess at low temperatures and pressures, and in near-neutral pH solutions where these species are important, because gibbsite solubilities under such conditions are of the same order of magnitude as the analytical detection limit for dissolved aluminum. Wesolowski and Palmer (1994), however, have developed a general model of aluminum's speciation behavior based on their experimental measurements of gibbsite solubilities at 50°C in 0.1 molal NaCl solutions containing acetate, bistris and tris pH-buffers, which is consistent with earlier studies by these

authors on gibbsite's solubility behavior in acidic and alkaline solutions (Palmer and Wesolowski, 1992; Wesolowski, 1992), and experimental measurements of boehmite solubilities at temperatures between 50 and 250°C at P_{sat} (Verdes, 1990; Bourcier *et al.*, 1993; Castet *et al.*, 1993). The model permits calculation of hydrolysis constants for $\text{Al}(\text{OH})_2^+$ and $\text{Al}(\text{OH})_3(\text{aq})$ at temperatures between 0 and 100°C at 1 bar that are consistent with all these experimental constraints.

These constants are compared with corresponding values calculated using SUPCRT and SPRONS.JNC in Figs. D.13 and D.14, respectively. As can be seen, the hydrolysis constants for $\text{Al}(\text{OH})_2^+$ determined by Wesolowski and Palmer (1994) are closely consistent with values calculated using SUPCRT and SPRONS.JNC over this temperature range (Fig. D.13). Hydrolysis constants for $\text{Al}(\text{OH})_3(\text{aq})$ from Wesolowski and Palmer (1994), however, agree with the value calculated using SUPCRT and SPRONS.JNC only at 100°C (Fig. D.14). The experimental uncertainty assigned to this constant is less than 0.2 log K units (Wesolowski and Palmer, 1994), which suggests that the discrepancies indicated in the figure are significant. Wesolowski and Palmer (1994) demonstrate that $\text{Al}(\text{OH})_3(\text{aq})$ does not contribute significantly to the aqueous speciation of aluminum between 0 and 100°C, however, which suggests that the discrepancies indicated in Fig. D.14 would not generate serious errors in geochemical calculations referring to these conditions.

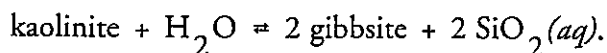
As noted in Section 6.3.1, the partial molal thermodynamic properties and HKF equation-of-state coefficients for AlO_2^- [i.e., $\text{Al}(\text{OH})_4^-$; see Section 5.2.5] are determined by Pokrovskii and Helgeson (1995) by regression of experimental data on gibbsite solubilities. Equilibrium constants for the corresponding hydrolysis reaction calculated using SUPCRT and SPRONS.JNC are compared with their experimental counterparts between 0 and 100°C at 1 bar in Fig. D.15, where it can be seen that the calculated constants are closely consistent with experimental values determined by Palmer and Wesolowski (1992) and Wesolowski and Palmer (1994), but that a discrepancy between 0.5 and 1 log K units is generated with respect to experimental constants determined by May *et al.* (1979). Palmer and Wesolowski (1992) suggest that the anomalously high solubilities determined by May *et al.* (1979) are due to some form of metastable material or extremely fine-grained gibbsite on the surface of the solid used in those experiments.

6.4 System $\text{SiO}_2\text{-Al}_2\text{O}_3\text{-H}_2\text{O}$

6.4.1 Kaolinite

The standard molal Gibbs free energy and enthalpy of formation, and standard entropy and volume, of kaolinite at the reference temperature and pressure in SPRONS.JNC are from Helgeson *et al.* (1978). These authors computed the standard Gibbs free energy of formation of gibbsite from the calorimetric value

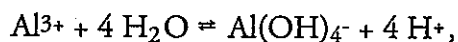
of its standard enthalpy of formation determined by Hemingway and Robie (1977), and used these data together with the compositions of mineral assemblages and coexisting interstitial waters in Jamaican bauxite deposits and weathered Hawaiian basalts to calculate the standard Gibbs free energy of formation of kaolinite consistent with the following reaction:



Helgeson *et al.* (1978) note that the ΔG_f° for kaolinite derived in this manner is consistent with solubility data from Kittrick (1966b) for coarsely crystalline kaolinite, but that the ΔH_f° of kaolinite is 2510 cal mol⁻¹ less negative than the calorimetric value determined by Hemingway and Robie (1977). The approach taken by Helgeson *et al.* (1978) is critically evaluated by Hemingway *et al.* (1982), who conclude that this enthalpy is incorrect.

The thermodynamic properties of kaolinite recommended by Helgeson *et al.* (1978) are therefore questionable. We note, however, that revisions to the thermodynamic properties of minerals and aqueous species in the system Al₂O₃-H₂O recommended by Pokrovskii and Helgeson (1995) are incorporated in SPRONS.JNC, and that the results of recent experimental studies on the solubility of kaolinite are now available. It is therefore possible to evaluate the accuracy of the thermodynamic data for kaolinite adopted in the present study by comparison of equilibrium constants for the kaolinite solubility reaction computed using SUPCRT and SPRONS.JNC with their experimental counterparts.

The solubility of kaolinite in dilute acidic solutions, where Al³⁺ is the dominant Al species in solution, has been experimentally determined at 25 and 80°C by May *et al.* (1986) and Nagy *et al.* (1991). Polzer and Hem (1965), Kittrick (1966b), Reesman and Keller (1968), Huang and Keller (1973), Devidal *et al.* (1996) and Zotov *et al.* (1998) have measured its solubility in alkaline solutions between 25 and 300°C at *P*_{sat}. Equilibrium constants calculated using SUPCRT and SPRONS.JNC for the kaolinite solubility reaction in acidic solutions are compared with the relevant experimental data in Fig. C.19. The experimental data referring to alkaline conditions from Devidal *et al.* (1996), Kittrick (1966b) and Zotov *et al.* (1998) are recalculated in the present study for conditions in acidic solutions using equilibrium constants for the reaction,



based on data for the aqueous aluminum species from Pokrovskii and Helgeson (1995) [see Fig. D.15]. As can be seen, the calculated solubility of kaolinite is in excellent agreement with the experimentally determined values reported by Kittrick (1966b) and May *et al.* (1986) at the reference temperature and pressure. There is considerable scatter, however, in the experimental solubility

data at higher temperatures and pressures. The calculated curve passes within this scatter, but calculated and experimental values at a given temperature differ by as much as 3 orders of magnitude. Zotov *et al.* (1998) evaluate the effects of grain size on kaolinite solubilities. As can be seen in Fig C.19, the solubility increases slightly with decreasing grain size, but the effects are too small to explain the scatter in experimental solubilities shown in this figure.

The equilibrium constant calculated using the data for kaolinite from Helgeson *et al.* (1978) and unrevised data for the Al^{3+} ion from Shock and Helgeson (1988) [inverted solid triangle in Fig. C.19] is nearly two orders of magnitude smaller at 25°C and 1 bar than the corresponding experimental data and the value of the equilibrium constant calculated in the present study (open circle). This suggests that the revised thermodynamic properties for Al^{3+} from Pokrovskii and Helgeson (1995), which are consistent with the thermodynamic properties of gibbsite determined by Robie *et al.* (1978), have resolved some of the confusion concerning kaolinite's solubility behavior.

Equilibrium constants for the kaolinite solubility reaction in alkaline solutions between 25 and 150°C at P_{sat} measured by Devidal (1996), and recalculated by these authors based on experimental results from Polzer and Hem (1965), Reesman and Keller (1968) and Huang and Keller (1973), are shown in Fig. C.20. As can be seen, there is also considerable scatter among these experimental data. The curve calculated using SUPCRT and SPRONS.JNC passes within this scatter, but calculated and experimental values differ by as much as 2 orders of magnitude.

Additional experimentation is clearly needed to resolve the discrepancies in kaolinite's solubility behavior at temperatures greater than 25°C. In the meantime, the thermodynamic properties of this mineral determined by Helgeson *et al.* (1978), and included in SPRONS.JNC, are accepted provisionally. It is important to note that questions raised by Hemingway *et al.* (1982) concerning the standard enthalpy of formation of kaolinite calculated by Helgeson *et al.* (1978) are still unresolved. Moreover, because kaolinite is used as a reference phase in simultaneous evaluations of aluminous mineral equilibria by Helgeson *et al.* (1978), the enthalpies of formation of these phases may also be questionable (see Section 4.3.4.3). Unfortunately, the scatter in experimental solubilities depicted in Figs. C.19 and C.20 preclude the use of these data to help clarify the thermodynamic properties of kaolinite.

6.4.2 Kaolinite, boehmite, diaspore, andalusite and pyrophyllite

Equilibria among various minerals in the Al_2O_3 - SiO_2 - H_2O system are examined in a series of experiments at high temperatures and pressures by Hemley *et al.* (1980), Huang (1993) and Zotov *et al.* (1998). The following diagrams document comparisons of equilibrium constants calculated using SUPCRT and SPRONS.JNC for the investigated reactions among the indicated minerals and

an aqueous phase:

- Fig C.21; kaolinite-boehmite,
- Fig C.22; kaolinite-diaspore,
- Fig C.23; pyrophyllite-kaolinite,
- Fig C.24; pyrophyllite-diaspore,
- Fig C.25; pyrophyllite-andalusite,
- Fig C.26; andalusite-corundum, and
- Fig C.27; andalusite-diaspore.

The equilibrium activity of $\text{SiO}_2(aq)$ at the experimental pressures and temperatures determines the value of the equilibrium constant for all these reactions. Silica activity is closely approximated by the total molality of dissolved silica because the solutions are sufficiently dilute and acidic that the concentrations of other silica species would be vanishingly small under such conditions, and because $\text{SiO}_2(aq)$ is uncharged. Dissolved silica is determined with a precision of $\pm 5\%$ of the measured value in the experiments of Hemley *et al.* (1980), and temperature is accurate to $\pm 5^\circ\text{C}$. Similar experimental uncertainties are assumed here for the experimental results reported by Huang (1993) and Zotov *et al.* (1998). As can be seen in these figures, the calculated equilibrium constants agree in some cases to within a few tenths of a log K unit with their experimental counterparts (Figs. C.23, C.24 and C.25), but in other cases discrepancies on the order of 0.5 to 1 log K units are apparent (Figs. C.21, C.22, C.26 and C.27).

Figures C.21-C.23 and C.25-C.27 are similar to Fig. 47(a-d) of Helgeson *et al.* (1978), who compared calculated equilibrium constants with unpublished experimental results from J. J. Hemley, which presumably are the same results later reported by Hemley *et al.* (1980). Comparison of our figures with Fig. 47 of Helgeson *et al.* (1978) indicates that our calculated curves in Figs. C.21, C.22, C.26 and C.27 are offset slightly from those generated by Helgeson *et al.* (1978), and that the latter curves agree significantly better with the experimental data. We note that the reactions considered in Figs. C.21, C.22, C.26 and C.27 involve minerals (boehmite, diaspore and corundum) whose thermodynamic properties are revised by Pokrovskii and Helgeson (1995) from values originally retrieved by Helgeson *et al.* (1978), based, in part, on the experimental results portrayed in their Fig. 47. This suggests that additional work may be needed to reconcile the discrepancies documented in Figs. C.21, C.22, C.26 and C.27 such that the resulting calculated curves are consistent with the experimental data of Hemley *et al.* (1980), as well as with the generally lower temperature and pressure experimental constraints evaluated by Pokrovskii and Helgeson (1995).

6.5 System $\text{Na}_2\text{O}-\text{K}_2\text{O}-\text{SiO}_2-\text{Al}_2\text{O}_3-\text{H}_2\text{O}$

Equilibria among various minerals in this system are examined in a series of experiments, generally at high temperatures and pressures, by Montoya and

Hemley (1975), Sverjensky *et al.* (1991), Burch (1993), Murphy *et al.* (1996) and Wilkin and Barnes (1998). The following diagrams document comparisons of equilibrium constants calculated using SUPCRT and SPRONS.JNC for the investigated reactions among the indicated minerals and an aqueous solution:

- Fig C.28; K-feldspar-muscovite-quartz (0.5 kb),
- Fig C.29; K-feldspar-muscovite-quartz (1.0 kb),
- Fig C.30; K-feldspar-muscovite-quartz (2.0 kb),
- Fig C.31; muscovite-andalusite-quartz,
- Fig C.32; muscovite-kaolinite,
- Fig C.33; muscovite-quartz-pyrophyllite,
- Fig C.34; albite-paragonite-quartz,
- Fig C.35; albite, and
- Fig C.36; analcime.

Equilibria among K-bearing minerals in this system at temperatures between 300 and 600°C and pressures from 0.5 to 2 kb were determined experimentally by Sverjensky *et al.* (1991). The results of this study are consistent with earlier investigations by Hemley (1959), Hemley and Jones (1964) and Montoya and Hemley (1975), which therefore are not considered further in the present study. Sverjensky *et al.* (1991) evaluated the thermodynamic properties of the following reactions (structural formulas of minerals as noted in Appendix A):

- $1.5 \text{ K-feldspar} + \text{H}^+ \rightleftharpoons 0.5 \text{ muscovite} + 3 \text{ quartz} + \text{K}^+$,
- $\text{K-feldspar} + \text{H}^+ \rightleftharpoons 0.5 \text{ andalusite} + 2.5 \text{ quartz} + \text{K}^+$,
- $\text{muscovite} + \text{H}^+ + 1.5 \text{ H}_2\text{O} \rightleftharpoons 1.5 \text{ kaolinite} + \text{K}^+$
- $\text{muscovite} + \text{H}^+ + 3 \text{ quartz} \rightleftharpoons 1.5 \text{ pyrophyllite} + \text{K}^+$, and
- $\text{muscovite} + \text{H}^+ \rightleftharpoons 1.5 \text{ andalusite} + 1.5 \text{ quartz} + 1.5 \text{ H}_2\text{O} + \text{K}^+$,

by measuring the total analytical concentrations of K^+ and H^+ in quenched solutions equilibrated with the mineral reactants and products. It was assumed as a first approximation that the ratio of these concentrations, $m_{\text{K}}/m_{\text{H}}$ (where m denotes molality), defines the equilibrium constant because the activity coefficients for K^+ and H^+ are approximately equal (and therefore cancel) and because the presence of other K- and H-bearing aqueous species are assumed to be present in negligible concentrations. Reactions analogous to the above were also considered by these authors to account for HCl(aq) and KCl(aq) complex formation, which it was assumed becomes significant at temperatures greater than about 300°C. Sverjensky *et al.* (1991) retrieved thermodynamic properties and HKF equation-of-state coefficients for HCl(aq) from the results of their experiments on K-feldspar-muscovite-quartz equilibrium, but these data are not included in SPRONS.JNC for reasons discussed in Section 5.2.3.

Equilibria among Na-bearing minerals in this system were determined experimentally by Montoya and Hemley (1975), Burch (1993), Murphy *et al.* (1996) and Wilkin and Barnes (1998). Montoya and Hemley (1975) used the

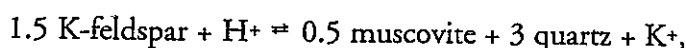
same experimental approach as Hemley (1959), Hemley and Jones (1964) and Sverjensky *et al.* (1991) to evaluate the thermodynamic properties of the following reactions between 300 and 500°C at 1 kb (structural formulas of minerals as noted in Appendix A):

- $3 \text{ albite} + 2 \text{ H}^+ \rightleftharpoons \text{paragonite} + 6 \text{ quartz} + 2 \text{ Na}^+$,
- $2 \text{ paragonite} + 2 \text{ H}^+ \rightleftharpoons 3 \text{ kaolinite} + 2 \text{ Na}^+$,
- $2 \text{ paragonite} + 6 \text{ quartz} + 2 \text{ H}^+ \rightleftharpoons 3 \text{ pyrophyllite} + 2 \text{ Na}^+$, and
- $2 \text{ paragonite} + 2 \text{ H}^+ \rightleftharpoons 3 \text{ andalusite} + 3 \text{ quartz} + 2 \text{ Na}^+$,

by measuring the total analytical concentrations of Na^+ and H^+ in quenched solutions equilibrated with the mineral reactants and products. It was assumed that the ratio of these concentrations, $m_{\text{t, Na}}/m_{\text{t, H}}$, defines the equilibrium constant. Burch (1993; see also Burch *et al.*, 1993) determined steady-state dissolution rates of albite at 80°C and pH 8.8 using a continuously-stirred flow-through reactor. The solubility of albite under these conditions was estimated from solution compositions sampled at steady-state in the dissolution and precipitation experiments nearest to equilibrium, as well as from near-equilibrium batch experiments. Murphy *et al.* (1996) and Wilkin and Barnes (1998) investigated the solubility of analcime. The experiments carried out by Murphy *et al.* (1996) reacted analcime (from Mt. St. Hilaire, Quebec) at 25°C in 0.1 molal NaCl - 0.01 molal NaHCO_3 solutions that were continuously equilibrated with atmospheric $\text{CO}_2(\text{g})$. Wilkin and Barnes (1998) reacted analcime (from Mt. St. Hilaire, Quebec and Wikieup, Arizona) in doubly distilled H_2O purged of CO_2 and O_2 in rocking autoclaves from 25 to 300°C at P_{sat} .

6.5.1 K-feldspar, muscovite, quartz

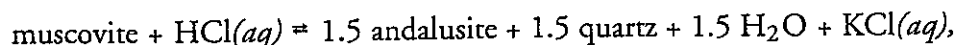
Equilibrium constants for the reaction



between 300 and 600°C at 0.5, 1 and 2 kb calculated using SUPCRT and SPRONS.JNC are compared with their experimental counterparts determined by Sverjensky *et al.* (1991) in Figs. C.28-C.30, respectively. Two calculated curves are shown in each figure representing solutions in which K^+ and H^+ , or $\text{KCl}(\text{aq})$ and $\text{HCl}(\text{aq})$, are assumed to be the dominant species in solution (reactions "1" and "2" as labelled in the figures, respectively). Experimental uncertainties ($\pm 0.15 \log K$ units) are estimated by Sverjensky *et al.* (1991). As can be seen, most of the experimentally determined equilibrium constants are consistent with calculated values for reaction 1 at 300°C, or reaction 2 at temperatures greater than 300°C. Discrepancies significantly exceeding experimental uncertainties are apparent, however, for the experimental datum at 500°C and 0.5 kb (Fig. C.28), and the experimental data at temperatures greater than 300°C at 2 kb (Fig. C.30).

6.5.2 *Muscovite, andalusite, quartz*

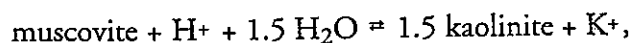
Equilibrium constants calculated using SUPCRT and SPRONS.JNC for the reaction



between 0 and 600°C at 1 and 2 kb are compared with their experimental counterparts determined by Sverjensky *et al.* (1991) in Fig. C.31. Experimental uncertainties ($\pm 0.15 \log K$ units) are omitted in this figure for clarity. As can be seen, the experimentally determined equilibrium constants are generally consistent with the calculated curves, but discrepancies slightly exceeding experimental uncertainties are apparent.

6.5.3 *Muscovite, kaolinite*

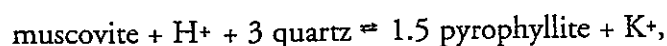
Equilibrium constants calculated using SUPCRT and SPRONS.JNC for the reaction



between 0 and 350°C at 1 kb are compared in Fig. C.32 with the experimental value at 300°C determined by Sverjensky *et al.* (1991). The calculated curve is consistent with this value and its associated uncertainty ($\pm 0.15 \log K$ units).

6.5.4 *Muscovite, quartz, pyrophyllite*

Equilibrium constants for the reaction

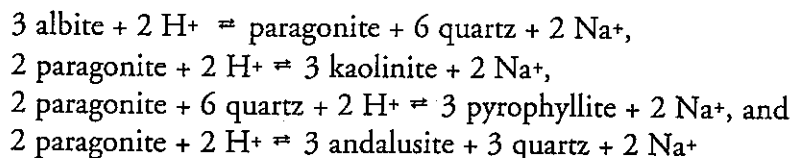


between 0 and 400°C at 1 kb calculated using SUPCRT and SPRONS.JNC are compared with their experimental counterparts determined by Sverjensky *et al.* (1991) in Fig. C.31. The two calculated curves in this figure represent solutions in which K^+ and H^+ , or $\text{KCl}(aq)$ and $\text{HCl}(aq)$, are assumed to be the dominant species in solution (reactions "1" and "2" as labelled in the figure, respectively). As can be seen, the calculated curve for reaction 1 is consistent with the experimental data and their associated uncertainties ($\pm 0.15 \log K$ units). The calculated curve for reaction 2 is inconsistent with these data. Complex formation of $\text{KCl}(aq)$ and $\text{HCl}(aq)$ is apparently unimportant at temperatures below about 300°C, although the exact temperature and pressure at which this transition in complexing behavior occurs is uncertain (Sverjensky *et al.*, 1991).

6.5.5 *Albite, paragonite, kaolinite, pyrophyllite, andalusite, quartz*

Helgeson *et al.* (1978) compare calculated equilibrium constants for the

reactions:



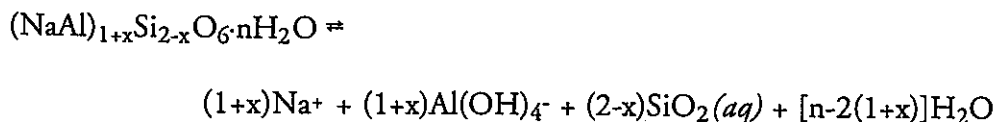
with their respective experimental counterparts determined by Montoya and Hemley (1975) (see Figs. 68a-d of Helgeson *et al.*, 1978, respectively). Calculated equilibrium constants are closely consistent with all the experimental values, except for the last reaction above where discrepancies on the order of 0.5 log K units are evident. Thermodynamic properties and equation-of-state parameters for H^+ and Na^+ from Helgeson and Kirkham (1976), and standard molal thermodynamic properties for the minerals in these reactions determined by Helgeson *et al.* (1978), were used in the calculations. These data are incorporated in SPRONS.JNC, and similar agreement between equilibrium constants computed using this database and experimental values determined by Montoya and Hemley (1975) is therefore expected. This is confirmed in Fig. C.34, which is similar to Fig 68a of Helgeson *et al.* (1978).

6.5.6 Albite

The equilibrium constant determined by Burch (1993) for the albite solubility reaction at 80°C is compared with values calculated using SUPCRT and SPRONS.JNC in Fig. C.35. The state of order/disorder of the albite used in the experiments is unknown, and two calculated curves are therefore shown representing equilibrium with respect to high albite (completely disordered) and low albite (completely ordered). As can be seen, the experimental solubility is nearly an order of magnitude more soluble than that calculated for the disordered phase, and nearly two orders of magnitude more soluble than that calculated for the ordered phase. Burch (1993) calculated the solubility of albite from steady-state solution compositions observed in dissolution and precipitation experiments that were "nearest" equilibrium. The equilibrium constant determined in this study is thus an approximation, but it is unclear how close the approximate value is to the actual equilibrium value.

6.5.7 Analcime

Equilibrium constants for the reaction



between 0 and 300°C at P_{sat} calculated using SUPCRT and SPRONS.JNC ($x = 1$, $n = 1$) are compared with their experimental counterparts determined by

Murphy *et al.* (1996) [Mt. St. Hilaire analcime: $x = 1.05$, $n = 1$] and Wilkin and Barnes (1998) [Mt. St. Hilaire analcime: $x = 1$, $n = 1$; Wikieup analcime: $x = 0.85$, $n = 1$] in Fig. C.36. As can be seen, the experimental values determined by Wilkin and Barnes (1998) for the Mt. St. Hilaire sample are remarkably consistent with calculated counterparts over the full range of temperatures considered, taking into account the experimental uncertainties determined by these authors. Significant discrepancies between calculated and observed solubilities are evident for the Wikieup analcime, however, presumably because this phase is significantly enriched in Si and deficient in Al relative to the idealized composition assumed for analcime in SPRONS.JNC. The solubility of Mt. St. Hilaire analcime determined by Murphy *et al.* (1996) at 25°C is more than an order of magnitude greater than that determined by Wilkin and Barnes (1998) and calculated using SUPCRT and SPRONS.JNC. The sample used by Murphy *et al.* (1996) is slightly richer in Si and poorer in Al than that used by Wilkin and Barnes (1998), but it is unlikely that this can explain the discrepancy between these experimental equilibrium constants.

6.6 System MgO-SiO₂-H₂O

Equilibrium relations among various minerals in this system have been characterized experimentally by Christ *et al.* (1973) and Hemley *et al.* (1977a, b). The latter study evaluated the thermodynamic properties of the following reactions (structural formulas of minerals as noted in Appendix A):

- $2 \text{ talc} = 3 \text{ forsterite} + 5 \text{ SiO}_2(\text{aq}) + 2 \text{ H}_2\text{O}$,
- $\text{talc} + \text{H}_2\text{O} = \text{chrysotile} + 2 \text{ SiO}_2(\text{aq})$
- $16 \text{ talc} + 15 \text{ H}_2\text{O} = \text{antigorite} + 30 \text{ SiO}_2(\text{aq})$,
- $7 \text{ talc} = 3 \text{ anthophyllite} + 4 \text{ SiO}_2(\text{aq}) + 4 \text{ H}_2\text{O}$,
- $\text{talc} = 3 \text{ enstatite} + \text{SiO}_2(\text{aq}) + \text{H}_2\text{O}$,
- $2 \text{ anthophyllite} = 7 \text{ forsterite} + 9 \text{ SiO}_2(\text{aq}) + 2 \text{ H}_2\text{O}$,
- $\text{anthophyllite} = 7 \text{ enstatite} + \text{SiO}_2(\text{aq}) + \text{H}_2\text{O}$, and
- $2 \text{ enstatite} = \text{forsterite} + \text{SiO}_2(\text{aq})$,

at relatively high temperatures and pressures by measuring the total analytical concentrations of $\text{SiO}_2(\text{aq})$ in quenched solutions equilibrated with the respective mineral reactants and products. The equilibrium activity of $\text{SiO}_2(\text{aq})$ (which is approximately equal to its molality, see Section 6.4.2) determines the value of the equilibrium constant for all these reactions. Christ *et al.* (1973) similarly evaluated the thermodynamic properties of the following reaction (structural formula of sepiolite as noted in Appendix A):

- $\text{sepiolite} + 8 \text{ H}^+ = 4 \text{ Mg}^{2+} + 6 \text{ SiO}_2(\text{aq}) + 11 \text{ H}_2\text{O}$,

between 50 and 90°C at 1 bar by analyzing equilibrated solutions for pH and for total concentrations of Mg^{2+} and $\text{SiO}_2(\text{aq})$ to estimate corresponding activities

of the aqueous reactants and products using the Debye-Hückel equation (e.g., Garrels and Christ, 1965).

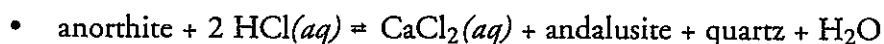
Helgeson *et al.* (1978) compare calculated equilibrium constants with the experimentally determined values for all the above reactions (Figs. 27-30 and Fig. 34 of Helgeson *et al.*, 1978, respectively). Calculated equilibrium constants agree closely with all the experimental data. Thermodynamic properties and equation-of-state parameters for Mg^{2+} from Helgeson and Kirkham (1976), similar data for $\text{SiO}_2(\text{aq})$ from Walther and Helgeson (1977), and standard molal thermodynamic properties for the minerals in these reactions from Helgeson *et al.* (1978) were used in the calculations. Most of these data are incorporated in SPRONS.JNC, but thermodynamic properties and equation-of-state coefficients for $\text{SiO}_2(\text{aq})$ are revised by Shock *et al.* (1989) from those determined by Walther and Helgeson (1977), as noted in Section 6.2.1. The revisions are relatively minor, however, and are not expected to significantly alter the consistency among experimental and calculated equilibrium constants for the minerals in this system, as documented by Helgeson *et al.* (1978). This is confirmed in Figs. C.37 and C.38, which exhibit the same close correspondence among calculated and experimental results as Figs. 27 and 34, respectively, of Helgeson *et al.* (1978).

6.6.1 Aqueous speciation

Experimental data for the hydrolysis of Mg^{2+} between 0 and 150°C at P_{sat} are compared in Fig. D.16 with equilibrium constants calculated for this reaction using SUPCRT and SPRONS.JNC. As noted by Shock *et al.* (1997a), the close agreement among most of the experimental data and the calculated curve is remarkable because the HKF equation-of-state parameters for MgOH^+ are estimated independently of experimental constraints, using the correlation algorithms developed by these authors.

6.7 System $\text{CaO-SiO}_2\text{-Al}_2\text{O}_3\text{-H}_2\text{O}$

There are few experimental data characterizing the solubilities of minerals in this system. Roselle and Baumgartner (1995) measured the solubility of the assemblage:



in supercritical chloride solutions from 400 to 600°C at 2 kb. Equilibrium constants retrieved by these authors are compared with equilibrium constants calculated using SUPCRT and SPRONS.JNC in Fig. C.39. As can be seen, the calculated curve is consistent with the experimental data at 500 and 600°C, but the discrepancy at 400°C is significantly greater than the experimental uncertainty.

6.7.1 Aqueous speciation

Experimental data for the hydrolysis of Ca^{2+} between 0 and 450°C at P_{sat} and at 0.5 and 1 kb are compared in Fig. D.17 with equilibrium constants calculated for this reaction using SUPCRT and SPRONS.JNC. As can be seen, the experimental data are closely consistent with the calculated curve over this range of temperatures and pressures. This supports the validity of correlations developed by Shock *et al.* (1997a), because equation-of-state coefficients for CaOH^+ were estimated using these correlations independently of the experimental data shown in the figure.

6.8 System $\text{CaO-MgO-FeO-BaO-SrO-H}_2\text{O-CO}_2$

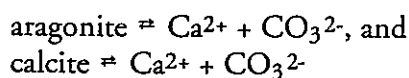
Plummer and Busenberg (1982), Busenberg *et al.* (1984) and Busenberg and Plummer (1986) carried out a series of experiments between 0 and 91°C at 1 bar to investigate mineral-solution and aqueous-speciation equilibria in this system. These authors determined the solubility of aragonite, calcite, vaterite, strontianite and witherite over this temperature range, and corresponding ion-pair formation constants for CaHCO_3^+ , $\text{CaCO}_3(\text{aq})$, SrHCO_3^+ , $\text{SrCO}_3(\text{aq})$, BaHCO_3^+ and $\text{BaCO}_3(\text{aq})$ that are internally consistent with the solubility data. The experimental data and thermodynamic properties retrieved from them by Plummer and Busenberg (1982), Busenberg *et al.* (1984) and Busenberg and Plummer (1986) supersede numerous earlier investigations, which are critically evaluated for the $\text{CaO-CO}_2\text{-H}_2\text{O}$ system by Langmuir (1968) and Jacobsen and Langmuir (1974), summarized by Busenberg *et al.* (1984) for the $\text{SrO-CO}_2\text{-H}_2\text{O}$ system, and summarized by Busenberg and Plummer (1986) for the $\text{BaO-CO}_2\text{-H}_2\text{O}$ system. The comprehensive experimental investigations of this chemical system by Plummer and Busenberg (1982), Busenberg *et al.* (1984) and Busenberg and Plummer (1986) did not include Fe, and corresponding thermodynamic properties of siderite and ferrous-carbonate complexes.

As noted in Section 5.2.2, Johnson *et al.* (1991) corrected the standard Gibbs free energy and enthalpy of formation of calcite and aragonite in their SPRONS database (SPRONS92.DAT) to be consistent with the solubility data for these phases determined by Plummer and Busenberg (1982). The data for strontianite and witherite were unrevised, however, and are therefore unconstrained by the experimental solubilities determined by Busenberg *et al.* (1984) and Busenberg and Plummer (1986), respectively. The thermodynamic properties of strontianite and witherite in SPRONS.JNC are thus the original data from Helgeson *et al.* (1978), which are constrained solely by calorimetric measurements (Table 9 of Helgeson *et al.*, 1978). The thermodynamic properties of siderite are similarly constrained only by calorimetric data. Helgeson *et al.* (1978) did not compare these calorimetric-based properties with other experimental observations, and notes that they are therefore subject to greater uncertainty.

Experimental data reported by Plummer and Busenberg (1982), Busenberg *et al.* (1984) and Busenberg and Plummer (1986) were regressed by Sverjensky *et al.* (1997) to determine partial molal thermodynamic properties and HKF equation-of-state coefficients for $\text{CaCO}_3(\text{aq})$, $\text{SrCO}_3(\text{aq})$ and $\text{BaCO}_3(\text{aq})$. These data are included in SPRONS.JNC. Similar data for the CaHCO_3^+ ion-pair in this database are attributed by Johnson *et al.* (1991) to unpublished work from Sverjensky (ref. 5, Appendix A), but these data are not documented by Sverjensky *et al.* (1997). Data for the SrHCO_3^+ and BaHCO_3^+ ion pairs also were not evaluated by these authors, and are not included in SPRONS.JNC.

6.8.1 Aragonite and calcite

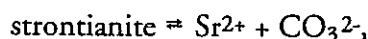
Equilibrium constants for the reactions



between 0 and 100°C at 1 bar calculated using SUPCRT and SPRONS.JNC are compared with their experimental counterparts determined by Plummer and Busenberg (1982) in Figs. C.40 and C.41, respectively. As can be seen, the calculated curves agree closely with all the experimental data. This is expected because these experimental data partially constrain the standard molal thermodynamic properties of aragonite and calcite in SPRONS.JNC.

6.8.2 Strontianite

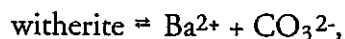
Equilibrium constants for the reaction



between 0 and 100°C at 1 bar calculated using SUPCRT and SPRONS.JNC are compared with their experimental counterparts determined by Busenberg *et al.* (1984) in Fig. C.42. The calculated curve underpredicts the experimentally determined solubility product of strontianite by more than an order of magnitude over this temperature range. As noted above, the thermodynamic properties of strontianite in SPRONS.JNC are constrained only by calorimetric measurements. The discrepancies in Fig. C.42 suggest these data should be revised to be consistent with both the calorimetric constraints and solubility data that are now available.

6.8.3 Witherite

Equilibrium constants for the reaction



between 0 and 100°C at 1 bar calculated using SUPCRT and SPRONS.JNC are compared with their experimental counterparts determined by Busenberg and Plummer (1986) in Fig. C.43. The calculated curve severely underpredicts the experimentally determined solubility product of witherite by 3 to 4 orders of magnitude over this temperature range. As noted above, the thermodynamic properties of witherite in SPRONS.JNC are constrained by calorimetric measurements alone. The discrepancies in Fig. C.43 suggest these data should be revised to be consistent with the available calorimetric and solubility data.

6.8.4 Siderite

Equilibrium constants for the reaction



between 0 and 100°C at 1 bar calculated using SUPCRT and SPRONS.JNC are compared with experimentally determined solubility products at 25°C in Fig. C.44. As can be seen, the curve falls within the scatter of the data, but there is considerable disagreement among the experimental values, and between most of these values and the calculated curve. Additional experimentation is needed to resolve experimental discrepancies in the solubility product of siderite at 25°C and at higher temperatures to better constrain siderite's standard molal thermodynamic properties.

6.8.5 Aqueous speciation

Dissociation constants for the ion pairs CaHCO_3^+ , $\text{CaCO}_3(\text{aq})$, $\text{SrCO}_3(\text{aq})$ and $\text{BaCO}_3(\text{aq})$ calculated using SUPCRT and SPRONS.JNC between 0 and 100°C are compared with experimentally determined dissociation constants in Figs. D.18-D.21, respectively. As can be seen in Fig. D.18, calculated dissociation constants for CaHCO_3^+ are reasonably consistent with experimental values determined by Plummer and Busenberg (1982), but discrepancies slightly exceeding experimental uncertainties are apparent at temperatures less than 65°C. Calculated dissociation constants for $\text{CaCO}_3(\text{aq})$, shown in Fig. D.19, are closely consistent with all but one of the experimentally determined values reported by Plummer and Busenberg (1982), as is expected because these experimental data were regressed by Sverjensky *et al.* (1997) to retrieve the partial molal thermodynamic properties and equation-of-state coefficients for this species. Experimental data determined by Reardon and Langmuir (1974) also lie within 0.1 to 0.2 log K units of calculated equilibrium constants for this reaction. The excellent agreement among calculated and experimental dissociation constants for $\text{SrCO}_3(\text{aq})$ and $\text{BaCO}_3(\text{aq})$, as can be seen in Figs. D.20 and D.21, respectively, is similarly due to the fact that experimental data for $\text{SrCO}_3(\text{aq})$ dissociation from Busenberg *et al.* (1984) and for $\text{BaCO}_3(\text{aq})$ dissociation from Busenberg and Plummer (1986) were used in the regression calculations by Sverjensky *et al.* (1997) to retrieve the thermodynamic properties

of these species. Experimental data on the dissociation constant of $\text{SrCO}_3(aq)$ at 25°C from Hasegawa *et al.* (1970) are clearly discrepant with respect to the calculated curve in Fig. D.20, and the experimental value determined by Busenberg *et al.* (1984). Experimental data on the dissociation constant of $\text{BaCO}_3(aq)$ at 25°C from Benes and Selecka (1973) are similarly discrepant with respect to the calculated curve in Fig. D.21, and the experimental value from Busenberg and Plummer (1986). The experimentally determined dissociation constant for $\text{BaCO}_3(aq)$ reported by Palmer and Van Eldik (1983) agrees closely with that determined by Busenberg and Plummer (1986).

Sverjensky *et al.* (1997) regressed experimental data from Reardon and Langmuir (1974) and Siebert and Hostettler (1977) on the dissociation constant of $\text{MgCO}_3(aq)$ between 10 and 90°C to retrieve partial molal thermodynamic properties and equation-of-state coefficients for this species. As can be seen in Fig. D.22, equilibrium constants for this reaction calculated using SUPCRT and SPRONS.JNC agree closely with the data from Siebert and Hostettler (1997), and lie within $0.2 \log K$ units of the values determined by Reardon and Langmuir (1974). There is considerable scatter among the other experimental data at 25°C , however.

Correlation algorithms developed by Shock *et al.* (1997a) were used by these authors to predict the standard partial molal properties of SrOH^+ . Association constants for this species calculated using SUPCRT and SPRONS.JNC are compared in Fig. D.23 with values determined experimentally between 5 and 50°C by Gimblett and Monk (1954). As can be seen, the calculated curve is closely consistent with all the experimental data, which supports the validity of the correlation approach used by Shock *et al.* (1997a).

6.9 System CaO-BaO-SrO-SO₃-H₂O

6.9.1 Anhydrite

The solubility of anhydrite between 0 and 350°C at P_{sat} determined experimentally by Yeatts and Marshall (1969) is compared in Fig. C.45 with calculated solubilities using SUPCRT and SPRONS.JNC. The standard molal thermodynamic properties of anhydrite in SPRONS.JNC are from Helgeson *et al.* (1978). Sverjensky *et al.* (1997) regressed the solubility data of Yeatts and Marshall (1969) to retrieve the standard partial molal thermodynamic properties and HKF equation-of-state coefficients for the $\text{CaSO}_4(aq)$ ion pair, and this is reflected by the close agreement among calculated and experimental solubilities shown in Fig. C.45.

6.9.2 Barite

The solubilities of barite between 0 and 300°C at P_{sat} determined experimental-

ly by Blount (1977), Strübel (1967), Melcher (1910), Templeton (1960), Malinin *et al.* (1969), Felmy *et al.* (1990) and Jiang (1996) are compared with equilibrium constants for this reaction calculated using SUPCRT and SPRONS.JNC in Fig. C.46. The standard molal thermodynamic properties of barite in SPRONS.JNC are from Helgeson *et al.* (1978), and are based solely on calorimetric measurements. The partial molal thermodynamic properties and equation-of-state coefficients of Ba^{2+} and SO_4^{2-} are from Shock and Helgeson (1988), and were retrieved independently of the solubility data for barite. The close correspondence among most of the calculated and experimental solubilities for this phase is therefore noteworthy, because it constitutes independent confirmation that the thermodynamic data for barite and the aqueous ions Ba^{2+} and SO_4^{2-} are reliable over this temperature range. Experimental equilibrium constants for this reaction reported by Strübel (1967) are clearly discrepant with respect to calculated counterparts and the other experimental data.

6.9.3 Celestite

The solubility of celestite at 25°C determined by Felmy *et al.* (1990) is compared with the equilibrium constant for this reaction calculated using SUPCRT and SPRONS.JNC in Fig. C.47. The standard molal thermodynamic properties of celestite in SPRONS.JNC are from Helgeson *et al.* (1978), and are constrained solely by calorimetric measurements. The partial molal thermodynamic properties and equation-of-state coefficients of Sr^{2+} and SO_4^{2-} are from Shock and Helgeson (1988), and were retrieved independently of the solubility data for celestite. As can be seen, this mineral is calculated to be nearly an order of magnitude more soluble at 25°C than measured by Felmy *et al.* (1990).

6.10 System FeO-S(rh)-H₂O-HCl

6.10.1 Pyrite

The solubility of pyrite in chloride solutions has been determined experimentally by Crerar *et al.* (1978) between 200 and 350°C at *Psat*. Calculated pyrite solubilities using SUPCRT and SPRONS.JNC in solutions equilibrated with FeCl^+ or $\text{FeCl}_2(\text{aq})$ are compared with experimental values determined by Crerar *et al.* (1978) in Figs. C.48 and C.49, respectively. The thermodynamic properties of pyrite in SPRONS.JNC are constrained only by calorimetric data (Helgeson *et al.*, 1978). Sverjensky *et al.* (1997) regressed experimental dissociation constants for FeCl^+ from Heinrich and Seward (1990) and Palmer and Hyde (1993), and total dissociation constants for $\text{FeCl}_2(\text{aq})$ reported by Wood and Crerar (1985), Heinrich and Seward (1990) and Fein *et al.* (1992), to retrieve the standard partial molal thermodynamic properties and HKF equation-of-state coefficients for these species. The close correspondence among the calculated and experimental solubilities for pyrite shown in Figs. C.48 and C.49 constitutes independent confirmation that the thermodynamic data for

pyrite and both ferrous chloride complexes are reliable over this temperature range.

6.10.2 Pyrrhotite

Crerar *et al.* (1978) measured the solubility of pyrrhotite in chloride solutions between 200 and 350°C at *Psat*. Berner (1967) also measured the solubility of a sample of nearly stoichiometric pyrrhotite ($\text{Fe}_{0.98}\text{S}$) at 25°C and 1 bar in dilute solutions equilibrated with $\text{H}_2\text{S}(g)$. Pyrrhotite solubilities calculated using SUPCRT and SPRONS.JNC are compared with the experimental value determined by Berner (1967) in Fig. C.50, and with solubilities measured by Crerar *et al.* (1978) in Figs. C.51 and C.52. The thermodynamic properties of pyrrhotite in SPRONS.JNC are constrained only by calorimetric data (Helgeson *et al.*, 1978). The standard partial molal thermodynamic properties and HKF equation-of-state coefficients for Fe^{2+} were revised from values recommended by Shock and Helgeson (1988) by Shock *et al.* (1997a). The standard partial molal thermodynamic properties and HKF equation-of-state coefficients for FeCl^+ and $\text{FeCl}_2(aq)$ were determined by Sverjensky *et al.* (1997), as noted above. As can be seen in Fig. C.50, pyrrhotite is calculated to be more than an order of magnitude more soluble at 25°C than measured by Berner (1967). A similar discrepancy is noted in Fig. C.51 for the solubility at 200°C, but better agreement is apparent in this figure, and in Fig. C.52, for the solubilities measured at higher temperatures. Additional experimentation is needed to better constrain the solubility behavior of pyrrhotite at temperatures between 0 and 300°C.

6.10.3 Magnetite

Two experimental datasets characterizing the solubility of magnetite and relative stabilities of ferrous hydroxide complexes between 25 and 300°C at *Psat* are reported by Sweeton and Baes (1970) and Tremaine and LeBlanc (1980). Both sets of data were obtained from measurements of magnetite solubility in solutions of varying pH, but equilibrium constants derived from these data for the hydroxide complexes differ by as much as two log units. Shock *et al.* (1997a) note that it is not evident that the dataset of Sweeton and Baes (1970) is intrinsically more accurate than that of Tremaine and LeBlanc (1980), but nevertheless select the data of Sweeton and Baes (1970) to retrieve by regression the partial molal thermodynamic properties and equation-of-state coefficients for ferrous hydroxide complexes that are incorporated in SPRONS.JNC. The rationale for this decision is that retrieved parameters, such as the standard partial molal entropies of these complexes, are more consistent with empirical correlations constructed by Shock *et al.* (1997a) than are similar parameters retrieved from the dataset of Tremaine and LeBlanc (1980). Calculated magnetite solubilities using SUPCRT and SPRONS.JNC between 0 and 300°C at *Psat* in solutions equilibrated with Fe^{2+} , FeOH^+ , $\text{Fe}(\text{OH})_2(aq)$ and $\text{Fe}(\text{OH})_3^-$ are compared with their experimental counterparts determined by Sweeton and

Baes (1970) in Fig. C.53. As can be seen, the agreement between calculated and experimental equilibrium constants is excellent over this temperature range, as expected.

6.10.4 Aqueous speciation

Association constants for the first, second and third hydrolysis constants of Fe(II) calculated using SUPCRT and SPRONS.JNC between 0 and 350°C at P_{sat} are compared with their experimental counterparts measured by Sweeton and Baes (1970) and Tremaine and LeBlanc (1980) in Figs. D.24, D.25 and D.26, respectively. As expected, there is close agreement between calculated values and the experimental constants determined by Sweeton and Baes (1970) because these experimental data constrain the partial molal thermodynamic properties and equation-of-state coefficients of the ferrous hydroxide complexes in SPRONS.JNC. Significant discrepancies are evident, however, between calculated association constants and the experimental values determined by Tremaine and LeBlanc (1980). As noted above, there is no reason to believe the experimental data of Sweeton and Baes (1970) are more accurate than those of Tremaine and LeBlanc (1980), other than the fact that thermodynamic properties of the hydroxide complexes retrieved from the former dataset are more consistent with correlations developed by Shock *et al.* (1997a) than are the latter. Additional experimentation is needed to resolve the discrepancies between these two experimental datasets.

Sverjensky *et al.* (1997) document comparisons of experimental and calculated dissociation constants for FeCl^+ and $\text{FeCl}_2(aq)$ over a wide range of temperatures and pressures (Figs. 4 and 17, respectively, of Sverjensky *et al.*, 1997). The excellent agreement between calculated and experimental dissociation constants documented by Sverjensky *et al.* (1997) would be expected in similar calculations using SUPCRT and SPRONS.JNC, because the partial molal thermodynamic properties and equation-of-state coefficients for FeCl^+ and $\text{FeCl}_2(aq)$ in this database are from Sverjensky *et al.* (1997). This is confirmed in Fig. D.46, which illustrates comparisons between calculated and experimentally determined dissociation constants for FeCl^+ between 0 and 600°C at P_{sat} and at 0.5, 1 and 2 kb.

6.11 Stability Relations Among Smectites, and the Solubility of Illite

As noted in Section 5.2.4, thermodynamic data in SPRONS.JNC and PHREEQE.JNC support different models of hydrolysis reactions involving non-stoichiometric solid solutions. The PHREEQE series of codes do not include solid-solution models in the general sense considered by Helgeson *et al.* (1978), although ion-exchange reactions can be simulated. Rather, solid-solution behavior is approximated using a *surrogate* stoichiometric phase to represent the solid solution. The stoichiometry of the surrogate phase is defined such that it is within the range of compositions exhibited by the solid solution it represents,

but the stoichiometry is not variable, as would otherwise be consistent with solid-solution behavior. For this reason, estimated thermodynamic properties for a limited number of such phases, including those representing di-octahedral smectites (Na-, K-, Ca- and Mg- nontronites, beidellites and montmorillonites) and illite, are included in SPRONS.JNC. These data are estimated by Wolery (1976), who used a re-calibrated and slightly modified version of the technique described by Tardy and Garrels (1974) to estimate Gibbs free energies of formation, and methods described by Helgeson *et al.* (1978) to estimate standard molal entropies and Maier-Kelly heat capacity coefficients.

6.11.1 Smectites

Experimental data on the solubilities of clay minerals whose compositions are identical to those of the surrogate phases in SPRONS.JNC are not available, and the accuracy of the estimated thermodynamic data for these phases is therefore difficult to evaluate quantitatively. It is possible, however, to assess whether the estimated data are consistent with phase relations observed in natural systems. In this approach, logarithmic activity diagrams are used to correlate the chemistry of aqueous solutions with the stabilities of coexisting minerals at equilibrium. The idealized stoichiometry assigned in the present study to the smectite clays dictates the intercepts and slopes of boundaries on such diagrams, which in turn define the stability fields of these minerals. The compositions of natural waters plotting within a given stability field, or along one of its boundaries, is necessary, but not sufficient, evidence that the solution has equilibrated with the corresponding minerals.

Stability relations among minerals, including smectites, in the systems

- $\text{Na}_2\text{O}-\text{MgO}-\text{Al}_2\text{O}_3-\text{SiO}_2-\text{H}_2\text{O}$,
- $\text{K}_2\text{O}-\text{MgO}-\text{Al}_2\text{O}_3-\text{SiO}_2-\text{H}_2\text{O}$,
- $\text{Na}_2\text{O}-\text{K}_2\text{O}-\text{MgO}-\text{Al}_2\text{O}_3-\text{SiO}_2-\text{H}_2\text{O}$, and
- $\text{CaO}-\text{MgO}-\text{Al}_2\text{O}_3-\text{SiO}_2-\text{H}_2\text{O}$

are correlated, for example, with the compositions of deep Japanese groundwaters in Figs. C.55 - C.58, respectively. These diagrams are constructed using the Geochemist's Workbench (Bethke, 1996) and its associated thermodynamic database, which was modified in the present study to ensure that the thermodynamic properties of hydrolysis reactions among minerals and aqueous species are consistent with the data in SPRONS.JNC. The groundwater chemistry database, compiled by JNC, includes analytical data for approximately 15,000 samples from throughout Japan. These data were screened on the basis of sampling depth and temperature, and to discriminate among samples obtained at depths greater than, or less than, 200 m, to derive a database consisting of approximately 1000 samples that characterize the chemistry of "deep" groundwaters in Japan. These data were used with PHREEQE and the H3 TDB (see section 5.3.3) to calculate the logarithms of

activity quotients represented by the axes in Figs. C.55 - C.58. Stability relations among the minerals in these figures are calculated assuming the activity of $\text{SiO}_2(aq)$ is controlled by amorphous silica solubility at 19°C, which is the average temperature of the deep groundwater dataset.

As can be seen in Figs. C.55 - C.58, the deep groundwater compositions plot rather consistently in the stability fields of the montmorillonite end members of the smectite clays. Discernable trends in the groundwater data are also revealed suggesting that ion-exchange reactions among these clay minerals control the aqueous activities of Na^+ , K^+ , Ca^{2+} and Mg^{2+} . These trends are roughly parallel to the slopes of stability-field boundaries between montmorillonite end members, and in some cases are coincident with them (Figs. C.57 and C.58). Small adjustments in $\log K$ values for the reactions that determine these boundaries in Figs. C.55 and C.56 ($< 0.5 \log K$ units) are found by trial and error to generate better agreement with the groundwater data, but these adjustments are arbitrary, and have therefore not been incorporated in SPRONS.JNC.

The observations in Figs C.55 - C.58 that most of the deep groundwaters plot in the stability fields of clay minerals is qualitatively reasonable because reactions among groundwaters and primary aluminosilicate minerals are known to precipitate clay minerals as reaction products at low temperatures (e.g. Drever, 1982; Langmuir, 1997). It is not known, however, if such reactions actually occur in the host rocks from which the groundwater samples were obtained, nor is information available characterizing the mineralogy of these clay minerals, if they exist. Thus, although the thermodynamic data in SPRONS.JNC for the smectites appear to be compatible with known stability relations among minerals in groundwater systems, additional experimentation and field studies are needed to confirm this conclusion. Alternative models of smectite-water equilibria that more realistically account for the solid-solution behavior of these minerals (e.g., Aagaard and Helgeson, 1983) should also be tested using the available data in SPRONS.JNC for end-member components of smectite [e.g., stoichiometrically equivalent to pyrophyllite, margarite, muscovite, paragonite and others (Aagaard and Helgeson, 1983; Ransom and Helgeson, 1993; 1994)].

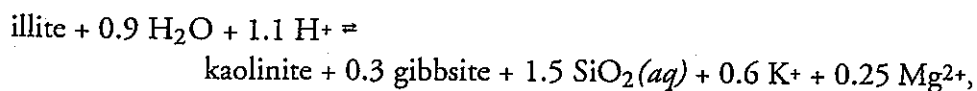
6.11.2 Illite

The solubility and stability of natural illites have been investigated experimentally by Routson and Kittrick (1971), Sass *et al.* (1987) (later reinterpreted by Aja, 1991), Aja *et al.* (1991a,b), and Yates and Rosenberg (1996). All the studies other than Aja *et al.* (1991b) fail to report Mg concentrations in the experimental solutions, however, and therefore cannot be used to test the reliability of thermodynamic data for illite in SPRONS.JNC.

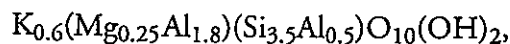
Aja *et al.* (1991b) measured the compositions of solutions coexisting with illite,

kaolinite, and gibbsite between 25 and 250°C at P_{sat} . Their interpretation of phase equilibria in this system is criticized by Essene and Peacor (1995), who assert that equilibrium was either not attained in the experiments, or that the composition of the illite solid-solution controlling aqueous compositions was mis-identified. Aja and Rosenberg (1996) argue that Essene and Peacor (1995) misinterpreted these experimental data, but this is disputed by Essene and Peacor (1997).

Despite these questions of interpretation, the basic experimental data generated by Aja *et al.* (1991b) can be used to test whether the thermodynamic data for illite in SPRONS.JNC are qualitatively reasonable. That is, although our assumption that illite's solid-solution behavior can be approximated using a stoichiometric surrogate phase is a gross simplification, solution compositions at, or near to, equilibrium with the illite surrogate, kaolinite and gibbsite should not conflict severely with solution compositions measured by Aja *et al.* (1991b). The relevant reaction is



where thermodynamic data for "stoichiometric" illite in SPRONS.JNC,



are estimated by Wolery (1976), as noted above.

Equilibrium constants calculated using SUPCRT and SPRONS.JNC for this reaction between 0 and 250°C at P_{sat} are compared in Fig. C.59 with equilibrium constants retrieved from experimental solution compositions by Aja *et al.* (1991b). The solid symbols refer to solutions sampled at the end of the experiments, and their open-symbol counterparts refer to solutions sampled at earlier times. As can be seen, the calculated equilibrium constants at relatively low temperatures are up to 3 log K units lower than their experimental counterparts, but these differences tend to decrease with increasing reaction time. At temperatures greater than 100°C, the calculated equilibrium constants are closely consistent with the experimental values determined at the end of the experiments. It is questionable whether gibbsite is stable, or metastable, in the experiments at temperatures greater than about 100°C (e.g., Pokrovskii and Helgeson, 1995), but accounting for boehmite rather than gibbsite in the above reaction (as shown by the dashed curve in Fig. C.59) does not significantly alter the close correspondence between calculated and experimental equilibrium constants at these temperatures, as can be seen in the figure.

The comparison of calculated and experimental equilibrium constants illustrated in Fig. C.59 suggests that the thermodynamic data for illite in SPRONS.JNC

are reasonably consistent with the limited amount of relevant experimental data over this range of temperatures and pressures. Although discrepancies at temperatures less than about 100°C are significant, they appear to diminish with increasing reaction time and temperature, which supports the contention of Essene and Peacor (1995) that equilibrium was not achieved in these experiments. We emphasize, however, that the reaction considered in this comparison is a gross approximation of illite's true solid-solution behavior, and that alternative modeling approaches that more realistically account for such behavior should be tested in similar comparisons with these experimental data.

6.12 Aqueous Dissociation Reactions

Experimental dissociation constants for Na^+ , K^+ , Mg^{2+} , Ca^{2+} , Ba^{2+} , Sr^{2+} , Fe^{2+} , and Mn^{2+} complexes with Cl^- , F^- , SO_4^{2-} , $\text{B}(\text{OH})_4^-$, H_3SiO_4^- or CH_3COO^- ligands are compared with corresponding dissociation constants calculated using SUPCRT and SPRONS.JNC in Figs. D.27 - D.50. Although most of these diagrams, and numerous others, are similarly described by Sverjensky *et al.* (1997), we include them in the present study because they incorporate experimental dissociation constants that were not considered by these authors. Comments summarized below address instances where these new and additional experimental data compare unfavorably with calculated values, although as can be seen in Figs. D.27 - D.50, calculated dissociation constants generally agree closely with their experimental counterparts.

Experimental dissociation constants for NaSO_4^- and $\text{NaB}(\text{OH})_4(\text{aq})$ determined by Pokrovski *et al.* (1995) between 0 and 300°C at P_{sat} are compared with values calculated using SUPCRT and SPRONS.JNC in Figs. D.29 and D.30, respectively. Partial molal thermodynamic properties and HKF equation-of-state coefficients for NaSO_4^- and $\text{NaB}(\text{OH})_4(\text{aq})$ in SPRONS.JNC were determined by regression of these experimental data, and the close correspondence between calculated and experimental dissociation constants in these figures is therefore expected. Dissociation constants for NaSO_4^- between 150 and 320°C at P_{sat} determined in flow-calorimetry experiments by Oscarson *et al.* (1988a) are generally consistent with their calculated counterparts, and with the experimental constants determined by Pokrovski *et al.* (1995).

Dissociation constants for the Na-acetate complex calculated using SUPCRT and SPRONS.JNC between 270 and 330°C at P_{sat} are compared with experimental values determined by Oscarson *et al.* (1988b) in Fig. D.31. The partial molal thermodynamic properties of this complex at 25°C and 1 bar in SPRONS.JNC are from Shock and Koretsky (1993), who used the correlation algorithm proposed by Shock and Helgeson (1988) to estimate corresponding equation-of-state coefficients. The poor agreement shown in the figure between calculated and experimental dissociation constants is attributed by Shock and Koretsky (1993) to incorrect dissociation constants for $\text{HCl}(\text{aq})$ used by

Oscarson *et al.* (1988b) to retrieve their dissociation constants for Na-acetate.

Dissociation constants for KSO_4^- calculated using SUPCRT and SPRONS.JNC between 0 and 200°C at P_{sat} are compared with experimental values for this reaction in Fig. D.33. Sverjensky *et al.* (1997) regressed experimental data from Truesdell and Hosterler (1968), Quist *et al.* (1963) and Bell and George (1953) (not shown) to retrieve partial molal thermodynamic properties and equation-of-state coefficients for this species. Dissociation constants determined by Righellato and Davies (1930) and Jenkins and Monk (1950) at low temperatures are consistent with the calculated curve in Fig. D.33, and the experimental data of Truesdell and Hosterler (1968) and Quist *et al.* (1963).

Dissociation constants for MgCl^+ calculated using SUPCRT and SPRONS.JNC are compared with their experimental counterparts from 0 to 300°C at P_{sat} in Fig. D.34, and from 300 to 600°C at higher pressures in Fig. D.35. Sverjensky *et al.* (1997) regressed experimental data from Majer and Stulik (1982) and Saccocia and Seyfried (1990) to retrieve partial molal thermodynamic properties and equation-of-state coefficients for this species. Dissociation constants determined by Gillespie *et al.* (1992) and Johnson and Pytkowicz (1978) are consistent with calculated values between 25 and 330°C at P_{sat} (Fig. D.34), but dissociation constants retrieved at higher temperatures and pressures from conductance measurements by Frantz and Marshall (1982) differ significantly (Fig. D.35).

Dissociation constants for MgF^+ between 0 and 300°C at P_{sat} calculated using SUPCRT and SPRONS.JNC are compared with experimental values in Fig. D.36. Sverjensky *et al.* (1997) regressed experimental data from Richardson and Holland (1979) to retrieve partial molal thermodynamic properties and equation-of-state coefficients for this species. Dissociation constants determined by Majer and Stulik (1982) are closely consistent with calculated values at low temperatures, but dissociation constants calculated by Richardson and Holland (1979) using experimental data from Cadek *et al.* (1971) and Tanner *et al.* (1968) are lower by less than 1 log K unit than dissociation constants calculated using SUPCRT and SPRONS.JNC.

Dissociation constants for CaCl^+ and $\text{CaCl}_2(aq)$ calculated using SUPCRT and SPRONS.JNC are compared with experimental values from 0 to 350°C at P_{sat} in Figs. D.37 and D.38, respectively, and from 400 to 600°C at higher pressures in Figs. D.39 and D.40, respectively. Sverjensky *et al.* (1997) regressed experimental data from Majer and Stulik (1982), Alakhverdov (1985), Simonson *et al.* (1985), Ramette (1986), Williams-Jones and Seward (1989) and Gillespie *et al.* (1992) to retrieve partial molal thermodynamic properties and equation-of-state coefficients for CaCl^+ . Experimental data from Ramette (1986) and Williams-Jones and Seward (1989) were used by Sverjensky *et al.* (1997) to retrieve partial molal thermodynamic properties and equation-of-state

coefficients for $\text{CaCl}_2(aq)$. Figures D.37 and D.38 confirm that dissociation constants for CaCl^+ and $\text{CaCl}_2(aq)$ determined in the respective experimental studies are consistent with calculated values between 25 and 360°C at P_{sat} , but dissociation constants retrieved at higher temperatures and pressures by Frantz and Marshall (1982) differ significantly (Fig. D.39 and D.40, respectively).

Dissociation constants for CaF^+ between 0 and 300°C at P_{sat} calculated using SUPCRT and SPRONS.JNC are compared with their experimental counterparts in Fig. D.41. Sverjensky *et al.* (1997) regressed experimental data from Richardson and Holland (1979) to retrieve partial molal thermodynamic properties and equation-of-state coefficients for this species. Dissociation constants determined Majer and Stulik (1982) are closely consistent with calculated values at low temperatures, but dissociation constants calculated by Richardson and Holland (1979) using experimental data from Tanner *et al.* (1968), Cadek *et al.* (1971) and Aziz and Lyle (1968) differ from values calculated using SUPCRT and SPRONS.JNC by less than 1 log K unit. Similar discrepancies are apparent in the solubility of fluorite between 0 and 260°C at P_{sat} determined by Richardson and Holland (1979) (Fig. C.54).

Dissociation constants for MnCl^+ between 0 and 300°C at P_{sat} calculated using SUPCRT and SPRONS.JNC are compared with experimental dissociation constants reported by Wheat and Carpenter (1988) and Gammons and Seward (1996) in Fig. D.49. Sverjensky *et al.* (1997) regressed these experimental data using the partial molal thermodynamic properties and HKF equation-of-state coefficients for Cl^- from Shock and Helgeson (1988) to retrieve partial molal thermodynamic properties and equation-of-state coefficients for MnCl^+ . The source of thermodynamic properties for Mn^{2+} used in this regression procedure is unclear, however, because these data were not determined by Shock and Helgeson (1988). Partial molal thermodynamic properties and equation-of-state parameters for Mn^{2+} from Shock *et al.* (1997a), which are included in SPRONS.JNC, apparently differ from the corresponding data used by Sverjensky *et al.* (1997) because the calculated curve in Fig. D.41 is displaced upward by 0.3 to 0.4 log K units relative to the corresponding curve depicted in Fig. 17 of Sverjensky *et al.* (1997). A similar displacement is noted in the calculated curve shown in Fig. D.50 for $\text{MnSO}_4(aq)$ dissociation compared with the corresponding curve in Fig. 14 of Sverjensky *et al.* (1997), who used the data from Wheat and Carpenter (1988) as experimental constraints. These observations suggest that the thermodynamic data for MnCl^+ and $\text{MnSO}_4(aq)$ in SPRONS.JNC should be revised by regression of the experimental data in Figs. D.49 and D.50, using partial molal thermodynamic properties and equation-of-state coefficients for Cl^- from Shock and Helgeson (1988), and for Mn^{2+} from Shock *et al.* (1997a).

6.13 Comments on Database Management

We conclude this section with comments on the conditions adopted for a reliable thermodynamic database (Section 3.1) with regard to SPRONS.JNC, and, hence, PHREEQE.JNC.

6.13.1 Internal consistency

SPRONS.JNC is *partially* consistent in the sense defined by Engi (1992), because all the data are compatible with basic thermodynamic definitions and functional relations comprising the SUPCRT model, and fundamental experimental and field observations that constrain these data are consistently evaluated within this modeling framework. The databases are not *fully* consistent, however, because the rationale for selecting certain parameters, such as standard molal or partial molal properties at the reference pressure and temperature, are not generally documented in sufficient detail in original data sources.

The internal consistency of SPRONS.JNC should be maintained as this database is revised. Revisions to the thermodynamic properties of boehmite, diaspore and corundum recommended by Pokrovskii and Helgeson (1995), for example, have resulted in less favorable agreement among calculated and experimental equilibria involving these minerals and various aluminosilicates than were achieved in the original evaluation of these data by Helgeson *et al.* (1978). Similar discrepancies can be expected if the revised partial molal properties of $\text{SiO}_2(\text{aq})$ determined by Rimstidt (1997) are incorporated in SPRONS.JNC. Resolving such discrepancies for a given reaction is reasonably straightforward using regression techniques, but matrix analytical methods may be required if multiple reactions are affected by the revised property.

6.13.2 Experimental basis

Additional experimental determination of mineral solubilities at relatively low pressures and temperatures (≈ 0 to 150°C at P_{sat}) are needed to improve confidence that the thermodynamic data in SPRONS.JNC are reliable for conditions that are most relevant to the Japanese repository concept. The thermodynamic properties of minerals in this database are constrained primarily by phase-equilibrium studies and a limited amount of calorimetric data. These properties should be integrated with those of aqueous species and gases, and results compared with independent experimental solubility equilibria. Recent studies of this kind, which are remarkably few in number, are described by Sverjensky *et al.* (1991) and Pokrovskii and Helgeson (1995), and provide a useful framework for similar work by JNC and others in the future. Based on the evaluations discussed in preceding sections, similar studies are recommended to revise standard molal thermodynamic properties in SPRONS.JNC for:

- albite,
- strontianite,
- witherite,
- siderite,
- celestite,
- pyrrhotite,
- magnetite, and
- fluorite.

Evaluations of data reliability such as are described in Sections 6.1 - 6.12 should also be undertaken routinely as the results of new solubility and aqueous-speciation studies become available.

Experimental investigations of mineral solubilities at relatively low temperatures are often complicated by sluggish reaction kinetics. To overcome this problem, experiments at higher temperatures are useful because reaction-rates generally increase with increasing temperature, and because thermodynamic relations incorporated in SUPCRT enable extrapolation of the high-temperature data to lower temperatures with reasonable confidence. We also recommend, however, that field investigations of phase relations in natural groundwater systems should be given greater priority for purposes of evaluating the reliability of thermodynamic data. The temperature in such systems is low, but reaction times greatly exceed those that can be accommodated experimentally. Detailed characterization of the mineralogy and groundwater chemistry comprising such systems is required to extract meaningful data that can be used to test the accuracy of thermodynamic data. Such characterization should take into account the effects of metastability, differences in crystallinity and the state of order/disorder in minerals formed under different conditions in different geologic environments, and the effects of solid-solution on mineral compositions.

The minerals in SPRONS.JNC are stoichiometric phases, but as noted above several non-stoichiometric minerals, such as the clays, zeolites and chlorites, are also important in geologic and near-field environments considered in the Japanese repository concept. The compositions of such minerals are variable within limits, and result from ideal or non-ideal mixing of thermodynamic components in solid solution. Helgeson *et al.* (1978), Stoessell (1979), Aagaard and Helgeson (1983) and Giggenbach (1985) describe models of this behavior in which the activities of components computed from site-mixing approximations are correlated with the compositions of the minerals and used as descriptive variables in activity diagrams. The approach is advantageous because it circumvents the need to assign values for the standard Gibbs free energies of formation to one or another solid-solution mineral with a specified composition. SPRONS.JNC includes data for several minerals that are stoichiometrically equivalent to thermodynamic components used in these models to constrain the thermodynamic behavior of solid-solution minerals. This suggests that the reliability of these data could be evaluated by comparing predicted mineral

stability relations in various chemical subsystems with phase relations observed in natural systems.

6.13.3 Documentation

The documentation criteria adopted in the present study are that the selected thermodynamic data are *traceable* to original experimental sources, and that these sources are *scientifically defensible* and *reproducible*. The traceability criterion is met by referencing sources from which the standard molal and partial molal properties in SPRONS.JNC were obtained, and these sources in turn identify sources of primary experimental results that were used to retrieve these data. The sources of thermodynamic data in SPRONS.JNC are summarized in Sections 5.2.2 and 5.2.3, and are referenced in Section 8. The sources of experimental constraints used to retrieve these data are identified in these references, and most are also documented in the present report. The scientific defensibility and reproducibility of the primary experimental results are documented in the original references.

6.13.4 Verification of accuracy

The accuracy of thermodynamic data in SPRONS.JNC is verified in the present study by comparing calculated thermodynamic quantities with their experimental counterparts, which are documented in the various diagrams comprising Appendixes C and D. Similar diagrams, and others, constructed using identical thermodynamic data are documented in the references summarized in Sections 5.2.2 and 5.2.3. It is recommended that a library of such diagrams should also be developed by JNC, and that these should be updated and revised as the results of new experimental investigations become available.

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Appendix A: Listing of the SPRONS.JNC Database

(see Tables 5.1_1 - 5.1_3 for formatting and unit conventions)

*** sprons.jnc -s[sequential access] pro[properties of] n[natural
 s[ubstances] database developed for the Japan Nuclear Cycle
 Development Institute (JNC).
 Sequential-access thermodynamic datafile to be used as input
 to program cprons.a code that converts sprons.jnc into its direct
 -access equivalent(dprons.jnc), which is used by program supcrt.
 *** NOTE: species can be added, deleted, or modified using program
 mprons.
 *** last updated: 20.Jan.98 by r. arthur
 *** 1.Apr.98: s2o3-2 data corrected (r. arthur)
 *** 14.Apr.98: corrected names of species HB02,aq, B02- and NaH2B03,aq to
 B(OH)3,aq B(OH)4- and NaB(OH)4,aq,respectively, in accordance
 ref 39 (r. arthur)

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minerals that do not undergo phase transitions

AKERMANITE	Ca ₂ MgSi ₂ O ₇			
AK	Ca(2)Mg(1)Si(2)O(7)			
ref:1,19	15.Mar.90			
	-879362.0	50.030	92.810	
	60.090000	11.400000	-11.400000	
	1700.00			
ALABANDITE	MnS			
Alb	Mn(1)S(1)			
ref:1	5.May.78			
	-51000.0	19.200	21.460	
	11.400000	1.800000	.000000	
	1803.00			
ALBITE, LOW	Na(AlSi ₃)O ₈			
Lo-Ab	Na(1)Al(1)Si(3)O(8)			
ref:1	5.May.78			
	-939680.0	49.510	100.070	
	61.700000	13.900000	-15.010000	
	1400.00			
ALUNITE	KAl ₃ (OH)6(SO ₄) ₂			
Alu	K(1)Al(3)S(2)O(14)H(6)			
ref:1	5.May.78			
	-1235600.0	78.400	293.600	
	153.450000	.000000	-54.950000	
	650.00			
AMESITE, 7A	Mg ₂ Al(AlSi) ₅ O ₅ (OH) ₄			
Ams-7A	Mg(2)Al(1)Si(1)O(9)H(4)			
ref:1	5.May.78			
	999999.0	52.000	103.000	
	81.030000	24.738000	-20.230000	
	1000.00			
AMORPHOUS-SILICA	SiO ₂ *NH ₂ O			
Amor-Si	Si(1)O(2)H(2N)O(N)			
ref:1	5.May.78			
	-202892.0	14.340	29.000	
	5.930000	47.200000	-22.780000	
	622.00			
ANALCIME	NaAlSi ₂ O ₆ *H ₂ O			
Anl	Na(1)Al(1)Si(2)O(7)H(2)			
ref:1	5.May.78			
	-738098.0	56.000	97.100	
	53.490000	24.140000	-8.880000	
	1000.00			
ANALCIME, DEHYDRATED	NaAlSi ₂ O ₆			
Dhyd-Anl	Na(1)Al(1)Si(2)O(6)			
ref:1	5.May.78			
	-674989.0	41.900	89.100	
	42.090000	24.140000	-8.880000	
	1000.00			
ANDALUSITE	Al ₂ SiO ₅			
And	Al(2)Si(1)O(5)			
ref:1	5.May.78			
	-580587.0	22.200	51.530	
	41.310833	6.292556	-12.392063	
	1043.00			
ANDRADITE	Ca ₃ Fe ₂ Si ₃ O ₁₂			
Adr	Ca(3)Fe(2)Si(3)O(12)			
ref:1,19	15.Mar.90			
	-1296819.0	70.130	131.850	
	113.532000	15.636000	-30.889000	
	1100.00			
ANGLESITE	PbSO ₄			
Ang	Pb(1)S(1)O(4)			
ref:1	5.May.78			
	-194353.0	35.510	47.950	
	10.960000	31.000000	4.200000	
	1100.00			
ANHYDRITE	CaSO ₄			

Anh	Ca(1)S(1)O(4)			
ref:1	5.May.78			
	-315925.0	25.500	45.940	
	16.780000	23.600000	.000000	
	1453.00			
ANNITE	KFe ₃ (AlSi ₃)O ₁₀ (OH) ₂			
Ann	K(1)Fe(3)Al(1)Si(3)O(12)H(2)			
ref:1	5.May.78			
	-1147156.0	95.200	154.320	
	106.430000	29.770000	-19.310000	
	1000.00			
ANORTHITE	Ca(Al ₂ Si ₂)O ₈			
An	Ca(1)Al(2)Si(2)O(8)			
ref:1,19	15.Mar.90			
	-954078.0	49.100	100.790	
	63.311000	14.794000	-15.440000	
	1700.00			
ARAGONITE	CaCO ₃			
Arg	Ca(1)C(1)O(3)			
ref:1,18	9.Mar.90			
	-269683.0	21.560	34.150	
	20.130000	10.240000	-3.340000	
	600.00			
ARTINITE	Mg ₂ (OH) ₂ (CO ₃)*3H ₂ O			
Art	Mg(2)C(1)O(8)H(8)			
ref:1	5.May.78			
	-613915.0	55.670	96.900	
	70.870000	27.660000	-7.430000	
	1000.00			
AZURITE	Cu ₃ (OH) ₂ (CO ₃) ₂			
Azr	Cu(3)C(2)O(8)H(2)			
ref:1	5.May.78			
	-334417.0	66.970	91.010	
	36.880000	77.440000	-.920000	
	780.00			
BARITE	BaSO ₄			
Brt	Ba(1)S(1)O(4)			
ref:1	5.May.78			
	-325563.0	31.600	52.100	
	33.800000	.000000	-8.430000	
	1422.00			
BEIDELLITE, CA	Ca _{0.165} Al _{2.33} Si _{3.67} O ₁₀ (OH) ₂			
Ca-Bei	Ca(0.165)Al(2.33)Si(3.67)O(12)H(2)			
ref:32	14.Jan.98			
	-1280909.0	58.300	129.530	
	82.191000	37.150000	-18.030000	
	600.00			
BEIDELLITE, K	K _{0.33} Al _{2.33} Si _{3.67} O ₁₀ (OH) ₂			
K-Bei	K(0.33)Al(2.33)Si(3.67)O(12)H(2)			
ref:32	14.Jan.98			
	-1281773.0	60.510	133.700	
	83.319000	38.400000	-17.920000	
	600.00			
BEIDELLITE, MG	Mg _{0.165} Al _{2.33} Si _{3.67} O ₁₀ (OH) ₂			
Mg-Bei	Mg(0.165)Al(2.33)Si(3.67)O(12)H(2)			
ref:32	14.Jan.98			
	-1277064.0	57.700	123.190	
	81.945000	37.260000	-18.020000	
	600.00			
BEIDELLITE, NA	Na _{0.33} Al _{2.33} Si _{3.67} O ₁₀ (OH) ₂			
Na-Bei	Na(0.33)Al(2.33)Si(3.67)O(12)H(2)			
ref:32	14.Jan.98			
	-1279688.0	59.620	130.540	
	83.277000	37.780000	-18.250000	
	600.00			
BERNDTITE	SnS ₂			
Brn	Sn(1)S(2)			
ref:12	5.May.78			
	-34750.0	20.900	40.960	
	15.510000	4.200000	.000000	

	1000.00			
BOEHMITE	AlO(OH)			
Bhm	Al(1)O(2)H(1)			
ref:20	29.Apr.97			
	-219599.0	8.890	19.535	
	12.902000	12.421000	-3.228000	
	900.00			
BROMELLITE	BeO			
Brm	Be(1)O(1)			
ref:1	5.May.78			
	999999.0	3.380	8.309	
	8.450000	4.000000	-3.170000	
	1200.00			
BRUCITE	Mg(OH) ₂			
Brc	Mg(1)O(2)H(2)			
ref:1	5.May.78			
	-199646.0	15.090	24.630	
	24.147000	-221390.0	-6.110000	
	900.00	4.012400		
CA-AL-PYROXENE	CaAl(AlSi) ₆			
Ca-Al-Px	Ca(1)Al(2)Si(1)O(6)			
ref:1,19	15.Mar.90			
	-742067.0	35.000	63.500	
	54.130000	-783793.0	-14.900000	
	1400.00	6.420000		
CALCITE	CaCO ₃			
Cal	Ca(1)C(1)O(3)			
ref:1,18	9.Mar.90			
	-269880.0	22.150	36.934	
	24.980000	-288552.0	-6.200000	
	1200.00	5.240000		
CASSITERITE	SnO ₂			
Cst	Sn(1)O(2)			
ref:12	5.May.78			
	-124260.0	12.500	21.550	
	17.245000	-138800.0	-4.900100	
	1500.00	2.802600		
CELADONITE	K(MgAl)Si ₄ O ₁₀ (OH) ₂			
Cln	K(1)Mg(1)Al(1)Si(4)O(12)H(2)			
ref:1,30	9.Jan.98			
	-1305895.0	74.900	157.100	
	80.250000	-1394899.0	-18.540000	
	1000.00	25.300000		
CELESTITE	SrSO ₄			
Cls	Sr(1)S(1)O(4)			
ref:1	5.May.78			
	-320435.0	28.000	46.250	
	21.800000	-347300.0	.000000	
	1500.00	13.300000		
CERUSSITE	PbCO ₃			
Car	Pb(1)C(1)O(3)			
ref:1	5.May.78			
	-150370.0	31.300	40.590	
	12.390000	-168000.0	.000000	
	800.00	28.600000		
CHABAZITE	Ca(Al ₂ Si ₄)O ₁₂ *6H ₂ O			
Cbz	Ca(1)Al(2)Si(4)O(18)H(12)			
ref:1	5.May.78			
	999999.0	152.900	247.760	
	146.000000	44.470000	-16.430000	
	1000.00			
CHALCEDONY	SiO ₂			
Cha	Si(1)O(2)			
ref:1	5.May.78			
	-204276.0	9.880	22.688	
	11.220000	-217282.0	-2.700000	
	848.00	8.200000		
CHAMOSITE, 7A	Fe ₂ Al(AlSi) ₅ O ₅ (OH) ₄			
Chm-7A	Fe(2)Al(2)Si(1)O(9)H(4)			
ref:1,30	9.Jan.98			

	-828026.0	-902403.0	64.700	106.200
	84.910000	25.400000	-18.770000	
	1000.00			
CHLORARGYRITE	AgCl			
Crg	Ag(1)Cl(1)			
ref:1,15	5.May.78			
	-26247.0	-30370.0	23.000	25.727
	14.330000	1.821000	-2.430000	
	728.00			
CHLORITOID	FeAl ₂ SiO ₅ (OH) ₂			
Cld	Fe(1)Al(2)Si(1)O(7)H(2)			
ref:1	5.May.78			
	999999.0	999999.0	42.100	69.800
	60.630000	17.130000	-13.630000	
	1000.00			
CHRYSSOTILE	Mg ₃ Si ₂ O ₅ (OH) ₄			
Ctl	Mg(3)Si(2)O(9)H(4)			
ref:1	5.May.78			
	-964871.0	-1043123.0	52.900	108.500
	75.820000	31.600000	-17.580000	
	1000.00			
CINNABAR	HgS			
Cin	Hg(1)S(1)			
ref:1	5.May.78			
	-10940.0	-12750.0	19.700	28.416
	10.460000	3.720000	.000000	
	618.00			
CLINOCHLORE, 14A	Mg ₅ Al(AlSi ₃)O ₁₀ (OH) ₈			
14A-Cnc	Mg(5)Al(2)Si(3)O(18)H(8)			
ref:1	5.May.78			
	-1961703.0	-2116964.0	111.200	207.110
	166.500000	42.100000	-37.470000	
	900.00			
CLINOZOISITE	Ca ₂ Al ₃ Si ₃ O ₁₂ (OH)			
Czo	Ca(2)Al(3)Si(3)O(13)H(1)			
ref:1,19	15.Mar.90			
	-1549240.0	-1643781.0	70.640	136.200
	106.118000	25.214000	-27.145000	
	700.00			
COPPER, NATIVE	Cu			
Cu	Cu(1)			
ref:1	5.May.78			
	.0	.0	7.923	7.113
	5.410000	1.500000	.000000	
	1357.00			
CORDIERITE	Mg ₂ Al ₃ (AlSi ₅)O ₁₈			
Crd	Mg(2)Al(4)Si(5)O(18)			
ref:1	5.May.78			
	-2061279.0	-2183199.0	97.330	233.220
	143.830000	25.800000	-38.600000	
	1700.00			
CORDIERITE, HYDROUS	Mg ₂ Al ₃ (AlSi ₅)O ₁₈ *H ₂ O			
Hyd-Crd	Mg(2)Al(4)Si(5)O(19)H(2)			
ref:1	5.May.78			
	-2121350.0	-2255676.0	111.430	241.220
	155.230000	25.800000	-38.600000	
	1700.00			
CORUNDUM	Al ₂ O ₃			
Crn	Al(2)O(3)			
ref:20	05.May.98			
	-378167.0	-400500.0	12.170	25.575
	29.321000	1.644000	-12.0070000	
	1200.00			
COVELLITE	CuS			
Cv	Cu(1)S(1)			
ref:1	5.May.78			
	-12612.0	-12500.0	15.900	20.420
	10.600000	2.640000	.000000	
	1273.00			
CRISTOBALITE, ALPHA	SiO ₂			
A-Crs	Si(1)O(2)			
ref:1	5.May.78			

	-203895.0	-216755.0	10.372	25.740
	13.980000	3.340000	-3.810000	
	1000.00			
CRISTOBALITE, BETA	SiO ₂			
B-Crs	Si(1)O(2)			
ref:1	5.May.78			
	-203290.0	-215675.0	11.963	27.380
	17.390000	.310000	-9.890000	
	2000.00			
CUMINGTONITE	Mg ₇ Si ₈ O ₂₂ (OH) ₂			
Cum	Mg(7)Si(8)O(24)H(2)			
ref:1	5.May.78			
	999999.0	999999.0	127.500	271.900
	185.240000	58.610000	-44.660000	
	800.00			
CUPRITE	Cu ₂ O			
Cpr	Cu(2)O(1)			
ref:1	5.May.78			
	-35384.0	-40830.0	22.080	23.437
	14.080000	5.880000	-.760000	
	1515.00			
DAPHNITE, 14A	Fe ₅ Al(AlSi ₃)O ₁₀ (OH) ₈			
Dph-14A	Fe(5)Al(2)Si(3)O(18)H(8)			
ref:1,29	9.Jan.98			
	-1551548.0	-1696591.0	142.500	213.420
	176.210000	43.760000	-33.820000	
	1000.00			
DIASPORE	AlO(OH)			
Dsp	Al(1)O(2)H(1)			
ref:20	29.Apr.97			
	-220150.0	-238924.0	8.446	17.760
	12.273000	13.070000	-2.903000	
	900.00			
DICKITE	Al ₂ Si ₂ O ₅ (OH) ₄			
Dck	Al(2)Si(2)O(9)H(4)			
ref:1,7	5.May.78			
	999999.0	999999.0	47.100	99.300
	72.770000	29.200000	-21.520000	
	1000.00			
DIOPSIDE	CaMg(SiO ₃) ₂			
Di	Ca(1)Mg(1)Si(2)O(6)			
ref:1,19	15.Mar.90			
	-723780.0	-765378.0	34.200	66.090
	52.870000	7.840000	-15.740000	
	1600.00			
DOLOMITE	CaMg(CO ₃) ₂			
Dol	Ca(1)Mg(1)C(2)O(6)			
ref:1,19	15.Mar.90			
	-517760.0	-556631.0	37.090	64.365
	41.557000	23.952000	-9.884000	
	1000.00			
DOLOMITE, DISORDERED	CaMg(CO ₃) ₂			
Dis-Dol	Ca(1)Mg(1)C(2)O(6)			
ref:1,19	15.Mar.90			
	-515653.0	-553704.0	39.840	64.390
	44.711000	17.779000	-10.948000	
	1000.00			
DOLOMITE, ORDERED	CaMg(CO ₃) ₂			
Ord-Dol	Ca(1)Mg(1)C(2)O(6)			
ref:1,19	15.Mar.90			
	-517760.0	-556631.0	37.090	64.340
	44.711000	17.779000	-10.948000	
	1000.00			
EPIDOTE	Ca ₂ FeAl ₂ Si ₃ O ₁₂ (OH)			
Ep	Ca(2)Fe(1)Al(2)Si(3)O(13)H(1)			
ref:1,19	15.Mar.90			
	-1450906.0	-1543992.0	75.280	139.200
	117.622000	12.816000	-31.864000	
	1100.00			
EPIDOTE, ORDERED	Ca ₂ FeAl ₂ Si ₃ O ₁₂ (OH)			
Ord-Ep	Ca(2)Fe(1)Al(2)Si(3)O(13)H(1)			
ref:1,19	15.Mar.90			

	-1450906.0	-1544016.0	75.200	139.200
	113.780000	14.690000	-28.920000	
	1100.00			
FAYALITE	Fe ₂ SiO ₄			
Fa	Fe(2)Si(1)O(4)			
ref:1	5.May.78			
	-330233.0	-354119.0	35.450	46.390
	36.510000	9.360000	-6.700000	
	1490.00			
FERROPARGASITE	Na(Ca ₂ Fe ₄ Al)(Al ₂ Si ₆)O ₂₂ (OH) ₂			
Fe-Prg	Na(1)Ca(2)Fe(4)Al(3)Si(6)O(24)H(2)			
ref:1	5.May.78			
	999999.0	999999.0	185.500	279.890
	213.570000	42.990000	-47.290000	
	1000.00			
FERROTREMOLITE	(Ca ₂ Fe ₅)Si ₈ O ₂₂ (OH) ₂			
Fe-Tr	Ca(2)Fe(5)Si(8)O(24)H(2)			
ref:1	5.May.78			
	999999.0	999999.0	163.500	282.800
	197.930000	58.950000	-41.170000	
	800.00			
FERROUS-OXIDE	FeO			
Frs-Ox	Fe(1)O(1)			
ref:1	5.May.78			
	-60097.0	-65020.0	14.520	12.000
	12.122000	2.072000	-.750000	
	1600.00			
FLUORITE	CaF ₂			
Fl	Ca(1)F(2)			
ref:1	5.May.78			
	-280493.0	-293000.0	16.390	24.542
	14.300000	7.280000	.470000	
	1424.00			
FLUORPHLOGOPITE	KMg ₃ (AlSi ₃)O ₁₀ (F) ₂			
F-Phl	K(1)Mg(3)Al(1)Si(3)O(10)F(2)			
ref:1,7	5.May.78			
	999999.0	999999.0	80.380	146.370
	96.130000	26.740000	-20.410000	
	1000.00			
FLUOTREMOLITE	(Ca ₂ Mg ₅)Si ₈ O ₂₂ (F) ₂			
F-Tr	Ca(2)Mg(5)Si(8)O(22)F(2)			
ref:1	5.May.78			
	999999.0	999999.0	129.400	270.450
	183.740000	55.250000	-43.730000	
	800.00			
FORSTERITE	Mg ₂ SiO ₄			
For	Mg(2)Si(1)O(4)			
ref:1	5.May.78			
	-491564.0	-520000.0	22.750	43.790
	35.810000	6.540000	-8.520000	
	1800.00			
GALENA	PbS			
Gn	Pb(1)S(1)			
ref:1	5.May.78			
	-23115.0	-23500.0	21.800	31.490
	11.170000	2.200000	.000000	
	1388.00			
GEHLENITE	Ca ₂ (Al ₂ Si) ₇ O ₇			
Gh	Ca(2)Al(2)Si(1)O(7)			
ref:1,19	15.Mar.90			
	-903148.0	-951225.0	48.100	90.240
	63.740000	8.000000	-15.120000	
	1800.00			
GIBBSITE	Al(OH) ₃			
Gbs	Al(1)O(3)H(3)			
ref:20	29.Apr.97			
	-276025.0	-309065.0	16.360	31.956
	13.073000	40.696000	-2.920000	
	600.00			
GLAUCOPHANE	Na ₂ (Mg ₃ Al ₂)Si ₈ O ₂₂ (OH) ₂			
Gln	Na(2)Mg(3)Al(2)Si(8)O(24)H(2)			
ref:1	5.May.78			

999999.0	999999.0	130.000	269.700
190.260000	59.360000	-50.010000	
800.00			
GOLD, NATIVE	Au		
ref:1	Au(1)		
	5.May.78		
		11.330	10.215
5.660000	1.240000	.000000	
1336.00			
GRAPHITE	C		
ref:1	C(1)		
	5.May.78		
		1.372	5.298
4.030000	1.140000	-2.040000	
2500.00			
GREENALITE	Fe3Si2O5(OH)4		
ref:1,30	Fe(3)Si(2)O(9)H(4)		
	9.Jan.98		
		72.600	115.000
-716501.0	-787774.0	-15.390000	
81.650000	32.600000		
1000.00			
GROSSULAR	Ca3Al2Si3O12		
ref:1,19	Ca(3)Al(2)Si(3)O(12)		
	15.Mar.90		
		60.870	125.300
-1496307.0	-1582737.0	-27.318000	
104.017000	17.013000		
1000.00			
GRUNERITE	Fe7Si8O22(OH)2		
ref:1	Fe(7)Si(8)O(24)H(2)		
	5.May.78		
		163.300	256.670
999999.0	999999.0	-39.550000	
198.830000	60.930000		
800.00			
HALITE	NaCl		
ref:1	Na(1)Cl(1)		
	5.May.78		
		17.240	27.015
-91807.0	-98260.0	.000000	
10.980000	3.900000		
1073.00			
HALLOYSITE	Al2Si2O5(OH)4		
ref:1,7,28	Al(2)Si(2)O(9)H(4)		
	9.Jan.98		
		48.520	99.300
-903628.2	-980294.7	-21.520000	
72.770000	29.200000		
1000.00			
HEDENBERGITE	CaFe(SiO3)2		
ref:1,19	Ca(1)Fe(1)Si(2)O(6)		
	5.Mar.90		
		40.700	68.270
-638998.0	-678276.0	-15.010000	
54.810000	8.170000		
1600.00			
HUNTITE	CaMg3(CO3)4		
ref:1	Ca(1)Mg(3)C(4)O(12)		
	5.May.78		
		71.590	122.900
-1004710.0	-1082600.0	-20.440000	
84.170000	42.860000		
1000.00			
HYDROMAGNESITE	Mg5(OH)2(CO3)4*4H2O		
ref:1	Mg(5)C(4)O(18)H(10)		
	5.May.78		
		129.380	208.800
-1401687.0	-1557090.0	-21.670000	
141.460000	65.280000		
1000.00			
ILLITE	K0.6Mg0.25Al1.8Al0.5Si3.5O10(OH)2		
ref:32	K(0.6)Mg(0.25)Al(2.3)Si(3.5)O(12)H(2)		
	14.Jan.98		
		63.590	138.940
-1303971.0	-1394709.0	-17.820000	
86.044000	38.570000		
600.00			
JADEITE	NaAl(SiO3)2		
ref:1	Na(1)Al(1)Si(2)O(6)		
	5.May.78		

-679445.0	-722116.0	31.900	60.400
48.160000	11.420000	-11.870000	
1400.00			
K-FELDSPAR	K(AlSi3)O8		
ref:1,31	K(1)Al(1)Si(3)O(8)		
	9.Jan.98		
		51.130	108.870
-895610.0	-949424.0	-29.945000	
76.617000	4.311000		
1400.00			
KAOLINITE	Al2Si2O5(OH)4		
ref:1	Al(2)Si(2)O(9)H(4)		
	5.May.78		
		48.530	99.520
-905614.0	-982221.0	-21.520000	
72.770000	29.200000		
1000.00			
KYANITE	Al2SiO5		
ref:1	Al(2)Si(1)O(5)		
	5.May.78		
		20.000	44.090
-580956.0	-616897.0	-12.882104	
41.393132	6.816463		
466.00			
LAUMONTITE	Ca(Al2Si4)O12*4H2O		
ref:1,19	Ca(1)Al(2)Si(4)O(16)H(8)		
	15.Mar.90		
		116.100	207.550
-1596823.0	-1728664.0	-16.430000	
123.200000	44.470000		
1000.00			
LEONHARDITE	Ca2(Al4Si8)O24*7H2O		
ref:1	Ca(2)Al(4)Si(8)O(31)H(14)		
	5.May.78		
		220.400	407.860
999999.0	999999.0	-32.860000	
235.000000	88.940000		
1000.00			
LIME	CaO		
ref:1	Ca(1)O(1)		
	5.May.78		
		9.500	16.764
-144366.0	-151790.0	-1.560000	
11.670000	1.080000		
2000.00			
MAGNESITE	MgCO3		
ref:1	Mg(1)C(1)O(3)		
	5.May.78		
		15.700	28.018
-245658.0	-265630.0	-4.748000	
19.731000	12.539000		
1000.00			
MALACHITE	Cu2(OH)2(CO3)		
ref:1	Cu(2)C(1)O(5)H(2)		
	5.May.78		
		44.500	54.860
-214204.0	-251900.0	-1.340000	
27.760000	43.780000		
780.00			
MANGANOSITE	MnO		
ref:1,6	Mn(1)O(1)		
	5.May.78		
		14.270	13.221
-86735.0	-92070.0	-8.80000	
11.110000	1.940000		
1800.00			
MERWINITE	Ca3Mg(SiO4)2		
ref:1,19	Ca(3)Mg(1)Si(2)O(8)		
	15.Mar.90		
		60.500	104.400
-1036526.0	-1090796.0	-14.440000	
72.970000	11.960000		
1700.00			
METACINNABAR	HgS		
ref:1	Hg(1)S(1)		
	5.May.78		
		21.200	30.169
-10437.0	-11800.0	.000000	
10.520000	3.630000		
1000.00			
MICROCLINE, MAXIMUM	K(AlSi3)O8		
ref:1,31	K(1)Al(1)Si(3)O(8)		
	9.Jan.98		

-895610.0	-949424.0	51.130	108.741
63.830000	12.900000	-17.050000	
1400.00			
MINNESOTAITE	Fe3Si4O10(OH)2		
ref:1,30	Fe(3)Si(4)O(12)H(2)		
	9.Jan.98		
		83.500	147.860
-1070048.0	-1153366.0	-11.150000	
88.310000	42.610000		
800.00			
MONTICELLITE	CaMgSiO4		
ref:1	Ca(1)Mg(1)Si(1)O(4)		
	5.May.78		
		26.400	51.470
-512829.0	-540800.0	-8.000000	
36.820000	5.340000		
1400.00			
MONTMORILLONITE, CA	Ca0.165Mg0.33Al1.67Si4O10(OH)2		
ref:34	Ca(0.165)Mg(0.33)Al(1.67)Si(4)O(12)H(2)		
	14.Jan.98		
		59.893	133.264
-1272340.0	-1361490.0	-16.650000	
80.180000	39.500000		
600.00			
MONTMORILLONITE, K	K0.33Mg0.33Al1.67Si4O10(OH)2		
ref:34	K(0.33)Mg(0.33)Al(1.67)Si(4)O(12)H(2)		
	15.Jan.98		
		62.103	137.432
-1273300.0	-1362823.0	-16.530000	
84.320000	40.750000		
600.00			
MONTMORILLONITE, MG	Mg0.495Al1.67Si4O10(OH)2		
ref:34	Mg(0.495)Al(1.67)Si(4)O(12)H(2)		
	15.Jan.98		
		59.131	131.247
-1268590.0	-1357862.0	-16.630000	
79.940000	39.610000		
600.00			
MONTMORILLONITE, NA	Na0.33Mg0.33Al1.67Si4O10(OH)2		
ref:34	Na(0.33)Mg(0.33)Al(1.67)Si(4)O(12)H(2)		
	15.Jan.98		
		61.213	134.271
-1271210.0	-1360684.0	-16.870000	
81.270000	40.130000		
600.00			
MUSCOVITE	KAl2(AlSi3)O10(OH)2		
ref:1,31	K(1)Al(3)Si(3)O(12)H(2)		
	9.Jan.98		
		70.000	140.710
-1335667.0	-1427644.0	-25.440000	
97.560000	26.380000		
1000.00			
NEPHELINE	Na(AlSi)O4		
ref:1	Na(1)Al(1)Si(1)O(4)		
	5.May.78		
		29.720	54.160
-472872.0	-500241.0	-7.327965	
35.908062	6.457918		
1500.00			
NONTRONITE, CA	Ca0.165Fe2Al0.33Si3.67O10(OH)2		
ref:32	Ca(0.165)Fe(2)Al(0.33)Si(3.67)O(12)H(2)		
	14.Jan.98		
		66.350	131.100
-1079492.0	-1166692.0	-13.200000	
78.191000	52.930000		
600.00			
NONTRONITE, K	K0.33Fe2Al0.33Si3.67O10(OH)2		
ref:32	K(0.33)Fe(2)Al(0.33)Si(3.67)O(12)H(2)		
	14.Jan.98		
		68.570	135.270
-1080356.0	-1167926.0	-13.090000	
79.319000	54.180000		
600.00			
NONTRONITE, MG	Mg0.165Fe2Al0.33Si3.67O10(OH)2		
ref:32	Mg(0.165)Fe(2)Al(0.33)Si(3.67)O(12)H(2)		
	14.Jan.98		
		65.740	129.760
-1075647.0	-1162924.0	-13.190000	
77.945000	53.040000		
600.00			
NONTRONITE, NA	Na0.33Fe2Al0.33Si3.67O10(OH)2		
ref:32	Na(0.33)Fe(2)Al(0.33)Si(3.67)O(12)H(2)		
	14.Jan.98		

-1078271.0 79.277000 600.00	-1165798.0 53.560000	67.660 132.110 -13.420000	
PALLADIUM Pd ref:38	Pd Pd(1) 20.Jan.98		
5.800000 973.00	1.380000	8.990 8.850 .000000	
Pd-OXIDE Pd-Ox ref:38	PdO Pd(1)O(1) 20.Jan.98		
-14400.0 3.300000 973.00	-20400.0 14.200000	13.200 14.070 .000000	
PARAGONITE Pg ref:1	NaAl2(AlSi3)O10(OH)2 Na(1)Al(3)Si(3)O(12)H(2) 5.May.78		
-1326012.0 97.430000 1000.00	-1416963.0 24.500000	66.400 132.530 -26.440000	
PARGASITE Prg ref:1,19	Na(Ca2Mg4Al)(Al2Si6)O22(OH)2 Na(1)Ca(2)Mg(4)Al(3)Si(6)O(24)H(2) 15.Mar.90		
-2846728.0 205.800000 1000.00	-3016624.0 41.660000	160.000 273.500 -50.210000	
PERICLASE Per ref:1	MgO Mg(1)O(1) 5.May.78		
-136086.0 10.180000 2100.00	-143800.0 1.740000	6.440 11.248 -1.480000	
PHLOGOPITE Phl ref:1,31	KMg3(AlSi3)O10(OH)2 K(1)Mg(3)Al(1)Si(3)O(12)H(2) 9.Jan.98		
-1396423.0 100.610000 1000.00	-1488303.0 28.780000	76.100 149.660 -21.500000	
POTASSIUM-OXIDE K-Ox ref:1	K2O K(2)O(1) 5.May.78		
-77056.0 18.510000 1100.00	-86800.0 8.650000	22.500 40.380 -1.880000	
PYRITE Py ref:1	FeS2 Fe(1)S(2) 5.May.78		
-38293.0 17.880000 1000.00	-41000.0 1.320000	12.650 23.940 -3.050000	
PYROPHYLLITE Pri ref:1	Al2Si4O10(OH)2 Al(2)Si(4)O(12)H(2) 5.May.78		
-1255997.0 79.432000 800.00	-1345313.0 39.214000	57.200 126.600 -17.282000	
RHODOCHROSITE Rds ref:1	MnCO3 Mn(1)C(1)O(3) 5.May.78		
-195045.0 21.990000 700.00	-212521.0 9.300000	23.900 31.075 -4.690000	
RICHTERITE Rct ref:1	Na2(CaMg5)Si8O22(OH)2 Na(2)Ca(1)Mg(5)Si(8)O(24)H(2) 5.May.78		
999999.0 194.800000 800.00	999999.0 61.100000	137.600 272.800 -46.150000	
ROMARCHITE Sn-Ox ref:12	SnO Sn(1)O(1) 5.May.78		

-61459.0 9.550000 1237.00	-68340.0 3.500000	13.660 20.895 .000000	
RUTILE Ru ref:7,9	TiO2 Ti(1)O(2) Mar.83		
-212883.0 15.014420 1800.00	-226101.0 2.720501	12.020 18.820 -2.365364	
SANIDINE,HIGH Hi-Sa ref:1,31	K(AlSi3)O8 K(1)Al(1)Si(3)O(8) 9.Jan.98		
-893974.0 63.830000 1400.00	-946774.0 12.900000	54.530 109.008 -17.050000	
SEPIOLITE Sep ref:1	Mg4Si6O15(OH)2(H2O)2*4H2O Mg(4)Si(6)O(23)H(14) 5.May.78		
-2211192.0 157.920000 800.00	-2418000.0 104.300000	146.600 285.600 -18.680000	
SIDERITE Sd ref:1,13	FeCO3 Fe(1)C(1)O(3) 5.May.78		
-162414.0 11.630000 885.00	-179173.0 26.800000	25.100 29.378 .000000	
SILLIMANITE Sil ref:1	Al2SiO5 Al(2)Si(1)O(5) 5.May.78		
-580091.0 40.023988 1200.00	-615099.0 7.390511	23.130 49.900 -11.674053	
SILVER,NATIVE Ag ref:1	Ag Ag(1) 5.May.78		
5.090000 1234.00	2.040000	10.170 10.272 .360000	
SMITHSONITE Smt ref:1	ZnCO3 Zn(1)C(1)O(3) 5.May.78		
-174850.0 9.300000 780.00	-194260.0 33.000000	19.700 28.275 .000000	
SODIUM-OXIDE Na-Ox ref:1	Na2O Na(2)O(1) 5.May.78		
-89883.0 18.250000 1000.00	-99140.0 4.890000	17.935 25.000 -2.890000	
SPHALERITE Sp ref:1	ZnS Zn(1)S(1) 5.May.78		
-47947.0 11.770000 1300.00	-49000.0 1.260000	14.019 23.830 -1.160000	
SPINEL Spn ref:1	MgAl2O4 Mg(1)Al(2)O(4) 5.May.78		
-517006.0 36.772808 2000.00	-546847.0 6.415456	19.270 39.710 -9.708792	
STAUFOLITE St ref:1	Fe2Al9Si4O23(OH) Fe(2)Al(9)Si(4)O(24)H(1) 5.May.78		
999999.0 207.120000 1000.00	999999.0 36.940000	117.100 224.400 -57.220000	
STRONTIANITE Str ref:1	SrCO3 Sr(1)C(1)O(3) 5.May.78		

-275470.0 23.520000 1197.00	-294600.0 6.320000	23.200 39.010 -5.080000	
SYLVITE Sy ref:1	KCl K(1)Cl(1) 5.May.78		
-97735.0 9.890000 1043.00	-104370.0 5.200000	19.730 37.524 .770000	
TALC Tcl ref:1	Mg3Si4O10(OH)2 Mg(3)Si(4)O(12)H(2) 5.May.78		
-1320188.0 82.482000 800.00	-1410920.0 41.614000	62.340 136.250 -13.342000	
TENORITE Tn ref:1	CuO Cu(1)O(1) 5.May.78		
-30568.0 11.530000 1600.00	-37200.0 1.880000	10.180 12.220 -1.760000	
TITANITE Ti ref:7,9,19	CaTiSiO5 Ca(1)Ti(1)Si(1)O(5) Mar.90		
-587129.0 42.234220 1670.00	-620960.0 5.703951	30.880 55.650 -9.525038	
TREMOLITE Tr ref:1,19	(Ca2Mg5)Si8O22(OH)2 Ca(2)Mg(5)Si(8)O(24)H(2) 15.Mar.90		
-2770245.0 188.222000 800.00	-2944038.0 57.294000	131.190 272.920 -44.822000	
WAIKAKITE Wai ref:1,19	Ca(Al2Si4)O12*2H2O Ca(1)Al(2)Si(4)O(14)H(4) 15.Mar.90		
-1477432.0 100.400000 1000.00	-1579333.0 44.470000	105.100 186.870 -16.430000	
URANINITE Uran ref:37	UO2 U(1)O(2) 18.Jan.98		
-246610.0 15.003000 573.15	-259320.0 7.586000	18.400 24.638 -1.839000	
WITHERITE Wth ref:1	BaCO3 Ba(1)C(1)O(3) 5.May.78		
-278400.0 21.500000 1079.00	-297500.0 11.060000	26.800 45.810 -3.910000	
WOLLASTONITE Wo ref:1,19	CaSiO3 Ca(1)Si(1)O(3) 15.Mar.90		
-369225.0 26.640000 1400.00	-389590.0 3.600000	19.600 39.930 -6.520000	
WURTZITE Wur ref:1	ZnS Zn(1)S(1) 5.May.78		
-44810.0 11.820000 1300.00	-45850.0 1.160000	14.064 23.846 -1.040000	
ZINCITE Znc ref:1	ZnO Zn(1)O(1) 5.May.78		
999999.0 11.710000 2000.00	999999.0 1.220000	10.430 14.338 -2.180000	
ZOISITE Zo ref:1,19	Ca2Al3Si3O12(OH) Ca(2)Al(3)Si(3)O(13)H(1) 15.Mar.90		

-1549179.0	-1643691.0	70.740	135.900
106.118000	25.214000	-27.145000	
700.00			

minerals that undergo one phase transition			

ALBITE	Na(AlSi3)O8		
Ab	Na(1)Al(1)Si(3)O(8)		
ref:1	5.May.78		
-886308.0	-939680.0	49.510	100.250
61.700000	13.900000	-15.010000	
473.00	999999.0	999999.000	999999.000
81.880000	3.554000	-50.154000	
1200.00			
ALBITE, HIGH	NaAlSi3O8		
Hi-Ab	Na(1)Al(1)Si(3)O(8)		
ref:1	5.May.78		
-884509.0	-937050.0	52.300	100.430
61.700000	13.900000	-15.010000	
623.00	999999.0	999999.000	999999.000
64.170000	13.900000	-15.010000	
1400.00			
ALMANDINE	Fe3Al2Si3O12		
Alm	Fe(3)Al(2)Si(3)O(12)		
ref:1	5.May.78		
999999.0	999999.0	75.600	115.280
97.520000	33.640000	-18.730000	
848.00	999999.0	999999.000	999999.000
107.090000	14.860000	-10.630000	
1600.00			
AMESITE, 14A	Mg4Al2(Al2Si2)O10(OH)8		
14A-Am	Mg(4)Al(4)Si(2)O(18)H(8)		
ref:1	5.May.78		
999999.0	999999.0	108.900	205.400
172.590000	34.980000	-41.670000	
848.00	999999.0	999999.000	999999.000
169.400000	41.240000	-44.370000	
1000.00			
ANTIGORITE	Mg48Si134085(OH)62		
Atg	Mg(48)Si(34)O(147)H(62)		
ref:1	5.May.78		
-15808020.0	-17070891.0	861.360	1749.130
1228.450000	513.760000	-286.680000	
848.00	999999.0	999999.000	999999.000
1234.830000	501.240000	-281.280000	
1000.00			
CLINOCHLORE, 7A	Mg5Al(AlSi3)O10(OH)8		
7A-Cnc	Mg(5)Al(2)Si(3)O(18)H(8)		
ref:1	5.May.78		
-1957101.0	-2113197.0	106.500	211.500
162.820000	50.620000	-40.880000	
848.00	999999.0	999999.000	999999.000
166.010000	44.360000	-38.180000	
900.00			
COESITE	SiO2		
Coz	Si(1)O(2)		
ref:1	5.May.78		
-203541.0	-216614.0	9.650	20.641
11.000000	8.200000	-2.700000	
848.00	999999.0	999999.000	999999.000
14.190000	1.940000	.000000	
2000.00			
CRISTOBALITE	SiO2		
Crs	Si(1)O(2)		
ref:1	5.May.78		
-203895.0	-216755.0	10.372	25.740
13.980000	3.340000	-3.810000	
543.00	321.0	999999.000	999999.000
17.390000	.310000	-9.890000	
2000.00			
DAPHNITE, 7A	Fe5Al(AlSi3)O10(OH)8		
7A-Dph	Fe(5)Al(2)Si(3)O(18)H(8)		

ref:1,30	9.May.98		
-1544331.0	-1689504.0	138.900	221.200
172.530000	52.280000	-37.230000	
848.00	999999.0	999999.000	999999.000
175.720000	46.020000	-34.630000	
1000.00			
EDENITE	Na(Ca2Mg5)(AlSi7)O22(OH2)		
Ed	Na(1)Ca(2)Mg(5)Al(1)Si(7)O(24)H(2)		
ref:1	5.May.78		
999999.0	999999.0	161.000	271.000
199.710000	48.780000	-46.010000	
848.00	999999.0	999999.000	999999.000
202.900000	42.520000	-43.310000	
1000.00			
EPISTILBITE	Ca(Al2Si6)O16*5H2O		
Eps	Ca(1)Al(2)Si(6)O(21)H(10)		
ref:1	5.May.78		
999999.0	999999.0	152.900	267.560
157.040000	60.870000	-21.830000	
848.00	999999.0	999999.000	999999.000
163.420000	48.350000	-16.430000	
1000.00			
FERROEDENITE	Na(Ca2Fe5)(AlSi7)O22(OH)2		
Fe-Ed	Na(1)Ca(2)Fe(5)Al(1)Si(7)O(24)H(2)		
ref:1	5.May.78		
999999.0	999999.0	193.700	280.900
209.420000	50.440000	-42.360000	
848.00	999999.0	999999.000	999999.000
212.610000	44.180000	-39.660000	
1000.00			
FERROSILITE	FeSiO3		
Fe	Fe(1)Si(1)O(3)		
ref:1,13	5.May.78		
-267588.0	-285658.0	21.630	32.952
26.490000	5.070000	-5.550000	
413.00	37.0	.056	67.000
23.865000	8.780000	-4.700000	
1400.00			
FLUOREDENITE	Na(Ca2Mg5)(AlSi7)O22(F)2		
Fl-Ed	Na(1)Ca(2)Mg(5)Al(1)Si(7)O(22)F(2)		
ref:1	5.May.78		
999999.0	999999.0	146.700	272.500
195.230000	46.740000	-44.920000	
848.00	999999.0	999999.000	999999.000
198.420000	40.480000	-42.220000	
1000.00			
HERZENBERGITE	SnS		
Hrz	Sn(1)S(1)		
ref:12	5.May.78		
-25023.0	-25464.0	18.360	29.010
8.530000	7.480000	.900000	
875.00	160.0	999999.000	999999.000
9.780000	3.740000	.000000	
1153.00			
HEULANDITE	Ca(Al2Si7)O18*6H2O		
Hul	Ca(1)Al(2)Si(7)O(24)H(12)		
ref:1	5.May.78		
999999.0	999999.0	182.400	316.370
179.660000	69.070000	-24.530000	
848.00	999999.0	999999.000	999999.000
189.230000	50.290000	-16.430000	
1000.00			
KALSILITE	K(AlSi)O4		
Kls	K(1)Al(1)Si(1)O(4)		
ref:1	5.May.78		
-481750.0	-509408.0	31.850	59.890
29.430000	17.360000	-5.320000	
810.00	160.0	999999.000	999999.000
42.500000	.000000	.000000	
1800.00			
LAWSONITE	CaAl2Si2O7(OH)2*H2O		
Lws	Ca(1)Al(2)Si(2)O(10)H(4)		

ref:1,19	15.Mar.90		
-1073408.0	-1158104.0	55.800	101.320
81.800000	23.360000	-16.260000	
848.00	999999.0	999999.000	999999.000
84.990000	17.100000	-13.560000	
1000.00			
MAGNETITE	Fe3O4		
Mag	Fe(3)O(4)		
ref:1	5.May.78		
-242574.0	-267250.0	34.830	44.524
21.880000	48.200000	.000000	
900.00	999999.0	999999.000	999999.000
48.000000	.000000	.000000	
1800.00			
MARGARITE	CaAl2(Al2Si2)O10(OH)2		
Mrg	Ca(1)Al(4)Si(2)O(12)H(2)		
ref:1,19	15.Mar.90		
-1394150.0	-1485803.0	63.800	129.400
102.500000	16.350000	-28.050000	
848.00	999999.0	999999.000	999999.000
99.310000	22.610000	-30.750000	
1000.00			
NATROLITE	Na2(Al2Si3)O10*2H2O		
Ntr	Na(2)Al(2)Si(3)O(12)H(4)		
ref:1	5.May.78		
999999.0	999999.0	101.600	169.720
95.760000	40.080000	-15.060000	
848.00	999999.0	999999.000	999999.000
92.570000	46.340000	-17.760000	
1000.00			
NESQUEHONITE	MgCO3*3H2O		
Nsh	Mg(1)C(1)O(6)H(6)		
ref:1	5.May.78		
-412035.0	-472576.0	46.760	74.790
-1574.804000	3899.173000	-417.325000	
306.50	184.0	999999.000	999999.000
25.246000	91.289000	-4.222000	
340.00			
NICKEL	Ni		
Ni	Ni(1)		
ref:1	5.May.78		
.0	.0	7.140	6.588
4.060000	7.040000	.000000	
633.00	.0	999999.000	999999.000
6.000000	1.800000	.000000	
1725.00			
PHILLIPSITE, CA	Ca(Al2Si5)O14*5H2O		
Ca-Php	Ca(1)Al(2)Si(5)O(19)H(10)		
ref:1	5.May.78		
999999.0	999999.0	166.600	265.000
145.820000	52.670000	-19.130000	
848.00	999999.0	999999.000	999999.000
149.010000	46.410000	-16.430000	
1000.00			
PHILLIPSITE, K	K2(Al2Si5)O14*5H2O		
K-Php	K(2)Al(2)Si(5)O(19)H(10)		
ref:1	5.May.78		
999999.0	999999.0	172.100	265.000
152.660000	60.240000	-18.450000	
848.00	999999.0	999999.000	999999.000
155.850000	53.980000	-15.750000	
1000.00			
PHILLIPSITE, NA	Na2(Al2Si5)O14*5H2O		
Na-Php	Na(2)Al(2)Si(5)O(19)H(10)		
ref:1	5.May.78		
999999.0	999999.0	172.400	265.000
152.400000	56.480000	-20.460000	
848.00	999999.0	999999.000	999999.000
155.590000	50.220000	-17.760000	
1000.00			
PREHNITE	Ca2Al2Si3O10(OH)2		
Prh	Ca(2)Al(2)Si(3)O(12)H(2)		

ref:1,19 15.Mar.90
 -1390097.0 -1481649.0 65.000 140.330
 91.600000 37.820000 -19.600000
 848.00 999999.0 999999.000 999999.000
 101.170000 19.040000 -11.500000
 1000.00

PYROPE
 Prp
 ref:1,7 Mg3Al2Si3O12
 Mg(3)Al(2)Si(3)O(12)
 5.May.78
 999999.0 999999.0 62.300 113.270
 91.690000 32.640000 -20.920000
 848.00 999999.0 999999.000 999999.000
 101.260000 13.860000 -12.820000
 1800.00

QUARTZ
 Qtz
 ref:1 SiO2
 Si(1)O(2)
 5.May.78
 -204646.0 -217650.0 9.880 22.688
 11.220000 8.200000 -2.700000
 848.00 290.0 .372 38.500
 14.410000 1.940000 .000000
 2000.00

QUICKSILVER
 Hg
 ref:1 Hg(1)
 5.May.78
 .0 .0 18.170 14.822
 6.440000 .000000 .190000
 629.00 14140.0 999999.000 999999.000
 4.970000 .000000 .000000
 3000.00

SPESSARTINE
 Sps
 ref:1 Mn3Al2Si3O12
 Mn(3)Al(2)Si(3)O(12)
 5.May.78
 999999.0 999999.0 74.500 117.900
 94.480000 33.240000 -19.120000
 848.00 999999.0 999999.000 999999.000
 104.050000 14.460000 -11.020000
 1800.00

STILBITE
 Stb
 ref:1 NaCa2(Al5Si11)O36·14H2O
 Na(1)Ca(2)Al(5)Si(13)O(50)H(28)
 5.May.78
 999999.0 999999.0 399.300 649.910
 390.550000 137.680000 -49.840000
 848.00 999999.0 999999.000 999999.000
 400.120000 118.900000 -41.740000
 1000.00

TIN,NATIVE
 Sn
 ref:12 Sn(1)
 Aug.85
 .0 .0 12.240 16.289
 4.420000 6.300000 .000000
 505.06 1680.0 999999.000 999999.000
 7.300000 .000000 .000000
 3000.00

 minerals that undergo two phase transitions

ACANTHITE
 Acn
 ref:1 Ag2S
 Ag(2)S(1)
 5.May.78
 -9446.0 -7550.0 34.300 34.200
 15.630000 8.600000 .000000
 450.00 950.0 999999.000 999999.000
 1.819000 53.000000 .000000
 620.00 600.0 999999.000 999999.000
 21.600000 .000000 .000000
 1000.00

AEGERINE
 Agt
 ref:1 NaFe(SiO3)2
 Na(1)Fe(1)Si(2)O(6)
 5.May.78
 999999.0 999999.0 36.700 64.370
 46.160000 19.310000 -9.460000

950.00 999999.0 999999.000 999999.000
 52.420000 10.010000 -7.680000
 1050.00 999999.0 999999.000 999999.000
 50.270000 10.890000 -7.680000
 1400.00

ANTHOPHYLLITE
 Ath
 ref:1 Mg7Si8O22(OH)2
 Mg(7)Si(8)O(24)H(2)
 5.May.78
 -2715430.0 -2888749.0 128.600 264.400
 180.682000 60.574000 -38.462000
 903.00 999999.0 999999.000 999999.000
 197.542000 41.614000 -13.342000
 1258.00 999999.0 999999.000 999999.000
 199.522000 41.614000 -13.342000
 1800.00

BORNITE
 Bn
 ref:1,16 Cu5FeS4
 Cu(5)Fe(1)S(4)
 5.May.78
 -86704.0 -79922.0 99.290 98.600
 49.760000 35.080000 -1.350000
 485.00 1430.0 999999.000 999999.000
 -34.310000 247.000000 .000000
 540.00 .0 999999.000 999999.000
 80.330000 -2.040000 .000000
 1200.00

BUNSENITE
 Bsn
 ref:1 NiO
 Ni(1)O(1)
 5.May.78
 -50573.0 -57300.0 9.080 10.970
 -4.990000 37.580000 3.890000
 525.00 .0 999999.000 999999.000
 13.880000 .000000 .000000
 565.00 .0 999999.000 999999.000
 11.180000 2.020000 .000000
 2000.00

CHALCOCITE
 Cc
 ref:1 Cu2S
 Cu(2)S(1)
 5.May.78
 -20626.0 -19000.0 28.900 27.480
 12.630000 18.820000 .000000
 376.00 920.0 999999.000 999999.000
 26.780000 -7.350000 .000000
 717.00 287.0 999999.000 999999.000
 20.320000 .000000 .000000
 1400.00

CHALCOPYRITE
 Ccp
 ref:1,16 CuFeS2
 Cu(1)Fe(1)S(2)
 5.May.78
 -44900.0 -44453.0 31.150 42.830
 20.790000 12.800000 -1.340000
 830.00 2405.0 999999.000 999999.000
 -141.400000 210.000000 .000000
 930.00 .0 999999.000 999999.000
 41.220000 .000000 .000000
 1200.00

CRONSTEDTITE, 7A
 7A-Csd
 ref:1 Fe2Fe(FeSi)O5(OH)4
 Fe(4)Si(1)O(9)H(4)
 5.May.78
 999999.0 999999.0 73.500 110.900
 84.790000 41.840000 -12.476000
 950.00 999999.0 999999.000 999999.000
 97.300000 23.240000 -8.930000
 1050.00 999999.0 999999.000 999999.000
 93.010000 25.000000 -8.930000
 1500.00

ENSTATITE
 En
 ref:1 MgSiO3
 Mg(1)Si(1)O(3)
 5.May.78
 -348930.0 -369686.0 16.200 31.276
 24.550000 4.740000 -6.280000
 903.00 166.0 .020 384.615
 28.765000 .000000 .000000

1258.00 390.0 1.090 11.900
 29.260000 .000000 .000000
 1800.00

FERROGEDRITE
 Fe-Ged
 ref:1 (Fe5Al2)(Al2Si6)O22(OH)2
 Fe(5)Al(4)Si(6)O(24)H(2)
 5.May.78
 999999.0 999999.0 153.500 265.900
 196.270000 58.390000 -39.010000
 903.00 999999.0 999999.000 999999.000
 204.700000 48.910000 -26.450000
 1258.00 999999.0 999999.000 999999.000
 205.690000 48.910000 -26.450000
 1500.00

HASTINGSITE
 Hs
 ref:1 Na(Ca2Fe4Fe)(Al2Si6)O22(OH)2
 Na(1)Ca(2)Fe(5)Al(2)Si(6)O(24)H(2)
 5.May.78
 999999.0 999999.0 189.200 280.300
 211.570000 50.880000 -44.880000
 950.00 999999.0 999999.000 999999.000
 217.820000 41.580000 -43.100000
 1050.00 999999.0 999999.000 999999.000
 215.680000 42.460000 -43.100000
 1500.00

HEMATITE
 Hem
 ref:1 Fe2O3
 Fe(2)O(3)
 5.May.78
 -178155.0 -197720.0 20.940 30.274
 23.490000 18.600000 -3.550000
 950.00 160.0 999999.000 999999.000
 36.000000 .000000 .000000
 1050.00 .0 999999.000 999999.000
 31.710000 1.760000 .000000
 1800.00

LARNITE
 Lrn
 ref:1,8 Ca2SiO4
 Ca(2)Si(1)O(4)
 5.May.78
 999999.0 999999.0 30.500 51.600
 34.870000 9.740000 -6.260000
 970.00 440.0 999999.000 999999.000
 32.160000 11.020000 .000000
 1710.00 3390.0 999999.000 999999.000
 49.000000 .000000 .000000
 2000.00

MAGNESIOHASTINGSITE
 Mg-Hs
 ref:1 Na(Ca2Mg4Fe)(Al2Si6)O22(OH)2
 Na(1)Ca(2)Mg(4)Fe(1)Al(2)Si(6)O(24)H(2)
 5.May.78
 999999.0 999999.0 163.800 273.800
 203.800000 49.550000 -47.800000
 950.00 999999.0 999999.000 999999.000
 210.060000 40.250000 -46.020000
 1050.00 999999.0 999999.000 999999.000
 207.910000 41.130000 -46.020000
 1500.00

MAGNESIOBIEBECKITE
 Mg-Rbk
 ref:1 Na2(Mg3Fe2)Si8O22(OH)2
 Na(2)Mg(3)Fe(2)Si(8)O(24)H(2)
 5.May.78
 999999.0 999999.0 138.000 271.300
 186.260000 75.140000 -45.180000
 950.00 999999.0 999999.000 999999.000
 198.770000 56.540000 -41.630000
 1050.00 999999.0 999999.000 999999.000
 194.480000 58.300000 -41.630000
 1500.00

PYRRHOTITE
 Po
 ref:1 FeS
 Fe(1)S(1)
 5.May.78
 -24084.0 -24000.0 14.410 18.200
 5.190000 26.400000 .000000
 411.00 570.0 999999.000 999999.000
 17.400000 .000000 .000000
 598.00 120.0 999999.000 999999.000
 12.200000 2.380000 .000000

1468.00			
RIEBECKITE	Na ₂ (Fe ₃ Fe ₂)Si ₈ O ₂₂ (OH) ₂		
Rbk	Na(2)Fe(5)Si(8)O(24)H(2)		
ref:1	5.May.78		
999999.0	999999.0	156.600	274.900
192.090000	76.140000	-42.990000	
950.00	999999.0	999999.000	999999.000
204.600000	57.540000	-39.440000	
1050.00	999999.0	999999.000	999999.000
200.310000	59.300000	-39.440000	
1500.00			

minerals that undergo three phase transitions			

IRON	Fe		
Fe	Fe(1)		
ref:7,8	5.May.78		
.0	.0	6.520	7.092
3.040000	7.580000	.600000	
1033.00	326.0	999999.000	999999.000
11.130000	.000000	.000000	
1183.00	215.0	999999.000	999999.000
5.800000	1.980000	.000000	
1673.00	165.0	999999.000	999999.000
6.740000	1.600000	.000000	
1812.00			
PD-OXYANNITE	KFe ₃ (AlSi ₃)O ₁₀ (OH)O-		
Pd-Oxn	K(1)Fe(3)Al(1)Si(3)O(12)H(1)		
ref:1	5.May.78		
999999.0	999999.0	68.500	143.200
88.340000	54.120000	-18.060000	
848.00	999999.0	999999.000	999999.000
97.910000	35.340000	-9.960000	
950.00	999999.0	999999.000	999999.000
116.680000	7.440000	-4.630000	
1050.00	999999.0	999999.000	999999.000
110.240000	10.080000	-4.630000	
1100.00			

gases			

Ar,g	ARGON		
Ar(g)	Ar(1)		
ref:6,8	27.Dec.89		
.0	.0	37.008	.000
4.968000	.000000	.000000	
8000.00			
CH ₄ ,g	METHANE		
CH ₄ (g)	C(1)H(4)		
ref:6,8	15.Dec.87		
-12122.4	-17880.0	44.518	.000
5.650000	11.440000	-4.600000	
1500.00			
CO ₂ ,g	CARBON-DIOXIDE		
CO ₂ (g)	C(1)O(2)		
ref:6,8	27.Dec.89		
-94254.0	-94051.0	51.085	.000
10.570000	2.100000	-2.060000	
2500.00			
H ₂ ,g	HYDROGEN		
H ₂ (g)	H(2)		
ref:6,8	27.Dec.89		
.0	.0	31.234	.000
6.520000	.780000	.120000	
3000.00			
H ₂ O,g	STEAM		
H ₂ O(g)	H(2)O		
ref:17	24.Mar.90		
-54524.8	-57935.1	44.763	.000
12.664700	-10.404100	.013078	
2523.15			
H ₂ S,g	HYDROGEN-SULFIDE		

H ₂ S(g)	H(2)S(1)		
ref:6,8	27.Dec.89		
-8021.0	-4931.0	49.185	.000
7.810000	2.960000	-4.460000	
2300.00			
He,g	HELIUM		
He(g)	He(1)		
ref:6,8	27.Dec.89		
.0	.0	30.151	.000
4.968000	.000000	.000000	
8000.00			
Kr,g	KRYPTON		
Kr(g)	Kr(1)		
ref:6,8	27.Dec.89		
.0	.0	39.217	.000
4.968000	.000000	.000000	
8000.00			
N ₂ ,g	NITROGEN		
N ₂ (g)	N(2)		
ref:6,8	27.Dec.89		
.0	.0	45.796	.000
6.830000	.900000	-1.200000	
3000.00			
NH ₃ ,g	AMMONIA		
NH ₃ (g)	N(1)H(3)		
ref:6,8	26.Jul.89		
-3932.0	-11021.0	46.000	.000
7.110000	6.000000	-3.700000	
1800.00			
Ne,g	NEON		
Ne(g)	Ne(1)		
ref:6,8	27.Dec.89		
.0	.0	34.973	.000
4.968000	.000000	.000000	
8000.00			
O ₂ ,g	OXYGEN		
O ₂ (g)	O(2)		
ref:6,8	27.Dec.89		
.0	.0	49.029	.000
7.160000	1.000000	-4.400000	
3000.00			
Rn,g	RADON		
Rn(g)	Rn(1)		
ref:6,8	27.Dec.89		
.0	.0	42.120	.000
4.968000	.000000	.000000	
3000.00			
S ₂ ,g	SULFUR		
S ₂ (g)	S(2)		
ref:6,8	27.Dec.89		
18953.0	30681.0	54.540	.000
8.720000	.160000	-1.900000	
3000.00			
SO ₂ ,g	SULFUR-DIOXIDE		
SO ₂ (g)	S(1)O(2)		
ref:6,8	27.Dec.89		
-71748.0	-70944.0	59.330	.000
11.040000	1.880000	-1.840000	
2000.00			
Xe,g	XENON		
Xe(g)	Xe(1)		
ref:6,8	27.Dec.89		
.0	.0	40.555	.000
4.968000	.000000	.000000	
8000.00			

aqueous species			

FORMIC-ACID,aq	H ₂ CO ₂		
FOAC(aq)	C(1)H(2)O(2)+(0)		
ref:26	17.Jun.92		

-88982.	-101680.	38.900	
6.3957	7.7713	2.8318	-3.1002
26.1000	3.1000	-0.3300	0.
FORMATE	HCO ₂		
FOR-	H(1)C(1)O(2)-(1)		
ref:26	28.Feb.92		
-83862.	-101680.	21.700	
5.7842	4.7242	7.3630	-2.9742
17.0000	-12.4000	1.3003	-1.
ACETIC-ACID,aq	C ₂ H ₄ O ₂		
ACAC(aq)	C(2)H(4)O(2)+(0)		
ref:26	6.Mar.92		
-94760.	-116100.	42.700	
11.6198	5.2180	2.5088	-2.9946
42.0760	-1.5417	-0.1500	0.
ACETATE	C ₂ H ₃ O ₂		
AE-	C(2)H(3)O(2)-(1)		
ref:26	28.Feb.92		
-88270.	-116160.	20.600	
7.7525	8.6996	7.5825	-3.1385
26.3000	-3.8600	1.3182	-1.
PROPANOIC-ACID,aq	C ₃ H ₆ O ₂		
PRAC(aq)	C(3)H(6)O(2)+(0)		
ref:26	6.Mar.92		
-93450.	-122470.	49.400	
11.0057	18.7077	-0.7792	-3.5523
62.9700	-1.1900	-0.1500	0.
PROPANOATE	C ₃ H ₅ O ₂		
PRO-	C(3)H(5)O(2)-(1)		
ref:26	28.Feb.92		
-86780.	-122630.	26.500	
9.6992	12.1344	9.0612	-3.2805
52.3000	-4.2000	1.2276	-1.
BUTANOIC-ACID,aq	C ₄ H ₈ O ₂		
BUAC(aq)	C(4)H(8)O(2)+(0)		
ref:26	6.Mar.92		
-91210.	-127950.	56.100	
13.2702	23.3487	-0.6985	-3.7441
72.4342	2.9924	-0.2155	0.
BUTANOATE	C ₄ H ₇ O ₂		
BUT-	C(4)H(7)O(2)-(1)		
ref:26	17.Jun.92		
-84650.	-128630.	31.800	
11.7724	17.0492	7.4458	-3.4837
62.3135	-3.5666	1.1469	-1.
PENTANOIC-ACID,aq	C ₅ H ₁₀ O ₂		
PEAC(aq)	C(5)H(10)O(2)+(0)		
ref:26	6.Mar.92		
-89240.	-133690.	62.800	
15.4596	28.3156	-1.8374	-3.9495
91.0659	5.2450	-0.1710	0.
PENTANOATE	C ₅ H ₉ O ₂		
N-PEN-	C(5)H(9)O(2)-(1)		
ref:26	17.Jun.92		
-82630.	-134380.	38.300	
13.9304	24.8456	-1.0403	-3.8060
86.5816	-3.2539	1.0496	-1.
GLYCOLIC-ACID,aq	C ₂ H ₄ O ₃		
glyac	C(2)H(4)O(3)+(0)		
ref:26	16.Jun.92		
-126400.	-154890.	43.100	
8.7445	12.1496	4.0222	-3.2812
43.6669	-0.4655	-0.3017	0.
GLYCOLATE	C ₂ H ₃ O ₃		
C ₂ H ₃ O ₃ -	C(2)H(3)O(3)-(1)		
ref:26	17.Jun.92		
-121170.	-154700.	26.200	
7.6349	11.0570	0.9834	-3.2360
26.0463	-4.0271	1.2334	-1.
LACTIC-ACID,aq	C ₃ H ₆ O ₃		
lacac	C(3)H(6)O(3)+(0)		
ref:26	16.Jun.92		

	-127800.	-164000.	49.800	
	11.1720	17.2432	3.8083	-3.4917
	61.0199	1.6290	-0.2572	0.
LACTATE		C3H5O3-		
lac8		C(3)H(5)O(3)-(1)		
ref:26		17.JUN.92		
	-122530.	-164070.	31.900	
	9.8498	12.8244	8.0990	-3.3091
	44.9518	-3.7823	1.1469	-1.
OXALIC-ACID,aq		C2H2O4		
OXAC		C(2)H(2)O(4)+(0)		
ref:26		28.FEB.92		
	-168640.	-194580.	44.000	
	8.4291	12.3793	1.7866	-3.2907
	25.4994	-2.7181	-0.2957	0.
H-OXALATE		C2HO4-		
Hoxa-		C(2)H(1)O(4)-(1)		
ref:26		17.JUN.92		
	-166907.	-195600.	36.000	
	7.9276	11.2406	2.0500	-3.2436
	6.5915	-6.3548	1.0831	-1.
OXALATE		C2O4-2		
Ox-2		C(2)O(4)-(2)		
ref:26		17.JUN.92		
	-161100.	-197200.	10.900	
	6.9414	9.6716	0.8669	-3.1787
	-10.7618	-18.6787	3.0480	-2.
SUCCINIC-ACID,aq		C4H5O4		
SUCAC		C(4)H(6)O(4)+(0)		
ref:26		28.FEB.92		
	-177800.	-218000.	62.300	
	12.9872	22.7109	-0.5623	-3.7178
	51.0740	0.3050	-0.1744	0.
H-SUCCINATE		C4H5O4-		
Hsuc-		C(4)H(5)O(4)-(1)		
ref:26		16.JUN.92		
	-172060.	-217350.	45.200	
	11.6617	19.7057	0.1226	-3.5935
	26.117	-3.9932	0.9446	-1.
SUCCINATE		C4H4O4-2		
Suc-2		C(4)H(4)O(4)-(2)		
ref:26		17.JUN.92		
	-164380.	-217350.	19.500	
	10.4577	15.8622	3.5716	-3.4346
	-14.0383	-4.7180	2.9170	-2.
GLYCINE,aq		C2H5NO2		
GLV(aq)		C(2)H(5)O(2)N(1)+(0)		
ref:4		4.Sep.87		
	-88618.	-122846.	37.840	
	7.6046	7.0825	10.9119	-3.0717
	14.1998	-3.4185	-2.330	0.
ALANINE,aq		C3H7NO2		
ALA(aq)		C(3)H(7)O(2)N(1)+(0)		
ref:4		4.Sep.87		
	-88800.	-132130.	40.000	
	9.9472	11.7629	11.3023	-3.2652
	34.9465	-1.7690	-2.2658	0.
Ac+2		Ac(+2)		
Ac+2		Ac(1)+(2)		
ref:36		16.JAN.98		
	-79200.	-73200.	4.000	
	0.1865	-7.3171	8.6064	-2.4764
	7.9173	-5.6215	0.9979	2.
Ac+3		Ac(+3)		
Ac+3		Ac(1)+(3)		
ref:36		16.JAN.98		
	-153200.	-155300.	-39.000	
	-2.5931	-14.1059	11.2780	-2.1958
	3.9756	-10.7344	2.1668	3.
Ag+		Ag(+)		
Ag+		Ag(1)+(1)		
ref:2		1.JUL.87		

	18427.	25275.	17.540	
	1.7285	-3.5608	7.1496	-2.6318
	12.7862	-1.4254	.2160	1.
AgOH,aq		AgOH(0)		
AgOH(aq)		Ag(1)O(1)H(1)+(0)		
ref:21		12.Dec.97		
	-21900.	-31800.	17.200	
	2.2885	-2.1912	6.6063	-2.6883
	-0.3221	-5.1937	-0.0300	0.
AgO-		AgO(-1)		
AgO-		Ag(1)O(1)-(1)		
ref:21		12.Dec.97		
	-5500.	-16300.	13.900	
	1.4640	-4.1994	7.3839	-2.6053
	0.8786	-9.4103	1.4171	-1.
AgCl,aq		AgCl(0)		
AgCl(aq)		Ag(1)Cl(1)+(0)		
ref:22		17.Sep.97		
	-17450.	-18270.	34.100	
	5.2088	4.9399	3.8015	-2.9832
	9.8168	-1.6698	-0.0300	0.
AgCl2-		AgCl2(-)		
AgCl2-		Ag(1)Cl(2)-(1)		
ref:40		20.JAN.98		
	-51350.	-55438.	49.780	
	7.1327	9.8065	1.8947	-3.1844
	4.8953	-6.7789	0.6667	-1.
AgCl3-2		AgCl3(-2)		
AgCl3-2		Ag(1)Cl(3)-(2)		
ref:22		17.Sep.97		
	-82710.	-105613.	44.500	
	14.5040	27.6363	-5.1192	-3.9214
	35.8339	-1.8570	2.5402	-2.
AgCl4-3		AgCl4(-3)		
AgCl4-3		Ag(1)Cl(4)-(3)		
ref:22		17.Sep.97		
	-112280.	-150616.	35.000	
	20.0386	41.1504	-10.4309	-4.4801
	55.0238	243	4.2796	-3.
AgF,aq		AgF(0)		
AgF(aq)		Ag(1)F(1)+(0)		
ref:22		6.JAN.98		
	-49459.	-54722.	16.740	
	2.5788	-1.4818	6.3255	-2.7177
	12.1829	-0.8218	-0.0380	0.
AgCO3-		AgCO3(-)		
AgCO3-		Ag(1)C(1)O(3)-(1)		
ref:22		17.Sep.97		
	-111430.	-142490.	0.000	
	2.2588	-2.2632	6.6326	-2.6854
	15.2147	-5.1123	1.6310	-1.
Ag(CO3)2-3		Ag(CO3)2(-3)		
Ag(CO3)2-3		Ag(1)C(2)O(6)-(3)		
ref:22		17.Sep.97		
	-236890.	-304200.	-19.000	
	3.7128	1.2872	5.2371	-2.8322
	36.5589	-8.7993	5.0992	-3.
AgNO3(aq)		AgNO3(0)		
AgNO3(aq)		Ag(1)N(1)O(3)+(0)		
ref:22		17.Sep.97		
	-7738.	-23320.	49.000	
	6.7660	8.7423	2.3069	-3.1404
	24.1485	3.3372	-0.0380	0.
Ag(HS)2-		Ag(HS)2(-)		
Ag(HS)2-		Ag(1)H(2)S(2)-(1)		
ref:22		6.JAN.98		
	0.	-8200.	44.680	
	10.3654	17.5311	-1.1475	-3.5037
	37.2753	4.7273	0.9527	-1.
Ag(Ae),aq		AgCH3COO		
Ag(Ae),aq		Ag(1)C(2)H(3)O(2)		
ref:33		17.AUG.92		

	-70840.	-91650.	38.900	
	8.3987	12.7257	0.7485	-3.3050
	48.3694	11.7300	-0.0300	0.
Ag(Ae)2-		Ag(CH3COO)2-		
Ag(Ae)2-		Ag(1)C(4)H(6)O(4)-(1)		
ref:33		17.AUG.92		
	-158990.	-210560.	49.900	
	16.2287	31.8418	-6.7592	-4.0952
	110.8373	30.5473	0.8741	-1.
Ag+2		Ag(+2)		
Ag+2		Ag(1)+(2)		
ref:2		16.FEB.88		
	64300.	64200.	-21.000	
	-4299	-8.8310	9.2210	-2.4139
	15.2224	-4.1142	1.3201	2.
Al+3		Al(+3)		
Al+3		Al(1)+(3)		
ref:20		29.APR.97		
	-116543.	-128681.	-80.800	
	-3.3404	-17.1108	14.9917	-2.0716
	10.7000	-8.0600	2.8711	3.
AlOH+2		AlOH(+2)		
AlOH+2		Al(1)O(1)H(1)+(2)		
ref:20		30.APR.97		
	-166468.	-185096.	-46.820	
	2580	-9.6965	9.6000	-2.5767
	38.0000	-2.6000	1.7666	2.
AlO+		AlO(+)		
AlO+		Al(1)O(1)+(1)		
ref:20		29.APR.97		
	-215465.	-241825.	-17.000	
	3.5000	-2.4600	6.3000	-2.8100
	46.7000	-1.5000	.8089	1.
AlO2-		AlO2(-)		
AlO2-		Al(1)O(2)-(1)		
ref:20		29.APR.97		
	-198700.	-222079.	-7.000	
	3.7280	3.9800	-1.5170	-2.9435
	19.1000	-6.2000	1.7595	-1.
Al(Ae)+2		AlCH3COO+2		
Al(Ae)+2		Al(1)C(2)H(3)O(2)+(2)		
ref:33,35		15.JAN.98		
	-208564.	-249130.	-32.770	
	2.1643	-2.4961	6.7288	-2.6757
	63.2437	11.8151	1.5579	2.
Al(Ae)2+		Al(CH3COO)2+		
Al(Ae)2+		Al(1)C(4)H(6)O(4)+(1)		
ref:33,35		15.JAN.98		
	-299359.	-372080.	8.520	
	8.5167	13.0156	0.6302	-3.3170
	109.3239	31.4722	0.4210	1.
Am+2		Am(+2)		
Am+2		Am(1)+(2)		
ref:36		16.JAN.98		
	-89600.	-85100.	-3.000	
	0.0294	-7.7042	8.7663	-2.4604
	9.6780	-5.3363	1.1000	2.
Am+3		Am(+3)		
Am+3		Am(1)+(3)		
ref:36		16.JAN.98		
	-142900.	-147400.	-48.900	
	-2.9531	-14.9848	11.6223	-2.1594
	6.5235	-10.3270	2.3161	3.
Am+4		Am(+4)		
Am+4		Am(1)+(4)		
ref:36		16.JAN.98		
	-82900.	-99200.	-104.000	
	-4.3164	-18.3144	12.9330	-2.0218
	40.0550	-3.2586	3.7484	4.
AmO2+		AmO2(+)		
AmO2+		Am(1)O(2)+(1)		
ref:36		18.JAN.98		

	-177100.	-192400.	-4.900	
	3.4121	0.5496	5.5350	-2.8016
	-2.3999	-8.0048	0.6223	1.
AmO2+2		AmO2(+2)		
AmO2+2		Am(1)O(2)+(2)		
ref:36		18.Jan.98		
	-140400.	-155800.	-20.800	
	3.1185	-0.1662	5.8139	-2.7720
	23.6816	-1.3438	1.3732	2.
Ar,aq		Ar(0)		
Ar(aq)		Ar(1)+(0)		
ref:3		26.Jun.87		
	3900.	-2870.	14.300	
	6.0003	6.8698	3.0499	-3.0630
	35.6471	7.3160	-3.073	0.
AsO2-		AsO2(-)		
AsO2-		As(1)O(2)-(1)		
ref:21		12.Dec.97		
	-83650.	-102540.	9.700	
	3.2101	0.0554	5.7305	-2.7812
	3.4104	-8.7381	1.4820	-1.
AsO4-3		AsO4(-3)		
AsO4-3		As(1)O(4)-(3)		
ref:21		12.Dec.97		
	-154970.	-212270.	-38.900	
	1.0308	-5.2609	7.8091	-2.5614
	-12.1352	-26.6841	5.3990	-3.
Au+		Au(+)		
Au+		Au(1)+(1)		
ref:21		5.Jan.98		
	39000.	47900.	25.600	
	3.5312	0.8428	5.4139	-2.8137
	7.5089	-3.0956	0.1648	1.
AuCl,aq		AuCl(0)		
AuCl(aq)		Au(1)Cl(1)+(0)		
ref:22		5.Jan.98		
	-3184.	-2140.	41.470	
	7.0357	9.4008	2.0481	-3.1676
	0.0551	-5.0371	-0.0380	0.
AuCl2-		AuCl2(-)		
AuCl2-		Au(1)Cl(2)-(1)		
ref:22		5.Jan.98		
	-36781.	-44725.	53.600	
	11.5192	20.3482	-2.2547	-3.6201
	-0.6764	-8.0300	0.8173	-1.
AuCl3-2		AuCl3(-2)		
AuCl3-2		Au(1)Cl(3)-(2)		
ref:22		5.Jan.98		
	-67811.	-86023.	61.390	
	16.6878	32.9687	-7.2151	-4.1419
	1.7155	-11.8910	2.2827	-2.
Au(HS)2-		Au(HS)2(-)		
Au(HS)2-		Au(1)H(2)S(2)-(1)		
ref:22		6.Jan.98		
	2429.	-2509.	56.770	
	12.3420	22.3572	-3.0443	-3.7032
	16.8038	-1.8007	0.7693	-1.
Au(Ae),aq		AuCH3COO		
Au(Ae),aq		Au(1)C(2)H(3)O(2)		
ref:33		10 Sept 92		
	-49870.	-68310.	48.000	
	10.2130	17.1576	-0.9969	-3.4882
	38.7432	8.3843	-0.0300	0.
Au(Ae)2-		Au(CH3COO)2-		
Au(Ae)2-		Au(1)C(4)H(6)O(4)-(1)		
ref:33		10 Sept 92		
	-138240.	-186750.	61.300	
	18.1917	36.6392	-8.6534	-4.2936
	90.4610	24.0193	0.7010	-1.
Au+3		Au(+3)		
Au+3		Au(1)+(3)		
ref:21		5.Jan.98		

	103600.	97800.	-54.800	
	-1.7167	-11.9654	10.4352	-2.2843
	23.5775	-4.7048	2.4115	3.
AuCl4-		AuCl4(-)		
AuCl4-		Au(1)Cl(4)-(1)		
ref:22		5.Jan.98		
	35332.	23573.	-32.720	
	-0.5361	-9.0876	9.3148	-2.4033
	19.7251	-3.3107	1.5579	-1.
B(OH)3,aq		BOH3(0)		
B(OH)3(aq)		B(1)O(3)H(3)+(0)		
ref:39		20.Jan.98		
	-231540.	-256820.	38.790	
	8.0000	7.5000	15.0000	-6.4400
	40.0000	-6.8500	0.0450	0.
B(OH)4-		BOH4(-)		
B(OH)4-		B(1)O(4)H(4)-(1)		
ref:39		20.Jan.98		
	-275610.	-321230.	24.500	
	5.5200	6.3300	3.8000	-3.2500
	45.0000	-19.3300	0.9000	-1.
BF4-		BF4(-)		
BF4-		B(1)F(4)-(1)		
ref:2		16.Feb.88		
	-355400.	-376400.	43.000	
	7.9796	11.7027	1.1504	-3.2628
	11.8941	-4.1753	.9779	-1.
Ba+2		Ba(+2)		
Ba+2		Ba(1)+(2)		
ref:2		1.Jul.87		
	-134030.	-128500.	2.300	
	2.7382	-10.0565	-.0470	-2.3633
	3.8000	-3.4500	.9850	2.
BaOH+		BaOH(+1)		
BaOH+		Ba(1)O(1)H(1)+(1)		
ref:21		12.Sep.97		
	-172300.	-175900.	27.500	
	3.0700	-.2825	5.8541	-2.7672
	-5.7064	-7.5974	.1362	1.
BaCl+		BaCl(+)		
BaCl+		Ba(1)Cl(1)+(1)		
ref:22		17.Sep.97		
	-164730.	-165323.	24.000	
	3.4534	.6537	5.4861	-2.8060
	11.8213	-1.6759	.1895	1.
BaF+		BaF(+)		
BaF+		Ba(1)F(1)+(1)		
ref:22(tab. 3)		15.Sep.97		
	-201120.	-206510.	5.500	
	.9652	-5.4218	7.8740	-2.5548
	16.7505	-.8529	.4675	1.
BaCO3,aq		BaCO3(0)		
BaCO3(aq)		Ba(1)C(1)O(3)+(0)		
ref:22		17.Sep.97		
	-263830.	-285850.	16.000	
	.1361	-7.4461	8.6697	-2.4711
	-14.3440	-10.0418	-.0380	0.
Ba(Ae)+		BaCH3COO+		
Ba(Ae)+		Ba(1)C(2)H(3)O(2)+(1)		
ref:33		17 Aug 92		
	-223650.	-242850.	33.500	
	6.6253	8.3926	2.4574	-3.1259
	48.9806	11.6995	0.0459	1.
Ba(Ae)2,aq		Ba(CH3COO)2		
Ba(Ae)2,aq		Ba(1)C(4)H(6)O(4)		
ref:33		10 Sept 92		
	-312620.	-358010.	59.700	
	13.9186	26.2066	-4.5556	-3.8623
	103.6344	30.9388	-0.0300	0.
Ba(Ala)+		Ba(C3H6NO2)+		
Ba(Ala)+		Ba(1)C(3)H(6)N(1)O(2)+(1)		
ref:27		14.Aug.93		

	-210013.	-243703.	56.097	
	9.4348	15.2528	-0.2391	-3.4095
	58.4325	16.0850	-0.2977	1.
Ba(Ala)2,aq		Ba(C3H6NO2)2		
Ba(Ala)2,aq		Ba(1)C(6)H(12)N(2)O(4)		
ref:27		14.Aug.93		
	-285287.	-359051.	107.026	
	20.1040	41.3073	-10.4865	-4.4865
	133.7647	41.4112	-0.0300	0.
Ba(But)+		BaCH3(CH2)2CO2+		
Ba(But)+		Ba(1)C(4)H(7)O(2)+(1)		
ref:27		14.Aug.93		
	-218640.	-253285.	44.647	
	10.6454	18.2150	-1.4160	-3.5319
	89.0694	26.1794	-0.1247	1.
Ba(But)2,aq		Ba(CH3CH2CH2CO2)2		
Ba(But)2,aq		Ba(1)C(8)H(14)O(4)		
ref:27		14.Aug.93		
	-302827.	-378066.	85.592	
	22.5399	47.2556	-12.8248	-4.7324
	203.1160	65.5159	-0.0300	0.
Ba(For)+		BaCHO2+		
Ba(For)+		Ba(1)C(1)H(1)O(2)+(1)		
ref:27		26.Aug.93		
	-219775.	-228918.	34.547	
	4.6569	3.5906	4.3358	-2.9273
	18.1411	1.0380	0.0279	1.
Ba(For)2,aq		Ba(CHO2)2		
Ba(For)2,aq		Ba(1)C(2)H(2)O(4)		
ref:27		26.Aug.93		
	-304769.	-329933.	62.275	
	9.7700	16.0719	-0.5623	-3.4433
	30.3870	5.4799	-0.0300	0.
Ba(Gly)+		Ba(C2H4NO2)+		
Ba(Gly)+		Ba(1)C(2)H(4)N(1)O(2)+(1)		
ref:27		27.Aug.93		
	-211341.	-235808.	54.327	
	6.9656	9.2295	2.1164	-3.1604
	32.1389	6.8602	-0.2709	1.
Ba(Gly)2,aq		Ba(C2H4NO2)2		
Ba(Gly)2,aq		Ba(1)C(4)H(8)N(2)O(4)		
ref:27		27.Aug.93		
	-287819.	-343302.	102.940	
	14.8647	28.5140	-5.4575	-3.9577
	70.3874	19.3830	-0.0300	0.
Ba(Glyc)+		BaCH3OCO2+		
Ba(Glyc)+		Ba(1)C(2)H(3)O(3)+(1)		
ref:27		30.Aug.93		
	-256564.	-282924.	34.000	
	6.5329	8.1729	2.5314	-3.1168
	49.2137	11.8129	0.0358	1.
Ba(Glyc)2,aq		Ba(CH3OCO2)2		
Ba(Glyc)2,aq		Ba(1)C(4)H(6)O(6)		
ref:27		30.Aug.93		
	-378689.	-436833.	66.059	
	13.7306	25.7444	-4.3677	-3.8432
	104.4137	31.2096	-0.0300	0.
Ba(Lac)+		BaCH3CH2OCO2+		
Ba(Lac)+		Ba(1)C(3)H(5)O(3)+(1)		
ref:27		30.Aug.93		
	-257433.	-291416.	41.000	
	8.7414	13.5626	0.4195	-3.3396
	70.2142	19.4498	-0.0697	1.
Ba(Lac)2,aq		Ba(CH3CH2OCO2)2		
Ba(Lac)2,aq		Ba(1)C(6)H(10)O(6)		
ref:27		30.Aug.93		
	-380522.	-453654.	80.919	
	18.4752	37.3316	-8.9265	-4.3222
	156.8818	49.4461	-0.0300	0.
Ba(Pent)+		BaCH3(CH2)3CO2+		
Ba(Pent)+		Ba(1)C(5)H(9)O(2)+(1)		
ref:27		31.Aug.93		

-216357.	-259492.	51.147	
12.8030	23.4822	-3.4843	-3.7497
125.4720	39.1471	-0.2231	1.
Ba(Pent)2,aq	Ba(CH ₃ CH ₂ CH ₂ CH ₂ CO ₂) ₂		
Ba(Pent)2,aq	Ba(1)C(10)H(18)O(4)		
ref:27	31.Aug.93		
-298289.	-389909.	100.598	
27.1717	58.5642	-17.2675	-5.1999
292.2076	96.4819	-0.0300	0.
Ba(Prop)+	BaCH ₃ CH ₂ CO ₂ +		
Ba(Prop)+	Ba(1)C(3)H(5)O(2)+(1)		
ref:27	15.Sep.93		
-221005.	-248190.	39.347	
8.5720	13.1515	0.5756	-3.3226
75.0141	21.0377	-0.0446	1.
Ba(Prop)2,aq	Ba(CH ₃ CH ₂ CO ₂) ₂		
Ba(Prop)2,aq	Ba(1)C(6)H(10)O(4)		
ref:27	15.Sep.93		
-307529.	-368336.	73.356	
18.0991	36.4138	-8.5671	-4.2842
167.7909	53.2379	-0.0300	0.
Be+2	Be(+2)		
Be+2	Be(1)+(2)		
ref:21	15.Sep.97		
-90550.	-91500.	-32.100	
-5718	-9.1742	9.3462	-2.3996
17.0245	-4.2160	1.5475	2.
BeOH+	BeOH(+1)		
BeOH+	Be(1)O(1)H(1)+(1)		
ref:21	16.Sep.97		
-139900.	-153200.	-17.900	
2.3302	-2.0945	6.5792	-2.6923
30.9714	2.9542	.8222	1.
BeO,aq	BeO(0)		
BeO(aq)	Be(1)O(1)+(0)		
ref:21	16.Sep.97		
-128600.	-144100.	-25.100	
1.8917	-3.1609	6.9887	-2.6482
13.5675	-3.3660	-0.0300	0.
BeO2-2	BeO2(-2)		
BeO2-2	Be(1)O(2)-(2)		
ref:21	16.Sep.97		
-152900.	-189700.	-41.000	
3.2657	.1906	5.6788	-2.7868
34.0366	-5.6215	3.8330	-2.
BeCl+	BeCl(+1)		
BeCl+	Be(1)Cl(1)+(1)		
ref:22	13.Sep.97		
-115316.	-124998.	92.310	
1.1974	-4.8548	7.6512	-2.5783
16.6475	3.3197	-0.8467	1.
BeCl2,aq	BeCl2(0)		
BeCl2(aq)	Be(1)Cl(2)+(0)		
ref:22	13.Sep.97		
-145521.	-156432.	130.440	
5.0297	4.5026	3.9733	-2.9651
39.6530	8.7262	-0.0380	0.
BeF+	BeF(+1)		
BeF+	Be(1)F(1)+(1)		
ref:22	15.Sep.97		
-158493.	-177322.	59.200	
-1.2603	-10.8560	10.0098	-2.3302
23.7740	4.1893	-3.447	1.
BeF2,aq	BeF2(0)		
BeF2(aq)	Be(1)F(2)+(0)		
ref:22	15.Sep.97		
-231386.	-258350.	71.720	
-5.5250	-9.0604	9.3042	-2.4044
45.6271	10.8026	-0.0380	0.
BeF3-	BeF3(-1)		
BeF3-	Be(1)F(3)-(1)		
ref:22	15.Sep.97		

-302532.	-331512.	104.780	
.2052	-7.2776	8.6034	-2.4781
62.8555	16.5406	.0401	-1.
BeF4-2	BeF4(-2)		
BeF4-2	Be(1)F(4)-(2)		
ref:22	15.Sep.97		
-370691.	-386096.	190.120	
1.0882	-5.1214	7.7560	-2.5672
79.5432	21.4033	.3329	-2.
Be(Ae)+	BeCH ₃ COO+		
Be(Ae)+	Be(1)C(2)H(3)O(2)+(1)		
ref:33	11.Mar.93		
-174020.	-213040.	-45.600	
5.0795	4.6238	3.9264	-2.9700
74.5470	16.7410	1.2465	1.
Be(Ae)2,aq	Be(CH ₃ COO) ₂		
Be(Ae)2,aq	Be(1)C(4)H(6)O(4)		
ref:33	11.Mar.93		
-263740.	-336230.	-43.800	
11.7442	20.8961	-2.4661	-3.6427
131.9355	40.7754	-0.0300	0.
Bi+3	Bi(+3)		
Bi+3	Bi(1)+(3)		
ref:21	11.Sep.97		
22880.	19360.	-45.000	
-1.0957	-10.4536	9.8506	-2.3467
27.5027	-2.8715	2.2651	3.
BiOH+2	BiOH(+2)		
BiOH+2	Bi(1)O(1)H(1)+(2)		
ref:21	16.Sep.97		
-32300.	-44800.	-19.500	
2.4693	-1.7539	6.4433	-2.7064
9.8610	-6.0900	1.3552	2.
BiO+	BiO(+)		
BiO+	Bi(1)O(1)+(1)		
ref:21	16.Sep.97		
-29300.	-30300.	19.100	
2.9487	-5.825	5.9810	-2.7548
-39.8159	-19.8601	.2634	1.
BiO2-	BiO2(-)		
BiO2-	Bi(1)O(2)-(1)		
ref:21	16.Sep.97		
-61700.	-71600.	45.100	
5.2319	4.9920	3.7909	-2.9853
-59.7858	-28.9859	.9456	-1.
Bi(Ae)+2	BiCH ₃ COO+2		
Bi(Ae)+2	Bi(1)C(2)H(3)O(2)+(2)		
ref:33	26.Sep.92		
-67750.	-96420.	-15.200	
5.1690	4.8386	3.8508	-2.9789
107.9465	28.2236	1.2860	2.
Bi(Ae)2+	Bi(CH ₃ COO) ₂ +		
Bi(Ae)2+	Bi(1)C(4)H(6)O(4)+(1)		
ref:33	26.Sep.92		
-157070.	-212380.	9.600	
11.9760	21.4619	-2.6877	-3.6661
201.6950	63.6303	0.4046	1.
Bi(Ae)3,aq	Bi(CH ₃ COO) ₃		
Bi(Ae)3,aq	Bi(1)C(6)H(9)O(6)		
ref:33	26.Sep.92		
-245540.	-329240.	28.600	
19.7454	40.4309	-10.1406	-4.4503
313.6160	103.9230	-0.0300	0.
Br-	Br(-)		
Br-	Br(1)-(1)		
ref:2	1.Jul.87		
-24870.	-29040.	19.800	
5.2690	6.5940	4.7450	-3.1430
-3.8000	-6.8110	1.3858	-1.
Br3-	Br3(-)		
Br3-	Br(3)-(1)		
ref:2	12.Jul.87		

-25590.	-31170.	51.500	
9.0992	14.4364	.0759	-3.3758
17.2705	-1.8939	.8490	-1.
BrO-	BrO(-)		
BrO-	Br(1)O(1)-(1)		
ref:21	11.Sep.97		
-8000.	-22500.	9.700	
2.3206	-2.1131	6.5758	-2.6915
3.4104	-8.7381	1.4820	-1.
BrO3-	BrO3(-)		
BrO3-	Br(1)O(3)-(1)		
ref:2	12.Jul.87		
4450.	-16030.	38.650	
6.9617	9.2173	2.1272	-3.1600
3.7059	-7.2308	1.0433	-1.
BrO4-	BrO4(-)		
BrO4-	Br(1)O(4)-(1)		
ref:21	30.Apr.97		
28200.	3100.	47.600	
8.7361	13.5515	.4200	-3.3391
11.4884	-4.0937	.9085	-1.
CH4,aq	CH4(0)		
CH4(aq)	C(1)H(4)+(0)		
ref:4	31.Aug.87		
-8234.	-21010.	20.990	
6.7617	8.7279	2.3212	-3.1397
42.0941	10.4707	-3.179	0.
CN-	CN(-1)		
CN-	C(1)N(1)-(1)		
ref:21	11.Sep.97		
41200.	36000.	22.500	
5.4714	5.5813	3.5497	-3.0096
7.6773	-6.6400	1.2900	-1.
CO2,aq	CO2(0)		
CO2(aq)	C(1)O(2)+(0)		
ref:3	26.Jun.87		
-92250.	-98900.	28.100	
6.2466	7.4711	2.8136	-3.0879
40.0325	8.8004	-0.0200	0.
CO3-2	CO3(-2)		
CO3-2	C(1)O(3)-(2)		
ref:2	1.Jul.87		
-126191.	-161385.	-11.950	
2.8524	-3.9844	6.4142	-2.6143
-3.3206	-17.1917	3.3914	-2.
Ca+2	Ca(+2)		
Ca+2	Ca(1)+(2)		
ref:2	3.Jun.87		
-132120.	-129800.	-13.500	
-1.1947	-7.2520	5.2966	-2.4792
9.0000	-2.5220	1.2366	2.
CaOH+	CaOH(+1)		
CaOH+	Ca(1)O(1)H(1)+(1)		
ref:21	12.Sep.97		
-171300.	-179600.	6.700	
2.7243	-1.1303	6.1958	-2.7322
11.1286	-2.7493	.4496	1.
CaCl+	CaCl(+)		
CaCl+	Ca(1)Cl(1)+(1)		
ref:22	17.Sep.97		
-163100.	-168607.	4.500	
2.7148	-1.1497	6.1949	-2.7314
20.8839	.5241	.4862	1.
CaCl2,aq	CaCl2(0)		
CaCl2(aq)	Ca(1)Cl(2)+(0)		
ref:22	17.Sep.97		
-194000.	-211060.	6.000	
6.2187	7.4058	2.8322	-3.0851
23.9610	3.2720	-0.0380	0.
CaF+	CaF(+)		
CaF+	Ca(1)F(1)+(1)		
ref:22	12.Sep.97		

-200390. 1568 30.1743	-208600. -7.3958 3.0968	-9.000 8.6499 6.911	-2.4732 1.
CaHCO ₃ + CaHCO ₃ + ref:5	CaHCO ₃ (+) Ca(1)H(1)C(1)O(3)+(1) 1.Jan.91	16.000 5.7459 3.084	-2.7794 1.
-273830. 3.1911 42.3545	-294350. 0.104 8.5559	16.000 5.7459 3.084	-2.7794 1.
CaCO ₃ ,aq CaCO ₃ (aq) ref:22	CaCO ₃ (0) Ca(1)C(1)O(3)+(0) 17.Sep.97	2.500 9.1753 -0.0380	-2.4179 0.
-262850. -3.907 -11.5309	-287390. -8.7325 -9.0641	2.500 9.1753 -0.0380	-2.4179 0.
CaSO ₄ ,aq CaSO ₄ (aq) ref:22	CaSO ₄ (0) Ca(1)S(1)O(4)+(0) 12.Sep.97	5.000 6.4895 -0.0010	-2.7004 0.
-312930. 2.4079 -8.4942	-345900. -1.8992 -8.1271	5.000 6.4895 -0.0010	-2.7004 0.
Ca(HSiO ₃)+ Ca(HSiO ₃)+ ref:22	Ca(HSiO ₃)(+1) Ca(1)H(1)Si(1)O(3)+(1) 15.Sep.97	-1.990 7.7785 0.5831	-2.5649 1.
-376299. 1.0647 30.8048	-403109. -5.1787 3.6619	-1.990 7.7785 0.5831	-2.5649 1.
Ca(Ae)+ Ca(Ae)+ ref:33	CaCH ₃ COO+ Ca(1)C(2)H(3)O(2)+(1) 5 Apr 93	12.500 3.1505 0.3636	-3.0527 1.
-221660. 5.9002 58.1976	-245620. 6.6232 13.8857	12.500 3.1505 0.3636	-3.0527 1.
Ca(Ae)2,aq Ca(Ae)2,aq ref:33	Ca(CH ₃ COO)2 Ca(1)C(4)H(6)O(4) 5 Apr 93	32.300 -3.6556 -0.0300	-3.7685 0.
-311570. 12.9911 115.9068	-362650. 23.9379 35.2043	32.300 -3.6556 -0.0300	-3.7685 0.
Ca(Ala)+ Ca(Ala)+ ref:27	Ca(C ₃ H ₆ NO ₂)2 Ca(1)C(3)H(6)O(2)N(1)+(1) 14.Aug.93	34.555 0.4540 0.0279	-3.3365 1.
-208472. 8.7124 67.7227	-247083. 13.4888 18.2713	34.555 0.4540 0.0279	-3.3365 1.
Ca(Ala)2,aq Ca(Ala)2,aq ref:27	Ca(C ₃ H ₆ NO ₂)2 Ca(1)C(6)H(12)N(2)O(4) 14.Aug.93	78.835 -9.6021 -0.0300	-4.3930 0.
-284046. 19.1765 146.0369	-364714. 39.0448 45.6768	78.835 -9.6021 -0.0300	-4.3930 0.
Ca(But)+ Ca(But)+ ref:27	CaCH ₃ (CH ₂)2CO ₂ + Ca(1)C(4)H(7)O(2)+(1) 14.Aug.93	23.105 -0.7166 0.2025	-3.4589 1.
-217467. 9.9235 98.3731	-257034. 16.4497 28.3657	23.105 -0.7166 0.2025	-3.4589 1.
Ca(But)2,aq Ca(But)2,aq ref:27	Ca(CH ₃ CH ₂ CH ₂ CO ₂)2 Ca(1)C(8)H(14)O(4) 14.Aug.93	57.401 -11.9404 -0.0300	-4.6389 0.
-302268. 21.6124 215.3882	-384411. 44.9931 69.7815	57.401 -11.9404 -0.0300	-4.6389 0.
Ca(For)+ Ca(For)+ ref:27	CaCHO ₂ + Ca(1)C(1)H(1)O(2)+(1) 16.Aug.93	13.005 5.0305 0.3539	-2.8544 1.
-217933. 3.9346 27.4342	-231998. 1.8264 3.2242	13.005 5.0305 0.3539	-2.8544 1.
Ca(For)2,aq Ca(For)2,aq ref:27	Ca(CHO ₂)2 Ca(1)C(2)H(2)O(4) 16.Aug.93		

-302982. 8.8425 42.6594	-335050. 13.8094 9.7454	34.084 0.3220 -0.0300	-3.3498 0.
Ca(Gly)+ Ca(Gly)+ ref:27	Ca(C ₂ H ₄ NO ₂)2 Ca(1)C(2)H(4)O(2)N(1)+(1) 16.Aug.93	32.785 2.8241 0.0543	-3.0873 1.
-209240. 6.2431 41.4246	-238629. 7.4595 9.0464	32.785 2.8241 0.0543	-3.0873 1.
Ca(Gly)2,aq Ca(Gly)2,aq ref:27	Ca(C ₂ H ₄ NO ₂)2 Ca(1)C(4)H(8)N(2)O(4) 16.Aug.93	74.749 -4.5733 -0.0300	-3.8641 0.
-285555. 13.9371 82.6597	-347942. 26.2517 23.6485	74.749 -4.5733 -0.0300	-3.8641 0.
Ca(Glyc)+ Ca(Glyc)+ ref:27	CaCH ₃ OCO ₂ + Ca(1)C(2)H(3)O(3)+(1) 16.Aug.93	17.505 3.2594 0.2873	-3.0411 1.
-255541. 5.7854 57.8209	-285318. 6.3436 13.9991	17.505 3.2594 0.2873	-3.0411 1.
Ca(Glyc)2,aq Ca(Glyc)2,aq ref:27	Ca(CH ₃ OCO ₂)2 Ca(1)C(4)H(6)O(6) 16.Aug.93	44.473 -3.4840 -0.0300	-3.7496 0.
-378403. 12.8031 116.6861	-441481. 23.4820 35.4751	44.473 -3.4840 -0.0300	-3.7496 0.
Ca(Lac)+ Ca(Lac)+ ref:27	CaCH ₃ CH ₂ OCO ₂ + Ca(1)C(3)H(5)O(3)+(1) 16.Aug.93	23.00 1.1355 0.2025	-3.2647 1.
-256587. 8.0009 79.0115	-294436. 11.7524 21.6361	23.00 1.1355 0.2025	-3.2647 1.
Ca(Lac)2,aq Ca(Lac)2,aq ref:27	Ca(CH ₃ CH ₂ OCO ₂)2 Ca(1)C(6)H(10)O(6) 16.Aug.93	57.364 -8.0266 -0.0300	-4.2284 0.
-380563. 17.5477 169.1540	-459217. 35.0629 53.7117	57.364 -8.0266 -0.0300	-4.2284 0.
Ca(Pent)+ Ca(Pent)+ ref:27	CaCH ₃ (CH ₂)3CO ₂ + Ca(1)C(5)H(9)O(2)+(1) 16.Aug.93	29.605 -2.7923 0.1023	-3.6767 1.
-215129. 12.0805 134.7595	-263187. 21.7185 41.3334	29.605 -2.7923 0.1023	-3.6767 1.
Ca(Pent)2,aq Ca(Pent)2,aq ref:27	Ca(CH ₃ CH ₂ CH ₂ CH ₂ CO ₂)2 Ca(1)C(10)H(18)O(4) 16.Aug.93	72.407 -16.3832 -0.0300	-5.1064 0.
-297634. 26.2442 304.4798	-396159. 56.3017 100.7475	72.407 -16.3832 -0.0300	-5.1064 0.
Ca(Prop)+ Ca(Prop)+ ref:27	CaCH ₃ CH ₂ CO ₂ + Ca(1)C(3)H(5)O(2)+(1) 15.Sep.93	17.805 1.2776 0.2832	-3.2496 1.
-219818. 7.8503 84.3239	-251925. 11.3857 23.2239	17.805 1.2776 0.2832	-3.2496 1.
Ca(Prop)2,aq Ca(Prop)2,aq ref:27	Ca(CH ₃ CH ₂ CO ₂)2 Ca(1)C(6)H(10)O(4) 15.Sep.93	45.165 -7.6668 -0.0300	-4.1905 0.
-306942. 17.1716 180.0633	-374653. 34.1451 57.5034	45.165 -7.6668 -0.0300	-4.1905 0.
Cd+2 Cd+2 ref:2	Cd(+2) Cd(1)+(2) 1.Jul.87		
-18560. 0.0537 15.6573	-18140. -10.7080 -3.7476	-17.400 16.5176 1.2528	-2.3363 2.
CdOH+ CdOH+ ref:21	CdOH(+) Cd(1)O(1)H(1)+(1) 12.Dec.97		

-61500. 0.1473 18.8886	-73400. -7.4141 -0.5290	-2.900 8.6471 0.5985	-2.4724 1.
CdO,aq CdO(aq) ref:21	CdO(0) Cd(1)O(1)+(0) 12.Dec.97	-4.200 8.9312 -0.0300	-2.4426 0.
-47500. -0.1471 -0.3806	-59750. -8.1341 -5.2141	-4.200 8.9312 -0.0300	-2.4426 0.
CdO2-2 CdO2-2 ref:21	CdO2(-2) Cd(1)O(2)-(2) 12.Dec.97	-20.300 8.0303 3.5190	-2.5375 -2.
-67300. 0.7922 3.4810	-101000. -5.8401 -15.2361	-20.300 8.0303 3.5190	-2.5375 -2.
CdCl+ CdCl+ ref:22	CdCl(+) Cd(1)Cl(1)+(1) 5.Jan.98	0.750 5.8262 0.5394	-2.7702 1.
-52627. 3.0990 26.6485	-59392. -0.2116 2.3572	0.750 5.8262 0.5394	-2.7702 1.
CdCl2,aq CdCl2(aq) ref:22	CdCl2(0) Cd(1)Cl(2)+(0) 5.Jan.98	10.310 2.4408 -0.0380	-3.1263 0.
-84851. 6.6265 34.2501	-101362. 8.4016 6.8482	10.310 2.4408 -0.0380	-3.1263 0.
CdCl3- CdCl3- ref:22	CdCl3(-1) Cd(1)Cl(3)-(1) 5.Jan.98	10.860 -2.0276 1.4662	-3.5963 -1.
-115971. 11.2826 56.4451	-144916. 19.7705 9.7459	10.860 -2.0276 1.4662	-3.5963 -1.
CdCl4-2 CdCl4-2 ref:22	CdCl4(-2) Cd(1)Cl(4)-(2) 5.Jan.98	0.230 -7.0271 3.2060	-4.1221 -2.
-146081. 16.4919 76.2256	-190793. 32.4903 11.0502	0.230 -7.0271 3.2060	-4.1221 -2.
CdF+ CdF+ ref:22	CdF(+) Cd(1)F(1)+(1) 6.Jan.98	-13.110 8.2799 0.7483	-2.5121 1.
-87373. 0.5423 31.0752	-97544. -6.4544 3.2268	-13.110 8.2799 0.7483	-2.5121 1.
CdF2,aq CdF2(aq) ref:22	CdF2(0) Cd(1)F(2)+(0) 6.Jan.98	-23.720 7.7717 -0.0380	-2.5656 0.
-155164. 1.0718 40.2242	-180369. -5.1614 8.9247	-23.720 7.7717 -0.0380	-2.5656 0.
Cd(Ae)+ Cd(Ae)+ ref:33	CdCH ₃ COO+ Cd(1)C(2)H(3)O(2)+(1) 17 Aug 92	6.600 2.7593 0.4496	-3.0935 1.
-109460. 6.3043 64.3035	-135920. 7.6112 15.7327	6.600 2.7593 0.4496	-3.0935 1.
Cd(Ae)2,aq Cd(Ae)2,aq ref:33	Cd(CH ₃ COO)2 Cd(1)C(4)H(6)O(4) 17 Aug 92	24.600 -4.0570 -0.0300	-3.8107 0.
-199400. 13.4090 126.2752	-254520. 24.9584 38.8081	24.600 -4.0570 -0.0300	-3.8107 0.
Cd(Ae)3- Cd(Ae)3- ref:33	Cd(CH ₃ COO)3- Cd(1)C(6)H(9)O(6)-(1) 10 Sep 92	31.900 -12.2199 1.1469	-4.6683 -1.
-288030. 21.9033 213.8505	-376010. 45.7036 65.4786	31.900 -12.2199 1.1469	-4.6683 -1.
Cd(Ala)+ Cd(Ala)+ ref:27	Cd(C ₃ H ₆ NO ₂)2+ Cd(1)C(3)H(6)O(2)N(1)+(1) 14.Aug.93		

	-100082.	-141016.	29.238	
	9.1151	14.4742	0.0630	-3.3773
	73.7914	20.1183	0.1098	1.
Cd(Ala)2,aq		Cd(C3H6NO2)2		
Cd(Ala)2,aq		Cd(1)(6)H(12)N(2)O(4)		
ref:27		14.Aug.93		
	-179940.	-263420.	71.877	
	19.5944	40.0653	-10.0034	-4.4352
	156.4053	49.2806	-0.0300	0.
Cd(But)+		CdCH3(CH2)2CO2+		
Cd(But)+		Cd(1)(4)H(7)O(2)+(1)		
ref:27		14.Aug.93		
	-105285.	-147174.	17.788	
	10.3259	17.4297	-1.0970	-3.4994
	104.4311	30.2127	0.2832	1.
Cd(But)2,aq		Cd(CH3CH2CH2CO2)2		
Cd(But)2,aq		Cd(1)(8)H(14)O(4)		
ref:27		14.Aug.93		
	-191463.	-276419.	50.443	
	22.0303	46.0136	-12.3417	-4.6811
	225.7566	73.3853	-0.0300	0.
Cd(For)+		CdCHO2+		
Cd(For)+		Cd(1)(1)H(1)O(2)+(1)		
ref:27		26.Aug.93		
	-104932.	-121320.	7.688	
	4.3381	2.8097	4.6479	-2.8951
	33.5224	5.0712	0.4379	1.
Cd(For)2,aq		Cd(CHO2)2		
Cd(For)2,aq		Cd(1)(2)H(2)O(4)		
ref:27		26.Aug.93		
	-191523.	-226403.	27.125	
	9.2604	14.8299	-0.0792	-3.3920
	53.0278	13.3492	-0.0300	0.
Cd(Gly)+		Cd(C2H4NO2)+		
Cd(Gly)+		Cd(1)(2)H(4)O(2)N(1)+(1)		
ref:27		27.Aug.93		
	-100236.	-132088.	27.000	
	6.6477	8.4525	2.4231	-3.1283
	47.5446	10.8935	0.1417	1.
Cd(Gly)2,aq		Cd(C2H4NO2)2		
Cd(Gly)2,aq		Cd(1)(4)H(8)N(2)O(4)		
ref:27		27.Aug.93		
	-180576.	-246607.	65.000	
	14.3550	27.2722	-4.9746	-3.9063
	93.0282	27.2523	-0.0300	0.
Cd(Glyc)+		CdCH3OCO2+		
Cd(Glyc)+		Cd(1)(2)H(3)O(3)+(1)		
ref:27		30.Aug.93		
	-142281.	-174381.	12.188	
	6.1880	7.3293	2.8656	-3.0819
	63.8836	15.8461	0.3686	1.
Cd(Glyc)2,aq		Cd(CH3OCO2)2		
Cd(Glyc)2,aq		Cd(1)(4)H(6)O(6)		
ref:27		30.Aug.93		
	-265388.	-331279.	37.514	
	13.2210	24.5024	-3.8851	-3.7918
	127.0546	39.0789	-0.0300	0.
Cd(Lac)+		CdCH3CH2OCO2+		
Cd(Lac)+		Cd(1)(3)H(5)O(3)+(1)		
ref:27		30.Aug.93		
	-143409.	-183519.	17.888	
	8.4019	12.7361	0.7390	-3.3054
	85.0322	23.4831	0.2792	1.
Cd(Lac)2,aq		Cd(CH3CH2OCO2)2		
Cd(Lac)2,aq		Cd(1)(6)H(10)O(6)		
ref:27		30.Aug.93		
	-267699.	-349085.	50.673	
	17.9656	36.0834	-8.4279	-4.2706
	179.5224	57.3155	-0.0300	0.
Cd(Pent)+		CdCH3(CH2)3CO2+		
Cd(Pent)+		Cd(1)(5)H(9)O(2)+(1)		
ref:carla		31.Aug.93		

	-103384.	-153764.	24.288	
	12.4829	22.6985	-3.1719	-3.7173
	140.8184	43.1804	0.1832	1.
Cd(Pent)2,aq		Cd(CH3CH2CH2CH2CO2)2		
Cd(Pent)2,aq		Cd(1)(10)H(18)O(4)		
ref:27		31.Aug.93		
	-187389.	-288726.	65.449	
	26.6621	57.3222	-16.7845	-5.1486
	314.8482	104.3513	-0.0300	0.
Cd(Prop)+		CdCH3CH2CO2+		
Cd(Prop)+		Cd(1)(3)H(5)O(2)+(1)		
ref:27		15.Sep.93		
	-107908.	-142338.	12.488	
	8.2526	12.3660	0.8957	-3.2901
	90.3786	25.0710	0.3636	1.
Cd(Prop)2,aq		Cd(CH3CH2CO2)2		
Cd(Prop)2,aq		Cd(1)(6)H(10)O(4)		
ref:27		15.Sep.93		
	-196520.	-267043.	38.207	
	17.5895	35.1655	-8.0682	-4.2326
	190.4317	61.1072	-0.0300	0.
Ce+2		Ce(+2)		
Ce+2		Ce(1)+(2)		
ref:21		9.Dec.97		
	-74900.	-70300.	1.400	
	0.1315	-7.4538	8.6653	-2.4708
	8.5755	-5.5196	1.0376	2.
Ce+3		Ce(+3)		
Ce+3		Ce(1)+(3)		
ref:21		10.Dec.97		
	-161600.	-167400.	-49.000	
	-3.4833	-16.2789	12.1302	-2.1059
	-0.3550	-12.7510	2.3265	3.
CeOH+2		CeOH+2		
CeOH+2		Ce(1)(1)H(1)+(2)		
ref:25		1.Dec.95		
	-206800.	-218200.	-2.7	
	2.7525	-1.0606	6.1673	-2.7351
	-3.6840	-9.9807	1.1000	2.
CeO+		CeO+		
CeO+		Ce(1)(1)O(1)+(1)		
ref:25		1.Dec.95		
	-195900.	-208900.	13.4	
	2.8409	-0.8430	6.0772	-2.7441
	-35.2162	-18.5361	0.3491	1.
CeO2-		CeO2-		
CeO2-		Ce(1)(2)O(2)-(1)		
ref:25		1.Dec.95		
	-222100.	-230300.	38.6	
	5.0465	4.5418	3.9626	-2.9667
	-51.0768	-26.2767	1.0449	-1.
CeCl+2		CeCl+2		
CeCl+2		Ce(1)(1)Cl(1)+(2)		
ref:25		1.Dec.95		
	-193400.	-203800.	-22.1	
	-0.1746	-8.1993	8.5925	-2.4399
	14.4780	-4.6011	1.3914	2.
CeCl2+		CeCl2+		
CeCl2+		Ce(1)(1)Cl(2)+(1)		
ref:25		1.Dec.95		
	-224400.	-242300.	-5.1	
	2.9511	-0.5763	5.9777	-2.7551
	2.5793	-6.3002	0.6305	1.
CeCl3,aq		CeCl3(aq)		
CeCl3(aq)		Ce(1)(1)Cl(3)		
ref:25		1.Dec.95		
	-255200.	-283500.	2.6	
	6.4972	8.0855	2.5659	-3.1132
	-31.3975	-15.9948	-0.0300	0.
CeCl4-		CeCl4-		
CeCl4-		Ce(1)(1)Cl(4)-(1)		
ref:25		1.Dec.95		

	-286100.	-327600.	0.3	
	11.2558	19.7046	-2.0003	-3.5935
	-67.0564	-33.6847	1.6238	-1.
CeF+2		CeF+2		
CeF+2		Ce(1)(1)F(1)+(2)		
ref:25		1.Dec.95		
	-234700.	-242000.	-14.3	
	-2.8281	-14.6820	11.5084	-2.1719
	15.8334	-3.7656	1.2776	2.
CeF2+		CeF2+		
CeF2+		Ce(1)(1)F(2)+(1)		
ref:25		1.Dec.95		
	-306200.	-324100.	-10.1	
	-2.5538	-14.0092	11.2373	-2.1998
	8.9628	-4.3051	0.7003	1.
CeF3,aq		CeF3(aq)		
CeF3(aq)		Ce(1)(1)F(3)		
ref:25		1.Dec.95		
	-376600.	-409300.	-19.7	
	-2.2777	-13.3348	10.9718	-2.2276
	-21.3883	-12.5158	-0.0300	0.
CeF4-		CeF4-		
CeF4-		Ce(1)(1)F(4)-(1)		
ref:25		1.Dec.95		
	-446500.	-498900.	-46.1	
	-0.8995	-9.9693	9.6488	-2.3668
	-45.3959	-28.3978	2.3239	-1.
CeBr+2		CeBr+2		
CeBr+2		Ce(1)(1)Br(1)+(2)		
ref:25		1.Dec.95		
	-186988.	-195709.	-25.01	
	0.8072	-5.8045	8.0178	-2.5389
	13.8428	-4.9716	1.4382	2.
CeHCO3+2		CeHCO3+2		
CeHCO3+2		Ce(1)(1)H(1)C(1)O(3)+(2)		
ref:25		1.Dec.95		
	-304500.	-330200.	-9.5	
	0.6941	-6.0823	8.1309	-2.5275
	35.5792	3.3307	1.2048	2.
CeCO3+		CeCO3+		
CeCO3+		Ce(1)(1)C(1)O(3)+(1)		
ref:25		1.Dec.95		
	-297887.	-309988.	-40.6	
	-0.9160	-10.0134	9.6736	-2.3649
	-21.4145	-16.3769	1.1729	1.
CeH2PO4+2		CeH2PO4+2		
CeH2PO4+2		Ce(1)(1)H(2)P(1)O(4)+(2)		
ref:25		1.Dec.95		
	-435100.	-480100.	-25.9	
	1.6930	-3.6450	7.1770	-2.6282
	39.4058	3.8826	1.4477	2.
CeNO3+2		CeNO3+2		
CeNO3+2		Ce(1)(1)N(1)O(3)+(2)		
ref:25		1.Dec.95		
	-189900.	-223200.	-32.2	
	1.4120	-4.3253	7.4319	-2.6001
	30.1006	0.3288	1.5475	2.
CeIO3+2		CeIO3+2		
CeIO3+2		Ce(1)(1)I(1)O(3)+(2)		
ref:25		1.Dec.95		
	-194792.	-225358.	-28.97	
	0.9706	-5.4027	7.8548	-2.5556
	29.8495	0.4044	1.4967	2.
CeClO4+2		CeClO4+2		
CeClO4+2		Ce(1)(1)Cl(1)O(4)+(2)		
ref:25		1.Dec.95		
	-166246.	-210026.	-36.06	
	3.5099	0.7863	5.4456	-2.8114
	42.4445	4.4497	1.6005	2.
CeSO4+		CeSO4+		
CeSO4+		Ce(1)(1)S(1)O(4)+(1)		
ref:25		1.Dec.95		

	-334500.	-380200.	-12.6	
	1.5220	-4.0622	7.3397	-2.6110
	-22.4285	-15.3380	0.7384	1.
Ce(Ae)+2		CeCH3COO+2		
Ce(Ae)+2		Ce(1)C(2)H(3)O(2)+(2)		
ref:33		10 Sep 92		
	-253590.	-286390.	-25.400	
	2.9487	-0.5825	5.9809	-2.7548
	53.6757	8.8731	1.4382	2.
Ce(Ae)2+		Ce(CH3COO)2+		
Ce(Ae)2+		Ce(1)C(4)H(6)O(4)+(1)		
ref:33		10 Sep 92		
	-344550.	-405090.	-4.300	
	9.5157	15.4527	-0.3219	-3.4177
	95.0022	25.8754	0.6143	1.
Ce(Ae)3,aq		Ce(CH3COO)3		
Ce(Ae)3,aq		Ce(1)C(6)H(9)O(6)		
ref:33		10 Sep 92		
	-434760.	-524960.	10.300	
	16.9252	33.5426	-7.4282	-4.1656
	130.0191	40.1094	-0.0300	0.
Ce+4		Ce(+4)		
Ce+4		Ce(1)+(4)		
ref:21		10.Dec.97		
	-121300.	-137700.	-100.100	
	-4.2792	-18.2247	12.8996	-2.0255
	40.2790	-3.0141	3.6964	4.
Cl-		Cl(-)		
Cl-		Cl(1)-(1)		
ref:2		11.Sep.97		
	-31379.	-39933.	13.560	
	4.0320	4.8010	5.5630	-2.8470
	-4.4000	-5.7140	1.4560	-1.
ClO-		ClO(-)		
ClO-		Cl(1)O(1)-(1)		
ref:21		11.Sep.97		
	-8800.	-25600.	10.000	
	2.3599	-2.0164	6.5356	-2.6955
	3.4786	-6.6974	1.4767	-1.
ClO2-		ClO2(-)		
ClO2-		Cl(1)O(2)-(1)		
ref:21		11.Sep.97		
	4100.	-15900.	24.200	
	5.2163	4.9580	3.7949	-2.9839
	6.4977	-6.9659	1.2637	-1.
ClO3-		ClO3(-)		
ClO3-		Cl(1)O(3)-(1)		
ref:2		12.Jul.87		
	-1900.	-24850.	38.800	
	7.1665	9.7172	1.9307	-3.1807
	8.5561	-5.5401	1.0418	-1.
ClO4-		ClO4(-)		
ClO4-		Cl(1)O(4)-(1)		
ref:21		11.Sep.97		
	-2040.	-30910.	43.500	
	8.1411	15.5654	-7.8077	-3.4224
	16.4500	-6.5700	9.699	-1.
Cm+2		Cm(+2)		
Cm+2		Cm(1)+(2)		
ref:36		16.Jan.98		
	-77100.	-74100.	-4.000	
	0.0094	-7.7516	8.7810	-2.4584
	9.9946	-5.2956	1.1216	2.
Cm+3		Cm(+3)		
Cm+3		Cm(1)+(3)		
ref:36		16.Jan.98		
	-141200.	-147000.	-49.200	
	-2.9633	-15.0142	11.6438	-2.1582
	6.6776	-10.3066	2.3265	3.
Cm+4		Cm(+4)		
Cm+4		Cm(1)+(4)		
ref:36		16.Jan.98		

	-70700.	-88200.	-104.000	
	-4.3164	-18.3144	12.9330	-2.0218
	40.0550	-3.2586	3.7484	4.
Co+2		Co(+2)		
Co+2		Co(1)+(2)		
ref:21		15.Sep.97		
	-13000.	-13900.	-27.000	
	-1.2252	-8.9356	5.3191	-2.4095
	15.2013	-4.6234	1.4769	2.
CoOH+		CoOH(+)		
CoOH+		Co(1)O(1)H(1)+(1)		
ref:21		11.Dec.97		
	-56025.	-68450.	-10.000	
	-0.2561	-8.4039	9.0457	-2.4315
	26.6839	1.8542	0.7003	1.
CoO,aq		CoO(0)		
CoO(aq)		Co(1)O(1)+(0)		
ref:21		11.Dec.97		
	-44000.	-58800.	-17.800	
	-1.5566	-11.5736	10.2791	-2.3004
	8.7031	-2.0567	-0.0300	0.
CoO2-2		CoO2(-2)		
CoO2-2		Co(1)O(2)-(2)		
ref:21		11.Dec.97		
	-63200.	-99000.	-32.700	
	-0.0590	-7.9219	8.8540	-2.4514
	21.8518	-9.4714	3.7127	-2.
CoCl+		CoCl(+)		
CoCl+		Co(1)Cl(1)+(1)		
ref:22		5.Jan.98		
	-45157.	-53965.	-11.270	
	1.8028	-3.3766	7.0702	-2.6394
	22.7656	0.4323	0.7191	1.
CoF+		CoF(+)		
CoF+		Co(1)F(1)+(1)		
ref:22		6.Jan.98		
	-81732.	-93191.	-22.600	
	-0.7659	-9.6486	9.5354	-2.3801
	26.8657	1.3018	0.8926	1.
Co(Ae)+		CoCH3COO+		
Co(Ae)+		Co(1)C(2)H(3)O(2)+(1)		
ref:33		17 Aug 92		
	-103260.	-132080.	-6.500	
	5.0294	4.4992	3.9806	-2.9649
	60.4543	13.7619	0.6472	1.
Co(Ae)2,aq		Co(CH3COO)2		
Co(Ae)2,aq		Co(1)C(4)H(6)O(4)		
ref:33		10 Sep 92		
	-192770.	-251460.	7.500	
	11.9141	21.3120	-2.6321	-3.6599
	115.2121	34.9629	-0.0300	0.
Co(Ae)3-		Co(CH3COO)3-		
Co(Ae)3-		Co(1)C(6)H(9)O(6)-(1)		
ref:33		10 Sep 92		
	-281890.	-373730.	10.400	
	20.3474	41.8989	-10.7127	-4.5110
	198.1420	58.9793	1.4714	-1.
Co(Ala)+		Co(C3H6NO2)+		
Co(Ala)+		Co(1)C(3)H(5)N(1)O(2)+(1)		
ref:27		14.Aug.93		
	-94099.	-136245.	20.000	
	7.8202	11.3148	1.2990	-3.2467
	69.3955	18.1475	0.2482	1.
Co(Ala)2,aq		Co(C3H6NO2)2		
Co(Ala)2,aq		Co(1)C(6)H(12)N(2)O(4)		
ref:27		14.Aug.93		
	-173793.	-259272.	60.000	
	18.0995	36.4127	-8.5626	-4.2842
	145.3423	45.4353	-0.0300	0.
Co(But)+		CoCH3(CH2)2CO2+		
Co(But)+		Co(1)C(4)H(7)O(2)+(1)		
ref:27		14.Aug.93		

	-99984.	-144234.	4.699	
	9.0507	14.3185	0.1200	-3.3708
	100.5727	28.2419	0.4799	1.
Co(But)2,aq		Co(CH3CH2CH2CO2)2		
Co(But)2,aq		Co(1)C(8)H(14)O(4)		
ref:27		14.Aug.93		
	-186135.	-274655.	33.314	
	20.5354	42.3610	-10.9009	-4.5301
	214.6936	69.5400	-0.0300	0.
Co(For)+		CoCHO2+		
Co(For)+		Co(1)C(1)H(1)O(2)+(1)		
ref:27		26.Aug.93		
	-99399.	-118148.	-5.401	
	3.0614	-0.3037	5.8638	-2.7663
	29.6264	3.1004	0.6305	1.
Co(For)2,aq		Co(CHO2)2		
Co(For)2,aq		Co(1)C(2)H(2)O(4)		
ref:27		26.Aug.93		
	-184926.	-223371.	9.997	
	7.7655	11.1773	1.3615	-3.2410
	41.9649	9.5039	-0.0300	0.
Co(Gly)+		Co(C2H4NO2)+		
Co(Gly)+		Co(1)C(2)H(4)N(1)O(2)+(1)		
ref:27		27.Aug.93		
	-95195.	-129082.	15.000	
	5.3683	5.3273	3.6539	-2.9991
	43.5722	8.9227	0.3260	1.
Co(Gly)2,aq		Co(C2H4NO2)2		
Co(Gly)2,aq		Co(1)C(4)H(8)N(2)O(4)		
ref:27		27.Aug.93		
	-175957.	-243427.	55.000	
	12.8601	23.6197	-3.5341	-3.7553
	81.9652	23.4070	-0.0300	0.
Co(Glyc)+		CoCH3OCO2+		
Co(Glyc)+		Co(1)C(2)H(3)O(3)+(1)		
ref:27		30.Aug.93		
	-136871.	-171331.	-0.901	
	4.9137	4.2158	4.0944	-2.9532
	60.0521	13.8754	0.5681	1.
Co(Glyc)2,aq		Co(CH3OCO2)2		
Co(Glyc)2,aq		Co(1)C(4)H(6)O(6)		
ref:27		30.Aug.93		
	-260101.	-329556.	20.386	
	11.7260	20.8501	-2.4444	-3.6408
	115.9914	35.2337	-0.0300	0.
Co(Lac)+		CoCH3CH2OCO2+		
Co(Lac)+		Co(1)C(3)H(5)O(3)+(1)		
ref:27		30.Aug.93		
	-137385.	-179856.	4.799	
	7.1281	9.6212	1.9722	-3.1766
	81.2111	21.5123	0.4799	1.
Co(Lac)2,aq		Co(CH3CH2OCO2)2		
Co(Lac)2,aq		Co(1)C(6)H(10)O(6)		
ref:27		30.Aug.93		
	-261457.	-346408.	33.545	
	16.4707	32.4370	-7.0030	-4.1198
	168.4594	53.4702	-0.0300	0.
Co(Pent)+		CoCH3(CH2)3CO2+		
Co(Pent)+		Co(1)C(5)H(9)O(2)+(1)		
ref:27		31.Aug.93		
	-97933.	-150673.	11.199	
	11.2090	19.5901	-1.9551	-3.5888
	136.9956	41.2096	0.3837	1.
Co(Pent)2,aq		Co(CH3CH2CH2CH2CO2)2		
Co(Pent)2,aq		Co(1)C(10)H(18)O(4)		
ref:27		31.Aug.93		
	-182034.	-286935.	48.320	
	25.1672	53.6696	-15.3437	-4.9976
	303.7853	100.5060	-0.0300	0.
Co(Prop)+		CoCH3CH2CO2+		
Co(Prop)+		Co(1)C(3)H(5)O(2)+(1)		
ref:27		15.Sep.93		

	-101557.	-138347.	-0.601	
	6.9775	9.2545	2.1152	-3.1615
	86.5252	23.1002	0.5608	1.
Co(Prop)2,aq		Co(CH ₃ CH ₂ CO ₂) ₂		
Co(Prop)2,aq		Co(1)C(6)H(10)O(4)		
ref:27		15.Sep.93		
	-189405.	-263492.	21.078	
	16.0946	31.5192	-6.6433	-4.0819
	179.3685	57.2620	-0.0300	0.
Co+3		Co(+3)		
Co+3		Co(1)+(3)		
ref:21		5.Jan.98		
	32000.	22000.	-73.000	
	-2.8678	-14.7777	11.5439	-2.1680
	16.4739	-8.0659	2.6901	3.
CoOH+2		CoOH(+2)		
CoOH+2		Co(1)O(1)H(1)+(2)		
ref:21		11.Dec.97		
	-23000.	-36100.	-27.800	
	-1.0884	-10.4353	9.8425	-2.3475
	16.5492	-4.1549	1.4769	2.
Cr+2		Cr(+2)		
Cr+2		Cr(1)+(2)		
ref:21		15.Sep.97		
	-39400.	-39000.	-24.200	
	-8036	-9.7400	9.5688	-2.3762
	15.0502	-4.5215	1.4287	2.
Cr+3		Cr(+3)		
Cr+3		Cr(1)+(3)		
ref:21		15.Sep.97		
	-49300.	-60000.	-77.000	
	-2.7824	-14.5709	11.4661	-2.1765
	17.9914	-7.6992	2.7403	3.
CrOH+2		CrOH(+2)		
CrOH+2		Cr(1)O(1)H(1)+(2)		
ref:21		16.Sep.97		
	-100500.	-118600.	-46.100	
	-2.6482	-14.2423	11.3347	-2.1901
	31.4127	.1024	1.7607	2.
CrO+		CrO(+)		
CrO+		Cr(1)O(1)+(1)		
ref:21		16.Sep.97		
	-92800.	-105000.	-26.300	
	-1.1728	-10.6385	9.9161	-2.3391
	-3.1901	-9.3085	.9437	1.
CrO2-		CrO2(-)		
CrO2-		Cr(1)O(2)-(1)		
ref:21		16.Sep.97		
	-125300.	-148300.	-6.700	
	1.2013	-4.8403	7.6350	-2.5788
	10.0021	-7.2511	1.7331	-1.
CrO4-2		CrO4(-2)		
CrO4-2		Cr(1)O(4)-(2)		
ref:21		11.Sep.97		
	-174800.	-210930.	13.800	
	5.4891	5.6223	3.5382	-3.0113
	-2.8608	-15.7861	3.0024	-2.
Cr2O7-2		Cr2O7(-2)		
Cr2O7-2		Cr(2)O(7)-(2)		
ref:21		11.Sep.97		
	-312970.	-355370.	72.000	
	12.4303	22.5680	-3.1161	-3.7119
	8.6580	-8.9622	2.1216	-2.
Cs+		Cs(+)		
Cs+		Cs(1)+(1)		
ref:2		1.Jul.87		
	-69710.	-61670.	31.750	
	6.1475	-1.1309	4.2094	-2.7736
	6.2700	-5.7360	.0974	1.
CsOH,aq		CsOH		
CsOH(aq)		Cs(1)O(1)H(1)+(0)		
ref:21		12.Sep.97		

	-105000.	-112300.	35.900	
	5.6685	6.0571	3.3740	-3.0293
	-12.8052	-9.5325	-0.0300	0.
CsCl,aq		CsCl(0)		
CsCl(aq)		Cs(1)Cl(1)+(0)		
ref:22		13.Sep.97		
	-100900.	-100950.	52.580	
	8.4010	12.7344	.7379	-3.3054
	-6.0563	-7.9234	.2000	0.
CsBr,aq		CsBr(0)		
CsBr(aq)		Cs(1)Br(1)+(0)		
ref:22		13.Sep.97		
	-94610.	-88490.	57.300	
	9.5646	15.5757	-.3789	-3.4228
	-9.4143	-8.4469	-.0010	0.
CsI,aq		CsI(0)		
CsI(aq)		Cs(1)I(1)+(0)		
ref:22		13.Sep.97		
	-83460.	-77820.	60.300	
	11.6513	20.6709	-2.3815	-3.6335
	-7.2120	-8.0048	.1000	0.
Cs(Ae),aq		CsCH ₃ COO		
Cs(Ae),aq		Cs(1)C(2)H(3)O(2)		
ref:33		19 Aug 92		
	-158010.	-176320.	57.500	
	11.7865	20.9974	-2.5020	-3.6469
	29.6579	5.2264	-0.0300	0.
Cs(Ae)2-		Cs(CH ₃ COO)2-		
Cs(Ae)2-		Cs(1)C(4)H(6)O(4)-(1)		
ref:33		10 Sep 92		
	-245900.	-293570.	73.200	
	19.8839	40.7670	-10.2671	-4.4642
	71.0739	17.8580	0.5208	-1.
Cu+		Cu(+)		
Cu+		Cu(1)+(1)		
ref:21		1.May.97		
	11950.	17132.	9.700	
	.8070	-5.8040	8.0165	-2.5390
	17.9233	-2.2438	.4046	1.
CuCl,aq		CuCl(0)		
CuCl(aq)		Cu(1)Cl(1)+(0)		
ref:22		14.Sep.97		
	-22608.	-26338.	22.060	
	4.1084	2.2530	4.8575	-2.8721
	17.3292	.9670	-.0380	0.
CuCl2-		CuCl2(-1)		
CuCl2-		Cu(1)Cl(2)-(1)		
ref:22		14.Sep.97		
	-58038.	-72903.	26.960	
	8.3943	12.7182	0.7442	-3.3047
	36.7555	3.6846	1.2219	-1.
CuCl3-2		CuCl3(-2)		
CuCl3-2		Cu(1)Cl(3)-(2)		
ref:22		14.Sep.97		
	-89963.	-118524.	23.280	
	13.2482	24.5700	-3.9140	-3.7947
	63.9808	7.9089	2.8579	-2.
Cu(Ae),aq		CuCH ₃ COO		
Cu(Ae),aq		Cu(1)C(2)H(3)O(2)		
ref:33		19 Aug 92		
	-76770.	-99970.	28.700	
	7.3009	10.0483	1.7946	-3.1943
	56.0175	14.3883	-0.0300	0.
Cu(Ae)2-		Cu(CH ₃ COO)2-		
Cu(Ae)2-		Cu(1)C(4)H(6)O(4)-(1)		
ref:33		10 Sep 92		
	-165000.	-219740.	37.000	
	15.0715	29.0205	-5.6592	-3.9786
	127.5564	35.7339	1.0691	-1.
Cu+2		Cu(+2)		
Cu+2		Cu(1)+(2)		
ref:2		1.Jul.87		

	15675.	15700.	-23.200	
	-1.1021	-10.4726	9.8662	-2.3461
	20.3000	-4.3900	1.4769	2.
CuOH+		CuOH(+1)		
CuOH+		Cu(1)O(1)H(1)+(1)		
ref:21		1.May.97		
	-30200.	-41700.	-6.100	
	-.0277	-7.8416	8.8142	-2.4547
	21.5062	.2246	.6472	1.
CuO,aq		CuO(0)		
CuO(aq)		Cu(1)O(1)+(0)		
ref:21		1.May.97		
	-20800.	-34200.	-12.400	
	-.3937	-8.7360	9.1675	-2.4178
	12.1490	-3.3197	.7384	0.
CuO2-2		CuO2(-2)		
CuO2-2		Cu(1)O(2)-(2)		
ref:21		1.May.97		
	-41200.	-74400.	-23.100	
	.6024	-6.3038	8.2129	-2.5183
	7.6532	-13.9325	3.5647	-2.
CuCl+		CuCl(+1)		
CuCl+		Cu(1)Cl(1)+(1)		
ref:22		14.Sep.97		
	-16250.	-23847.	-6.510	
	1.7480	-3.5103	7.1228	-2.6338
	29.3560	2.9530	.6472	1.
CuCl2,aq		CuCl2(0)		
CuCl2(aq)		Cu(1)Cl(2)+(0)		
ref:22		14.Sep.97		
	-46142.	-64160.	0.780	
	5.0806	4.6270	3.9244	-2.9702
	37.5948	8.0108	-.0380	0.
CuCl3-		CuCl3(-1)		
CuCl3-		Cu(1)Cl(3)-(1)		
ref:22		14.Sep.97		
	-75338.	-106846.	-2.210	
	9.6259	15.7252	-.4377	-3.4290
	63.8881	11.7108	1.6605	-1.
CuCl4-2		CuCl4(-2)		
CuCl4-2		Cu(1)Cl(4)-(2)		
ref:22		14.Sep.97		
	-103579.	-152534.	-18.480	
	14.6696	28.0407	-5.2782	-3.9382
	87.5037	14.0532	3.4923	-2.
CuF+		CuF(+1)		
CuF+		Cu(1)F(1)+(1)		
ref:22		15.Sep.97		
	-53739.	-64285.	-18.840	
	-.8164	-9.7719	9.5838	-2.3750
	33.5733	3.8226	.8334	1.
Cu(Ae)+		CuCH ₃ COO+		
Cu(Ae)+		Cu(1)C(2)H(3)O(2)+(1)		
ref:33		19 Aug 92		
	-75640.	-103120.	-1.300	
	4.9722	4.3620	4.0290	-2.9592
	62.4950	14.7244	0.5681	1.
Cu(Ae)2,aq		Cu(CH ₃ COO)2		
Cu(Ae)2,aq		Cu(1)C(4)H(6)O(4)		
ref:33		19 Aug 92		
	-165820.	-222690.	14.200	
	11.8801	21.2264	-2.5925	-3.6564
	120.6150	36.8408	-0.0300	0.
Cu(Ae)3-		Cu(CH ₃ COO)3-		
Cu(Ae)3-		Cu(1)C(6)H(9)O(6)-(1)		
ref:33		10 Sep 92		
	-255810.	-345320.	18.990	
	20.2654	41.7019	-10.6422	-4.5029
	206.0700	62.1534	1.3408	-1.
Cu(Ala)+		Cu(C ₃ H ₆ NO ₂)+		
Cu(Ala)+		Cu(1)C(3)H(6)N(1)O(2)+(1)		
ref:27		16.Aug.93		

-70595.	-109970.	25.000	
7.7387	11.1165	1.3762	-3.2385
70.7759	19.1100	0.0974	1.
Cu(Ala)2,aq	Cu(C3H6NO2)2		
Cu(Ala)2,aq	Cu(1)C(6)H(12)N(2)O(4)		
ref:27	16.Aug.93		
-154654.	-237360.	65.000	
18.0655	36.3271	-8.5230	-4.2807
150.7452	47.3132	-0.0300	0.
Cu(But)+	CuCH3(CH2)2CO2+		
Cu(But)+	Cu(1)C(4)H(7)O(2)+(1)		
ref:27	14.Aug.93		
-71854.	-114768.	9.880	
8.9947	14.1844	0.1682	-3.3653
102.6486	29.2044	0.4046	1.
Cu(But)2,aq	Cu(CH3CH2CH2CO2)2		
Cu(But)2,aq	Cu(1)C(8)H(14)O(4)		
ref:27	14.Aug.93		
-158470.	-245176.	40.094	
20.5014	42.2754	-10.8613	-4.5266
220.0963	71.4180	-0.0300	0.
Cu(For)+	CuCHO2+		
Cu(For)+	Cu(1)C(1)H(1)O(2)+(1)		
ref:27	16.Aug.93		
-70888.	-88300.	-0.220	
3.0050	-0.4430	5.9211	-2.7606
31.6869	4.0629	0.5536	1.
Cu(For)2,aq	Cu(CHO2)2		
Cu(For)2,aq	Cu(1)C(2)H(2)O(4)		
ref:27	16.Aug.93		
-156551.	-193183.	16.777	
7.7315	11.0979	1.3851	-3.2377
47.3676	11.3819	-0.0300	0.
Cu(Gly)+	Cu(C2H4NO2)+		
Cu(Gly)+	Cu(1)C(2)H(4)N(1)O(2)+(1)		
ref:27	16.Aug.93		
-71295.	-102408.	25.000	
5.2864	5.1238	3.7413	-2.9907
44.9395	9.8852	0.1739	1.
Cu(Gly)2,aq	Cu(C2H4NO2)2		
Cu(Gly)2,aq	Cu(1)C(4)H(8)N(2)O(4)		
ref:27	16.Aug.93		
-156477.	-221770.	63.000	
12.8262	23.5338	-3.4944	-3.7518
87.3679	25.2850	-0.0300	0.
Cu(Glyc)+	CuCH3OCO2+		
Cu(Glyc)+	Cu(1)C(2)H(3)O(3)+(1)		
ref:27	16.Aug.93		
-109438.	-142561.	4.280	
4.8555	4.0750	4.1470	-2.9474
62.0660	14.8378	0.4862	1.
Cu(Glyc)2,aq	Cu(CH3OCO2)2		
Cu(Glyc)2,aq	Cu(1)C(4)H(6)O(6)		
ref:27	16.Aug.93		
-233022.	-300664.	27.166	
11.6921	20.7705	-2.4205	-3.6376
121.3943	37.1116	-0.0300	0.
Cu(Lac)+	CuCH3CH2OCO2+		
Cu(Lac)+	Cu(1)C(3)H(5)O(3)+(1)		
ref:27	16.Aug.93		
-110347.	-151481.	9.980	
7.0703	9.4856	2.0447	-3.1710
83.2379	22.4748	0.3993	1.
Cu(Lac)2,aq	Cu(CH3CH2OCO2)2		
Cu(Lac)2,aq	Cu(1)C(6)H(10)O(6)		
ref:27	16.Aug.93		
-235047.	-318184.	40.325	
16.4367	32.3514	-6.9634	-4.1163
173.8623	55.3481	-0.0300	0.
Cu(Pent)+	CuCH3(CH2)3CO2+		
Cu(Pent)+	Cu(1)C(5)H(9)O(2)+(1)		
ref:27	16.Aug.93		

-69817.	-121221.	16.380	
11.1516	19.4472	-1.8933	-3.5828
139.0314	42.1721	0.3041	1.
Cu(Pent)2,aq	Cu(CH3CH2CH2CH2CO2)2		
Cu(Pent)2,aq	Cu(1)C(10)H(18)O(4)		
ref:27	16.Aug.93		
-154382.	-257470.	55.100	
25.1332	53.5902	-15.3200	-4.9943
309.1882	102.3839	-0.0300	0.
Cu(Prop)+	CuCH3CH2CO2+		
Cu(Prop)+	Cu(1)C(3)H(5)O(2)+(1)		
ref:27	16.Aug.93		
-74124.	-109577.	4.580	
6.9197	9.1126	2.1719	-3.1556
88.5486	24.0627	0.4799	1.
Cu(Prop)2,aq	Cu(CH3CH2CO2)2		
Cu(Prop)2,aq	Cu(1)C(6)H(10)O(4)		
ref:27	16.Aug.93		
-162995.	-235268.	27.858	
16.0606	31.4336	-6.6036	-4.0784
184.7714	59.1399	-0.0300	0.
Dy+2	Dy(+2)		
Dy+2	Dy(1)+(2)		
ref:21	10.Dec.97		
-102800.	-99900.	-3.600	
0.0206	-7.7248	8.7724	-2.4596
9.8690	-5.3159	1.1144	2.
Dy+3	Dy(+3)		
Dy+3	Dy(1)+(3)		
ref:2	13.Jul.87		
-158700.	-166500.	-55.200	
-3.0003	-15.1074	11.6879	-2.1545
9.5076	-9.4919	2.3792	3.
DyOH+2	DyOH+2		
DyOH+2	Dy(1)O(1)H(1)+(2)		
ref:25	1.Dec.95		
-204700.	-216500.	-10.9	
2.6154	-1.3941	6.2950	-2.7213
2.8769	-8.0863	1.2206	2.
DyO+	DyO+		
DyO+	Dy(1)O(1)+(1)		
ref:25	1.Dec.95		
-193400.	-209000.	4.6	
2.6935	-1.2056	6.2255	-2.7291
-28.1511	-16.4991	0.4799	1.
DyO2-	DyO2-		
DyO2-	Dy(1)O(2)-(1)		
ref:25	1.Dec.95		
-226400.	-237700.	28.6	
4.7414	3.7936	4.2636	-2.9357
-37.5629	-22.0601	1.1949	-1.
DyCl+2	DyCl+2		
DyCl+2	Dy(1)Cl(1)+(2)		
ref:25	1.Dec.95		
-190400.	-203200.	-29.5	
-0.2729	-8.4410	9.0516	-2.4299
24.6385	-1.4386	1.5067	2.
DyCl2+	DyCl2+		
DyCl2+	Dy(1)Cl(2)+(1)		
ref:25	1.Dec.95		
-221400.	-242200.	-14.4	
2.8449	-0.8347	6.0770	-2.7444
21.6041	-0.1300	0.7686	1.
DyCl3,aq	DyCl3(aq)		
DyCl3(aq)	Dy(1)Cl(3)		
ref:25	1.Dec.95		
-252200.	-284200.	-9.6	
6.3268	7.6646	2.7424	-3.0958
-1.3925	-5.5657	-0.0300	0.
DyCl4-	DyCl4-		
DyCl4-	Dy(1)Cl(4)-(1)		
ref:25	1.Dec.95		

-283000.	-329600.	-16.4	
11.1518	19.4467	-1.8909	-3.5828
-18.8614	-17.7460	1.8776	-1.
DyF+2	DyF+2		
DyF+2	Dy(1)F(1)+(2)		
ref:25	1.Dec.95		
-232400.	-241100.	-18.4	
-2.9451	-14.9683	11.6229	-2.1601
25.4842	-0.6031	1.3375	2.
DyF2+	DyF2+		
DyF2+	Dy(1)F(2)+(1)		
ref:25	1.Dec.95		
-304500.	-323800.	-14.1	
-2.6837	-14.3302	11.3726	-2.1865
27.3446	1.8651	0.7686	1.
DyF3,aq	DyF3(aq)		
DyF3(aq)	Dy(1)F(3)		
ref:25	1.Dec.95		
-375400.	-409800.	-25.0	
-2.4481	-13.7557	11.1483	-2.2102
8.6168	-2.0868	-0.0300	0.
DyF4-	DyF4-		
DyF4-	Dy(1)F(4)-(1)		
ref:25	1.Dec.95		
-445600.	-500800.	-54.9	
-1.0402	-10.3182	9.7972	-2.3523
1.7998	-12.4591	2.4692	-1.
DyHCO3+2	DyHCO3+2		
DyHCO3+2	Dy(1)H(1)C(1)O(3)+(2)		
ref:25	1.Dec.95		
-301300.	-329700.	-18.2	
0.6017	-6.3082	8.2207	-2.5181
45.9011	6.4931	1.3375	2.
DyCO3+	DyCO3+		
DyCO3+	Dy(1)C(1)O(3)+(1)		
ref:25	1.Dec.95		
-295800.	-310100.	-48.1	
-0.9340	-10.0536	9.6821	-2.3633
-16.3311	-14.9713	1.2858	1.
DyH2PO4+2	DyH2PO4+2		
DyH2PO4+2	Dy(1)H(2)P(1)O(4)+(2)		
ref:25	1.Dec.95		
-431800.	-479700.	-34.8	
1.6038	-3.8668	7.2721	-2.6190
49.8123	7.0451	1.5897	2.
DyNO3+2	DyNO3+2		
DyNO3+2	Dy(1)N(1)O(3)+(2)		
ref:25	1.Dec.95		
-185400.	-223200.	-43.8	
1.3347	-4.5159	7.5104	-2.5922
40.8317	3.4913	1.7247	2.
DySO4+	DySO4+		
DySO4+	Dy(1)S(1)O(4)+(1)		
ref:25	1.Dec.95		
-341600.	-379000.	-17.9	
-1.4956	-4.1268	7.3650	-2.6083
-17.6132	-13.9325	0.8222	1.
Dy(Ae)+2	DyCH3COO+2		
Dy(Ae)+2	Dy(1)C(2)H(3)O(2)+(2)		
ref:33	11.Sep.92		
-250590.	-286150.	-34.200	
2.8591	-0.7970	6.0561	-2.7460
64.0720	12.0356	1.5790	2.
Dy(Ae)2+	Dy(CH3COO)2+		
Dy(Ae)2+	Dy(1)C(4)H(6)O(4)+(1)		
ref:33	11.Sep.92		
-341450.	-405710.	-16.300	
9.4257	15.2330	-0.2363	-3.4086
114.4678	32.0457	0.8002	1.
Dy(Ae)3,aq	Dy(CH3COO)3		
Dy(Ae)3,aq	Dy(1)C(6)H(9)O(6)		
ref:33	11.Sep.92		

	-431590.	-526620.	-5.400	-4.1484	
	16.7546	33.1279	-7.2677		0.
	160.0243	50.5384	-0.0300		
Dy+4		Dy(+4)			
Dy+4		Dy(1)+(4)			
ref:21		10.Dec.97			
	-55700.	-73400.	-104.000		
	-4.3164	-18.3144	12.9330	-2.0218	
	40.0550	-3.2586	3.7484		4.
Er+2		Er(+2)			
Er+2		Er(1)+(2)			
ref:21		10.Dec.97			
	-91600.	-89150.	-5.500		
	-0.0104	-7.7996	8.7981	-2.4565	
	10.3150	-5.2548	1.1437		2.
Er+3		Er(+3)			
Er+3		Er(1)+(3)			
ref:2		13.Jul.87			
	-159900.	-168500.	-58.300		
	-3.3041	-15.8492	11.9794	-2.1238	
	8.2815	-10.0215	2.4115		3.
ErOH+2		ErOH+2			
ErOH+2		Er(1)O(1)H(1)+(2)			
ref:25		1.Dec.95			
	-206000.	-219000.	-14.6		
	2.5526	-1.5478	6.3556	-2.7149	
	5.8637	-7.2307	1.2776		2.
ErO+		ErO+			
ErO+		Er(1)O(1)+(1)			
ref:25		1.Dec.95			
	-194800.	-211600.	0.6		
	2.6315	-1.3551	6.2803	-2.7229	
	-24.9068	-15.5621	0.5394		1.
ErO2-		ErO2-			
ErO2-		Er(1)O(2)-(1)			
ref:25		1.Dec.95			
	-228800.	-241500.	24.0		
	4.6152	3.4877	4.3792	-2.9231	
	-31.3910	-20.1453	1.2669		-1.
ErCl+2		ErCl+2			
ErCl+2		Er(1)Cl(1)+(2)			
ref:25		1.Dec.95			
	-191700.	-205400.	-33.2		
	-0.6062	-9.2588	9.3810	-2.3961	
	21.6825	-2.6303	1.5579		2.
ErCl2+		ErCl2+			
ErCl2+		Er(1)Cl(2)+(1)			
ref:25		1.Dec.95			
	-222600.	-244700.	-19.0		
	2.4799	-1.7255	6.4270	-2.7076	
	15.6173	-2.4550	0.8449		1.
ErCl3,aq		ErCl3(aq)			
ErCl3(aq)		Er(1)Cl(3)			
ref:25		1.Dec.95			
	-253400.	-287100.	-15.6		
	5.8915	6.6023	3.1580	-3.0518	
	-12.6987	-9.4955	-0.0300		0.
ErCl4-		ErCl4-			
ErCl4-		Er(1)Cl(4)-(1)			
ref:25		1.Dec.95			
	-284200.	-333200.	-24.7		
	10.7108	18.3744	-1.4789	-3.5385	
	-34.9400	-23.7519	2.0080		-1.
ErF+2		ErF+2			
ErF+2		Er(1)F(1)+(2)			
ref:25		1.Dec.95			
	-233700.	-242900.	-20.4		
	-3.2867	-15.8018	11.9486	-2.1257	
	22.3012	-1.7948	1.3642		2.
ErF2+		ErF2+			
ErF2+		Er(1)F(2)+(1)			
ref:25		1.Dec.95			

	-305900.	-325700.	-16.1		
	-3.0637	-15.2566	11.7333	-2.1482	
	20.9464	-0.4598	0.8002		1.
ErF3,aq		ErF3(aq)			
ErF3(aq)		Er(1)F(3)			
ref:25		1.Dec.95			
	-376800.	-411900.	-27.7		
	-2.8834	-14.8179	11.5640	-2.1663	
	-2.6893	-6.0166	-0.0300		0.
ErF4-		ErF4-			
ErF4-		Er(1)F(4)-(1)			
ref:25		1.Dec.95			
	-447100.	-503500.	-59.3		
	-1.5016	-11.4431	10.2350	-2.3058	
	-14.8367	-18.4650	2.5390		-1.
ErHCO3+2		ErHCO3+2			
ErHCO3+2		Er(1)H(1)C(1)O(3)+(2)			
ref:25		1.Dec.95			
	-302600.	-332200.	-22.6		
	0.2724	-7.1116	8.5345	-2.4849	
	43.0537	5.3015	1.4006		2.
ErCO3+		ErCO3+			
ErCO3+		Er(1)C(1)O(3)+(1)			
ref:25		1.Dec.95			
	-297200.	-312600.	-51.9		
	-1.0539	-10.3510	9.8086	-2.3510	
	-17.2819	-15.5009	1.3480		1.
ErH2PO4+2		ErH2PO4+2			
ErH2PO4+2		Er(1)H(2)P(1)O(4)+(2)			
ref:25		1.Dec.95			
	-433100.	-482200.	-39.3		
	1.2754	-4.6600	7.5655	-2.5863	
	46.9907	5.8534	1.6556		2.
ErNO3+2		ErNO3+2			
ErNO3+2		Er(1)N(1)O(3)+(2)			
ref:25		1.Dec.95			
	-186600.	-226000.	-49.7		
	1.0128	-5.3011	7.8182	-2.5598	
	38.1891	2.2996	1.8100		2.
ErSO4+		ErSO4+			
ErSO4+		Er(1)S(1)O(4)+(1)			
ref:25		1.Dec.95			
	-342700.	-381048.	-21.1		
	1.3743	-4.4228	7.4814	-2.5961	
	-18.7119	-14.4621	0.8683		1.
Er(Ae)+2		ErCH3COO+2			
Er(Ae)+2		Er(1)C(2)H(3)O(2)+(2)			
ref:33		11 Sept 92			
	-251770.	-288520.	-38.600		
	2.5305	-1.6022	6.3787	-2.7127	
	61.2455	10.8439	1.6444		2.
Er(Ae)2+		Er(CH3COO)2+			
Er(Ae)2+		Er(1)C(4)H(6)O(4)+(1)			
ref:33		11 Sept 92			
	-342620.	-408540.	-22.300		
	9.0662	14.3528	0.1146	-3.3722	
	108.6293	29.7206	0.8926		1.
Er(Ae)3,aq		Er(CH3COO)3			
Er(Ae)3,aq		Er(1)C(6)H(9)O(6)			
ref:33		11 Sept 92			
	-432750.	-529990.	-13.200		
	16.3195	32.0657	-6.8520	-4.1045	
	148.7181	46.6086	-0.0300		0.
Er+4		Er(+4)			
Er+4		Er(1)+(4)			
ref:21		10.Dec.97			
	-28100.	-45900.	-104.700		
	-4.3256	-18.3402	12.9496	-2.0207	
	40.0003	-3.3197	3.7615		4.
Eu+2		Eu(+2)			
Eu+2		Eu(1)+(2)			
ref:21		10.Dec.97			

	-129200.	-126150.	-2.400		
	0.0407	-7.6776	8.7578	-2.4615	
	9.5539	-5.3567	1.0929		2.
EuCl+		EuCl+			
EuCl+		Eu(1)Cl(1)+(1)			
ref:25		1.Dec.95			
	-161000.	-164000.	19.5		
	5.1742	4.8499	3.8487	-2.9794	
	36.2777	6.6125	0.2557		1.
EuCl2,aq		EuCl2(aq)			
EuCl2(aq)		Eu(1)Cl(2)			
ref:25		1.Dec.95			
	-193600.	-204600.	35.00		
	9.1152	14.4740	0.0641	-3.3773	
	58.1447	15.1277	-0.0300		0.
EuCl3-		EuCl3-			
EuCl3-		Eu(1)Cl(3)-(1)			
ref:25		1.Dec.95			
	-226000.	-246800.	44.7		
	13.9460	26.2721	-4.5790	-3.8650	
	91.9571	23.7331	0.9527		-1.
EuCl4-2		EuCl4-2			
EuCl4-2		Eu(1)Cl(4)-(2)			
ref:25		1.Dec.95			
	-258500.	-290600.	48.6		
	19.4730	39.7656	-9.8784	-4.4228	
	131.0045	32.4289	2.4755		-2.
EuF+		EuF+			
EuF+		Eu(1)F(1)+(1)			
ref:25		1.Dec.95			
	-194600.	-202200.	1.7		
	2.6504	-1.3110	6.2669	-2.7247	
	41.1681	7.4480	0.5256		1.
EuF2,aq		EuF2(aq)			
EuF2(aq)		Eu(1)F(1)			
ref:25		1.Dec.95			
	-261000.	-282200.	-4.1		
	3.5866	0.9784	5.3596	-2.8193	
	63.8849	17.1229	-0.0300		0.
EuF3-		EuF3-			
EuF3-		Eu(1)F(3)-(1)			
ref:25		1.Dec.95			
	-327700.	-365700.	-20.5		
	5.5026	5.6556	3.5244	-3.0127	
	111.0055	27.2121	1.9339		-1.
EuF4-2		EuF4-2			
EuF4-2		Eu(1)F(4)-(2)			
ref:25		1.Dec.95			
	-394600.	-455700.	-58.1		
	7.6287	10.8431	1.4936	-3.2272	
	161.1451	37.7158	4.0961		-2.
Eu(Ala)+		Eu(C3H6NO2)+			
Eu(Ala)+		Eu(1)C(3)H(6)N(1)O(2)+(1)			
ref:27		14.Aug.93			
	-205330.	-242060.	49.7		
	11.0532	19.2058	-1.7947	-3.5729	
	83.4588	24.4724	-0.2006		1.
Eu(Ala)2,aq		Eu(C3H6NO2)2			
Eu(Ala)2,aq		Eu(1)C(6)H(12)N(2)O(4)			
ref:27		14.Aug.93			
	-281120.	-358510.	98.6		
	21.8707	45.6205	-12.1799	-4.6649	
	180.8470	57.7759	-0.0300		0.
Eu(But)+		EuCH3(CH2)2CO2+			
Eu(But)+		Eu(1)C(4)H(7)O(2)+(1)			
ref:27		14.Aug.93			
	-214119.	-251804.	38.239		
	12.2638	22.1618	-2.9570	-3.6951	
	114.0923	34.5668	-0.0278		1.
Eu(But)2,aq		Eu(CH3CH2CH2CO2)2			
Eu(But)2,aq		Eu(1)C(8)H(14)O(4)			
ref:27		14.Aug.93			

-298525.	-377392.	77.206	
24.3067	51.5686	-14.5188	-4.9107
250.1983	81.8806	-0.0300	0.
Eu(For)+	EuCH2O+		
Eu(For)+	Eu(1)C(1)H(1)O(2)+(1)		
ref:27	26.Aug.93		
-214872.	-227054.	28.139	
6.2755	7.5431	2.7814	-3.0907
43.1702	9.4253	0.1254	1.
Eu(For)2,aq	Eu(CH2O)2		
Eu(For)2,aq	Eu(1)C(2)H(2)O(4)		
ref:27	26.Aug.93		
-300521.	-329314.	53.889	
11.5367	20.3852	-2.2565	-3.6216
77.4695	21.8445	-0.0300	0.
Eu(Gly)+	Eu(C2H4NO2)+		
Eu(Gly)+	Eu(1)C(2)H(2)N(1)O(2)+(1)		
ref:27	27.Aug.93		
-206629.	-234136.	47.919	
8.5841	13.1760	0.5767	-3.3236
57.1645	15.2476	-0.1738	1.
Eu(Gly)2,aq	Eu(C2H4NO2)2		
Eu(Gly)2,aq	Eu(1)C(4)H(8)N(2)O(4)		
ref:27	27.Aug.93		
-283817.	-342929.	94.554	
16.6314	32.8272	-7.1514	-4.1360
117.4699	35.7476	-0.0300	0.
Eu(Glyc)+	EuCH3OCO2+		
Eu(Glyc)+	Eu(1)C(2)H(3)O(3)+(1)		
ref:27	30.Aug.93		
-252044.	-279938.	32.639	
8.1255	12.0562	1.0166	-3.2773
73.5346	20.2002	0.0564	1.
Eu(Glyc)2,aq	Eu(CH3OCO)2		
Eu(Glyc)2,aq	Eu(1)C(4)H(6)O(6)		
ref:27	30.Aug.93		
-374046.	-433849.	64.278	
15.4973	30.0577	-6.0619	-4.0215
151.4961	47.5743	-0.0300	0.
Eu(Lac)+	Eu(CH3CH2OCO2)+		
Eu(Lac)+	Eu(1)C(3)H(5)O(3)+(1)		
ref:27	30.Aug.93		
-252899.	-288803.	38.339	
10.3407	17.4658	-1.1113	-3.5009
94.7163	27.8372	-0.0294	1.
Eu(Lac)2,aq	Eu(CH3CH2OCO)2		
Eu(Lac)2,aq	Eu(1)C(6)H(10)O(6)		
ref:27	30.Aug.93		
-376425.	-451723.	77.437	
20.2419	41.6449	-10.6207	-4.5005
203.9641	65.8108	-0.0300	0.
Eu(Pent)+	EuCH3(CH2)3CO2+		
Eu(Pent)+	Eu(1)C(5)H(9)O(2)+(1)		
ref:27	31.Aug.93		
-211714.	-257888.	44.739	
14.4215	27.4350	-5.0392	-3.9131
150.5000	47.5345	-0.1257	1.
Eu(Prop)+	EuCH3CH2CO2+		
Eu(Prop)+	Eu(1)C(3)H(5)O(2)+(1)		
ref:27	15.Sep.93		
-216648.	-246872.	32.939	
10.1903	17.0986	-0.9659	-3.4858
100.0361	29.4251	0.0521	1.
Eu(Prop)2,aq	Eu(CH3CH2CO)2		
Eu(Prop)2,aq	Eu(1)C(6)H(10)O(4)		
ref:27	15.Sep.93		
-303186.	-367621.	64.970	
19.8658	40.7270	-10.2606	-4.4626
214.8734	69.6025	-0.0300	0.
Eu+3	Eu(+3)		
Eu+3	Eu(1)+(+3)		
ref:2	13.Jul.87		

-137300.	-144700.	-53.000	
-3.1037	-15.3599	11.7871	-2.1440
6.0548	-10.4900	2.3161	3.
EuOH+2	EuOH+2		
EuOH+2	Eu(1)O(1)H(1)+(2)		
ref:25	1.Dec.95		
-183200.	-194373.	-8.0	
2.6569	-1.2969	6.2659	-2.7253
0.5828	-8.7585	1.1815	2.
EuO+	EuO+		
EuO+	Eu(1)O(1)+(1)		
ref:25	1.Dec.95		
-171700.	-186500.	7.7	
2.7458	-1.0743	6.1663	-2.7345
-30.6415	-17.2120	0.4322	1.
EuO2-	EuO2-		
EuO2-	Eu(1)O(2)-(1)		
ref:25	1.Dec.95		
-203600.	-214100.	32.1	
4.8468	4.0541	4.1548	-2.9465
-42.3249	-23.5471	1.1424	-1.
EuCl+2	EuCl+2		
EuCl+2	Eu(1)Cl(1)+(2)		
ref:25	1.Dec.95		
-169100.	-181300.	-26.9	
-0.3777	-8.6968	9.1514	-2.4194
17.8123	-3.6844	1.4671	2.
EuCl2+	EuCl2+		
EuCl2+	Eu(1)Cl(2)+(1)		
ref:25	1.Dec.95		
-200000.	-220100.	-11.1	
2.7262	-1.122	6.1870	-2.7325
8.5409	-4.5117	0.7191	1.
EuCl3,aq	EuCl3(aq)		
EuCl3(aq)	Eu(1)Cl(3)		
ref:25	1.Dec.95		
-230800.	-261800.	-5.3	
6.2132	7.3881	2.8493	-3.0843
-22.7003	-12.9719	-0.0300	0.
EuCl4-	EuCl4-		
EuCl4-	Eu(1)Cl(4)-(1)		
ref:25	1.Dec.95		
-261600.	-306800.	-10.5	
10.9946	19.0660	-1.7473	-3.5671
-52.2612	-29.0648	1.7870	-1.
EuF+2	EuF+2		
EuF+2	Eu(1)F(1)+(2)		
ref:25	1.Dec.95		
-210700.	-219200.	-17.0	
-3.0453	-15.2112	11.7137	-2.1501
18.7828	-2.8489	1.3115	2.
EuF2+	EuF2+		
EuF2+	Eu(1)F(2)+(1)		
ref:25	1.Dec.95		
-282500.	-301700.	-12.7	
-2.7958	-14.6044	11.4813	-2.1752
14.4597	-2.5166	0.7384	1.
EuF3,aq	EuF3(aq)		
EuF3(aq)	Eu(1)F(3)		
ref:25	1.Dec.95		
-353200.	-387300.	-23.1	
-2.5616	-14.0324	11.2555	-2.1988
-12.6911	-9.4929	-0.0300	0.
EuF4-	EuF4-		
EuF4-	Eu(1)F(4)-(1)		
ref:25	1.Dec.95		
-423200.	-477800.	-51.8	
-1.1887	-10.6780	9.9326	-2.3375
-31.3648	-23.7779	2.4042	-1.
EuSO4+	EuSO4+		
EuSO4+	Eu(1)S(1)O(4)+(1)		
ref:25	1.Dec.95		

-320200.	-347200.	-15.3	
-1.4399	-4.2627	7.4184	-2.6027
-20.8829	-14.9306	0.7790	1.
EuHCO3+2	EuHCO3+2		
EuHCO3+2	Eu(1)H(1)C(1)O(3)+(2)		
ref:25	1.Dec.95		
-279800.	-307500.	-15.2	
0.4928	-6.5720	8.3198	-2.5072
38.9648	4.2473	1.2860	2.
EuCO3+	EuCO3+		
EuCO3+	Eu(1)C(1)O(3)+(1)		
ref:25	1.Dec.95		
-274300.	-287900.	-45.4	
-0.9842	-10.1779	9.7343	-2.3581
-19.5642	-15.9695	1.2465	1.
EuH2PO4+2	EuH2PO4+2		
EuH2PO4+2	Eu(1)H(2)P(1)O(4)+(2)		
ref:25	1.Dec.95		
-410500.	-457600.	-31.7	
1.4946	-4.1236	7.3517	-2.6084
42.8673	4.7993	1.5372	2.
EuNO3+2	EuNO3+2		
EuNO3+2	Eu(1)N(1)O(3)+(2)		
ref:25	1.Dec.95		
-165000.	-201800.	-39.7	
1.2198	-4.7951	7.6178	-2.5807
33.7330	1.2455	1.6556	2.
Eu(Ae)+2	EuCH3COO+2		
Eu(Ae)+2	Eu(1)C(2)H(3)O(2)+(2)		
ref:33	10.Sep.92		
-229390.	-264280.	-31.100	
2.7500	-1.0666	6.1690	-2.7348
57.1309	9.7898	1.5269	2.
Eu(Ae)2+	Eu(CH3COO)2+		
Eu(Ae)2+	Eu(1)C(4)H(6)O(4)+(1)		
ref:33	10.Sep.92		
-320420.	-383670.	-12.000	
9.3029	14.9307	-0.1123	-3.3961
101.2917	27.6639	0.7384	1.
Eu(Ae)3,aq	Eu(CH3COO)3		
Eu(Ae)3,aq	Eu(1)C(6)H(9)O(6)		
ref:33	10.Sep.92		
-410690.	-504320.	0.200	
16.6413	32.8512	-7.1605	-4.1370
138.7162	43.1323	-0.0300	0.
Eu(But)+2	EuCH3(CH2)2CO2+2		
Eu(But)+2	Eu(1)C(4)H(7)O(2)+(2)		
ref:27	14.Aug.93		
-225716.	-276036.	-19.856	
6.7698	8.7503	2.3069	-3.1406
97.2089	24.2698	1.3552	2.
Eu(But)2+	Eu(CH3CH2CH2CO2)2+		
Eu(But)2+	Eu(1)C(8)H(14)O(4)+(1)		
ref:27	14.Aug.93		
-313068.	-405964.	14.439	
17.7879	35.6504	-8.2592	-4.2527
197.0574	62.2411	0.3351	1.
Eu(For)+2	EuCH2O+2		
Eu(For)+2	Eu(1)C(1)H(1)O(2)+(2)		
ref:27	26.Aug.93		
-224968.	-249786.	-29.956	
0.7809	-5.8667	8.0384	-2.5364
26.2700	-0.8716	1.5067	2.
Eu(For)2+	Eu(CH2O)2+		
Eu(For)2+	Eu(1)C(2)H(2)O(4)+(1)		
ref:27	26.Aug.93		
-311559.	-354544.	-9.431	
5.1383	4.7630	3.8812	-2.9758
27.6085	2.2050	0.6911	1.
Eu(Pent)+2	EuCH3(CH2)3CO2+2		
Eu(Pent)+2	Eu(1)C(5)H(9)O(2)+(2)		
ref:27	1.Sep.93		

-223706.	-282516.	-13.356	
8.9288	14.0201	0.2395	-3.3585
133.6500	37.2375	1.2610	2.
Eu(Pent)2+	Eu(CH ₃ CH ₂ CH ₂ CH ₂ CO ₂) ₂ +		
Eu(Pent)2+	Eu(1)(C(10)H(18)O(4)+(1)		
ref:27	1.Sep.93		
-309035.	-418206.	29.802	
22.3402	46.7679	-12.6335	-4.7123
283.9817	93.2071	0.0999	1.
Eu(Prop)+2	EuCH ₃ CH ₂ CO ₂ +2		
Eu(Prop)+2	Eu(1)(C(3)H(5)O(2)+(2)		
ref:27	15.Sep.93		
-227972.	-270831.	-25.156	
4.6974	3.6905	4.2941	-2.9315
83.1800	19.1280	1.4382	2.
Eu(Prop)2+	Eu(CH ₃ CH ₂ CO ₂) ₂ +		
Eu(Prop)2+	Eu(1)(C(6)H(10)O(4)+(1)		
ref:27	15.Sep.93		
-317566.	-396115.	1.913	
13.4114	24.9646	-4.0601	-3.8109
163.4871	49.9630	0.5256	1.
Eu+4	Eu(+4)		
Eu+4	Eu(1)+(4)		
ref:21	10.Dec.97		
5000.	-12700.	-103.300	
-4.3071	-18.2951	12.9322	-2.0226
40.0515	-3.2178	3.7353	4.
F-	F(-)		
F-	F(1)-(1)		
ref:2	1.Jul.87		
-67340.	-80150.	-3.150	
6.870	1.3588	7.6033	-2.8352
4.4600	-7.4880	1.7870	-1.
Fe+2	Fe(+2)		
Fe+2	Fe(1)+(2)		
ref:21	11.Sep.97		
-21870.	-22050.	-25.300	
-7.867	-9.6969	9.5479	-2.3780
14.7860	-4.6437	1.4382	2.
FeOH+	FeOH(+1)		
FeOH+	Fe(1)O(1)H(1)+(1)		
ref:21	12.Sep.97		
-65850.	-78080.	-10.000	
-2.561	-8.4039	9.0457	-2.4315
21.4093	.0209	7.003	1.
FeO,aq	FeO		
FeO(aq)	Fe(1)O(1)+(0)		
ref:21	1.May.97		
-50720.	-62950.	-10.000	
-5.029	-9.0053	9.2791	-2.4066
5.8901	-3.0345	-0.0300	0.
FeCl+	FeCl(+)		
FeCl+	Fe(1)Cl(1)+(1)		
ref:22	12.Sep.97		
-53030.	-61260.	-10.060	
2.1468	-2.5367	6.7401	-2.6741
24.6912	1.1617	7.003	1.
FeCl2,aq	FeCl2(0)		
FeCl2(aq)	Fe(1)Cl(2)+(0)		
ref:22	17.Sep.97		
-73480.	-78490.	43.000	
5.5057	5.6650	3.5164	-3.0131
22.9295	2.9135	-0.0380	0.
FeF+	FeF(+1)		
FeF+	Fe(1)F(1)+(1)		
ref:22	15.Sep.97		
-91157.	-101452.	-20.920	
-4.4234	-8.8124	9.2067	-2.4146
26.3785	1.2102	.8683	1.
Fe(Ae)+	Fe(CH ₃ COO)(+)		
Fe(Ae)+	Fe(1)(C(2)H(3)O(2)+(1)		
ref:22	12.Sep.97		

-111900.	-139060.	-1.500	
5.2246	4.9785	3.7863	-2.9848
59.1160	13.5263	.5756	1.
Fe(Ae)2,aq	Fe(CH ₃ COO)2(0)		
Fe(Ae)2(aq)	Fe(1)(C(4)H(6)O(4)+(0)		
ref:22	12.Sep.97		
-201800.	-259100.	11.530	
12.1698	21.9370	-2.8791	-3.6858
113.7691	34.4870	-0.0380	0.
Fe(Ala)+	Fe(C ₃ H ₆ NO ₂)+		
Fe(Ala)+	Fe(1)(C(3)H(6)N(1)O(2)+(1)		
ref:27	14.Aug.93		
-103252.	-145225.	18.058	
8.1661	12.1558	0.9760	-3.2814
69.2856	18.0100	0.2792	1.
Fe(Ala)2,aq	Fe(C ₃ H ₆ NO ₂) ₂		
Fe(Ala)2,aq	Fe(1)(C(6)H(12)N(2)O(4)		
ref:27	14.Aug.93		
-182987.	-268535.	57.246	
18.4732	37.3244	-8.9189	-4.3219
144.5703	45.1671	-0.0300	0.
Fe(But)+	FeCH ₃ (CH ₂) ₂ CO ₂ +		
Fe(But)+	Fe(1)(C(4)H(7)O(2)+(1)		
ref:27	14.Aug.93		
-108714.	-151642.	6.608	
9.3758	15.1141	-0.1954	-3.4037
99.8978	28.1044	0.4496	1.
Fe(But)2,aq	Fe(CH ₃ CH ₂ CH ₂ CO ₂) ₂		
Fe(But)2,aq	Fe(1)(C(8)H(14)O(4)		
ref:27	14.Aug.93		
-194743.	-281765.	35.812	
20.9091	43.2727	-11.2572	-4.5678
213.9216	69.2718	-0.0300	0.
Fe(For)+	FeCHO ₂ +		
Fe(For)+	Fe(1)(C(1)H(1)O(2)+(1)		
ref:27	16.Aug.93		
-108225.	-125651.	-3.492	
3.3887	0.4923	5.5575	-2.7993
29.0080	2.9629	0.6063	1.
Fe(For)2,aq	Fe(CHO ₂) ₂		
Fe(For)2,aq	Fe(1)(C(2)H(2)O(4)		
ref:27	16.Aug.93		
-193710.	-230658.	12.495	
8.1392	12.0952	0.9892	-3.2789
41.1929	9.2357	-0.0300	0.
Fe(Gly)+	Fe(C ₂ H ₄ NO ₂)+		
Fe(Gly)+	Fe(1)(C(2)H(4)N(1)O(2)+(1)		
ref:27	16.Aug.93		
-103038.	-134682.	20.000	
5.6774	6.0837	3.3533	-3.0304
42.4593	8.7852	0.2482	1.
Fe(Gly)2,aq	Fe(C ₂ H ₄ NO ₂) ₂		
Fe(Gly)2,aq	Fe(1)(C(4)H(8)N(2)O(4)		
ref:27	16.Aug.93		
-182709.	-248527.	58.018	
13.2339	24.5311	-3.8903	-3.7930
81.1932	23.1388	-0.0300	0.
Fe(Glyc)+	FeCH ₃ OCO ₂ +		
Fe(Glyc)+	Fe(1)(C(2)H(3)O(3)+(1)		
ref:27	16.Aug.93		
-146010.	-179149.	1.008	
7.4836	10.4932	1.6217	-3.2127
59.3916	13.7379	0.5394	1.
Fe(Glyc)2,aq	Fe(CH ₃ OCO ₂) ₂		
Fe(Glyc)2,aq	Fe(1)(C(4)H(6)O(6)		
ref:27	16.Aug.93		
-269459.	-337416.	22.883	
16.8444	33.3487	-7.3593	-4.1575
115.2196	34.9654	-0.0300	0.
Fe(Lac)+	FeCH ₃ CH ₂ CO ₂ +		
Fe(Lac)+	Fe(1)(C(3)H(5)O(3)+(1)		
ref:27	30.Aug.93		

-147289.	-188437.	6.708	
7.4532	10.4168	1.6567	-3.2095
80.5362	21.3748	0.4496	1.
Fe(Lac)2,aq	Fe(CH ₃ CH ₂ COO) ₂		
Fe(Lac)2,aq	Fe(1)(C(6)H(10)O(6)		
ref:27	30.Aug.93		
-272042.	-355495.	36.043	
16.8444	33.3487	-7.3593	-4.1575
167.6874	53.2020	-0.0300	0.
Fe(Pent)+	FeCH ₃ (CH ₂) ₃ CO ₂ +		
Fe(Pent)+	Fe(1)(C(5)H(9)O(2)+(1)		
ref:27	1.Sep.93		
-106635.	-158054.	13.108	
11.5343	20.3852	-2.2687	-3.6216
136.3259	41.0721	0.3539	1.
Fe(Pent)2,aq	Fe(CH ₃ CH ₂ CH ₂ CH ₂ CO ₂) ₂		
Fe(Pent)2,aq	Fe(1)(C(10)H(18)O(4)		
ref:27	16.Aug.93		
-190586.	-293990.	50.8	
25.5409	54.5813	-15.7000	-5.0353
303.0133	100.2378	-0.0300	0.
Fe(Prop)+	FeCH ₃ CH ₂ CO ₂ +		
Fe(Prop)+	Fe(1)(C(3)H(5)O(2)+(1)		
ref:27	16.Aug.93		
-110833.	-146301.	1.308	
7.3034	10.0542	1.7914	-3.1945
85.8682	22.9627	0.5324	1.
Fe(Prop)2,aq	Fe(CH ₃ CH ₂ CO ₂) ₂		
Fe(Prop)2,aq	Fe(1)(C(6)H(10)O(4)		
ref:27	16.Aug.93		
-199008.	-271598.	23.576	
16.4683	32.4309	-6.9996	-4.1196
178.5967	56.9937	-0.0300	0.
Fe+3	Fe(+3)		
Fe+3	Fe(1)+(3)		
ref:21	11.Sep.97		
-4120.	-11850.	-66.300	
-2.4256	-13.6961	11.1141	-2.2127
19.0459	-6.8233	2.5812	3.
FeOH+2	FeOH(+2)		
FeOH+2	Fe(1)O(1)H(1)+(2)		
ref:21	1.May.97		
-57800.	-70000.	-25.400	
-1.1562	-10.6009	9.9077	-2.3407
14.6102	-4.7048	1.4382	2.
FeO+	FeO(+1)		
FeO+	Fe(1)O(1)+(1)		
ref:21	1.May.97		
-53100.	-61000.	-11.100	
-3.7118	-16.8408	12.3595	-2.0827
-15.3982	-12.8325	.7191	1.
FeO2-	FeO2(-1)		
FeO2-	Fe(1)O(2)-(-1)		
ref:21	1.May.97		
-88000.	-110700.	10.600	
2.3837	-1.9602	6.5182	-2.6979
-13.3207	-14.5028	1.4662	-1.
FeCl+2	FeCl(+2)		
FeCl+2	Fe(1)Cl(1)+(2)		
ref:22	14.Sep.97		
-37518.	-50820.	-42.740	
-7.1164	-9.5277	9.4878	-2.3851
23.8149	-2.3482	1.7013	2.
FeF+2	FeF(+2)		
FeF+2	Fe(1)F(1)+(2)		
ref:22	15.Sep.97		
-79645.	-87141.	-25.700	
-3.4294	-16.1520	12.0915	-2.1112
23.9805	-1.4787	1.4477	2.
Fr+	Fr(+)		
Fr+	Fr(1)+(1)		
ref:21	12.Dec.97		

	-70500.	-62500.	35.100	
	4.7139	3.7284	4.2849	-2.9330
	2.5442	-4.3586	0.0203	1.
Ga+3		Ga(+3)		
Ga+3		Ga(1)+(3)		
ref:21		15.Sep.97		
	-38000.	-50600.	-79.000	
	-2.8378	-14.7065	11.5197	-2.1709
	16.2356	-8.4326	2.7787	3.
GaOH+		GaOH(+)		
GaOH+		Ga(1)O(1)H(1)+(1)		
ref:21		10.Dec.97		
	-91100.	-116700.	-67.300	
	1.6572	-3.7321	7.2107	-2.6246
	48.3870	5.0115	2.0700	1.
GaO+		GaO(+)		
GaO+		Ga(1)O(1)+(1)		
ref:21		10.Dec.97		
	-86600.	-100900.	-29.200	
	2.1299	-2.5806	6.7629	-2.6722
	-0.7503	-8.6363	0.9986	1.
GaO2-		GaO2(-)		
GaO2-		Ga(1)O(2)-(1)		
ref:21		10.Dec.97		
	-128700.	-153900.	-10.000	
	3.6247	1.0721	5.3221	-2.8232
	14.3978	-5.8659	1.7777	-1.
Gd+2		Gd(+2)		
Gd+2		Gd(1)+(2)		
ref:21		10.Dec.97		
	-70600.	-67400.	-4.300	
	0.0094	-7.7516	8.7810	-2.4584
	9.9946	-5.2956	1.1216	2.
Gd+3		Gd(+3)		
Gd+3		Gd(1)+(3)		
ref:2		13.Jul.87		
	-158600.	-164200.	-49.200	
	-2.9771	-15.0506	11.6656	-2.1568
	6.5606	-10.3474	2.3265	3.
GdOH+2		GdOH+2		
GdOH+2		Gd(1)O(1)H(1)+(2)		
ref:25		1.Dec.95		
	-204500.	-213400.	-2.9	
	2.7389	-1.0936	6.1786	-2.7237
	-3.5082	-9.9196	1.1000	2.
GdO+		GdO+		
GdO+		Gd(1)O(1)+(1)		
ref:25		1.Dec.95		
	-193000.	-205500.	13.1	
	2.8425	-0.8409	6.0801	-2.7441
	-34.9963	-18.4750	0.3539	1.
GdO2-		GdO2-		
GdO2-		Gd(1)O(2)-(1)		
ref:25		1.Dec.95		
	-225000.	-233000.	38.3	
	5.0344	4.5111	3.9769	-2.9654
	-50.6235	-26.1341	1.0495	-1.
GdCl+2		GdCl+2		
GdCl+2		Gd(1)Cl(1)+(2)		
ref:25		1.Dec.95		
	-190400.	-200600.	-22.3	
	-0.2630	-8.4170	9.0425	-2.4309
	18.1232	-3.3636	1.4006	2.
GdCl2+		GdCl2+		
GdCl2+		Gd(1)Cl(2)+(1)		
ref:25		1.Dec.95		
	-221300.	-239000.	-5.4	
	2.8492	-0.8272	6.0803	-2.7447
	9.5260	-3.8858	0.6305	1.
GdCl3, aq		GdCl3(aq)		
GdCl3(aq)		Gd(1)Cl(3)		
ref:25		1.Dec.95		

	-252100.	-280200.	2.2	
	6.3836	7.8028	2.6888	-3.1015
	-19.6565	-11.9138	-0.0300	0.
GdCl4-		GdCl4-		
GdCl4-		Gd(1)Cl(4)-(1)		
ref:25		1.Dec.95		
	-282900.	-324300.	-0.2	
	11.1317	19.3995	-1.8761	-3.5809
	-49.0467	-27.4478	1.6310	-1.
GdF+2		GdF+2		
GdF+2		Gd(1)F(1)+(2)		
ref:25		1.Dec.95		
	-232200.	-239300.	-14.4	
	-2.9196	-14.9041	11.5935	-2.1628
	19.3937	-2.5281	1.2776	2.
GdF2+		GdF2+		
GdF2+		Gd(1)F(2)+(1)		
ref:25		1.Dec.95		
	-304100.	-321800.	-10.2	
	-2.6526	-14.2557	11.3457	-2.1896
	15.9952	-1.8906	0.7096	1.
GdF3, aq		GdF3(aq)		
GdF3(aq)		Gd(1)F(3)		
ref:25		1.Dec.95		
	-374900.	-407400.	-19.9	
	-2.3913	-13.6174	11.0949	-2.2160
	-9.6471	-8.4349	-0.0300	0.
GdF4-		GdF4-		
GdF4-		Gd(1)F(4)-(1)		
ref:25		1.Dec.95		
	-445000.	-497300.	-46.4	
	-1.0260	-10.2805	9.7756	-2.3539
	-27.4519	-22.1609	2.3239	-1.
GdHCO3+2		GdHCO3+2		
GdHCO3+2		Gd(1)H(1)C(1)O(3)+(2)		
ref:25		1.Dec.95		
	-301200.	-326700.	-9.7	
	0.6026	-6.3043	8.2153	-2.5183
	39.1396	4.5681	1.2048	2.
GdCO3+		GdCO3+		
GdCO3+		Gd(1)C(1)O(3)+(1)		
ref:25		1.Dec.95		
	-295500.	-307600.	-40.8	
	-0.9530	-10.1036	9.7095	-2.3612
	-19.8322	-15.8269	1.1729	1.
GdH2PO4+2		GdH2PO4+2		
GdH2PO4+2		Gd(1)H(2)P(1)O(4)+(2)		
ref:25		1.Dec.95		
	-431800.	-476600.	-26.1	
	1.6048	-3.8632	7.2686	-2.6192
	43.0549	5.1201	1.4574	2.
GdNO3+2		GdNO3+2		
GdNO3+2		Gd(1)N(1)O(3)+(2)		
ref:25		1.Dec.95		
	-185700.	-219800.	-32.5	
	1.3205	-4.5535	7.5323	-2.5907
	33.6609	1.5663	1.5475	2.
GdSO4+		GdSO4+		
GdSO4+		Gd(1)S(1)O(4)+(1)		
ref:25		1.Dec.95		
	-331500.	-376800.	-12.0	
	-1.4776	-4.1705	7.3822	-2.6065
	-20.9360	-14.7880	0.7287	1.
Gd(Ae)+2		GdCH3COO+2		
Gd(Ae)+2		Gd(1)C(2)H(3)O(2)+(2)		
ref:33		10.Sep.92		
	-250490.	-283100.	-25.700	
	2.8492	-0.7945	6.0567	-2.7461
	57.3241	10.1106	1.4477	2.
Gd(Ae)2+		Gd(CH3COO)2+		
Gd(Ae)2+		Gd(1)C(4)H(6)O(4)+(1)		
ref:33		10.Sep.92		

	-341350.	-401740.	-4.700	
	9.4165	15.2134	-0.2342	-3.4078
	102.0230	28.2899	0.6223	1.
Gd(Ae)3, aq		Gd(CH3COO)3		
Gd(Ae)3, aq		Gd(1)C(6)H(9)O(6)		
ref:33		10.Sep.92		
	-431490.	-521580.	9.800	
	16.8116	33.2662	-7.3215	-4.1541
	141.7603	44.1903	-0.0300	0.
Gd(But)+2		GdCH3(CH2)2CO2+2		
Gd(But)+2		Gd(1)C(4)H(7)O(2)+(2)		
ref:27		16.Aug.93		
	-246839.	-294884.	-14.467	
	6.8808	9.0211	2.2007	-3.1518
	97.4167	24.5906	1.2776	2.
Gd(But)2+		Gd(CH3CH2CH2CO2)2+		
Gd(But)2+		Gd(1)C(8)H(14)O(4)+(1)		
ref:27		16.Aug.93		
	-334041.	-424078.	21.787	
	17.9029	35.9294	-8.3657	-4.2642
	197.8246	62.8670	0.2229	1.
Gd(For)+2		GdCHO2+2		
Gd(For)+2		Gd(1)C(1)H(1)O(2)+(2)		
ref:27		16.Aug.93		
	-246091.	-268634.	-24.567	
	0.8917	-5.5954	7.9307	-2.5476
	26.4744	-0.5508	1.4287	2.
Gd(For)2+		Gd(CHO2)2+		
Gd(For)2+		Gd(1)C(2)H(2)O(4)+(1)		
ref:27		16.Aug.93		
	-332531.	-372659.	-2.084	
	5.2546	5.0511	3.7601	-2.9877
	28.4144	2.8309	0.5831	1.
Gd(Pent)+2		GdCH3(CH2)3CO2+2		
Gd(Pent)+2		Gd(1)C(5)H(9)O(2)+(2)		
ref:27		16.Aug.93		
	-244829.	-301364.	-7.967	
	9.0391	14.2927	0.1259	-3.3698
	133.8403	37.5583	1.1815	2.
Gd(Pent)2+		Gd(CH3CH2CH2CH2CO2)2+		
Gd(Pent)2+		Gd(1)C(10)H(18)O(4)+(1)		
ref:27		16.Aug.93		
	-330007.	-436320.	37.150	
	22.4554	47.0464	-12.7369	-4.7238
	284.7556	93.8330	-0.0116	1.
Gd(Prop)+2		GdCH3CH2CO2+2		
Gd(Prop)+2		Gd(1)C(3)H(5)O(2)+(2)		
ref:27		16.Aug.93		
	-249081.	-289666.	-19.767	
	4.8065	3.9537	4.1983	-2.9423
	83.3388	19.4489	1.3552	2.
Gd(Prop)2+		Gd(CH3CH2CO2)2+		
Gd(Prop)2+		Gd(1)C(6)H(10)O(4)+(1)		
ref:27		16.Aug.93		
	-338525.	-414216.	9.261	
	13.5253	25.2465	-4.1796	-3.8226
	164.2235	50.5890	0.4100	1.
Gd+4		Gd(+4)		
Gd+4		Gd(1)+(4)		
ref:21		10.Dec.97		
	12600.	-5700.	-107.600	
	-4.3532	-18.4052	12.9700	-2.0180
	38.8394	-3.5030	3.8013	4.
H+		H(+)		
H+		H(1)+(1)		
ref:2		17.Jul.89		
	0.	0.	0.000	
	0.0000	0.0000	0.0000	0.0000
	0.0000	0.0000	0.0000	1.
H2, aq		H2(O)		
H2(aq)		H(2)+(O)		
ref:3		13.Jul.87		

	4236.	-1000.	13.800	
	5.1427	4.7758	3.8729	-2.9764
	27.6251	5.0930	-2.090	0.
H2O2, aq		H2O2		
H2O2(aq)		H(2)O(2)+(0)		
ref:21		11.Sep.97		
	-32030.	-45690.	34.400	
	6.1922	7.3370	2.8686	-3.0822
	10.3600	-1.9975	-1.809	0.
HO2-		HO2(-)		
HO2-		H(1)O(2)-(1)		
ref:21		11.Sep.97		
	-16090.	-38320.	5.700	
	2.6566	-1.2961	6.2619	-2.7253
	-1.5190	-10.6529	1.5449	-1.
HA1O2, aq		HA1O2(0)		
HA1O2(aq)		H(1)Al(1)O(2)+(0)		
ref:20		29.Apr.97		
	-207500.	-230730.	-6.500	
	3.6808	1.2058	5.2761	-2.8288
	39.0000	2.5000	3.000	0.
H2AsO3-		H2AsO3(-)		
H2AsO3-		H(2)As(1)O(3)-(1)		
ref:2		3.Jul.87		
	-140330.	-170840.	26.400	
	5.7934	6.3646	3.2485	-3.0421
	15.8032	-3.6253	1.2305	-1.
HAsO2, aq		HAsO2(0)		
HAsO2(aq)		H(1)As(1)O(2)+(0)		
ref:21		12.Dec.97		
	-96240.	-109110.	30.100	
	5.5984	5.8913	3.4280	-3.0224
	6.7990	-2.4438	-0.1158	0.
H3AsO4, aq		H3AsO4(0)		
H3AsO4(aq)		H(3)As(1)O(4)+(0)		
ref:21		12.Dec.97		
	-183100.	-215700.	44.000	
	7.9357	11.5925	1.1990	-3.2581
	18.2805	2.2209	-0.3263	0.
H2AsO4-		H2AsO4(-)		
H2AsO4-		H(2)As(1)O(4)-(1)		
ref:21		12.Dec.97		
	-180010.	-217390.	28.000	
	6.7429	8.6835	2.3351	-3.1379
	16.9206	-3.1567	1.2055	-1.
HAsO4-2		HAsO4(-2)		
HAsO4-2		H(1)As(1)O(4)-(2)		
ref:21		12.Dec.97		
	-170790.	-216620.	-4.400	
	4.3994	2.9611	4.5853	-2.9013
	7.9908	-12.7102	3.2197	-2.
HBeO2-		HBeO2(-1)		
HBeO2-		H(1)Be(1)O(2)-(1)		
ref:21		16.Sep.97		
	-172200.	-206700.	-33.300	
	2.9518	-5.719	5.9708	-2.7553
	56.2537	7.5374	2.1351	-1.
HBiO2, aq		HBiO2(0)		
HBiO2(aq)		H(1)Bi(1)O(2)+(0)		
ref:21		16.Sep.97		
	-79300.	-86300.	54.600	
	5.2306	4.9892	3.7912	-2.9852
	-81.1398	-33.2840	-0.0300	0.
HBrO, aq		HBrO		
HBrO(aq)		H(1)Br(1)O(1)+(0)		
ref:21		11.Sep.97		
	-19700.	-27000.	33.800	
	5.5795	5.8410	3.4565	-3.0204
	9.8576	-1.2012	-1.1719	0.
HCO3-		HCO3(-)		
HCO3-		H(1)C(1)O(3)-(1)		
ref:2		1.Jul.87		

	-140282.	-164898.	23.530	
	7.5621	1.1505	1.2346	-2.8266
	12.9395	-4.7579	1.2733	-1.
HCdO2-		HCdO2(-)		
HCdO2-		H(1)Cd(1)O(2)-(1)		
ref:21		12.Dec.97		
	-86500.	-117400.	-10.900	
	0.9080	-5.5570	7.9176	-2.5492
	23.2446	-2.8512	1.7965	-1.
HCeO2, aq		HCeO2(aq)		
HCeO2(aq)		H(1)Ce(1)O(2)		
ref:25		1.Dec.95		
	-239300.	-249500.	48.5	
	5.0117	4.4582	3.9917	-2.9632
	-72.9937	-30.4525	-0.0300	0.
HCl, aq		HCl(0)		
HCl(aq)		H(1)Cl(1)+(0)		
ref:40		20.Jan.98		
	-30411.	-30285.	0.421	
	16.1573	-11.4311	-46.1866	-2.3036
	46.4716	-5.2811	0.0000	0.
HC1O, aq		HC1O		
HC1O(aq)		H(1)Cl(1)O(1)+(0)		
ref:21		11.Sep.97		
	-19100.	-28900.	33.900	
	5.5927	5.8751	3.4387	-3.0218
	9.9023	-1.1808	-1.734	0.
HC1O2, aq		HC1O2		
HC1O2(aq)		H(1)Cl(1)O(2)+(0)		
ref:21		11.Sep.97		
	1400.	-12400.	45.000	
	7.6706	10.9455	1.4527	-3.2314
	19.1373	2.5672	-3.415	0.
HCoO2-		HCoO2(-1)		
HCoO2-		H(1)Co(1)O(2)-(1)		
ref:21		11.Dec.97		
	-83400.	-117100.	-25.500	
	-0.0971	-8.0155	8.8926	-2.4475
	44.7712	3.9116	2.0211	-1.
HCrO2, aq		HCrO2(0)		
HCrO2(aq)		H(1)Cr(1)O(2)+(0)		
ref:21		16.Sep.97		
	-138100.	-157300.	6.000	
	1.5359	-4.0259	7.3211	-2.6125
	-16.2043	-10.7140	-0.0300	0.
HCrO4-		HCrO4(-)		
HCrO4-		H(1)Cr(1)O(4)-(1)		
ref:21		11.Sep.97		
	-183700.	-210000.	46.600	
	8.2211	12.2925	0.9174	-3.2871
	15.5492	-2.7290	9.230	-1.
HCuO2-		HCuO2(-1)		
HCuO2-		H(1)Cu(1)O(2)-(1)		
ref:21		1.May.97		
	-60100.	-86300.	300	
	1.3202	-4.5527	7.5282	-2.5907
	-3.2749	-8.0659	5.464	-1.
HDyO2, aq		HDyO2(aq)		
HDyO2(aq)		H(1)Dy(1)O(2)		
ref:25		1.Dec.95		
	-238200.	-251100.	39.1	
	4.6969	3.6855	4.3051	-2.9313
	-60.3933	-26.0730	-0.0300	0.
HEuO2, aq		HEuO2(aq)		
HEuO2(aq)		HEu(1)O(2)(1)		
ref:25		1.Dec.95		
	-216000.	-228200.	42.4	
	4.8064	3.9540	4.1968	-2.9424
	-64.9060	-27.6415	-0.3000	0.
HErO2, aq		HErO2(aq)		
HErO2(aq)		H(1)Er(1)O(2)		
ref:25		1.Dec.95		

	-240100.	-254300.	34.8	
	4.5464	3.3186	4.4476	-2.9161
	-54.7085	-24.0971	-0.0300	0.
HF2-		HF2(-)		
HF2-		H(1)F(2)-(1)		
ref:2		12.Jul.87		
	-138160.	-155340.	22.100	
	5.2263	4.9797	3.7928	-2.9849
	-1.3751	-9.7974	1.2934	-1.
HF, aq		HF(0)		
HF(aq)		H(1)F(1)+(0)		
ref:3		7.Feb.89		
	-71662.	-76835.	22.500	
	3.4753	.7042	5.4732	-2.8081
	14.3647	-1.828	-0.0007	0.
HFeO2-		HFeO2(-1)		
HFeO2-		H(1)Fe(1)O(2)-(1)		
ref:21		1.May.97		
	-95400.	-125700.	-15.000	
	.6272	-6.2440	8.1905	-2.5208
	36.1624	1.4468	1.8564	-1.
HFeO2, aq		HFeO2		
HFeO2(aq)		H(1)Fe(1)O(2)+(0)		
ref:21		1.May.97		
	-101100.	-120300.	22.200	
	2.7401	-1.0905	6.1776	-2.7338
	-37.8300	-18.2305	-0.0300	0.
HGaO2, aq		HGaO2		
HGaO2(aq)		H(1)Ga(1)O(2)+(0)		
ref:21		10.Dec.97		
	-137300.	-158600.	2.900	
	3.4470	0.6328	5.5060	-2.8051
	-12.1716	-9.2677	-0.0439	0.
HGdO2, aq		HGdO2(aq)		
HGdO2(aq)		H(1)Gd(1)O(2)		
ref:25		1.Dec.95		
	-237500.	-247200.	48.2	
	5.0117	4.4582	3.9917	-2.9632
	-72.5835	-30.3099	-0.0300	0.
HHfO2+		HHfO2(+)		
HHfO2+		H(1)Hf(1)O(2)+(1)		
ref:21		12.Dec.97		
	-237800.	-264200.	-51.000	
	2.0629	-2.7421	6.8221	-2.6655
	72.4834	15.7669	1.3268	1.
HHfO3-		HHfO3(-)		
HHfO3-		H(1)Hf(1)O(3)-(1)		
ref:21		12.Dec.97		
	-279200.	-331800.	-34.300	
	3.4908	0.7427	5.4561	-2.8096
	57.7445	8.0059	2.1506	-1.
HHgO2-		HHgO2(-1)		
HHgO2-		H(1)Hg(1)O(2)-(1)		
ref:21		12.Dec.97		
	-45200.	-73800.	2.300	
	1.8118	-3.3572	7.0689	-2.6401
	3.7588	-8.9825	1.5962	-1.
HHoO2, aq		HHoO2(aq)		
HHoO2(aq)		H(1)Ho(1)O(2)		
ref:25		1.Dec.95		
	-241300.	-253800.	40.5	
	4.7380	3.7900	4.2544	-2.9356
	-62.2688	-26.7248	-0.0300	0.
HIO, aq		HIO		
HIO(aq)		H(1)I(1)O(1)+(0)		
ref:21		11.Sep.97		
	-23700.	-33000.	22.800	
	3.8979	1.7354	5.0694	-2.8506
	.7260	-4.9085	-0.0053	0.
HI03, aq		HI03		
HI03(aq)		H(1)I(1)O(3)+(0)		
ref:21		11.Sep.97		

	-31700.	-50500.	39.900	
	6.8346	8.9050	2.2530	-3.1470
	14.8674	.8357	-2.2642	0.
HInO2, aq		HInO2		
HInO2(aq)		H(1)In(1)O(2)+(0)		
ref:21		11.Dec.97		
	-119800.	-135100.	27.200	
	4.2864	2.6831	4.6987	-2.8898
	-44.5111	-20.5527	-0.0300	0.
HLuO2, aq		HLuO2(aq)		
HLuO2(aq)		H(1)Lu(1)O(2)		
ref:25		1.Dec.95		
	-240200.	-255100.	27.1	
	4.2864	2.6831	4.6987	-2.8898
	-44.3938	-20.5120	-0.0300	0.
HMnO2-		HMnO2(-1)		
HMnO2-		H(1)Mn(1)O(2)-(-1)		
ref:21		12.Sep.97		
	-121000.	-149900.	-9.100	
	1.0354	-5.2480	7.8003	-2.5619
	20.5255	-3.7067	1.7685	-1.
HMoO4-		HMoO4(-1)		
HMoO4-		H(1)Mo(1)O(4)-(-1)		
ref:21		5.Jan.98		
	-206400.	-238500.	32.500	
	7.7320	11.0966	1.3908	-3.2376
	33.3930	2.7912	1.1359	-1.
HNO2, aq		HNO2		
HNO2(aq)		H(1)N(1)O(2)+(0)		
ref:21		11.Sep.97		
	-121100.	-28500.	32.400	
	5.9151	6.6590	3.1378	-3.0542
	8.7050	-1.6697	-1.507	0.
HNO3, aq		HNO3(0)		
HNO3(aq)		H(1)N(1)O(3)+(0)		
ref:21		11.Sep.97		
	-24730.	-45410.	42.700	
	7.1623	9.7063	1.9367	-3.1802
	13.8907	.6320	-3.066	0.
HN2O2-		HN2O2(-1)		
HN2O2-		H(1)N(2)O(2)-(-1)		
ref:21		15.Sep.97		
	18200.	-9400.	34.000	
	6.5748	8.2753	2.4914	-3.1210
	21.8724	-1.1401	1.1132	-1.
H2N2O2, aq		H2N2O2(0)		
H2N2O2(aq)		H(2)N(2)O(2)+(0)		
ref:21		15.Sep.97		
	8600.	-13700.	52.000	
	8.6337	13.3019	.5163	-3.3288
	24.9004	4.9097	-4.475	0.
HNbO3, aq		HNbO3(0)		
HNbO3(aq)		H(1)Nb(1)O(3)+(0)		
ref:21		5.Jan.98		
	-236800.	-257950.	31.100	
	5.7110	6.1642	3.3252	-3.0337
	20.2299	3.3616	-0.4710	0.
HNdO2, aq		HNdO2(aq)		
HNdO2(aq)		H(1)Nd(1)O(2)		
ref:25		1.Dec.95		
	-238000.	-248000.	47.8	
	4.9843	4.3865	4.0304	-2.9602
	-71.9974	-30.1062	-0.0300	0.
HNiO2-		HNiO2(-1)		
HNiO2-		H(1)Ni(1)O(2)-(-1)		
ref:21		16.Sep.97		
	-81980.	-118760.	-36.000	
	-9.8143	-9.7619	9.5688	-2.3753
	54.2564	6.7430	2.1664	-1.
H3PO2, aq		H3PO2(0)		
H3PO2(aq)		H(3)P(1)O(2)+(0)		
ref:21		15.Sep.97		

	-125100.	-145600.	36.900	
	6.5352	8.1731	2.5430	-3.1168
	12.4140	-1.623	-2.2188	0.
H2PO2-		H2PO2(-)		
H2PO2-		H(2)P(1)O(2)-(-1)		
ref:21		15.Sep.97		
	-122400.	-146700.	24.100	
	5.2026	4.9190	3.8223	-2.9823
	13.7063	-4.4604	1.2637	-1.
H3PO3, aq		H3PO3(0)		
H3PO3(aq)		H(3)P(1)O(3)+(0)		
ref:21		15.Sep.97		
	-204800.	-230600.	43.600	
	4.2704	2.6439	4.7145	-2.8882
	17.9261	2.0783	-3.3203	0.
H2PO3-		H2PO3(-)		
H2PO3-		H(2)P(1)O(3)-(-1)		
ref:21		15.Sep.97		
	-202350.	-231700.	31.700	
	4.1649	2.3899	4.8065	-2.8777
	19.9762	-1.9142	1.1491	-1.
HPO3-2		HPO3(-2)		
HPO3-2		H(1)P(1)O(3)-(-2)		
ref:21		15.Sep.97		
	-194000.	-231600.	4.000	
	3.4463	.6347	5.4974	-2.8051
	11.6523	-11.2232	3.1527	-2.
H4P2O7, aq		H4P2O7(0)		
H4P2O7(aq)		H(4)P(2)O(7)+(0)		
ref:21		15.Sep.97		
	-485700.	-542200.	64.000	
	9.2975	14.9199	-1.130	-3.3957
	34.8301	8.9430	-6.292	0.
H3P2O7-		H3P2O7(-)		
H3P2O7-		H(3)P(2)O(7)-(-1)		
ref:21		15.Sep.97		
	-483600.	-544100.	51.000	
	9.1292	14.5122	.0398	-3.3788
	35.9785	4.5838	.8568	-1.
H2P2O7-2		H2P2O7(-2)		
H2P2O7-2		H(2)P(2)O(7)-(-2)		
ref:21		15.Sep.97		
	-480400.	-544600.	39.000	
	9.0963	14.4299	.0760	-3.3754
	40.6359	.5505	2.6218	-2.
HP2O7-3		HP2O7(-3)		
HP2O7-3		H(1)P(2)O(7)-(-3)		
ref:21		15.Sep.97		
	-471400.	-543700.	-11.000	
	8.3302	12.5558	0.8208	-3.2980
	32.1587	-8.8807	4.6470	-3.
H3PO4, aq		H3PO4(0)		
H3PO4(aq)		H(3)P(1)O(4)+(0)		
ref:3		11.Jul.87		
	-273100.	-307920.	38.000	
	8.2727	12.4182	.8691	-3.2924
	17.9708	1.7727	-2.200	0.
H2PO4-		H2PO4(-)		
H2PO4-		H(2)P(1)O(4)-(-1)		
ref:2		3.Jul.87		
	-270140.	-309820.	21.600	
	6.4875	8.0594	2.5823	-3.1122
	14.0435	-4.4605	1.3003	-1.
HPO4-2		HPO4(-2)		
HPO4-2		H(1)P(1)O(4)-(-2)		
ref:2		3.Jul.87		
	-260310.	-308815.	-8.000	
	3.6149	1.0857	5.3233	-2.8239
	2.7357	-14.9103	3.3363	-2.
HPbO2-		HPbO2(-1)		
HPbO2-		H(1)Pb(1)O(2)-(-1)		
ref:21		16.Sep.97		

	-81000.	-104200.	18.000	
	4.4403	3.0599	4.5482	-2.9054
	-22.3600	-17.2935	1.3566	-1.
HPrO2, aq		HPrO2(aq)		
HPrO2(aq)		H(1)Pr(1)O(2)		
ref:25		1.Dec.95		
	-239700.	-250100.	47.0	
	4.9706	4.3537	4.0423	-2.9589
	-74.0012	-29.7599	-0.0300	0.
H2S, aq		H2S(0)		
H2S(aq)		H(2)S(1)+(0)		
ref:3		26.Jun.87		
	-6673.	-9001.	30.000	
	6.5097	6.7724	5.9646	-3.0590
	32.3000	4.7300	-1.000	0.
HS-		HS(-)		
HS-		H(1)S(1)-(-1)		
ref:2		1.Jul.87		
	2860.	-3850.	16.300	
	5.0119	4.9799	3.4765	-2.9849
	3.4200	-6.2700	1.4410	-1.
H2S2O3, aq		H2S2O3		
H2S2O3(aq)		H(2)S(2)O(3)+(0)		
ref:21		11.Sep.97		
	-128000.	-150400.	45.000	
	7.4242	10.3493	1.6762	-3.2067
	19.1373	2.5672	-3.415	0.
HS2O3-		HS2O3(-1)		
HS2O3-		H(1)S(2)O(3)-(-1)		
ref:21		11.Sep.97		
	-127200.	-153900.	30.500	
	6.1964	7.3510	2.8549	-3.0828
	18.9745	-2.3216	1.1676	-1.
H2S2O4, aq		H2S2O4		
H2S2O4(aq)		H(2)S(2)O(4)+(0)		
ref:21		11.Sep.97		
	-147400.	-175300.	51.000	
	8.6525	13.3462	.5023	-3.3306
	24.1023	4.5838	-4.323	0.
HS2O4-		HS2O4(-1)		
HS2O4-		H(1)S(2)O(4)-(-1)		
ref:21		11.Sep.97		
	-146900.	-179100.	36.500	
	7.6159	10.8142	1.5005	-3.2260
	23.9326	-3.3049	1.0760	-1.
HSO3-		HSO3(-)		
HSO3-		H(1)S(1)O(3)-(-1)		
ref:2		3.Jul.87		
	-126130.	-149670.	33.400	
	6.7014	8.5816	2.3771	-3.1338
	15.6949	-3.3198	1.1233	-1.
HSO4-		HSO4(-)		
HSO4-		H(1)S(1)O(4)-(-1)		
ref:2		3.Jul.87		
	-180630.	-212500.	30.000	
	6.9788	9.2590	2.1108	-3.1618
	20.0961	-1.9550	1.1748	-1.
HSO5-		HSO5(-)		
HSO5-		H(1)S(1)O(5)-(-1)		
ref:21		11.Sep.97		
	-152370.	-185380.	50.700	
	8.9391	14.0430	.2349	-3.3594
	35.7254	4.4819	.8611	-1.
HSbO2, aq		HSbO2(0)		
HSbO2(aq)		H(1)Sb(1)O(2)+(0)		
ref:21		11.Sep.97		
	-97400.	-116600.	11.100	
	2.9726	-0.5203	5.9475	-2.7574
	-8.9525	-8.8400	0.1719	0.
HScO2, aq		HScO2		
HScO2(aq)		H(1)Sc(1)O(2)+(0)		
ref:21		11.Dec.97		

	-231600.	-244300.	30.300	
	3.3422	0.3770	5.6066	-2.7945
HSe-	-47.9689	-21.7545	-0.0300	0.
HSe-		HSe(-)		
ref:21		H(1)Se(1)-(1)		
	10500.	3800.	19.000	
	4.9823	4.3855	4.0227	-2.9602
	6.9147	-7.0678	1.3400	-1.
H2SeO3, aq		H2SeO3		
H2SeO3 (aq)		H(2)Se(1)O(3)+(0)		
ref:21		11.Sep.97		
	-101850.	-121290.	49.700	
	7.9475	11.6240	1.1819	-3.2594
	22.9944	4.1356	-4.127	0.
HSeO3-		HSeO3(-)		
HSeO3-		H(1)Se(1)O(3)-(1)		
ref:21		11.Sep.97		
	-98340.	-122980.	32.300	
	6.4061	7.8624	2.6551	-3.1039
	20.4808	-1.7105	1.1402	-1.
HSeO4-		HSeO4(-)		
HSeO4-		H(1)Se(1)O(4)-(1)		
ref:21		11.Sep.97		
	-108100.	-139000.	35.700	
	7.5244	10.5922	1.5837	-3.2168
	23.2855	-5.697	1.0885	-1.
HSiO3-		HSiO3(-)		
HSiO3-		H(1)Si(1)O(3)-(1)		
ref:22		13.Sep.97		
	-242801.	-273872.	5.000	
	2.9735	-0.5181	5.9467	-2.7575
	8.1489	-7.3123	1.5511	-1.
HSnO2, aq		HSnO2(aq)		
HSnO2(aq)		H(1)Sn(1)O(2)		
ref:25		1.Dec.95		
	-237100.	-247700.	45.9	
	4.9296	4.2552	4.0768	-2.9548
	-69.5359	-29.2507	-0.3000	0.
HSnO2-		HSnO2(-1)		
HSnO2-		H(1)Sn(1)O(2)-(1)		
ref:21		12.Sep.97		
	-97300.	-122100.	9.400	
	4.1971	2.4678	4.7779	-2.8809
	-6.7372	-12.2825	1.4874	-1.
HTbO2, aq		HTbO2(aq)		
HTbO2(aq)		H(1)Tb(1)O(2)		
ref:25		1.Dec.95		
	-238700.	-251000.	40.9	
	4.7517	3.8228	4.2425	-2.9369
	-62.8548	-26.9285	-0.0300	0.
HTlO2, aq		HTlO2		
HTlO2(aq)		H(1)Tl(1)O(2)+(0)		
ref:21		11.Dec.97		
	-57600.	-65600.	53.100	
	5.1759	4.8579	3.8372	-2.9797
	-79.1473	-32.5914	-0.0300	0.
HTmO2, aq		HTmO2(aq)		
HTmO2(aq)		H(1)Tm(1)O(2)		
ref:25		1.Dec.95		
	-240300.	-254500.	34.7	
	4.5464	3.3186	4.4476	-2.9161
	-55.5327	-24.0360	-0.0300	0.
HUO2, aq		HUO2(0)		
HUO2(aq)		U(1)U(1)O(2)+(0)		
ref:37		20.Jan.98		
	-198000.	-205100.	52.900	
	5.1622	4.8252	3.8489	-2.9784
	-78.8543	-32.4895	-0.0300	0.
HUO2+		HUO2(+)		
HUO2+		H(1)U(1)O(2)+(1)		
ref:37		20.Jan.98		

	-233200.	-254800.	-11.400	
	3.2291	0.1055	5.7039	-2.7833
	14.0517	-2.6271	0.7287	1.
HUO3-		HUO3(-)		
HUO3-		H(1)U(1)O(3)-(1)		
ref:37		20.Jan.98		
	-274100.	-321200.	-41.300	
	3.1824	-0.0092	5.7491	-2.7785
	68.0400	11.2651	2.2503	-1.
HUO4-		HUO4(-)		
HUO4-		H(1)U(1)O(4)-(1)		
ref:37		20.Jan.98		
	-314800.	-362900.	-20.100	
	4.9364	4.2748	4.0637	-2.9556
	36.9351	1.4672	1.9339	-1.
H3VO4, aq		H3VO4(0)		
H3VO4(aq)		H(3)V(1)O(4)+(0)		
ref:21		5.Jan.98		
	-242900.	-284600.	37.100	
	6.9310	9.1392	2.1640	-3.1567
	15.3166	0.8561	-0.2218	0.
H2VO4-		H2VO4(-)		
H2VO4-		H(2)V(1)O(4)-(1)		
ref:21		5.Jan.98		
	-243990.	-280620.	29.000	
	6.4502	7.9653	2.6252	-3.1082
	38.1088	4.2579	1.1898	-1.
HVO4-2		HVO4(-2)		
HVO4-2		H(1)V(1)O(4)-(2)		
ref:21		5.Jan.98		
	-233000.	-277000.	4.000	
	4.5410	3.3077	4.4469	-2.9156
	86.3755	14.7484	3.1527	-2.
HWO4-		HWO4(-1)		
HWO4-		H(1)W(1)O(4)-(1)		
ref:21		15.Sep.97		
	-223400.	-255000.	31.200	
	7.7661	11.1820	1.3532	-3.2412
	34.7493	3.1986	1.1559	-1.
HYO2, aq		HYO2		
HYO2(aq)		H(1)Y(1)O(2)+(0)		
ref:21		12.Dec.97		
	-241700.	-254600.	31.800	
	4.4506	3.0828	4.5443	-2.9063
	-50.6646	-22.6916	-0.0300	0.
HYbO2, aq		HYbO2(aq)		
HYbO2(aq)		H(1)Yb(1)O(2)		
ref:25		1.Dec.95		
	-233800.	-246500.	36.5	
	4.6011	3.4560	4.3859	-2.9218
	-58.9357	-24.8711	-0.0300	0.
HZnO2-		HZnO2(-1)		
HZnO2-		H(1)Zn(1)O(2)-(1)		
ref:21		17.Sep.97		
	-110720.	-142370.	-16.000	
	5.523	-6.4047	8.2588	-2.5141
	35.0874	1.0394	1.8669	-1.
HZrO2+		HZrO2(+1)		
HZrO2+		H(1)Zr(1)O(2)+(1)		
ref:21		1.May.97		
	-239700.	-265300.	-49.700	
	2.1106	-2.6299	6.7878	-2.6702
	70.5343	15.1558	1.3060	1.
HZrO3-		HZrO3(-1)		
HZrO3-		H(1)Zr(1)O(3)-(1)		
ref:21		1.May.97		
	-281500.	-333300.	-32.700	
	3.5625	9.168	5.3901	-2.8168
	55.3517	7.2726	2.1199	-1.
He, aq		He(0)		
He(aq)		He(1)+(0)		
ref:3		13.Jul.87		

	4658.	-150.	14.020	
	3.4722	.6967	5.4762	-2.8078
	26.0707	4.5634	-2.123	0.
Hf+4		Hf(+4)		
Hf+4		Hf(1)+(4)		
ref:21		12.Dec.97		
	-132600.	-150250.	-111.200	
	-4.3761	-18.4638	12.9996	-2.0156
	39.6914	-3.7271	3.8552	4.
HfOH+3		HfOH(+3)		
HfOH+3		Hf(1)O(1)H(1)+(3)		
ref:21		12.Dec.97		
	-189000.	-211800.	-72.800	
	0.9047	-5.5670	7.9267	-2.5488
	57.6774	6.2949	2.6777	3.
HfO+2		HfO(+2)		
HfO+2		Hf(1)O(1)+(2)		
ref:21		12.Dec.97		
	-186000.	-200900.	-46.200	
	2.0179	-2.8551	6.8743	-2.6609
	17.6401	-4.6845	1.7607	2.
HfO2, aq		HfO2		
HfO2(aq)		Hf(1)O(2)+(0)		
ref:21		12.Dec.97		
	-231400.	-262600.	-45.100	
	1.8096	-3.3639	7.0739	-2.6398
	40.1163	8.8615	-0.0300	0.
Hg2+2		Hg2(+2)		
Hg2+2		Hg(2)+(2)		
ref:21		10.Dec.97		
	36710.	39880.	15.710	
	4.0263	2.0479	4.9491	-2.8636
	21.8090	-0.2234	.8201	2.
Hg+2		Hg(+2)		
Hg+2		Hg(1)+(2)		
ref:2		1.Jul.87		
	39360.	40670.	-8.680	
	-5280	-9.0707	9.3152	-2.4040
	18.0613	-2.5865	1.1512	2.
HgOH+		HgOH(+)		
HgOH+		Hg(1)O(1)H(1)+(1)		
ref:21		12.Dec.97		
	-12700.	-20400.	16.700	
	1.2505	-4.7196	7.5872	-2.5838
	36.1808	6.4374	0.2998	1.
HgO, aq		HgO		
HgO(aq)		Hg(1)O(1)+(0)		
ref:21		12.Dec.97		
	-8900.	-19200.	8.200	
	0.6191	-6.2664	8.2051	-2.5198
	-8.6440	-8.0863	-0.0300	0.
HgCl+		HgCl(+)		
HgCl+		Hg(1)Cl(1)+(1)		
ref:22		6.Jan.98		
	-2155.	-7374.	11.670	
	2.4941	-1.6886	6.4067	-2.7091
	32.5048	4.9328	0.3735	1.
HgCl2, aq		HgCl2(0)		
HgCl2(aq)		Hg(1)Cl(2)+(0)		
ref:22		6.Jan.98		
	-42838.	-56787.	24.640	
	6.0150	6.9084	3.0277	-3.0645
	48.6578	11.8560	-0.0380	0.
HgCl3-		HgCl3(-)		
HgCl3-		Hg(1)Cl(3)-(1)		
ref:22		6.Jan.98		
	-74504.	-99299.	30.510	
	10.5003	17.8603	-1.2769	-3.5173
	78.0458	18.2101	1.1676	-1.
HgCl4-2		HgCl4(-2)		
HgCl4-2		Hg(1)Cl(4)-(2)		
ref:22		6.Jan.98		

	-106988.	-145023.	28.360	
	15.5903	30.2888	-6.1618	-4.0311
	109.5610	23.9860	2.7846	-2.
HgF+		HgF(+)		
HgF+		Hg(1)F(1)+(1)		
ref:22		6.Jan.98		
	-30204.	-39512.	-4.480	
	-0.0492	-7.8986	8.8475	-2.4524
	37.2987	5.7934	0.6223	1.
Hg(Ae)+		HgCH3COO+		
Hg(Ae)+		Hg(1)C(2)H(3)O(2)+(1)		
ref:33		17 Aug 92		
	-54760.	-79390.	18.500	
	5.6341	5.9787	3.3935	-3.0261
	70.1764	18.3452	0.2712	1.
Hg(Ae)2,aq		Hg(CH3COO)2		
Hg(Ae)2,aq		Hg(1)C(4)H(6)O(4)		
ref:33		17 Aug 92		
	-146580.	-198780.	40.100	
	12.7295	23.3004	-3.4080	-3.7421
	140.9402	43.9053	-0.0300	0.
Hg(Ae)3-		Hg(CH3COO)3-		
Hg(Ae)3-		Hg(1)C(6)H(9)O(6)-(1)		
ref:33		10 Sep 92		
	-239020.	-321900.	51.400	
	21.0462	43.6061	-11.3861	-4.5816
	235.9096	74.0939	0.8508	-1.
Ho+2		Ho(+2)		
Ho+2		Ho(1)+(2)		
ref:21		10.Dec.97		
	-96900.	-94200.	-4.300	
	0.0094	-7.7516	8.7810	-2.4584
	9.9946	-5.2956	1.1216	2.
Ho+3		Ho(+3)		
Ho+3		Ho(1)+(3)		
ref:2		13.Jul.87		
	-161400.	-169000.	-54.300	
	-3.1198	-15.3992	11.8026	-2.1424
	8.6686	-9.8178	2.3899	3.
HoOH+2		HoOH+2		
HoOH+2		Ho(1)O(1)H(1)+(2)		
ref:25		1.Dec.95		
	-207500.	-219000.	-9.7	
	2.6374	-1.3394	6.2715	-2.7235
	1.9108	-8.3714	1.2048	2.
HoO+		HoO+		
HoO+		Ho(1)O(1)+(1)		
ref:25		1.Dec.95		
	-196200.	-211400.	5.9	
	2.7146	-1.1548	6.2077	-2.7312
	-29.1412	-16.7843	0.4615	1.
HoO2-		HoO2-		
HoO2-		Ho(1)O(2)-(1)		
ref:25		1.Dec.95		
	-229100.	-240000.	30.0	
	4.7894	3.9112	4.2160	-2.9406
	-39.5064	-22.6712	1.1748	-1.
HoCl+2		HoCl+2		
HoCl+2		Ho(1)Cl(1)+(2)		
ref:25		1.Dec.95		
	-193100.	-205600.	-28.4	
	-0.4168	-8.7940	9.1934	-2.4154
	22.3450	-2.1720	1.4867	2.
HoCl2+		HoCl2+		
HoCl2+		Ho(1)Cl(2)+(1)		
ref:25		1.Dec.95		
	-224100.	-244600.	-13.1	
	2.6851	-1.2272	6.2371	-2.7282
	17.3007	-1.5607	0.7483	1.
HoCl3,aq		HoCl3(aq)		
HoCl3(aq)		Ho(1)Cl(3)		
ref:25		1.Dec.95		

	-254900.	-286400.	-7.8	
	6.1565	7.2496	2.9030	-3.0786
	-8.3500	-7.9841	-0.0300	0.
HoCl4-		HoCl4-		
HoCl4-		Ho(1)Cl(4)-(1)		
ref:25		1.Dec.95		
	-285700.	-331700.	-14.0	
	10.9513	18.9611	-1.7081	-3.5628
	-29.7863	-21.4419	1.8460	-1.
HoF+2		HoF+2		
HoF+2		Ho(1)F(1)+(2)		
ref:25		1.Dec.95		
	-235200.	-243800.	-17.8	
	-3.0852	-15.3064	11.7462	-2.1461
	23.2937	-1.3364	1.3288	2.
HoF2+		HoF2+		
HoF2+		Ho(1)F(2)+(1)		
ref:25		1.Dec.95		
	-307400.	-326500.	-13.5	
	-2.8400	-14.7070	11.5098	-2.1709
	23.1340	0.4343	0.7584	1.
HoF3,aq		HoF3(aq)		
HoF3(aq)		Ho(1)F(3)		
ref:25		1.Dec.95		
	-378300.	-412500.	-24.2	
	-2.6184	-14.1706	11.3090	-2.1931
	1.6591	-4.5051	-0.0300	0.
HoF4-		HoF4-		
HoF4-		Ho(1)F(4)-(1)		
ref:25		1.Dec.95		
	-448500.	-503400.	-53.7	
	-1.2375	-10.7997	9.9860	-2.3324
	-9.0382	-16.1550	2.4471	-1.
HoHCO3+2		HoHCO3+2		
HoHCO3+2		Ho(1)H(1)C(1)O(3)+(2)		
ref:25		1.Dec.95		
	-304000.	-332100.	-17.0	
	0.4557	-6.6620	8.3538	-2.5035
	43.5512	5.7598	1.3115	2.
HoCO3+		HoCO3+		
HoCO3+		Ho(1)C(1)O(3)+(1)		
ref:25		1.Dec.95		
	-298600.	-312600.	-47.0	
	-0.9954	-10.2047	9.7433	-2.3570
	-17.4516	-15.2972	1.2659	1.
HoH2PO4+2		HoH2PO4+2		
HoH2PO4+2		Ho(1)H(2)P(1)O(4)+(2)		
ref:25		1.Dec.95		
	-434600.	-482100.	-33.5	
	1.4594	-4.2123	7.3930	-2.6048
	47.5067	6.3118	1.5684	2.
HoNO3+2		HoNO3+2		
HoNO3+2		Ho(1)N(1)O(3)+(2)		
ref:25		1.Dec.95		
	-188200.	-225600.	-42.1	
	1.1895	-4.8717	7.6537	-2.5775
	38.5060	2.7579	1.7013	2.
HoSO4+		HoSO4+		
HoSO4+		Ho(1)S(1)O(4)+(1)		
ref:25		1.Dec.95		
	-344200.	-381500.	-17.0	
	-1.4371	-4.2696	7.4212	-2.6024
	-18.6528	-14.2584	0.8111	1.
Ho(Ae)+2		HoCH3COO+2		
Ho(Ae)+2		Ho(1)C(2)H(3)O(2)+(2)		
ref:33		11.Sep.92		
	-253270.	-288520.	-32.900	
	2.7148	-1.1553	6.2089	-2.7311
	61.7679	11.3023	1.5579	2.
Ho(Ae)2+		Ho(CH3COO)2+		
Ho(Ae)2+		Ho(1)C(4)H(6)O(4)+(1)		
ref:33		11 Sept 92		

	-344120.	-407930.	-14.600	
	9.2621	14.8317	-0.0744	-3.3920
	110.0599	30.6149	0.7686	1.
Ho(Ae)3,aq		Ho(CH3COO)3		
Ho(Ae)3,aq		Ho(1)C(6)H(9)O(6)		
ref:33		11 Sept 92		
	-434250.	-528670.	-3.100	
	16.5845	32.7130	-7.1070	-4.1313
	153.0665	48.1201	-0.0300	0.
Ho+4		Ho(+4)		
Ho+4		Ho(1)+(4)		
ref:21		10.Dec.97		
	-30500.	-48300.	-104.200	
	-4.3164	-18.3144	12.9330	-2.0218
	39.9963	-3.2789	3.7484	4.
I-		I(-)		
I-		I(1)-(1)		
ref:2		1.Jul.87		
	-12410.	-13600.	25.500	
	7.7623	8.2762	1.4609	-3.1211
	-6.2700	-4.9440	1.2934	-1.
I3-		I3(-)		
I3-		I(3)-(1)		
ref:2		12.Jul.87		
	-12300.	-12300.	57.200	
	9.8500	16.2697	-6.6447	-3.4516
	20.8712	-3.6661	.7628	-1.
IO-		IO(-)		
IO-		I(1)O(1)-(1)		
ref:21		11.Sep.97		
	-9200.	-25700.	-1.400	
	.6131	-6.2757	8.1983	-2.5195
	1.1172	-10.0825	1.6530	-1.
IO3-		IO3(-)		
IO3-		I(1)O(3)-(1)		
ref:2		12.Jul.87		
	-30600.	-52900.	28.300	
	5.7148	6.1725	3.3240	-3.0342
	7.7293	-6.3345	1.2002	-1.
IO4-		IO4(-)		
IO4-		I(1)O(4)-(1)		
ref:21		30.Apr.97		
	-14000.	-36200.	53.000	
	9.2831	14.8827	-0.0933	-3.3942
	12.6081	-3.4419	.8264	-1.
In+3		In(+3)		
In+3		In(1)+(3)		
ref:21		12.Dec.97		
	-23400.	-25000.	-63.000	
	-2.2361	-13.2378	10.9433	-2.2316
	20.3744	-6.2122	2.5346	3.
InOH+2		InOH(+2)		
InOH+2		In(1)O(1)H(1)+(2)		
ref:21		11.Dec.97		
	-74600.	-94700.	-44.700	
	2.0372	-2.8058	6.8493	-2.6629
	30.2531	-0.2234	1.7366	2.
InO+		InO(+)		
InO+		In(1)O(1)+(1)		
ref:21		11.Dec.97		
	-69400.	-76900.	-2.600	
	2.5804	-1.4833	6.3385	-2.7176
	-22.2654	-14.8084	0.5907	1.
InO2-		InO2(-)		
InO2-		In(1)O(2)-(1)		
ref:21		11.Dec.97		
	-106700.	-125300.	15.900	
	4.3832	2.9224	4.5983	-2.8997
	-20.4685	-16.7435	1.3901	-1.
InCl+2		InCl(+2)		
InCl+2		In(1)Cl(1)+(2)		
ref:22		6.Jan.98		

	-59226.	-66209.	-38.800	
	-0.6441	-9.3513	9.4185	-2.3924
	22.2354	-2.7149	1.6444	2.
InF+2		InF(+2)		
InF+2		In(1)F(1)+(2)		
ref:22		6.Jan.98		
	-97016.	-98718.	-23.530	
	-3.3507	-15.9600	12.0160	-2.1192
	22.5771	-1.8453	1.4099	2.
K+		K(+)		
K+		K(1)+(1)		
ref:2		3.Jun.87		
	-67510.	-60270.	24.150	
	3.5590	-1.4730	5.4350	-2.7120
	7.4000	-1.7910	.1927	1.
KOH, aq		KOH		
KOH(aq)		K(1)O(1)H(1)+(0)		
ref:24		17.Sep.97		
	-103877.	-112104.	28.000	
	3.6808	1.2058	5.2761	-2.8288
	-4.8880	.6520	-.0500	0.
KCl, aq		KCl(0)		
KCl(aq)		K(1)Cl(1)+(0)		
ref:23		15.Sep.97		
	-96500.	-97400.	39.100	
	7.2386	9.8928	1.8616	-3.1880
	-1.434	6.0310	-.0300	0.
KBr, aq		KBr(0)		
KBr(aq)		K(1)Br(1)+(0)		
ref:22		13.Sep.97		
	-90010.	-86320.	47.500	
	8.0580	11.8970	1.0670	-3.2708
	1.4496	-4.6580	-.0050	0.
KI, aq		KI(0)		
KI(aq)		K(1)I(1)+(0)		
ref:22		13.Sep.97		
	-77740.	-71829.	49.200	
	9.8369	16.2406	-.6402	-3.4503
	2.7214	-4.2160	-.0050	0.
KAlO ₂ , aq		KAlO ₂ (0)		
KAlO ₂ (aq)		K(1)Al(1)O(2)+(0)		
ref:24		17.Sep.97		
	-264381.	-274982.	35.700	
	6.0070	6.8858	3.0436	-3.0637
	1.3940	1.2120	-.0500	0.
KHSO ₄ , aq		KHSO ₄ (0)		
KHSO ₄ (aq)		K(1)H(1)S(1)O(4)+(0)		
ref:22		17.Sep.97		
	-243400.	-267400.	55.000	
	9.1226	14.4964	.0453	-3.3782
	39.9849	8.7230	-.0010	0.
KSO ₄ -		KSO ₄ (-)		
KSO ₄ -		K(1)S(1)O(4)-(1)		
ref:22		12.Sep.97		
	-246640.	-276980.	35.000	
	5.9408	6.7274	3.0989	-3.0571
	9.9089	-5.2549	1.0996	-1.
K(Ae), aq		KCH ₃ COO		
K(Ae), aq		K(1)C(2)H(3)O(2)		
ref:33		17 Aug 92		
	-155420.	-175220.	47.600	
	9.9020	16.3937	-0.6874	-3.4566
	40.5631	9.0167	-0.0300	0.
K(Ae)2-		K(CH ₃ COO)2-		
K(Ae)2-		K(1)C(4)H(6)O(4)-(1)		
ref:33		10 Sep 92		
	-242990.	-292900.	60.700	
	17.8481	35.7984	-8.3193	-4.2588
	94.0916	25.2533	0.7097	-1.
K(But), aq		KCH ₃ (CH ₂) ₂ CO ₂		
K(But), aq		K(1)C(4)H(7)O(2)		
ref:27		26.Aug.93		

	-152161.	-187401.	58.768	
	13.9798	26.3520	-4.6040	-3.8683
	82.2231	82.4967	-0.0300	0.
K(But)2-		K(CH ₃ CH ₂ CH ₂ CO ₂)2-		
K(But)2-		K(1)C(8)H(14)O(4)-(1)		
ref:27		26.Aug.93		
	-236375.	-316310.	85.956	
	26.3402	56.5369	-16.4778	-5.1161
	190.0492	59.8305	0.3272	-1.
K(For), aq		KCHO ₂		
K(For), aq		K(1)C(1)H(1)O(2)		
ref:27		26.Aug.93		
	-151413.	-161151.	48.668	
	7.9397	11.6070	1.1839	-3.2587
	9.8889	-1.6447	-0.0300	0.
K(For)2-		K(CHO ₂)2-		
K(For)2-		K(1)C(2)H(2)O(4)-(1)		
ref:27		26.Aug.93		
	-234866.	-264561.	63.193	
	13.6868	25.6408	-4.3345	-3.8389
	20.4973	-0.2055	0.6721	-1.
K(Glyc), aq		KCH ₃ COO ₂		
K(Glyc), aq		K(1)C(2)H(3)O(3)		
ref:27		30.Aug.93		
	-188721.	-214171.	53.168	
	9.8131	16.1774	-0.6044	-3.4477
	40.8892	9.1302	-0.0300	0.
K(Glyc)2-		K(CH ₃ COO ₂)2-		
K(Glyc)2-		K(1)C(4)H(6)O(6)-(1)		
ref:27		30.Aug.93		
	-309741.	-370519.	73.335	
	17.5955	35.1810	-8.0761	-4.2333
	93.1092	25.5242	0.5185	-1.
K(Lac), aq		KCH ₃ CH ₂ COO ₂		
K(Lac), aq		K(1)C(3)H(5)O(3)		
ref:27		30.Aug.93		
	-190081.	-223541.	58.868	
	12.0572	21.6610	-2.7676	-3.6744
	62.8615	16.7671	-0.0300	0.
K(Lac)2-		K(CH ₃ CH ₂ COO ₂)2-		
K(Lac)2-		K(1)C(6)H(10)O(6)-(1)		
ref:27		30.Aug.93		
	-312475.	-388842.	86.182	
	22.2742	46.6039	-12.5625	-4.7055
	143.7822	43.7607	0.3236	-1.
K(Pent), aq		KCH ₃ (CH ₂) ₃ CO ₂		
K(Pent), aq		K(1)C(5)H(9)O(2)		
ref:27		1.Sep.93		
	-150151.	-193881.	65.268	
	16.1707	31.7005	-6.7050	-4.0894
	119.5323	36.4644	-0.0300	0.
K(Pent)2-		K(CH ₃ CH ₂ CH ₂ CH ₂ CO ₂)2-		
K(Pent)2-		K(1)C(10)H(18)O(4)-(1)		
ref:27		1.Sep.93		
	-232342.	-328765.	100.606	
	30.8965	67.6627	-20.8507	-5.5761
	277.0852	90.7965	0.1041	-1.
K(Prop), aq		KCH ₃ CH ₂ CO ₂		
K(Prop), aq		K(1)C(3)H(5)O(2)		
ref:27		15.Sep.93		
	-154321.	-182101.	53.468	
	11.8793	21.2223	-2.5860	-3.6562
	67.4299	18.3550	-0.0300	0.
K(Prop)2-		K(CH ₃ CH ₂ CO ₂)2-		
K(Prop)2-		K(1)C(6)H(10)O(4)-(1)		
ref:27		15.Sep.93		
	-240709.	-306125.	74.011	
	21.9605	45.8408	-12.2686	-4.6740
	156.3920	47.5525	0.5082	-1.
Kr, aq		Kr(0)		
Kr(aq)		Kr(1)+(0)		
ref:3		13.Jul.87		

	3554.	-3650.	15.060	
	6.2721	7.5333	2.7891	-3.0904
	37.8226	8.6985	-.2281	0.
La+2		La(+2)		
La+2		La(1)+(2)		
ref:21		9.Dec.97		
	-77700.	-72200.	0.700	
	0.1064	-7.5128	8.6825	-2.4683
	8.7551	-5.4789	1.0444	2.
La+3		La(+3)		
La+3		La(1)+(3)		
ref:2		13.Jul.87		
	-164000.	-169600.	-52.000	
	-2.7880	-14.3824	10.9602	-2.1844
	4.2394	-10.6122	2.1572	3.
LaOH+2		LaOH+2		
LaOH+2		La(1)O(1)H(1)+(2)		
ref:25		1.Dec.95		
	-208900.	-218200.	-6.6	
	2.6766	-1.2486	6.2460	-2.7273
	-0.5065	-9.0640	1.1586	2.
LaO+		La(1)O(1)+(1)		
LaO+		1.Dec.95		
ref:25		-195900.	-208900.	9.2
	2.7794	-0.9938	6.1373	-2.7378
	-31.7836	-17.5379	0.4100	1.
LaO ₂ H, aq		LaO ₂ H(aq)		
LaO ₂ H(aq)		La(1)O(2)H(1)		
ref:25		1.Dec.95		
	-239300.	-249500.	44.0	
	4.8611	4.0853	4.1503	-2.9478
	-66.8987	-28.3340	-0.0300	0.
LaO ₂ -		LaO ₂ -		
LaO ₂ -		La(1)O(2)-(1)		
ref:25		1.Dec.95		
	-221700.	-230200.	33.8	
	4.8931	4.1637	4.1178	-2.9510
	-44.4914	-24.2193	1.1172	-1.
LaCl+2		LaCl+2		
LaCl+2		La(1)Cl(1)+(2)		
ref:25		1.Dec.95		
	-195800.	-206100.	-25.7	
	0.0273	-7.7112	8.7724	-2.4601
	16.8429	-3.9594	1.4477	2.
LaCl ₂ +		LaCl ₂ +		
LaCl ₂ +		La(1)Cl(2)+(1)		
ref:25		1.Dec.95		
	-226700.	-244900.	-9.6	
	3.1785	-0.0177	5.7515	-2.7782
	6.8244	-5.0483	0.7003	1.
LaCl ₃ , aq		LaCl ₃ (aq)		
LaCl ₃ (aq)		La(1)Cl(3)		
ref:25		1.Dec.95		
	-257600.	-286400.	-3.3	
	6.7243	8.6386	2.3517	-3.1360
	-25.3096	-13.8787	-0.0300	0.
LaCl ₄ -		LaCl ₄ -		
LaCl ₄ -		La(1)Cl(4)-(1)		
ref:25		1.Dec.95		
	-288400.	-331200.	-7.8	
	11.5517	20.4270	-2.2839	-3.6234
	-56.5845	-30.4508	1.7506	-1.
LaF+2		LaF+2		
LaF+2		La(1)F(1)+(2)		
ref:25		1.Dec.95		
	-236600.	-243400.	-16.3	
	-2.6366	-14.2103	11.3144	-2.1914
	17.9127	-3.1239	1.3029	2.
LaF ₂ +		LaF ₂ +		
LaF ₂ +		La(1)F(2)+(1)		
ref:25		1.Dec.95		

	-307800.	-325200.	-12.0	
	-2.3404	-13.4887	11.0343	-2.2213
	12.8261	-3.0531	0.7287	1.
LaF3, aq		LaF3(aq)		
LaF3(aq)		La(1)F(3)		
ref:25		1. Dec. 95		
	-377900.	-410200.	-22.3	
	-2.0506	-12.7816	10.7577	-2.2505
	-15.3002	-10.3998	-0.0300	0.
LaF4-		LaF4-		
LaF4-		La(1)F(4)-(1)		
ref:25		1. Dec. 95		
	-447500.	-500100.	-50.4	
	-0.6263	-9.3060	9.3958	-2.3942
	-35.5432	-25.1639	2.3834	-1.
LaHCO3+2		LaHCO3+2		
LaHCO3+2		La(1)H(1)C(1)O(3)+(2)		
ref:25		1. Dec. 95		
	-307000.	-332900.	-13.7	
	0.8988	-5.5828	7.9358	-2.5481
	38.0197	3.9723	1.2693	2.
LaCO3+		LaCO3+		
LaCO3+		La(1)C(1)O(3)+(1)		
ref:25		1. Dec. 95		
	-299900.	-315100.	-44.2	
	-0.8237	-9.7872	9.5838	-2.3743
	-20.0908	-16.0917	1.2275	1.
LaNO3+2		LaNO3+2		
LaNO3+2		La(1)N(1)O(3)+(2)		
ref:25		1. Dec. 95		
	-191300.	-226000.	-37.8	
	1.6239	-3.8134	7.2430	-2.6213
	32.7363	0.9705	1.6332	2.
LaH2PO4+2		LaH2PO4+2		
LaH2PO4+2		La(1)H(2)P(1)O(4)+(2)		
ref:25		1. Dec. 95		
	-437600.	-482800.	-30.2	
	1.8993	-3.1434	6.9845	-2.6490
	41.8879	4.5243	1.5168	2.
LaSO4+		LaSO4+		
LaSO4+		La(1)S(1)O(4)+(1)		
ref:25		1. Dec. 95		
	-346900.	-382600.	-16.0	
	1.6077	-3.8530	7.2584	-2.6197
	-27.1374	-15.0529	0.7895	1.
La(Ae)+2		LaCH3COO+2		
La(Ae)+2		La(1)C(2)H(3)O(2)+(2)		
ref:33		17 Aug. 92		
	-255750.	-288710.	-29.600	
	3.1548	-0.0804	5.7863	-2.7756
	56.1530	9.5148	1.5067	2.
La(Ae)2+		La(CH3COO)2+		
La(Ae)2+		La(1)C(4)H(6)O(4)+(1)		
ref:33		17 Aug. 92		
	-346160.	-407330.	-10.100	
	9.7487	16.0216	-0.5464	-3.4412
	99.3969	27.1273	0.7003	1.
La(Ae)3, aq		La(CH3COO)3		
La(Ae)3, aq		La(1)C(6)H(9)O(6)		
ref:33		10 Sep. 92		
	-436550.	-527920.	2.800	
	17.1523	34.1020	-7.6583	-4.1887
	136.1072	42.2254	-0.0300	0.
La(But)+2		LaCH3(CH2)2CO2+2		
La(But)+2		La(1)C(4)H(7)O(2)+(2)		
ref:27		26 Aug. 93		
	-252198.	-300593.	-18.438	
	7.1755	9.7404	1.9188	-3.1816
	96.2547	23.9948	1.3375	2.
La(But)2+		La(CH3CH2CH2CO2)2+		
La(But)2+		La(1)C(8)H(14)O(4)+(1)		
ref:27		26 Aug. 93		

	-339359.	-430176.	16.373	
	18.2361	36.7475	-8.6968	-4.2980
	195.2278	61.7045	0.3041	1.
La(For)+2		LaCHO2+2		
La(For)+2		La(1)C(1)H(1)O(2)+(2)		
ref:27		26 Aug. 93		
	-251450.	-274343.	-28.538	
	1.1858	-4.8808	7.6567	-2.5771
	25.2949	-1.1466	1.4867	2.
La(For)2+		La(CHO2)2+		
La(For)2+		La(1)C(2)H(2)O(4)+(1)		
ref:27		26 Aug. 93		
	-337849.	-378757.	-7.498	
	5.5879	5.8627	3.4454	-3.0213
	25.8186	1.6684	0.6644	1.
La(Pent)+2		LaCH3(CH2)3CO2+2		
La(Pent)+2		La(1)C(5)H(9)O(2)+(2)		
ref:27		1. Sep. 93		
	-250188.	-307073.	-11.938	
	9.3322	15.0037	-0.1439	-3.3992
	132.6336	36.9625	1.2366	2.
La(Pent)2+		La(CH3CH2CH2CH2CO2)2+		
La(Pent)2+		La(1)C(10)H(18)O(4)+(1)		
ref:27		1. Sep. 93		
	-335325.	-442418.	31.735	
	22.7873	47.8627	-13.0596	-4.7575
	282.1786	92.6705	0.0717	1.
La(Prop)+2		LaCH3CH2CO2+2		
La(Prop)+2		La(1)C(3)H(5)O(2)+(2)		
ref:27		15. Sep. 93		
	-254208.	-295142.	-23.738	
	5.1026	4.6756	3.9168	-2.9722
	82.2145	18.8530	1.4192	2.
La(Prop)2+		La(CH3CH2CO2)2+		
La(Prop)2+		La(1)C(6)H(10)O(4)+(1)		
ref:27		15. Sep. 93		
	-343420.	-419891.	3.847	
	13.8589	26.0573	-4.4896	-3.8561
	161.6388	49.4265	0.4925	1.
Li+		Li(+)		
Li+		Li(1)+(1)		
ref:2		2. Jun. 87		
	-69933.	-66552.	2.700	
	-0.0237	-0.0690	11.5800	-2.7761
	19.2000	-2.400	4.862	1.
LiOH, aq		LiOH		
LiOH(aq)		Li(1)O(1)H(1)+(0)		
ref:21		12. Sep. 97		
	-108000.	-121500.	1.900	
	2.2749	-2.2242	6.6181	-2.6870
	9.8752	-1.6493	-0.0300	0.
LiCl, aq		LiCl(0)		
LiCl(aq)		Li(1)Cl(1)+(0)		
ref:22		13. Sep. 97		
	-99250.	-105680.	13.140	
	5.5837	5.8554	3.4416	-3.0210
	17.7136	1.1006	-0.0380	0.
Li(Ae), aq		LiCH3COO		
Li(Ae), aq		Li(1)C(2)H(3)O(2)		
ref:33		17 Aug. 92		
	-158610.	-184240.	19.500	
	8.3880	12.6976	0.7639	-3.3038
	56.6767	14.6175	-0.0300	0.
Li(Ae)2-		Li(CH3COO)2-		
Li(Ae)2-		Li(1)C(4)H(6)O(4)-(1)		
ref:33		10 Sept. 92		
	-246810.	-304670.	25.500	
	16.3412	32.1211	-6.8785	-4.1068
	130.4373	36.1810	1.2422	-1.
Lu+3		Lu(+3)		
Lu+3		Lu(1)+(3)		
ref:2		13. Jul. 87		

	-159400.	-167900.	-63.100	
	-3.5630	-16.4812	12.2279	-2.0977
	9.5650	-9.7160	2.4554	3.
LuOH+2		LuOH+2		
LuOH+2		Lu(1)O(1)H(1)+(2)		
ref:25		1. Dec. 95		
	-205700.	-219000.	-21.3	
	2.4374	-1.8325	6.4754	-2.7031
	11.3407	-5.6622	1.3823	2.
LuO+		LuO+		
LuO+		Lu(1)O(1)+(1)		
ref:25		1. Dec. 95		
	-195200.	-212400.	-6.6	
	2.5066	-1.6582	6.3958	-2.7104
	-19.0294	-13.8917	0.6557	1.
LuO2-		LuO2-		
LuO2-		Lu(1)O(2)-(1)		
ref:25		1. Dec. 95		
	-229200.	-242700.	15.8	
	4.3695	2.8897	4.6099	-2.8984
	-20.2927	-16.6824	1.3901	-1.
LuHCO3+2		LuHCO3+2		
LuHCO3+2		Lu(1)H(1)C(1)O(3)+(2)		
ref:25		1. Dec. 95		
	-302300.	-332400.	-29.5	
	0.0033	-7.7669	8.7884	-2.4578
	47.0638	6.3556	1.5067	2.
LuCO3+		LuCO3+		
LuCO3+		Lu(1)C(1)O(3)+(1)		
ref:25		1. Dec. 95		
	-296900.	-314100.	-57.7	
	-1.1546	-10.5926	9.8949	-2.3410
	-15.3205	-15.0324	1.4146	1.
LuCl+2		LuCl+2		
LuCl+2		Lu(1)Cl(1)+(2)		
ref:25		1. Dec. 95		
	-190700.	-204600.	-38.9	
	-0.8820	-9.9277	9.6353	-2.3685
	25.5113	-1.5761	1.6444	2.
LuCl2+		LuCl2+		
LuCl2+		Lu(1)Cl(2)+(1)		
ref:25		1. Dec. 95		
	-221300.	-244000.	-26.2	
	2.1736	-2.4767	6.7285	-2.6765
	22.4451	-0.3982	0.9437	1.
LuCl3, aq		LuCl3(aq)		
LuCl3(aq)		Lu(1)Cl(3)		
ref:25		1. Dec. 95		
	-251900.	-286846.	-25.1	
	5.5129	5.6785	3.5203	-3.0136
	-2.6969	-6.0192	-0.0300	0.
LuCl4-		LuCl4-		
LuCl4-		Lu(1)Cl(4)-(1)		
ref:25		1. Dec. 95		
	-282500.	-333800.	-37.7	
	10.3535	17.5009	-1.1323	-3.5024
	-17.8947	-18.4390	2.1990	-1.
LuF+2		LuF+2		
LuF+2		Lu(1)F(1)+(2)		
ref:25		1. Dec. 95		
	-233300.	-241900.	-23.6	
	-3.5730	-16.4994	12.2192	-2.0968
	25.8414	-0.7406	1.4192	2.
LuF2+		LuF2+		
LuF2+		Lu(1)F(2)+(1)		
ref:25		1. Dec. 95		
	-305600.	-324800.	-19.2	
	-3.3883	-16.0473	12.0392	-2.1155
	27.2751	1.5969	0.8449	1.
LuF3, aq		LuF3(aq)		
LuF3(aq)		Lu(1)F(3)		
ref:25		1. Dec. 95		

	-376563.	-411300.	-31.8	
	-3.2619	-15.7420	11.9264	-2.1281
	7.3122	-2.5402	-0.0300	0.
LuF4-		LuF4-		
LuF4-		Lu(1)F(4)-(1)		
ref:25		1.Dec.95		
	-446900.	-503800.	-66.1	
	-1.8892	-12.3870	10.6014	-2.2668
	1.3820	-13.1521	2.6403	-1.
LuH2PO4+2		LuH2PO4+2		
LuH2PO4+2		Lu(1)H(1)F(1)O(4)+(2)		
ref:25		1.Dec.95		
	-432800.	-482400.	-46.3	
	1.0060	-5.3207	7.8310	-2.5589
	50.9918	6.9076	1.7607	2.
LuNO3+2		LuNO3+2		
LuNO3+2		Lu(1)N(1)O(3)+(2)		
ref:25		1.Dec.95		
	-186700.	-227300.	-58.7	
	0.7523	-5.9353	8.0622	-2.5335
	42.4310	3.3538	1.9413	2.
LuSO4+		LuSO4+		
LuSO4+		Lu(1)S(1)O(4)+(1)		
ref:25		1.Dec.95		
	-342200.	-380630.	-26.6	
	1.2811	-4.6503	7.5708	-2.5867
	-16.5465	-13.9936	0.9570	1.
Lu(Ae)+2		LuCH3COO+2		
Lu(Ae)+2		Lu(1)C(2)H(3)O(2)+(2)		
ref:33		10 Sep 92		
	-251290.	-288534.	-45.400	
	2.2608	-2.2622	6.6412	-2.6854
	65.2386	11.8981	1.7486	2.
Lu(Ae)2+		Lu(CH3COO)2+		
Lu(Ae)2+		Lu(1)C(4)H(6)O(4)+(1)		
ref:33		10 Sep 92		
	-342150.	-409310.	-31.600	
	8.7721	13.6382	0.3879	-3.3427
	115.7906	31.7774	1.0276	1.
Lu(Ae)3,aq		Lu(CH3COO)3		
Lu(Ae)3,aq		Lu(1)C(6)H(9)O(6)		
ref:33		10 Sep 92		
	-432290.	-531620.	-25.300	
	15.9409	31.1419	-6.4897	-4.0663
	158.7197	50.0850	-0.0300	0.
Lu+4		Lu(+4)		
Lu+4		Lu(1)(+4)		
ref:21		10.Dec.97		
	27000.	9800.	-108.000	
	-4.3487	-18.3921	12.9608	-2.0186
	39.9042	-3.5234	3.8147	4.
Mg+2		Mg(+2)		
Mg+2		Mg(1)(+2)		
ref:2		1.Jul.87		
	-108505.	-111367.	-33.000	
	-8.217	-8.5990	8.3900	-2.3900
	20.8000	-5.8920	1.5372	2.
MgOH+		MgOH(+1)		
MgOH+		Mg(1)O(1)H(1)+(1)		
ref:21		12.Sep.97		
	-149255.	-152955.	-19.100	
	2.3105	-2.1365	6.5827	-2.6906
	32.0008	3.2394	8.449	1.
MgCl+		MgCl(+)		
MgCl+		Mg(1)Cl(1)+(1)		
ref:22		12.Sep.97		
	-139700.	-150933.	-19.000	
	2.2230	-2.3505	6.6669	-2.6818
	28.6016	2.058	8.449	1.
MgF+		MgF(+)		
MgF+		Mg(1)F(1)+(1)		
ref:22		12.Sep.97		

	-177690.	-190950.	-28.070	
	-2975	-8.5049	9.0858	-2.4274
	38.0239	4.9301	9706	1.
MgHCO3+		MgHCO3(+)		
MgHCO3+		Mg(1)H(1)C(1)O(3)+(1)		
ref:5		JAN.91		
	-250200.	-275750.	-3.000	
	2.7171	-1.1469	6.2008	-2.7316
	49.0065	9.9391	5985	1.
MgCO3,aq		MgCO3(0)		
MgCO3(aq)		Mg(1)C(1)O(3)+(0)		
ref:22		17.Sep.97		
	-238760.	-270570.	-24.000	
	-7355	-9.5745	9.5062	-2.3831
	-10.2416	-8.6159	-0.0380	0.
Mg(HSiO3)+		Mg(HSiO3)(+1)		
Mg(HSiO3)+		Mg(1)H(1)Si(1)O(3)+(1)		
ref:22		15.Sep.97		
	-353025.	-389700.	-23.780	
	6.289	-6.2428	8.1967	-2.5209
	36.7882	4.6702	9.177	1.
Mg(Ae)+		MgCH3COO+		
Mg(Ae)+		Mg(1)C(2)H(3)O(2)+(1)		
ref:33		17 Mar 93		
	-198520.	-229480.	-13.100	
	5.4981	5.6424	3.5341	-3.0122
	64.6297	14.8894	0.7483	1.
Mg(Ae)2,aq		Mg(CH3COO)2		
Mg(Ae)2,aq		Mg(1)C(4)H(6)O(4)		
ref:33		17 Mar 93		
	-287830.	-349260.	-1.200	
	12.3982	22.4898	-3.0853	-3.7086
	121.5413	37.1627	-0.0300	0.
Mg(Ala)+		Mg(C3H6NO2)+		
Mg(Ala)+		Mg(1)C(3)H(6)N(1)O(2)+(1)		
ref:27		14.Aug.93		
	-185839.	-231745.	7.968	
	8.3169	12.5282	0.8210	-3.2968
	74.3351	19.2750	0.4322	1.
Mg(Ala)2,aq		Mg(C3H6NO2)2		
Mg(Ala)2,aq		Mg(1)C(6)H(12)N(2)O(4)		
ref:27		14.Aug.93		
	-262231.	-352641.	44.042	
	18.5836	37.5967	-9.0318	-4.3331
	151.6713	47.6352	-0.0300	0.
Mg(But)+		MgCH3(CH2)2CO2+		
Mg(But)+		Mg(1)C(4)H(7)O(2)+(1)		
ref:27		26.Aug.93		
	-193879.	-240741.	-3.482	
	9.5279	15.4831	-0.3354	-3.4190
	104.9814	29.3694	0.6063	1.
Mg(But)2,aq		Mg(CH3CH2CH2CO2)2		
Mg(But)2,aq		Mg(1)C(8)H(14)O(4)		
ref:27		26.Aug.93		
	-278694.	-370578.	22.608	
	21.0196	43.5448	-11.3703	-4.5790
	221.0226	71.7399	-0.0300	0.
Mg(For)+		MgCHO2+		
Mg(For)+		Mg(1)C(1)H(1)O(2)+(1)		
ref:27		26.Aug.93		
	-194318.	-215678.	-13.582	
	3.5392	0.8593	5.4140	-2.8144
	34.0484	4.2279	0.7584	1.
Mg(For)2,aq		Mg(CHO2)2		
Mg(For)2,aq		Mg(1)C(2)H(2)O(4)		
ref:27		26.Aug.93		
	-279367.	-321177.	-0.709	
	8.2496	12.3613	0.8923	-3.2899
	48.2938	11.7038	-0.0300	0.
Mg(Gly)+		Mg(C2H4NO2)+		
Mg(Gly)+		Mg(1)C(2)H(4)N(1)O(2)+(1)		
ref:27		30.Aug.93		

	-188490.	-225174.	6.198	
	5.8466	6.4953	3.1944	-3.0474
	48.0086	10.0502	0.4555	1.
Mg(Gly)2,aq		Mg(C2H4NO2)2		
Mg(Gly)2,aq		Mg(1)C(4)H(8)N(2)O(4)		
ref:27		30.Aug.93		
	-267874.	-340003.	39.956	
	13.3443	24.8034	-4.0028	-3.8043
	88.2942	25.6069	-0.0300	0.
Mg(Glyc)+		MgCH3OCO2+		
Mg(Glyc)+		Mg(1)C(2)H(3)O(3)+(1)		
ref:27		30.Aug.93		
	-231489.	-266450.	-2.000	
	5.3533	5.2917	3.6660	-2.9977
	63.4339	15.0028	0.5831	1.
Mg(Glyc)2,aq		Mg(CH3OCO2)2		
Mg(Glyc)2,aq		Mg(1)C(4)H(6)O(6)		
ref:27		30.Aug.93		
	-353983.	-424040.	18.948	
	12.2102	22.0339	-2.9133	-3.6898
	122.3206	37.4335	-0.0300	0.
Mg(Lac)+		MgCH3CH2OCO2+		
Mg(Lac)+		Mg(1)C(3)H(5)O(3)+(1)		
ref:27		30.Aug.93		
	-232904.	-274593.	8.000	
	7.5465	10.6466	1.5614	-3.2190
	84.0159	22.6398	0.4322	1.
Mg(Lac)2,aq		Mg(CH3CH2OCO2)2		
Mg(Lac)2,aq		Mg(1)C(6)H(10)O(6)		
ref:27		30.Aug.93		
	-356826.	-440700.	37.734	
	16.9548	33.6148	-7.4563	-4.1685
	174.7884	55.6701	-0.0300	0.
Mg(Pent)+		MgCH3(CH2)3CO2+1		
Mg(Pent)+		Mg(1)C(5)H(9)O(2)+(1)		
ref:27		7.Sep.93		
	-191528.	-246880.	3.018	
	11.6846	20.7465	-2.3979	-3.6366
	141.3617	42.3371	0.5055	1.
Mg(Pent)2,aq		Mg(CH3CH2CH2CH2CO2)2		
Mg(Pent)2,aq		Mg(1)C(10)H(18)O(4)		
ref:27		7.Sep.93		
	-274046.	-382313.	37.615	
	25.6513	54.8536	-15.8129	-5.0465
	310.1143	102.7059	-0.0300	0.
Mg(Prop)+		MgCH3CH2CO2+		
Mg(Prop)+		Mg(1)C(3)H(5)O(2)+(1)		
ref:27		15.Sep.93		
	-196257.	-235660.	-8.782	
	7.4530	10.4173	1.6547	-3.2096
	90.8863	24.2277	0.6821	1.
Mg(Prop)2,aq		Mg(CH3CH2CO2)2		
Mg(Prop)2,aq		Mg(1)C(6)H(10)O(4)		
ref:27		15.Sep.93		
	-283437.	-360889.	10.373	
	16.5787	32.6970	-7.0965	-4.1306
	185.6977	59.4618	-0.0300	0.
Mn+2		Mn(+2)		
Mn+2		Mn(1)(+2)		
ref:21		30.Apr.97		
	-55100.	-52900.	-16.200	
	-1.1015	-8.0040	8.8400	-2.4480
	16.6672	-3.8697	1.4006	2.
MnOH+		MnOH(+)		
MnOH+		Mn(1)O(1)H(1)+(1)		
ref:21		12.Sep.97		
	-97300.	-106800.	300	
	3213	-6.9901	8.4821	-2.4899
	16.2994	-1.2623	5.464	1.
MnO,aq		MnO		
MnO(aq)		Mn(1)O(1)+(0)		
ref:21		12.Sep.97		

	-81500.	-99100.	-2.500	
	-0.0376	-7.8656	8.8229	-2.4537
	-1.5528	-5.6215	-0.0300	0.
MnO2-2		MnO2(-2)		
MnO2-2		Mn(1)O(2)-(2)		
ref:21		12.Sep.97		
	-102600.	-133300.	-15.200	
	1.1489	-4.9713	7.6937	-2.5734
	-4.0375	-17.5991	3.4408	-2.
MnCl+		MnCl(+)		
MnCl+		Mn(1)Cl(1)+(1)		
ref:22		17.Sep.97		
	-86290.	-88280.	12.000	
	2.8119	-9.126	6.1017	-2.7412
	24.2838	2.0824	3.686	1.
MnF+		MnF(+)		
MnF+		Mn(1)F(1)+(1)		
ref:22		15.Sep.97		
	-123641.	-132454.	-13.300	
	.3136	-7.0128	8.4994	-2.4890
	30.2840	2.9518	.7483	1.
MnSO4, aq		MnSO4(0)		
MnSO4(aq)		Mn(1)S(1)O(4)+(0)		
ref:22		17.Sep.97		
	-235640.	-266750.	5.000	
	2.1354	-2.5645	6.7510	-2.6729
	-6.2564	-7.2308	-0.0380	0.
Mn(Ae)+		MnCH3COO+		
Mn(Ae)+		Mn(1)C(2)H(3)O(2)+(1)		
ref:33		17 Aug 92		
	-144430.	-169560.	6.300	
	6.0776	7.0570	2.9786	-3.0706
	63.5668	15.4577	0.4555	1.
Mn(Ae)2, aq		Mn(CH3COO)2		
Mn(Ae)2, aq		Mn(1)C(4)H(6)O(4)		
ref:33		10 Sept 92		
	-233850.	-287670.	24.200	
	13.1542	24.3405	-3.8236	-3.7851
	124.7315	38.2716	-0.0300	0.
Mn(Ae)3-		Mn(CH3COO)3-		
Mn(Ae)3-		Mn(1)C(6)H(9)O(6)-(1)		
ref:33		10 Sept 92		
	-322580.	-408280.	31.500	
	21.6217	45.0124	-11.9409	-4.6397
	211.3036	64.5717	1.1536	-1.
Mn(Ala)+		Mn(C3H6NO2)+		
Mn(Ala)+		Mn(1)C(3)H(6)N(1)O(2)+(1)		
ref:27		14.Aug.93		
	-133580.	-173180.	28.965	
	8.8873	13.9167	0.2842	-3.3542
	73.0237	19.8433	0.1124	1.
Mn(Ala)2, aq		Mn(C3H6NO2)2		
Mn(Ala)2, aq		Mn(1)C(6)H(12)N(2)O(4)		
ref:27		14.Aug.93		
	-212074.	-294245.	71.520	
	19.3395	39.4414	-9.7539	-4.4094
	154.8617	48.7440	-0.0300	0.
Mn(But)+		MnCH3(CH2)2CO2+		
Mn(But)+		Mn(1)C(4)H(7)O(2)+(1)		
ref:27		26.Aug.93		
	-140788.	-181344.	17.515	
	10.0986	16.8770	-0.8847	-3.4766
	103.6776	29.9377	0.2873	1.
Mn(But)2, aq		Mn(CH3CH2CH2CO2)2		
Mn(But)2, aq		Mn(1)C(8)H(14)O(4)		
ref:27		26.Aug.93		
	-226367.	-310012.	50.086	
	21.7755	45.3895	-12.0924	-4.6553
	224.2130	72.8487	-0.0300	0.
Mn(For)+		MnCHO2+		
Mn(For)+		Mn(1)C(1)H(1)O(2)+(1)		
ref:27		26.Aug.93		

	-140681.	-155735.	7.415	
	4.1094	2.2545	4.8590	-2.8721
	32.7312	4.7962	0.4379	1.
Mn(For)2, aq		Mn(CHO2)2		
Mn(For)2, aq		Mn(1)C(2)H(2)O(4)		
ref:27		26.Aug.93		
	-226030.	-259601.	26.769	
	9.0055	14.2060	0.1701	-3.3662
	51.4842	12.8126	-0.0300	0.
Mn(Gly)+		Mn(C2H4NO2)+		
Mn(Gly)+		Mn(1)C(2)H(4)N(1)O(2)+(1)		
ref:27		27.Aug.93		
	-134771.	-165803.	25.000	
	6.4299	7.9185	2.6372	-3.1063
	47.0493	10.6185	0.1739	1.
Mn(Gly)2, aq		Mn(C2H4NO2)2		
Mn(Gly)2, aq		Mn(1)C(4)H(8)N(2)O(4)		
ref:27		27.Aug.93		
	-214101.	-278847.	64.561	
	14.1002	26.6481	-4.7253	-3.8805
	91.4846	26.7157	-0.0300	0.
Mn(Glyc)+		MnCH3OCO2+		
Mn(Glyc)+		Mn(1)C(2)H(3)O(3)+(1)		
ref:27		30.Aug.93		
	-177828.	-208594.	11.915	
	5.9609	6.7761	3.0811	-3.0590
	63.1381	15.5712	0.3735	1.
Mn(Glyc)2, aq		Mn(CH3OCO2)2		
Mn(Glyc)2, aq		Mn(1)C(4)H(6)O(6)		
ref:27		30.Aug.93		
	-300155.	-364736.	37.157	
	12.9567	23.8786	-3.6359	-3.7660
	125.5108	38.5424	-0.0300	0.
Mn(Lac)+		MnCH3CH2OCO2+		
Mn(Lac)+		Mn(1)C(3)H(5)O(3)+(1)		
ref:27		30.Aug.93		
	-178981.	-217756.	17.615	
	8.1746	12.1772	0.9667	-3.2823
	84.2763	23.2081	0.2832	1.
Mn(Lac)2, aq		Mn(CH3CH2OCO2)2		
Mn(Lac)2, aq		Mn(1)C(6)H(10)O(6)		
ref:27		30.Aug.93		
	-302971.	-383047.	50.317	
	17.7108	35.4655	-8.1945	-4.2450
	177.9788	56.7789	-0.0300	0.
Mn(Pent)+		MnCH3(CH2)3CO2+		
Mn(Pent)+		Mn(1)C(5)H(9)O(2)+(1)		
ref:27		9.Sep.93		
	-138601.	-187646.	24.015	
	12.2563	22.1440	-2.9511	-3.6943
	140.0855	42.9054	0.1895	1.
Mn(Pent)2, aq		Mn(CH3CH2CH2CH2CO2)2		
Mn(Pent)2, aq		Mn(1)C(10)H(18)O(4)		
ref:27		9.Sep.93		
	-222006.	-322033.	65.092	
	26.4073	56.6981	-16.5352	-5.1228
	313.3047	103.8147	-0.0300	0.
Mn(Prop)+		MnCH3CH2CO2+		
Mn(Prop)+		Mn(1)C(3)H(5)O(2)+(1)		
ref:27		15.Sep.93		
	-143016.	-176112.	12.215	
	8.0255	11.8128	1.1111	-3.2672
	89.6329	24.7960	0.3686	1.
Mn(Prop)2, aq		Mn(CH3CH2CO2)2		
Mn(Prop)2, aq		Mn(1)C(6)H(10)O(4)		
ref:27		15.Sep.93		
	-230823.	-300037.	37.850	
	17.3347	34.5476	-7.8345	-4.2071
	188.8879	60.5707	-0.0300	0.
Mn+3		Mn(+3)		
Mn+3		Mn(1)+(3)		
ref:21		11.Sep.97		

	-20300.	-30700.	-74.000	
	-2.9320	-14.9340	11.6041	-2.1615
	16.0611	-8.2492	2.7025	3.
MnO4-2		MnO4(-2)		
MnO4-2		Mn(1)O(4)-(2)		
ref:21		11.Sep.97		
	-120400.	-156600.	15.500	
	5.6596	6.0368	3.3786	-3.0285
	-5.2910	-16.5602	2.9803	-2.
MnO4-		MnO4(-)		
MnO4-		Mn(1)O(4)-(1)		
ref:21		30.Apr.97		
	-107600.	-129900.	46.500	
	7.8248	11.3277	1.2912	-3.2472
	13.6317	-3.4012	.9248	-1.
MoO4-2		MoO4(-2)		
MoO4-2		Mo(1)O(4)-(2)		
ref:21		11.Sep.97		
	-200400.	-238300.	9.000	
	6.9651	2.7095	18.6617	-2.8909
	6.6829	-12.7102	3.0777	-2.
N2, aq		N2(0)		
N2(aq)		N(2)+(0)		
ref:3		7.Feb.89		
	4347.	-2495.	22.900	
	6.2046	7.3685	2.8539	-3.0836
	35.7911	8.3726	-3.468	0.
N2H5+		N2H5(+)		
N2H5+		N(2)H(5)+(1)		
ref:21		15.Sep.97		
	19700.	-1800.	36.100	
	5.5711	5.8193	3.4672	-3.0195
	16.0644	.3876	.0057	1.
N2H6+2		N2H6(+2)		
N2H6+2		N(2)H(6)+(2)		
ref:21		15.Sep.97		
	21100.	-4000.	24.000	
	4.4762	3.1468	4.5160	-2.9090
	21.7574	.1635	.6936	2.
NH3, aq		NH3(0)		
NH3(aq)		N(1)H(3)+(0)		
ref:3		13.Jul.87		
	-6383.	-19440.	25.770	
	5.0911	2.7970	8.6248	-2.8946
	20.3000	-1.1700	-0.0500	0.
NH4+		NH4(+)		
NH4+		N(1)H(4)+(1)		
ref:2		1.Jul.87		
	-18990.	-31850.	26.570	
	3.8763	2.3448	8.5605	-2.8759
	17.4500	-.0210	.1502	1.
NH4(Ae), aq		NH4CH3COO		
NH4(Ae), aq		N(1)C(2)H(7)O(2)		
ref:33		10.Sep.92		
	-107550.	-147230.	50.700	
	11.2849	19.7719	-2.0187	-3.5963
	58.7075	15.3233	-0.0300	0.
NH4(Ae)2-		NH4(CH3COO)2-		
NH4(Ae)2-		N(1)C(4)H(10)O(4)-(1)		
ref:33		10.Sep.92		
	-195640.	-265200.	64.700	
	19.3685	39.5090	-9.7736	-4.4122
	128.9386	37.5581	0.6495	-1.
N2O2-2		N2O2(-2)		
N2O2-2		N(2)O(2)-(2)		
ref:21		15.Sep.97		
	33200.	-2590.	6.600	
	3.3102	.2986	5.6380	-2.7912
	.1067	-15.1139	3.1145	-2.
NO2-		NO2(-)		
NO2-		N(1)O(2)-(1)		
ref:2		3.Jul.87		

	-7700.	-25000.	29.400	
	5.5864	5.8590	3.4472	-3.0212
	3.4260	-7.7808	1.1847	-1.
NO3-		NO3(-)		
NO3-		N(1)O(3)-(1)		
ref:2		1.Jul.87		
	-26507.	-49429.	35.120	
	7.3161	6.7824	-4.6838	-3.0594
	7.7000	-6.7250	1.0977	-1.
Na+		Na(+)		
Na+		Na(1)+(1)		
ref:20		12.Sep.97		
	-62591.	-57433.	13.960	
	1.8390	-2.2850	3.2560	-2.7260
	18.1800	-2.9810	.3306	1.
NaOH,aq		NaOH		
NaOH(aq)		Na(1)O(1)H(1)+(0)		
ref:20		12.Sep.97		
	-99095.	-112927.	6.000	
	1.4675	-4.1982	7.4001	-2.6054
	22.0000	3.0000	-7.200	0.
NaCl,aq		NaCl(0)		
NaCl(aq)		Na(1)Cl(1)+(0)		
ref:20		29.Apr.97		
	-92910.	-96160.	28.000	
	5.0363	4.7365	3.4154	-2.9748
	10.8000	-1.3000	-.0380	0.
NaF,aq		NaF(0)		
NaF(aq)		Na(1)F(1)+(0)		
ref:22		13.Sep.97		
	-128570.	-135860.	12.000	
	2.5336	-1.5922	6.3689	-2.7131
	12.3804	-7.7531	-.0380	0.
NaBr,aq		NaBr(0)		
NaBr(aq)		Na(1)Br(1)+(0)		
ref:22		13.Sep.97		
	-85610.	-84830.	34.100	
	6.1257	7.1789	2.9214	-3.0757
	10.1926	-1.4111	-.0700	0.
NaI,aq		NaI(0)		
NaI(aq)		Na(1)I(1)+(0)		
ref:22		13.Sep.97		
	-72900.	-69101.	37.700	
	7.6488	10.8979	1.4597	-3.2295
	12.1001	-9.690	-.0010	0.
NaAlO2,aq		NaAlO2		
NaAlO2(aq)		Na(1)Al(1)O(2)+(0)		
ref:20		29.Apr.97		
	-260277.	-277259.	11.100	
	4.7640	4.850	12.6690	-4.8975
	32.4300	3.8000	-.1000	0.
NaB(OH)4,aq		NaBOH4(0)		
NaB(OH)4		Na(1)B(1)O(4)H(4)+(0)		
ref:39		20.Jan.98		
	-338580.	-378810.	39.200	
	6.2600	7.5100	2.8000	-3.0900
	73.5000	-18.9000	0.0000	0.
NaSO4-		NaSO4(-)		
NaSO4-		Na(1)S(1)O(4)-(1)		
ref:39		20.Jan.98		
	-241780.	-275480.	20.440	
	4.7500	3.8300	4.2500	-2.7300
	21.0500	-2.0200	1.3100	-1.
NaHSiO3,aq		NaHSiO3(0)		
NaHSiO3(aq)		Na(1)H(1)Si(1)O(3)+(0)		
ref:22		17.Sep.97		
	-307890.	-333894.	10.000	
	3.4928	.7500	5.4483	-2.8100
	20.2395	1.9785	-.0380	0.
Na(Ae),aq		NaCH3COO		
Na(Ae),aq		Na(1)C(2)H(3)O(2)		
ref:33		17 Aug 92		

	-150720.	-173540.	34.200	
	8.3514	12.6125	0.7884	-3.3003
	49.8989	12.2617	-0.0300	0.
Na(Ae)2-		Na(CH3COO)2-		
Na(Ae)2-		Na(1)C(4)H(6)O(4)-(1)		
ref:33		10 Sept 92		
	-238470.	-292400.	44.000	
	16.2062	31.7884	-6.7416	-4.0930
	114.6437	11.5846	0.9633	-1.
Na(But),aq		NaCH3(CH2)2CO2		
Na(But),aq		Na(1)C(4)H(7)O(2)		
ref:27		26.Aug.93		
	-147269.	-185529.	45.433	
	12.4293	22.5644	-3.1127	-3.7117
	91.5590	26.7417	-0.0300	0.
Na(But)2-		Na(CH3CH2CH2CO2)2-		
Na(But)2-		Na(1)C(8)H(14)O(4)-(1)		
ref:27		26.Aug.93		
	-231511.	-315475.	69.236	
	24.6982	52.5272	-14.9010	-4.9504
	210.5987	66.1618	0.5806	-1.
Na(For),aq		NaCHO2		
Na(For),aq		Na(1)C(1)H(1)O(2)		
ref:27		16.Aug.93		
	-146521.	-159279.	35.333	
	6.3891	7.8196	2.6757	-3.1022
	19.2249	1.6002	-0.0300	0.
Na(For)2-		Na(CHO2)2-		
Na(For)2-		Na(1)C(2)H(2)O(4)-(1)		
ref:27		16.Aug.93		
	-230001.	-263725.	45.473	
	12.0449	21.6308	-2.7566	-3.6731
	41.0498	6.1257	0.9257	-1.
Na(Glyc),aq		Na(CH3COO)2		
Na(Glyc),aq		Na(1)C(2)H(3)O(3)		
ref:27		16.Aug.93		
	-183829.	-212299.	39.833	
	8.2625	12.3962	0.8714	-3.2914
	50.2253	12.3751	-0.0300	0.
Na(Glyc)2-		Na(CH3COO)2-		
Na(Glyc)2-		Na(1)C(4)H(6)O(6)-(1)		
ref:27		16.Aug.93		
	-304876.	-369684.	56.615	
	15.9533	31.1718	-6.5007	-4.0675
	113.6563	31.8554	0.7716	-1.
Na(Lac),aq		NaCH3CH2COO2		
Na(Lac),aq		Na(1)C(3)H(5)O(3)		
ref:27		16.Aug.93		
	-185189.	-221669.	45.533	
	10.5067	17.8733	-1.2760	-3.5178
	72.1974	20.0121	-0.0300	0.
Na(Lac)2-		Na(CH3CH2COO2)2-		
Na(Lac)2-		Na(1)C(6)H(10)O(6)-(1)		
ref:27		16.Aug.93		
	-307610.	-388006.	69.462	
	20.6323	42.5939	-10.9852	-4.5397
	164.3333	50.0919	0.5771	-1.
Na(Pent),aq		NaCH3(CH2)3CO2		
Na(Pent),aq		Na(1)C(5)H(9)O(2)		
ref:27		16.Aug.93		
	-145259.	-192009.	51.933	
	14.6201	27.9193	-5.2291	-3.9331
	128.8681	39.7094	-0.0300	0.
Na(Pent)2-		Na(CH3CH2CH2CH2CO2)2-		
Na(Pent)2-		Na(1)C(10)H(18)O(4)-(1)		
ref:27		16.Aug.93		
	-227477.	-327929.	83.886	
	29.2549	63.6519	-19.2699	-5.4103
	297.6442	97.1277	0.3584	-1.
Na(Prop),aq		NaCH3CH2CO2		
Na(Prop),aq		Na(1)C(3)H(5)O(2)		
ref:27		16.Aug.93		

	-149429.	-180229.	40.133	
	10.3288	17.4408	-1.1100	-3.4999
	76.7660	21.5999	-0.0300	0.
Na(Prop)2-		Na(CH3CH2CO2)2-		
Na(Prop)2-		Na(1)C(6)H(10)O(4)-(1)		
ref:27		16.Aug.93		
	-235845.	-305289.	57.291	
	20.3185	41.8311	-10.6923	-4.5082
	176.9413	53.8837	0.7616	-1.
NbO3-		NbO3(-1)		
NbO3-		Nb(1)O(3)-(1)		
ref:21		5.Jan.98		
	-227100.	-255300.	3.200	
	5.0915	4.6485	3.9273	-2.9711
	-3.3376	-11.4066	1.5829	-1.
Nd+2		Nd(+2)		
Nd+2		Nd(1)+(2)		
ref:21		10.Dec.97		
	-100200.	-96100.	-0.500	
	0.0860	-7.5654	8.7097	-2.4661
	9.0616	-5.4382	1.0649	2.
Nd+3		Nd(+3)		
Nd+3		Nd(1)+(3)		
ref:2		13.Jul.87		
	-160600.	-166500.	-49.500	
	-3.3707	-14.5452	8.3211	-2.1777
	1.6236	-11.8344	2.2550	3.
NdOH+2		NdOH+2		
NdOH+2		Nd(1)O(1)H(1)+(2)		
ref:25		1.Dec.95		
	-206200.	-215500.	-3.3	
	2.7413	-1.0874	6.1755	-2.7339
	-3.2077	-9.8381	1.1072	2.
NdO+		NdO+		
NdO+		Nd(1)O(1)+(1)		
ref:25		1.Dec.95		
	-194000.	-207000.	12.7	
	2.8304	-0.8716	6.0955	-2.7429
	-34.6589	-18.3731	0.3587	1.
NdO2-		NdO2-		
NdO2-		Nd(1)O(2)-(1)		
ref:25		1.Dec.95		
	-223400.	-231700.	37.8	
	5.0229	4.4850	3.9825	-2.9643
	-49.9788	-25.9304	1.0559	-1.
NdCl+2		NdCl+2		
NdCl+2		Nd(1)Cl(1)+(2)		
ref:25		1.Dec.95		
	-192400.	-203000.	-22.7	
	-0.6746	-9.4228	9.439	-2.3894
	8.4972	-6.7094	1.4006	2.
NdCl2+		NdCl2+		
NdCl2+		Nd(1)Cl(2)+(1)		
ref:25		1.Dec.95		
	-223400.	-241500.	-5.9	
	2.3933	-1.9354	6.5057	-2.6989
	-9.1790	-10.4138	0.6388	1.
NdCl3,aq		NdCl3(aq)		
NdCl3(aq)		Nd(1)Cl(3)		
ref:25		1.Dec.95		
	-254300.	-282700.	1.6	
	5.8726	6.5582	3.1706	-3.0500
	-51.4009	-22.9475	-0.0300	0.
NdCl4-		NdCl4-		
NdCl4-		Nd(1)Cl(4)-(1)		
ref:25		1.Dec.95		
	-285100.	-327000.	-1.1	
	10.5672	18.0205	-1.3324	-3.5239
	-97.4278	-44.3105	1.6455	-1.
NdF+2		NdF+2		
NdF+2		Nd(1)F(1)+(2)		
ref:25		1.Dec.95		

	-233900.	-241200.	-14.6	
	-3.3312	-15.9098	11.9894	-2.1212
	9.7675	-5.8738	1.2776	2.
NdF2+		NdF2+		
NdF2+		Nd(1)F(2)+(1)		
ref:25		1.Dec.95		
	-305600.	-323500.	-10.4	
	-3.1113	-15.3752	11.7849	-2.1433
	-2.7862	-8.4186	0.7096	1.
NdF3,aq		NdF3(aq)		
NdF3(aq)		Nd(1)F(3)		
ref:25		1.Dec.95		
	-376100.	-408900.	-20.1	
	-2.9023	-14.8620	11.5767	-2.1645
	-41.3917	-19.4685	-0.0300	0.
NdF4-		NdF4-		
NdF4-		Nd(1)F(4)-(1)		
ref:25		1.Dec.95		
	-446100.	-498700.	-46.8	
	-1.5889	-11.6574	10.3228	-2.2970
	-75.7884	-39.0236	2.3433	-1.
NdHCO3+2		NdHCO3+2		
NdHCO3+2		Nd(1)H(1)O(3)+(2)		
ref:25		1.Dec.95		
	-303400.	-329200.	-10.2	
	0.1936	-7.3045	8.6115	-2.4769
	29.5859	1.2224	1.2126	2.
NdCO3+		NdCO3+		
NdCO3+		Nd(1)C(1)O(3)+(1)		
ref:25		1.Dec.95		
	-297300.	-309500.	-41.2	
	-1.1199	-10.5088	9.8634	-2.3445
	-24.1105	-17.3139	1.1729	1.
NdH2PO4+2		NdH2PO4+2		
NdH2PO4+2		Nd(1)H(2)P(1)O(4)+(2)		
ref:25		1.Dec.95		
	-433891.	-479076.	-26.57	
	1.1931	-4.8625	7.6492	-2.5779
	33.4289	1.7743	1.4574	2.
NdNO3+2		NdNO3+2		
NdNO3+2		Nd(1)N(1)O(3)+(2)		
ref:25		1.Dec.95		
	-188185.	-222586.	-33.09	
	0.9124	-5.5498	7.9230	-2.5495
	24.1306	-1.7794	1.5579	2.
NdSO4+		NdSO4+		
NdSO4+		Nd(1)S(1)O(4)+(1)		
ref:25		1.Dec.95		
	-343500.	-379100.	-12.3	
	1.3167	-4.5634	7.5367	-2.5903
	-25.1244	-16.2751	0.7484	1.
Nd(Ae)+2		NdCH3COO+2		
Nd(Ae)+2		Nd(1)C(2)H(3)O(2)+(2)		
ref:33		17 Aug 92		
	-252510.	-285470.	-26.100	
	2.4521	-1.7961	6.4594	-2.7046
	47.7869	6.7648	1.4574	2.
Nd(Ae)2+		Nd(CH3COO)2+		
Nd(Ae)2+		Nd(1)C(4)H(6)O(4)+(1)		
ref:33		17 Aug 92		
	-343330.	-404110.	-5.300	
	8.9606	14.0990	0.2066	-3.3618
	83.3168	21.7619	0.6305	1.
Nd(Ae)3,aq		Nd(CH3COO)3		
Nd(Ae)3,aq		Nd(1)C(6)H(9)O(6)		
ref:33		10 Sep 92		
	-433560.	-524090.	9.100	
	16.3006	32.0216	-6.8393	-4.1027
	110.0156	33.1567	-0.0300	0.
Nd+4		Nd(+4)		
Nd+4		Nd(1)+4		
ref:21		10.Dec.97		

	-47000.	-62900.	-98.700	
	-4.2742	-18.2128	12.8958	-2.0260
	40.2772	-2.9327	3.6707	4.
Ne,aq		Ne(0)		
Ne(aq)		Ne(1)+(0)		
ref:3		13.Jul.87		
	4565.	-870.	16.740	
	4.4709	3.1352	4.5177	-2.9086
	27.0977	5.0523	-2.535	0.
Ni+2		Ni(+2)		
Ni+2		Ni(1)+(2)		
ref:2		1.Jul.87		
	-10900.	-12900.	-30.800	
	-1.6942	-11.9181	10.4344	-2.2863
	13.1905	-5.4179	1.5067	2.
NiOH+		NiOH(+)		
NiOH+		Ni(1)O(1)H(1)+(1)		
ref:21		16.Sep.97		
	-52850.	-67650.	-18.000	
	-1.7075	-9.5051	9.4760	-2.3860
	31.9091	3.2801	.8222	1.
NiO,aq		NiO(0)		
NiO(aq)		Ni(1)O(1)+(0)		
ref:21		16.Sep.97		
	-39340.	-56230.	25.000	
	-1.4334	-11.2786	10.1751	-2.3126
	11.7507	-9.975	-0.0300	0.
NiO2-2		NiO2(-2)		
NiO2-2		Ni(1)O(2)-(2)		
ref:21		16.Sep.97		
	-64200.	-101800.	-38.900	
	-1.4814	-8.9493	9.2495	-2.4089
	30.9114	-6.5993	3.7992	-2.
NiCl+		NiCl(+)		
NiCl+		Ni(1)Cl(1)+(1)		
ref:22		12.Sep.97		
	-40920.	-51400.	-17.000	
	1.1319	-5.0147	7.7140	-2.5716
	18.3863	-1.3846	8.111	1.
NiF+		NiF(+)		
NiF+		Ni(1)F(1)+(1)		
ref:22		15.Sep.97		
	-79768.	-92315.	-26.360	
	-1.4455	-11.3080	10.1875	-2.3115
	22.3170	-4.857	.9570	1.
Ni(Ae)+		NiCH3COO+		
Ni(Ae)+		Ni(1)C(2)H(3)O(2)		
ref:33		17 Aug 92		
	-101120.	-131450.	-11.700	
	4.3556	2.8512	4.6343	-2.8968
	56.0621	11.9745	0.7287	1.
Ni(Ae)2,aq		Ni(CH3COO)2		
Ni(Ae)2,aq		Ni(1)C(4)H(6)O(4)		
ref:33		10 Sep 92		
	-190610.	-251280.	0.700	
	11.1327	19.4031	-1.8801	-3.5810
	105.1782	31.4753	-0.0300	0.
Ni(Ae)3-		Ni(CH3COO)3-		
Ni(Ae)3-		Ni(1)C(6)H(9)O(6)-(1)		
ref:33		10 Sep 92		
	-279690.	-374030.	1.900	
	19.5212	39.8827	-9.9226	-4.4277
	182.3942	53.0847	1.6030	-1.
Ni(Ala)+		Ni(C3H6NO2)+		
Ni(Ala)+		Ni(1)C(3)H(5)N(1)O(2)+(1)		
ref:27		14.Aug.93		
	-93527.	-137131.	15.000	
	7.1451	9.6640	1.9536	-3.1784
	64.9703	16.3600	0.3260	1.
Ni(Ala)2,aq		Ni(C3H6NO2)2		
Ni(Ala)2,aq		Ni(1)C(6)H(12)N(2)O(4)		
ref:27		14.Aug.93		

	-174545.	-262972.	50.000	
	17.3180	34.5040	-7.8105	-4.2053
	135.3083	41.9478	-0.0300	0.
Ni(But)+		NiCH3(CH2)2CO2+		
Ni(But)+		Ni(1)C(4)H(7)O(2)+(1)		
ref:27		26.Aug.93		
	-97925.	-143687.	-0.482	
	8.3767	12.6711	0.7713	-3.3027
	96.1758	26.4544	0.5608	1.
Ni(But)2,aq		Ni(CH3CH2CH2CO2)2		
Ni(But)2,aq		Ni(1)C(8)H(14)O(4)		
ref:27		26.Aug.93		
	-184117.	-274625.	26.534	
	19.7540	40.4521	-10.1489	-4.4512
	204.6596	66.0525	-0.0300	0.
Ni(For)+		NiCHO2+		
Ni(For)+		Ni(1)C(1)H(1)O(2)+(1)		
ref:27		26.Aug.93		
	-97313.	-117573.	-10.582	
	2.3868	-1.9558	6.5236	-2.6980
	25.2125	1.3130	0.7096	1.
Ni(For)2,aq		Ni(CHO2)2		
Ni(For)2,aq		Ni(1)C(2)H(2)O(4)		
ref:27		26.Aug.93		
	-182853.	-223287.	3.217	
	6.9840	9.2687	2.1133	-3.1621
	31.9308	6.0164	-0.0300	0.
Ni(Gly)+		Ni(C2H4NO2)+		
Ni(Gly)+		Ni(1)C(2)H(4)N(1)O(2)+(1)		
ref:27		27.Aug.93		
	-94541.	-129289.	12.000	
	4.6813	3.6515	4.3097	-2.9299
	38.8214	7.1352	0.3686	1.
Ni(Gly)2,aq		Ni(C2H4NO2)2		
Ni(Gly)2,aq		Ni(1)C(4)H(8)N(2)O(4)		
ref:27		27.Aug.93		
	-176531.	-246055.	48.000	
	12.0787	21.7107	-2.7819	-3.6764
	71.9311	19.9195	-0.0300	0.
Ni(Glyc)+		NiCH3OCO2+		
Ni(Glyc)+		Ni(1)C(2)H(3)O(3)+(1)		
ref:27		30.Aug.93		
	-135153.	-171125.	-6.082	
	4.2391	2.5699	4.7380	-2.8851
	55.6378	12.0879	0.6472	1.
Ni(Glyc)2,aq		Ni(CH3OCO2)2		
Ni(Glyc)2,aq		Ni(1)C(4)H(6)O(6)		
ref:27		30.Aug.93		
	-258711.	-330154.	13.605	
	10.9446	18.9412	-1.6924	-3.5619
	105.9573	31.7462	-0.0300	0.
Ni(Lac)+		NiCH3CH2OCO2+		
Ni(Lac)+		Ni(1)C(3)H(5)O(3)+(1)		
ref:27		30.Aug.93		
	-135599.	-179581.	-0.382	
	6.4541	7.9800	2.6079	-3.1088
	76.8142	19.7248	0.5608	1.
Ni(Lac)2,aq		Ni(CH3CH2OCO2)2		
Ni(Lac)2,aq		Ni(1)C(6)H(10)O(6)		
ref:27		30.Aug.93		
	-259957.	-346896.	26.765	
	15.6892	30.5284	-6.2512	-4.0409
	158.4254	49.9827	-0.0300	0.
Ni(Pent)+		NiCH3(CH2)3CO2+		
Ni(Pent)+		Ni(1)C(5)H(9)O(2)+(1)		
ref:27		7.Sep.93		
	-95874.	-150126.	6.018	
	10.5339	17.9393	-1.3007	-3.5205
	132.5696	39.4221	0.4615	1.
Ni(Pent)2,aq		Ni(CH3CH2CH2CH2CO2)2		
Ni(Pent)2,aq		Ni(1)C(10)H(18)O(4)		
ref:27		7.Sep.93		

-180002.	-286892.	41.540	
24.3858	51.7607	-14.5917	-4.9187
293.7512	97.0185	-0.0300	0.
Ni (Prop)+	NiCH3CH2CO2+		
Ni (Prop)+	Ni (1)C(3)H(5)O(2)+(1)		
ref:27	15.Sep.93		
-99635.	-137936.	-5.782	
6.3025	7.6097	2.7546	-3.0935
82.1010	21.3127	0.6388	1.
Ni (Prop)2,aq	Ni (CH3CH2CO2)2		
Ni (Prop)2,aq	Ni (1)C(6)H(10)O(4)		
ref:27	15.Sep.93		
-187632.	-263708.	14.298	
15.3131	29.6105	-5.8911	-4.0030
169.3347	53.7744	-0.0300	0.
Np+2	Np(+2)		
Np+2	Np(1)+(2)		
ref:36	16.Jan.98		
-57700.	-52100.	-0.400	
0.0860	-7.5654	8.7097	-2.4661
9.0615	-5.4382	1.0649	2.
Np+3	Np(+3)		
Np+3	Np(1)+(3)		
ref:36	16.Jan.98		
-122500.	-126000.	-46.400	
-2.8540	-14.7452	11.5327	-2.1693
5.9466	-10.4288	2.2853	3.
Np+4	Np(+4)		
Np+4	Np(1)+(4)		
ref:36	16.Jan.98		
-117800.	-133000.	-101.400	
-4.2886	-18.2500	12.9154	-2.0244
40.1636	-3.0956	3.7093	4.
NpO2+	NpO2(+)		
NpO2+	Np(1)O(2)+(1)		
ref:36	18.Jan.98		
-218700.	-233800.	-5.200	
3.4012	0.5220	5.5478	-2.8005
-1.9730	-7.8826	0.6305	1.
NpO2+2	NpO2(+2)		
NpO2+2	Np(1)O(2)+(2)		
ref:36	18.Jan.98		
-190200.	-205700.	-22.100	
3.0837	-0.2498	5.8429	-2.7686
25.4318	-0.7938	1.3914	2.
O2,aq	O(2)		
O2(aq)	O(2)+(0)		
ref:3	26.Jun.87		
3954.	-2900.	26.040	
5.7889	6.3536	3.2528	-3.0417
35.3530	8.3726	-3.943	0.
OCN-	OCN(-1)		
OCN-	O(1)C(1)N(1)-(1)		
ref:21	11.Sep.97		
-23280.	-34900.	25.500	
5.7591	6.2815	3.2790	-3.0386
8.3509	-6.2530	1.2422	-1.
OH-	OH(-)		
OH-	O(1)H(1)-(1)		
ref:2	29.Apr.97		
-37595.	-54977.	-2.560	
1.2527	.0738	1.8423	-2.7821
4.1500	-10.3460	1.7246	-1.
PO4-3	PO4(-3)		
PO4-3	P(1)O(4)-(3)		
ref:21	11.Sep.97		
-243500.	-305300.	-53.000	
-5258	-9.0576	9.2927	-2.4045
-15.1599	-28.4155	5.6114	-3.
P2O7-4	P2O7(-4)		
P2O7-4	P(2)O(7)-(4)		
ref:21	15.Sep.97		

-458700.	-542800.	-28.000	
7.0687	9.4773	2.0273	-3.1707
-11.7223	-31.3692	6.9069	-4.
Pa+2	Pa(+2)		
Pa+2	Pa(1)+(2)		
ref:36	16.Jan.98		
-46700.	-40600.	1.600	
0.1292	-7.4602	8.6697	-2.4705
8.5137	-5.5196	1.0309	2.
Pa+3	Pa(+3)		
Pa+3	Pa(1)+(3)		
ref:36	16.Jan.98		
-104600.	-107700.	-44.300	
-2.7822	-14.5714	11.4684	-2.1765
5.4329	-10.5103	2.2550	3.
Pa+4	Pa(+4)		
Pa+4	Pa(1)+(4)		
ref:36	16.Jan.98		
-144300.	-158800.	-98.800	
-4.2742	-18.2128	12.8958	-2.0260
40.2772	-2.9327	3.6707	4.
Pb+2	Pb(+2)		
Pb+2	Pb(1)+(2)		
ref:2	1.Jul.87		
-5710.	220.	4.200	
-0.0051	-7.7939	8.8134	-2.4568
8.6624	-5.6216	1.0788	2.
PbOH+	PbOH(+)		
PbOH+	Pb(1)O(1)H(1)+(1)		
ref:21	16.Sep.97		
-53950.	-68900.	-10.000	
2.4669	-1.7600	6.4461	-2.7061
.8970	-7.1085	.7003	1.
PbO,aq	PbO(0)		
PbO(aq)	Pb(1)O(1)+(0)		
ref:21	16.Sep.97		
-39350.	-44700.	22.000	
2.8906	-7.236	6.0351	-2.7490
-18.7244	-11.5899	-0.0300	0.
PbCl+	PbCl(+)		
PbCl+	Pb(1)Cl(1)+(1)		
ref:22	17.Sep.97		
-39050.	-38630.	28.000	
2.9756	-5.130	5.9446	-2.7577
10.2182	-2.0364	.1281	1.
PbCl2,aq	PbCl2(0)		
PbCl2(aq)	Pb(1)Cl(2)+(0)		
ref:22	17.Sep.97		
-71200.	-77700.	47.000	
6.6429	8.4416	2.4251	-3.1279
6.9886	-2.6272	-0.0380	0.
PbCl3-	PbCl3(-)		
PbCl3-	Pb(1)Cl(3)-(1)		
ref:22	17.Sep.97		
-102150.	-117700.	59.000	
11.0547	19.2141	-1.8089	-3.5733
5.9698	-5.4586	.7356	-1.
PbCl4-2	PbCl4(-2)		
PbCl4-2	Pb(1)Cl(4)-(2)		
ref:22	17.Sep.97		
-133177.	-159209.	66.000	
16.1777	31.7230	-6.7255	-4.0904
5.9335	-10.2007	2.2126	-2.
PbF+	PbF(+)		
PbF+	Pb(1)F(1)+(1)		
ref:22	6.Jan.98		
-75860.	-80591.	8.260	
0.4488	-6.6827	8.3696	-2.5027
15.9794	-0.9898	0.4266	1.
PbF2,aq	PbF2(0)		
PbF2(aq)	Pb(1)F(2)+(0)		
ref:22	6.Jan.98		

-145056.	-162783.	4.480	
1.0888	-5.1199	7.7554	-2.5673
16.5543	0.6976	-0.0380	0.
Pb(HS)2,aq	Pb(HS)2(0)		
Pb(HS)2(aq)	Pb(1)H(2)S(2)+(0)		
ref:22	6.Jan.98		
-20058.	-22851.	52.560	
7.4825	10.4917	1.6193	-3.2127
28.5025	4.8505	-0.0380	0.
Pb(HS)3-	Pb(HS)3(-)		
Pb(HS)3-	Pb(1)H(3)S(3)-(1)		
ref:22	6.Jan.98		
-18971.	-28704.	68.080	
12.3397	22.3518	-3.0422	-3.7030
39.6902	6.7025	0.5980	-1.
Pb(Ae)+	PbCH3COO+		
Pb(Ae)+	Pb(1)C(2)H(3)O(2)+(1)		
ref:33	16.Feb.91		
-97250.	-115880.	36.000	
6.1543	7.2429	2.9082	-3.0783
48.0824	11.5161	0.0057	1.
Pb(Ae)2,aq	Pb(CH3COO)2		
Pb(Ae)2,aq	Pb(1)C(4)H(6)O(4)		
ref:33	16.Feb.91		
-186890.	-229460.	69.800	
13.4090	24.9584	-4.0570	-3.8107
102.6053	30.5811	-0.0300	0.
Pb(Ae)3-	Pb(CH3COO)3-		
Pb(Ae)3-	Pb(1)C(6)H(9)O(6)-(1)		
ref:33	08.Apr.93		
-277750.	-348760.	88.600	
21.6128	44.9921	-11.9361	-4.6389
165.9232	51.5732	0.2871	-1.
Pb(Ala)+	Pb(C3H6NO2)+		
Pb(Ala)+	Pb(1)C(3)H(6)O(2)N(1)+(1)		
ref:27	14.Aug.93		
-87191.	-120275.	58.688	
8.9640	14.1088	0.1992	-3.3622
57.5413	15.9017	-0.3372	1.
Pb(Ala)2,aq	Pb(C3H6NO2)2		
Pb(Ala)2,aq	Pb(1)C(6)H(12)N(2)O(4)		
ref:27	14.Aug.93		
-166271.	-239191.	110.416	
19.5944	40.0653	-10.0034	-4.4352
132.7356	41.0535	-0.0300	0.
Pb(But)+	PbCH3(CH2)2CO2+		
Pb(But)+	Pb(1)C(4)H(7)O(2)+(1)		
ref:27	16.Aug.93		
-92817.	-126856.	47.238	
10.1748	17.0643	-0.9612	-3.4843
88.1802	25.9961	-0.1639	1.
Pb(But)2,aq	Pb(CH3CH2CH2CO2)2		
Pb(But)2,aq	Pb(1)C(8)H(14)O(4)		
ref:27	16.Aug.93		
-179077.	-253472.	88.982	
22.0303	46.0136	-12.3417	-4.6811
202.0867	65.1583	-0.0300	0.
Pb(For)+	PbCHO2+		
Pb(For)+	Pb(1)C(1)H(1)O(2)+(1)		
ref:27	16.Aug.93		
-92150.	-100688.	37.138	
4.1867	2.4389	4.7962	-2.8797
17.2647	0.8546	-0.0100	1.
Pb(For)2,aq	Pb(CHO2)2		
Pb(For)2,aq	Pb(1)C(2)H(2)O(4)		
ref:27	16.Aug.93		
-177718.	-202038.	65.665	
9.2604	14.8299	-0.0792	-3.3920
29.3579	5.1222	-0.0300	0.
Pb(Gly)+	Pb(C2H4NO2)+		
Pb(Gly)+	Pb(1)C(2)H(4)O(2)N(1)+(1)		
ref:27	16.Aug.93		

-88450.	-112312.	56.918	
6.4949	8.0790	2.5712	-3.1129
31.2483	6.6769	-0.3103	1.
Pb(Glyc)2,aq	Pb(C2H4NO2)2		
Pb(Glyc)2,aq	Pb(1)C(4)H(8)N(2)O(4)		
ref:27	16.Aug.93		
-168353.	-222992.	106.330	
14.3550	27.2722	-4.9746	-3.9063
69.3582	19.0253	-0.0300	0.
Pb(Glyc)+	PbCH3OCO2+		
Pb(Glyc)+	Pb(1)C(2)H(3)O(3)+(1)		
ref:27	16.Aug.93		
-130018.	-154267.	41.638	
8.2809	12.4416	0.8529	-3.2932
47.6306	11.6296	-0.0788	1.
Pb(Glyc)2,aq	Pb(CH3OCO2)2		
Pb(Glyc)2,aq	Pb(1)C(4)H(6)O(6)		
ref:27	16.Aug.93		
-253616.	-308946.	76.1054	
17.9656	36.0834	-8.4279	-4.2706
103.3846	30.8519	-0.0300	0.
Pb(Lac)+	PbCH3CH2OCO2+		
Pb(Lac)+	Pb(1)C(3)H(5)O(3)+(1)		
ref:27	16.Aug.93		
-131350.	-163610.	47.338	
8.2516	12.3686	0.8843	-3.2902
68.8031	19.2665	-0.1656	1.
Pb(Lac)2,aq	Pb(CH3CH2OCO2)2		
Pb(Lac)2,aq	Pb(1)C(6)H(10)O(6)		
ref:27	16.Aug.93		
-256295.	-327120.	89.213	
17.9656	36.0834	-8.4279	-4.2706
155.8527	49.0884	-0.0300	0.
Pb(Pent)+	PbCH3(CH2)3CO2+		
Pb(Pent)+	Pb(1)C(5)H(9)O(2)+(1)		
ref:27	7.Sep.93		
-90793.	-133322.	53.738	
12.3324	22.3315	-3.0281	-3.7021
124.5859	38.9638	-0.2620	1.
Pb(Pent)2,aq	Pb(CH3CH2CH2CH2CO2)2		
Pb(Pent)2,aq	Pb(1)C(10)H(18)O(4)		
ref:27	7.Sep.93		
-175016.	-265793.	103.988	
26.6621	57.3222	-16.7845	-5.1486
291.1785	96.1242	-0.0300	0.
Pb(Prop)+	PbCH3CH2CO2+		
Pb(Prop)+	Pb(1)C(3)H(5)O(2)+(1)		
ref:27	16.Aug.93		
-95672.	-122252.	41.938	
8.1013	12.0011	1.0303	-3.2750
74.1249	20.8544	-0.0838	1.
Pb(Prop)2,aq	Pb(CH3CH2CO2)2		
Pb(Prop)2,aq	Pb(1)C(6)H(10)O(4)		
ref:27	16.Aug.93		
-184202.	-244164.	76.746	
17.5895	35.1655	-8.0682	-4.2326
166.7618	52.8802	-0.0300	0.
Pd+2	Pd(+2)		
Pd+2	Pd(1)+(2)		
ref:2	1.Jul.87		
42200.	42080.	-22.600	
-6079	-9.2658	9.3919	-2.3960
15.3780	-4.3179	1.4006	2.
PdOH+	PdOH(+)		
PdOH+	Pd(1)O(1)H(1)+(1)		
ref:38	20.Jan.98		
-13000.	-27000.	-13.100	
0.8484	-5.7100	7.9943	-2.5429
26.0312	1.4735	0.7483	1.
PdO,aq	PdO(0)		
PdO(aq)	Pd(1)O(1)+(0)		
ref:38	20.Jan.98		

-11500.	-24700.	-9.900	
-0.2867	-8.4815	9.0836	-2.4284
1.8779	-3.7056	-0.2560	0.
PdCl+	PdCl(+)		
PdCl+	Pd(1)Cl(1)+(1)		
ref:38	20.Jan.98		
2500.	-5500.	-6.800	
2.1187	-2.6082	6.7752	-2.6712
28.4586	2.6138	0.6557	1.
PdCl2,aq	PdCl2(0)		
PdCl2(aq)	Pd(1)Cl(2)+(0)		
ref:38	20.Jan.98		
-35200.	-53600.	0.400	
4.8050	3.9511	4.1971	-2.9423
39.4445	9.1990	-0.2083	0.
PdCl3-	PdCl3(-)		
PdCl3-	Pd(1)Cl(3)-(1)		
ref:38	20.Jan.98		
-69800.	-102000.	-3.300	
8.4201	12.7782	0.7276	-3.3073
67.2220	12.8203	1.6759	-1.
PdCl4-2	PdCl4(-2)		
PdCl4-2	Pd(1)Cl(4)-(2)		
ref:38	20.Jan.98		
-104000.	-154000.	-21.200	
12.0275	21.5864	-2.7344	-3.6714
84.6986	12.9349	3.5371	-2.
Pm+2	Pm(+2)		
Pm+2	Pm(1)+(2)		
ref:21	10.Dec.97		
-92700.	-88400.	-1.900	
0.0520	-7.6511	8.7496	-2.4626
9.4303	-5.3771	1.0858	2.
Pm+3	Pm(+3)		
Pm+3	Pm(1)+(3)		
ref:21	10.Dec.97		
-158000.	-164000.	-50.000	
-3.2471	-15.7059	11.9125	-2.1296
1.5577	-12.1195	2.3369	3.
Pm+4	Pm(+4)		
Pm+4	Pm(1)+(4)		
ref:21	10.Dec.97		
-34200.	-50400.	-99.700	
-4.2836	-18.2319	12.8955	-2.0252
40.2193	-2.9938	3.6835	4.
Pr+2	Pr(+2)		
Pr+2	Pr(1)+(2)		
ref:21	10.Dec.97		
-92700.	-88400.	0.700	
0.1064	-7.5128	8.6835	-2.4683
8.7551	-5.4789	1.0444	2.
Pr+3	Pr(+3)		
Pr+3	Pr(1)+(3)		
ref:21	15.Sep.97		
-162600.	-168800.	-50.000	
-3.2061	-14.2894	8.5328	-2.1882
3.9607	-11.2844	2.3369	3.
PrOH+2	PrOH+2		
PrOH+2	Pr(1)O(1)H(1)+(2)		
ref:25	1.Dec.95		
-208000.	-217700.	-4.0	
2.7325	-1.1081	6.1822	-2.7331
-2.6058	-9.6751	1.1216	2.
PrO+	PrO+		
PrO+	Pr(1)O(1)+(1)		
ref:25	1.Dec.95		
-195700.	-209000.	12.0	
2.8201	-0.8945	6.0986	-2.7419
-34.0998	-18.2101	0.3685	1.
PrO2-	PrO2-		
PrO2-	Pr(1)O(2)-(1)		
ref:25	1.Dec.95		

-224700.	-233400.	37.0	
5.0000	4.4265	4.0111	-2.9619
-48.8607	-25.5841	1.0691	-1.
PrHCO3+2	PrHCO3+2		
PrHCO3+2	Pr(1)H(1)C(1)O(3)+(2)		
ref:25	1.Dec.95		
-305500.	-336800.	-28.3	
0.4387	-6.7049	8.3735	-2.5017
35.6716	2.4598	1.4867	2.
PrCO3+	PrCO3+		
PrCO3+	Pr(1)C(1)O(3)+(1)		
ref:25	1.Dec.95		
-299100.	-311600.	-41.8	
-1.0523	-10.3427	9.7957	-2.3513
-22.3638	-16.7639	1.1907	1.
PrCl+2	PrCl+2		
PrCl+2	Pr(1)Cl(1)+(2)		
ref:25	1.Dec.95		
-194400.	-205300.	-23.3	
-0.5190	-9.0442	9.2930	-2.4050
12.1429	-5.4719	1.4099	2.
PrCl2+	PrCl2+		
PrCl2+	Pr(1)Cl(2)+(1)		
ref:25	1.Dec.95		
-225400.	-243800.	-6.6	
2.5689	-1.5094	6.3442	-2.7165
-2.0764	-7.9993	0.6557	1.
PrCl3,aq	PrCl3(aq)		
PrCl3(aq)	Pr(1)Cl(3)		
ref:25	1.Dec.95		
-256300.	-285200.	0.6	
6.0618	7.0173	2.9974	-3.0690
-39.6599	-18.8665	-0.0300	0.
PrCl4-	PrCl4-		
PrCl4-	Pr(1)Cl(4)-(1)		
ref:25	1.Dec.95		
-287100.	-329500.	-2.4	
10.7857	18.5526	-1.5382	-3.5459
-79.2757	-38.0736	1.6681	-1.
PrF+2	PrF+2		
PrF+2	Pr(1)F(1)+(2)		
ref:25	1.Dec.95		
-235700.	-243400.	-15.0	
-3.1759	-15.5305	11.8402	-2.1369
13.4052	-4.6364	1.2860	2.
PrF2+	PrF2+		
PrF2+	Pr(1)F(2)+(1)		
ref:25	1.Dec.95		
-307300.	-325600.	-10.7	
-2.9414	-14.9592	11.6187	-2.1605
4.1602	-6.0041	0.7096	1.
PrF3,aq	PrF3(aq)		
PrF3(aq)	Pr(1)F(3)		
ref:25	1.Dec.95		
-377800.	-410800.	-20.5	
-2.7130	-14.4032	11.4035	-2.1835
-29.6505	-15.3876	-0.0300	0.
PrF4-	PrF4-		
PrF4-	Pr(1)F(4)-(1)		
ref:25	1.Dec.95		
-447700.	-500700.	-47.6	
-1.3780	-11.1430	10.1215	-2.3182
-57.8444	-32.7867	2.3433	-1.
PrH2PO4+2	PrH2PO4+2		
PrH2PO4+2	Pr(1)H(2)P(1)O(4)+(2)		
ref:25	1.Dec.95		
-436000.	-481500.	-27.3	
1.3489	-4.4844	7.5048	-2.5935
37.0786	3.0118	1.4671	2.
PrNO3+2	PrNO3+2		
PrNO3+2	Pr(1)N(1)O(3)+(2)		
ref:25	1.Dec.95		

	-190000.	-224900.	-34.0	
	1.0684	-5.1660	7.7658	-2.5653
	27.7877	-0.5419	1.5684	2.
PrSO4+		PrSO4+		
PrSO4+		Pr(1)S(1)O(4)+(1)		
ref:25		1.Dec.95		
	-335500.	-381500.	-13.1	
	-1.3885	-4.3882	7.4678	-2.5975
	-23.4508	-15.7251	0.7483	1.
Pr(Ae)+2		PrCH3COO+2		
Pr(Ae)+2		Pr(1)C(2)H(3)O(2)+(2)		
ref:33		10 Sep 92		
	-254570.	-287880.	-26.800	
	2.6078	-1.4116	6.2996	-2.7205
	41.9424	4.7024	1.4671	2.
Pr(Ae)2+		Pr(CH3COO)2+		
Pr(Ae)2+		Pr(1)C(4)H(6)O(4)+(1)		
ref:33		10 Sep 92		
	-345500.	-406710.	-6.200	
	9.1362	14.5250	0.0441	-3.3794
	71.8931	17.7378	0.6472	1.
Pr(Ae)3,aq		Pr(CH3COO)3		
Pr(Ae)3,aq		Pr(1)C(6)H(9)O(6)		
ref:33		10 Sep 92		
	-435690.	-526750.	7.800	
	16.4898	32.4806	-7.0127	-4.1216
	90.4470	26.3552	-0.0300	0.
Pr+4		Pr(+4)		
Pr+4		Pr(1)+(4)		
ref:21		10.Dec.97		
	-72700.	-88700.	-98.500	
	-4.2742	-18.2128	12.8958	-2.0260
	40.3357	-2.9123	3.6707	4.
Pu+2		Pu(+2)		
Pu+2		Pu(1)+(2)		
ref:36		16.Jan.98		
	-73500.	-68800.	-2.000	
	0.0520	-7.6511	8.7496	-2.4626
	9.4303	-5.3771	1.0858	2.
Pu+3		Pu(+3)		
Pu+3		Pu(1)+(3)		
ref:36		16.Jan.98		
	-137400.	-141500.	-47.100	
	-2.8780	-14.8009	11.5494	-2.1670
	6.0992	-10.4084	2.2955	3.
Pu+4		Pu(+4)		
Pu+4		Pu(1)+(4)		
ref:36		16.Jan.98		
	-114200.	-128200.	-100.300	
	-4.2792	-18.2247	12.8996	-2.0255
	40.2204	-3.0345	3.6964	4.
PuO2+		PuO2(+)		
PuO2+		Pu(1)O(2)+(1)		
ref:36		18.Jan.98		
	-203100.	-218600.	-5.200	
	3.4012	0.5220	5.5478	-2.8005
	-1.9730	-7.8826	0.6305	1.
PuO2+2		PuO2(+2)		
PuO2+2		Pu(1)O(2)+(2)		
ref:36		18.Jan.98		
	-180900.	-196500.	-21.100	
	3.1049	-0.1993	5.8259	-2.7707
	24.0918	-1.2012	1.3732	2.
Ra+2		Ra(+2)		
Ra+2		Ra(1)+(2)		
ref:21		11.Sep.97		
	-134200.	-126100.	12.900	
	3740	-6.8598	8.4279	-2.4953
	5.7496	-5.9474	8.6644	2.
Ra(Ae)+		RaCH3COO+		
Ra(Ae)+		Ra(1)C(2)H(3)O(2)+(1)		
ref:33		11 Mar 93		

	-223900.	-239380.	48.000	
	6.8554	8.9566	2.2315	-3.1492
	44.1726	10.7370	-0.1754	1.
Ra(Ae)2,aq		Ra(CH3COO)2		
Ra(Ae)2,aq		Ra(1)C(4)H(6)O(4)		
ref:33		11 Mar 93		
	-312940.	-353260.	78.800	
	14.2584	27.0324	-4.8725	-3.8964
	98.2315	29.0609	-0.0300	0.
Rb+		Rb(+)		
Rb+		Rb(1)+(1)		
ref:2		29.Jun.87		
	-67800.	-60020.	28.800	
	4.2913	-9.041	7.4070	-2.7416
	5.7923	-3.6457	1.502	1.
RbOH,aq		RbOH(0)		
RbOH(aq)		Rb(1)O(1)H(1)+(0)		
ref:21		12.Sep.97		
	-105100.	-113000.	32.000	
	4.5875	3.4168	4.4132	-2.9202
	-10.2264	-8.6363	-0.0300	0.
RbCl,aq		RbCl(0)		
RbCl(aq)		Rb(1)Cl(1)+(0)		
ref:22		13.Sep.97		
	-97870.	-96800.	48.560	
	7.4542	10.4227	1.6465	-3.2098
	-5.6468	-7.1086	-0.0100	0.
RbF,aq		RbF(0)		
RbF(aq)		Rb(1)F(1)+(0)		
ref:22		13.Sep.97		
	-136450.	-139710.	31.600	
	4.3647	2.8788	4.6115	-2.8980
	-3.2196	-6.2938	-0.0010	0.
RbBr,aq		RbBr(0)		
RbBr(aq)		Rb(1)Br(1)+(0)		
ref:22		13.Sep.97		
	-91010.	-85730.	54.200	
	8.4668	12.8952	6.747	-3.3120
	-5.1720	-6.9436	-0.0100	0.
RbI,aq		RbI(0)		
RbI(aq)		Rb(1)I(1)+(0)		
ref:22		13.Sep.97		
	-78900.	-71720.	56.300	
	10.5225	17.9145	-1.2982	-3.5195
	-3.8174	-6.5015	-0.0010	0.
Rb(Ae),aq		RbCH3COO		
Rb(Ae),aq		Rb(1)C(2)H(3)O(2)		
ref:33		19 Aug 92		
	-156110.	-174950.	53.700	
	10.6948	18.3352	-1.4623	-3.5369
	33.9961	6.7343	-0.0300	0.
Rb(Ae)2-		Rb(CH3COO)2-		
Rb(Ae)2-		Rb(1)C(4)H(6)O(4)-(1)		
ref:33		10 Sep 92		
	-244000.	-292490.	68.300	
	18.6924	37.8608	-9.1318	-4.3441
	60.2139	20.8000	0.5941	-1.
ReO4-		ReO4(-)		
ReO4-		Re(1)O(4)-(1)		
ref:21		10.Dec.97		
	-166000.	-188200.	48.100	
	8.6103	20.0581	-16.7541	-3.6081
	18.9300	-7.0800	9.004	-1.
Rh+3		Rh(+3)		
Rh+3		Rh(1)+(3)		
ref:21		15.Sep.97		
	52450.	42830.	-71.600	
	-3.2866	-15.8020	11.9492	-2.1256
	13.0231	-9.1862	2.6654	3.
Rn,aq		Rn(0)		
Rn(aq)		Rn(1)+(0)		
ref:3		13.Jul.87		

	2790.	-5000.	16.000	
	9.0862	14.4045	0.884	-3.3745
	52.5774	13.8725	-2.423	0.
RuO4-2		RuO4(-2)		
RuO4-2		Ru(1)O(4)-(2)		
ref:21		5.Jan.98		
	-71600.	-110200.	6.600	
	6.9502	9.1887	2.1387	-3.1588
	6.6707	-12.8325	3.1145	-2.
S2-2		S2(-2)		
S2-2		S(2)-(2)		
ref:2		3.Jul.87		
	19000.	7200.	6.800	
	5.5797	5.8426	3.4536	-3.0205
	-3.3496	-16.2955	3.1083	-2.
S2O3-2		S2O3(-2)		
S2O3-2		S(2)O(3)-(2)		
ref:2		3.Jul.87		
	-124900.	-155000.	16.000	
	6.6685	12.4951	-7.7281	-3.2955
	-0.0577	-14.7066	2.9694	-2.
S2O4-2		S2O4(-2)		
S2O4-2		S(2)O(4)-(2)		
ref:21		11.Sep.97		
	-143500.	-180100.	22.000	
	6.6784	8.5280	2.3917	-3.1314
	3.3115	-13.2399	2.8772	-2.
S2O5-2		S2O5(-2)		
S2O5-2		S(2)O(5)-(2)		
ref:21		11.Sep.97		
	-189000.	-232000.	25.000	
	7.3618	10.1945	1.7414	-3.2003
	3.9715	-12.8732	2.8343	-2.
S2O6-2		S2O6(-2)		
S2O6-2		S(2)O(6)-(2)		
ref:21		11.Sep.97		
	-231000.	-280400.	30.000	
	8.2257	12.3054	9.087	-3.2876
	5.0331	-12.2621	2.7587	-2.
S2O8-2		S2O8(-2)		
S2O8-2		S(2)O(8)-(2)		
ref:2		3.Jul.87		
	-266500.	-321400.	58.400	
	13.3622	24.8454	-4.0153	-3.8061
	12.9632	-8.1271	2.3281	-2.
S3-2		S3(-2)		
S3-2		S(3)-(2)		
ref:2		3.Jul.87		
	17600.	6200.	15.800	
	6.7661	8.7396	2.3150	-3.1403
	-3.3595	-14.8288	2.9749	-2.
S3O6-2		S3O6(-2)		
S3O6-2		S(3)O(6)-(2)		
ref:21		11.Sep.97		
	-229000.	-279000.	33.000	
	8.4155	12.7691	7.268	-3.3068
	5.6683	-11.8954	2.7131	-2.
S4-2		S4(-2)		
S4-2		S(4)-(2)		
ref:2		3.Jul.87		
	16500.	5500.	24.700	
	7.9381	11.6012	1.1902	-3.2586
	2.6081	-13.3622	2.8390	-2.
S4O6-2		S4O6(-2)		
S4O6-2		S(4)O(6)-(2)		
ref:21		11.Sep.97		
	-248700.	-292600.	61.500	
	10.2672	17.2902	-1.0502	-3.4937
	11.7042	-8.4122	2.2805	-2.
S5-2		S5(-2)		
S5-2		S(5)-(2)		
ref:2		3.Jul.87		

	15700.	5100.	33.600	
	9.1107	14.4645	0.0649	-3.3770
	5.5361	-11.9159	2.7051	-2.
SS06-2		S506(-2)		
SS06-2		S(5)O(6)-(2)		
ref:21		11.Sep.97		
	-229000.	-281000.	40.000	
	8.8725	13.8806	2.9886	-3.3527
	7.1576	-11.0399	2.6076	-2.
SO2,aq		SO2(0)		
SO2(aq)		S(1)O(2)+(0)		
ref:3		13.Jul.87		
	-71980.	-77194.	38.700	
	6.9502	9.1890	2.1383	-3.1589
	31.2101	6.4578	-2.461	0.
SO3-2		SO3(-2)		
SO3-2		S(1)O(3)-(2)		
ref:21		11.Sep.97		
	-116300.	-151900.	-7.000	
	2.4632	-1.7691	6.4494	-2.7058
	-2.7967	-16.7843	3.3210	-2.
SO4-2		SO4(-2)		
SO4-2		S(1)O(4)-(2)		
ref:2		3.Jun.87		
	-177930.	-217400.	4.500	
	8.3014	-1.9846	-6.2122	-2.6970
	1.6400	-17.9980	3.1463	-2.
SCN-		SCN(-1)		
SCN-		S(1)C(1)N(1)-(1)		
ref:21		11.Sep.97		
	22160.	18270.	34.500	
	7.0244	9.3687	2.0708	-3.1662
	10.7414	-4.9900	1.1073	-1.
SbO2-		SbO2(-1)		
SbO2-		Sb(1)O(2)-(1)		
ref:21		11.Sep.97		
	-82400.	-110000.	-17.100	
	-0.5113	-9.0269	9.2893	-2.4057
	-2.2802	-12.0177	1.8885	-1.
Sc+3		Sc(+3)		
Sc+3		Sc(1)+(3)		
ref:21		5.Jan.98		
	-140200.	-146800.	-61.000	
	-2.1109	-12.9294	10.8170	-2.2444
	21.1134	-5.8456	2.5003	3.
ScOH+2		ScOH(+2)		
ScOH+2		Sc(1)O(1)H(1)+(2)		
ref:21		11.Dec.97		
	-19100.	-200700.	-15.700	
	-0.4111	-8.7778	9.1827	-2.4160
	6.7806	-6.9659	1.2944	2.
ScO+		ScO(+)		
ScO+		Sc(1)O(1)+(1)		
ref:21		11.Dec.97		
	-183600.	-189800.	-3.600	
	0.1089	-7.5069	8.6817	-2.4686
	-21.4773	-14.5843	0.6063	1.
ScO2-		ScO2(-)		
ScO2-		Sc(1)O(2)-(1)		
ref:21		11.Dec.97		
	-218100.	-234100.	19.200	
	2.9831	-0.4979	5.9465	-2.7583
	-24.9439	-18.1287	1.3369	-1.
Sc(Ae)+2		ScCH3COO+2		
Sc(Ae)+2		Sc(1)C(2)H(3)O(2)+(2)		
ref:33		10.Sep.92		
	-233010.	-268100.	-42.400	
	2.7175	-1.1437	6.1937	-2.7316
	60.3195	10.3398	1.7013	2.
Sc(Ae)2+		Sc(CH3COO)2+		
Sc(Ae)2+		Sc(1)C(4)H(6)O(4)+(1)		
ref:33		10.Sep.92		

	-324640.	-389320.	-27.500	
	9.2794	14.8737	-0.0899	-3.3938
	106.5183	28.7370	0.9706	1.
Sc(Ae)3,aq		Sc(CH3COO)3		
Sc(Ae)3,aq		Sc(1)C(6)H(9)O(6)		
ref:33		10.Sep.92		
	-415370.	-511840.	-20.000	
	16.5277	32.5748	-7.0539	-4.1255
	143.9346	44.9460	-0.0300	0.
SeO3-2		SeO3(-2)		
SeO3-2		Se(1)O(3)-(2)		
ref:21		11.Sep.97		
	-88400.	-121700.	3.100	
	3.6149	1.0478	5.3323	-2.8222
	-6.517	-15.5417	3.1658	-2.
SeO4-2		SeO4(-2)		
SeO4-2		Se(1)O(4)-(2)		
ref:21		11.Sep.97		
	-105500.	-143200.	12.900	
	5.7548	6.2678	3.2916	-3.0380
	1.3971	-14.3602	3.0192	-2.
SeCN-		SeCN(-1)		
SeCN-		Se(1)C(1)N(1)-(1)		
ref:21		11.Sep.97		
	19500.	18530.	46.600	
	8.8752	13.8922	2.829	-3.3532
	13.0877	-3.5845	9.230	-1.
SiF6-2		SiF6(-2)		
SiF6-2		Si(1)F(6)-(2)		
ref:2		16.Feb.88		
	-525700.	-571000.	29.200	
	8.5311	13.0492	6.211	-3.3185
	4.0970	-12.6289	2.7716	-2.
SiO2,aq		SiO2(0)		
SiO2(aq)		Si(1)O(2)+(0)		
ref:3		13.Jan.89		
	-199190.	-209775.	18.000	
	1.9000	1.7000	20.0000	-2.7000
	29.1000	-51.2000	1.291	0.
Sm+2		Sm(+2)		
Sm+2		Sm(1)+(2)		
ref:21		10.Dec.97		
	-123000.	-120500.	-6.200	
	-0.0353	-7.8592	8.8194	-2.4540
	10.5009	-5.2141	1.1512	2.
Sm+3		Sm(+3)		
Sm+3		Sm(1)+(3)		
ref:2		13.Jul.87		
	-159100.	-165200.	-50.700	
	-3.2065	-15.6108	11.8857	-2.1337
	1.9385	-11.8548	2.2955	3.
SmOH+2		SmOH+2		
SmOH+2		Sm(1)O(1)H(1)+(2)		
ref:25		1.Dec.95		
	-204900.	-214600.	-4.9	
	2.7076	-1.1676	6.2027	-2.7306
	-1.9523	-9.4714	1.1289	2.
SmO+		SmO+		
SmO+		Sm(1)O(1)+(1)		
ref:25		1.Dec.95		
	-193300.	-206500.	11.0	
	2.8115	-0.9157	6.1076	-2.7410
	-33.2571	-17.9657	0.3837	1.
SmO2-		SmO2-		
SmO2-		Sm(1)O(2)-(1)		
ref:25		1.Dec.95		
	-224700.	-233500.	35.9	
	4.9642	4.3993	4.0456	-2.9583
	-47.3681	-25.1156	1.0848	-1.
SmCl+2		SmCl+2		
SmCl+2		Sm(1)Cl(1)+(2)		
ref:25		1.Dec.95		

	-190900.	-201700.	-24.1	
	-0.5006	-8.9988	9.2743	-2.4069
	8.5369	-6.7552	1.4192	2.
SmCl2+		SmCl2+		
SmCl2+		Sm(1)Cl(2)+(1)		
ref:25		1.Dec.95		
	-221800.	-240300.	-7.7	
	2.5888	-1.4617	6.3276	-2.7185
	-9.2005	-10.5032	0.6644	1.
SmCl3,aq		SmCl3(aq)		
SmCl3(aq)		Sm(1)Cl(3)		
ref:25		1.Dec.95		
	-252700.	-281700.	-0.7	
	6.0808	7.0673	2.9692	-3.0711
	-51.8359	-23.0986	-0.0300	0.
SmCl4-		SmCl4-		
SmCl4-		Sm(1)Cl(4)-(1)		
ref:25		1.Dec.95		
	-283500.	-326200.	-4.3	
	10.8148	18.6261	-1.5732	-3.5489
	-97.6675	-44.5415	1.6917	-1.
SmF+2		SmF+2		
SmF+2		Sm(1)F(1)+(2)		
ref:25		1.Dec.95		
	-232400.	-239900.	-15.4	
	-3.1578	-15.4843	11.8177	-2.1388
	9.7908	-5.9197	1.2944	2.
SmF2+		SmF2+		
SmF2+		Sm(1)F(2)+(1)		
ref:25		1.Dec.95		
	-304200.	-322200.	-11.2	
	-2.9212	-14.9061	11.5895	-2.1627
	-2.9564	-8.5080	0.7191	1.
SmF3,aq		SmF3(aq)		
SmF3(aq)		Sm(1)F(3)		
ref:25		1.Dec.95		
	-374800.	-407700.	-21.2	
	-2.6941	-14.3529	11.3752	-2.1856
	-41.8267	-19.6196	-0.0300	0.
SmF4-		SmF4-		
SmF4-		Sm(1)F(4)-(1)		
ref:25		1.Dec.95		
	-444800.	-497700.	-48.5	
	-1.3502	-11.0723	10.0878	-2.3212
	-76.2703	-39.2546	2.3632	-1.
SmHCO3+2		SmHCO3+2		
SmHCO3+2		Sm(1)H(1)C(1)O(3)+(2)		
ref:25		1.Dec.95		
	-301800.	-327900.	-11.9	
	0.3694	-6.8727	8.4365	-2.4948
	29.6747	1.1765	1.2366	2.
SmCO3+		SmCO3+		
SmCO3+		Sm(1)C(1)O(3)+(1)		
ref:25		1.Dec.95		
	-296000.	-308800.	-42.6	
	-1.0455	-10.3293	9.7980	-2.3519
	-24.0049	-17.3342	1.1907	1.
SmH2PO4+2		SmH2PO4+2		
SmH2PO4+2		Sm(1)H(2)P(1)O(4)+(2)		
ref:25		1.Dec.95		
	-432300.	-477800.	-28.3	
	1.3708	-4.4295	7.4801	-2.5958
	33.5675	1.7285	1.4867	2.
SmNO3+2		SmNO3+2		
SmNO3+2		Sm(1)N(1)O(3)+(2)		
ref:25		1.Dec.95		
	-186700.	-221600.	-35.4	
	1.0908	-5.1124	7.7478	-2.5676
	24.2912	-1.8252	1.5897	2.
SmSO4+		SmSO4+		
SmSO4+		Sm(1)S(1)O(4)+(1)		
ref:25		1.Dec.95		

	-342000.	-377800.	-13.3	
	-1.3885	-4.3882	7.4678	-2.5975
	-25.0918	-16.2954	0.7483	1.
Sm(Ae)+2		SmCH3COO+2		
Sm(Ae)+2		Sm(1)C(2)H(3)O(2)+(2)		
ref:33		17 Aug 92		
	-251240.	-284550.	-27.800	
	2.6264	-1.3667	6.2827	-2.7224
	47.8346	6.7190	1.4769	2.
Sm(Ae)2+		Sm(CH3COO)2+		
Sm(Ae)2+		Sm(1)C(4)H(6)O(4)+(1)		
ref:33		17 Aug 92		
	-342190.	-403500.	-7.600	
	9.1590	14.5839	0.0138	-3.3818
	83.3717	21.6725	0.6644	1.
Sm(Ae)3,aq		Sm(CH3COO)3		
Sm(Ae)3,aq		Sm(1)C(6)H(9)O(6)		
ref:33		10 Sep 92		
	-432630.	-523910.	6.000	
	16.5088	32.5307	-7.0412	-4.1237
	109.5807	33.0056	-0.0300	0.
Sm+4		Sm(+4)		
Sm+4		Sm(1)+(4)		
ref:21		10.Dec.97		
	-39900.	-56400.	-101.100	
	-4.2886	-18.2500	12.9154	-2.0244
	40.2222	-3.0752	3.7093	4.
Sn+2		Sn(+2)		
Sn+2		Sn(1)+(2)		
ref:21		11.Sep.97		
	-6570.	-2100.	-4.000	
	.0094	-7.7516	8.7810	-2.4584
	7.4160	-6.1919	1.1216	2.
SnO,aq		SnO		
SnO(aq)		Sn(1)O(1)+(0)		
ref:21		12.Sep.97		
	-53600.	-60100.	14.800	
	2.7401	-1.0905	6.1776	-2.7338
	-13.0981	-9.6344	-0.0300	0.
SnOH+		SnOH(+1)		
SnOH+		Sn(1)O(1)H(1)+(1)		
ref:21		12.Sep.97		
	-58600.	-63800.	19.200	
	2.9337	-6.6181	5.9928	-2.7533
	.9969	-5.6622	.2595	1.
Sr+2		Sr(+2)		
Sr+2		Sr(1)+(2)		
ref:2		1.Jul.87		
	-134760.	-131670.	-7.530	
	.7071	-10.1508	7.0027	-2.3594
	10.7452	-5.0818	1.1363	2.
SrOH+		SrOH(+1)		
SrOH+		Sr(1)O(1)H(1)+(1)		
ref:21		12.Sep.97		
	-173300.	-180000.	14.600	
	2.8620	-7.922	6.0586	-2.7462
	4.7576	-4.5826	.3306	1.
SrCl+		SrCl(+)		
SrCl+		Sr(1)Cl(1)+(1)		
ref:22		17.Sep.97		
	-165800.	-169790.	11.000	
	2.7719	-1.0104	6.1402	-2.7372
	16.6402	-6.6227	.3837	1.
SrF+		SrF(+)		
SrF+		Sr(1)F(1)+(1)		
ref:22		17.Sep.97		
	-202290.	-210670.	-6.200	
	.2336	-7.2081	8.5761	-2.4810
	21.5706	.2471	.6472	1.
SrCO3,aq		SrCO3(0)		
SrCO3(aq)		Sr(1)C(1)O(3)+(0)		
ref:22		17.Sep.97		

	-264860.	-288620.	8.500	
	-3.332	-8.5922	9.1201	-2.4237
	-12.9961	-9.5733	-0.0380	0.
Sr(Ae)+		SrCH3COO+		
Sr(Ae)+		Sr(1)C(2)H(3)O(2)+(1)		
ref:33		17 Aug 92		
	-224510.	-247220.	20.000	
	5.9602	6.7718	3.0884	-3.0588
	53.8109	12.7307	0.2482	1.
Sr(Ae)2,aq		Sr(CH3COO)2		
Sr(Ae)2,aq		Sr(1)C(4)H(6)O(4)		
ref:33		10 Sept 92		
	-313610.	-363740.	42.200	
	13.1015	24.2102	-3.7685	-3.7797
	109.4234	32.9508	-0.0300	0.
Sr(Ala)+		Sr(C3H6NO2)+		
Sr(Ala)+		Sr(1)C(3)H(6)N(1)O(2)+(1)		
ref:27		14.Aug.93		
	-210430.	-247624.	42.695	
	8.7700	13.6312	0.3947	-3.3424
	63.2691	17.1163	-0.0948	1.
Sr(Ala)2,aq		Sr(C3H6NO2)2		
Sr(Ala)2,aq		Sr(1)C(6)H(12)N(2)O(4)		
ref:27		14.Aug.93		
	-285431.	-363933.	89.487	
	19.2869	39.3109	-9.6990	-4.4040
	139.5534	43.4233	-0.0300	0.
Sr(But)+		SrCH3(CH2)2CO2+		
Sr(But)+		Sr(1)C(4)H(7)O(2)+(1)		
ref:27		26.Aug.93		
	-219575.	-257725.	31.245	
	9.9807	16.5869	-0.7651	-3.4646
	93.9079	27.2107	0.0785	1.
Sr(But)2,aq		Sr(CH3CH2CH2CO2)2		
Sr(But)2,aq		Sr(1)C(8)H(17)O(4)		
ref:27		26.Aug.93		
	-303925.	-383903.	68.053	
	21.7228	45.2592	-12.0373	-4.6499
	208.9047	67.5280	-0.0300	0.
Sr(For)+		SrCHO2+		
Sr(For)+		Sr(1)C(1)H(1)O(2)+(1)		
ref:27		26.Aug.93		
	-220518.	-233167.	21.145	
	3.9918	1.9636	4.9821	-2.8601
	22.9691	2.0692	0.2299	1.
Sr(For)2,aq		Sr(CHO2)2		
Sr(For)2,aq		Sr(1)C(2)H(2)O(4)		
ref:27		26.Aug.93		
	-305513.	-335415.	44.736	
	8.9529	14.0817	0.2091	-3.3610
	36.1759	7.4919	-0.0300	0.
Sr(Gly)+		Sr(C2H4NO2)+		
Sr(Gly)+		Sr(1)C(2)H(4)N(1)O(2)+(1)		
ref:27		27.Aug.93		
	-211334.	-239307.	40.925	
	6.3007	7.6019	2.7652	-3.0932
	36.9719	7.8914	-0.0684	1.
Sr(Gly)2,aq		Sr(C2H4NO2)2		
Sr(Gly)2,aq		Sr(1)C(4)H(8)N(2)O(4)		
ref:27		27.Aug.93		
	-287199.	-347420.	85.401	
	14.0476	26.5176	-4.6704	-3.8751
	76.1763	21.3950	-0.0300	0.
Sr(Glyc)+		SrCH3OCO2+		
Sr(Glyc)+		Sr(1)C(2)H(3)O(3)+(1)		
ref:27		30.Aug.93		
	-257717.	-286078.	25.645	
	5.8421	6.4821	3.2052	-3.0469
	53.3420	12.8441	0.1618	1.
Sr(Glyc)2,aq		Sr(CH3OCO2)2		
Sr(Glyc)2,aq		Sr(1)C(4)H(6)O(6)		
ref:27		30.Aug.93		

	-380197.	-441109.	55.125	
	12.9135	23.7481	-3.5809	-3.7606
	110.2025	33.2217	-0.0300	0.
Sr(Lac)+		SrCH3CH2COO2+		
Sr(Lac)+		Sr(1)C(3)H(5)O(3)+(1)		
ref:27		30.Aug.93		
	-258627.	-295697.	29.000	
	8.0696	11.9219	1.0649	-3.2718
	74.8585	20.4811	0.1124	1.
Sr(Lac)2,aq		Sr(CH3CH2COO2)2		
Sr(Lac)2,aq		Sr(1)C(6)H(10)O(6)		
ref:27		30.Aug.93		
	-382098.	-459421.	65.215	
	17.6581	35.3352	-8.1390	-4.2397
	162.6705	51.4582	-0.0300	0.
Sr(Pent)+		SrCH3(CH2)3CO2+		
Sr(Pent)+		Sr(1)C(5)H(9)O(2)+(1)		
ref:27		7.Sep.93		
	-217114.	-263755.	37.745	
	12.1383	21.8541	-2.8337	-3.6823
	130.3108	40.1784	-0.0199	1.
Sr(Pent)2,aq		Sr(CH3CH2CH2CH2CO2)2		
Sr(Pent)2,aq		Sr(1)C(10)H(18)O(4)		
ref:27		7.Sep.93		
	-299073.	-395432.	83.059	
	26.5546	56.5678	-16.4801	-5.1174
	297.9963	98.4940	-0.0300	0.
Sr(Prop)+		SrCH3CH2CO2+		
Sr(Prop)+		Sr(1)C(3)H(5)O(2)+(1)		
ref:27		15.Sep.93		
	-221857.	-252548.	25.945	
	7.9074	11.5294	1.2116	-3.2555
	79.8554	22.0690	0.1589	1.
Sr(Prop)2,aq		Sr(CH3CH2CO2)2		
Sr(Prop)2,aq		Sr(1)C(6)H(10)O(4)		
ref:27		15.Sep.93		
	-308491.	-374036.	55.817	
	17.2820	34.4174	-7.7796	-4.2017
	173.5798	55.2499	-0.0300	0.
Tb+2		Tb(+2)		
Tb+2		Tb(1)+(2)		
ref:21		10.Dec.97		
	-79400.	-76200.	-2.900	
	0.0294	-7.7042	8.7663	-2.4604
	9.6780	-5.3363	1.1000	2.
Tb+3		Tb(+3)		
Tb+3		Tb(1)+(3)		
ref:21		15.Sep.97		
	-159500.	-166900.	-54.000	
	-2.9245	-14.9162	11.5979	-2.1623
	7.1853	-10.3677	2.4007	3.
TbOH+2		TbOH+2		
TbOH+2		Tb(1)O(1)H(1)+(2)		
ref:25		1.Dec.95		
	-205500.	-216700.	-9.3	
	2.6348	-1.3451	6.2720	-2.7233
	1.6043	-8.4529	1.1969	2.
TbO+		TbO+		
TbO+		Tb(1)O(1)+(1)		
ref:25		1.Dec.95		
	-194100.	-209000.	6.3	
	2.7263	-1.1230	6.1877	-2.7325
	-29.4895	-16.8861	0.4555	1.
TbO2-		TbO2-		
TbO2-		Tb(1)O(2)-(1)		
ref:25		1.Dec.95		
	-226200.	-236900.	30.5	
	4.8006	3.9380	4.2071	-2.9417
	-40.1590	-22.8749	1.1676	-1.
TbCl+2		TbCl+2		
TbCl+2		Tb(1)Cl(1)+(2)		
ref:25		1.Dec.95		

	-191200.	-203500.	-28.1	
	-0.2034	-8.2736	8.9909	-2.4369
	18.7845	-3.4094	1.4867	2.
TbCl2+		TbCl2+		
TbCl2+		Tb(1)Cl(2)+(1)		
ref:25		1.Dec.95		
	-222200.	-242400.	-12.6	
	2.9196	-0.6498	5.9993	-2.7520
	10.2630	-3.9752	0.7384	1.
TbCl3,aq		TbCl3(aq)		
TbCl3(aq)		Tb(1)Cl(3)		
ref:25		1.Dec.95		
	-253000.	-284300.	-7.2	
	6.4214	7.8972	2.6479	-3.1054
	-20.0913	-12.0650	-0.0300	0.
TbCl4-		TbCl4-		
TbCl4-		Tb(1)Cl(4)-(1)		
ref:25		1.Dec.95		
	-283800.	-329400.	-13.2	
	11.2397	19.6657	-1.9861	-3.5919
	-47.9171	-27.6788	1.6257	-1.
TbF+2		TbF+2		
TbF+2		Tb(1)F(1)+(2)		
ref:25		1.Dec.95		
	-233200.	-241600.	-17.6	
	-2.8718	-14.7859	11.5438	-2.1677
	19.7334	-2.5739	1.3288	2.
TbF2+		TbF2+		
TbF2+		Tb(1)F(2)+(1)		
ref:25		1.Dec.95		
	-305300.	-324300.	-13.3	
	-2.6056	-14.1355	11.2872	-2.1945
	16.0944	-1.9800	0.7483	1.
TbF3,aq		TbF3(aq)		
TbF3(aq)		Tb(1)F(3)		
ref:25		1.Dec.95		
	-376100.	-410200.	-24.0	
	-2.3534	-13.5233	11.0536	-2.2198
	-10.0820	-5.5860	-0.0300	0.
TbF4-		TbF4-		
TbF4-		Tb(1)F(4)-(1)		
ref:25		1.Dec.95		
	-446300.	-500900.	-53.2	
	-0.9496	-10.0938	9.7025	-2.3616
	-27.1819	-22.3919	2.4254	-1.
TbHCO3+2		TbHCO3+2		
TbHCO3+2		Tb(1)H(1)C(1)O(3)+(2)		
ref:25		1.Dec.95		
	-302100.	-335300.	-34.3	
	0.7596	-5.9232	8.0699	-2.5340
	42.4558	4.5223	1.5790	2.
TbCO3+		TbCO3+		
TbCO3+		Tb(1)C(1)O(3)+(1)		
ref:25		1.Dec.95		
	-296500.	-310400.	-46.7	
	1.5196	-4.0643	7.3324	-2.6109
	-20.3354	-14.8084	0.8002	1.
TbH2PO4+2		TbH2PO4+2		
TbH2PO4+2		Tb(1)H(2)P(1)O(4)+(2)		
ref:25		1.Dec.95		
	-432600.	-479900.	-33.1	
	1.6693	-3.7077	7.2115	-2.6256
	43.8497	5.0743	1.5579	2.
TbNO3+2		TbNO3+2		
TbNO3+2		Tb(1)N(1)O(3)+(2)		
ref:25		1.Dec.95		
	-186700.	-223800.	-41.6	
	1.3991	-4.1601	7.4524	-2.5987
	34.8389	1.5205	1.6897	2.
TbSO4+		TbSO4+		
TbSO4+		Tb(1)S(1)O(4)+(1)		
ref:25		1.Dec.95		

	-342400.	-379600.	-17.1	
	1.5192	-4.0691	7.3424	-2.6107
	-20.2352	-14.8084	0.8111	1.
Tb(Ae)+2		TbCH3COO+2		
Tb(Ae)+2		Tb(1)C(2)H(3)O(2)+(2)		
ref:33		10 Sep 92		
	-251390.	-286400.	-32.500	
	2.9247	-0.6382	5.9963	-2.7525
	52.1779	8.0023	1.5475	2.
Tb(Ae)+2		Tb(CH3COO)+2		
Tb(Ae)+2		Tb(1)C(4)H(6)O(4)+(1)		
ref:33		10 Sep 92		
	-342250.	-405780.	-14.000	
	9.4965	15.4094	-0.3132	-3.4159
	91.4417	24.1763	0.7584	1.
Tb(Ae)+3,aq		Tb(CH3COO)+3		
Tb(Ae)+3,aq		Tb(1)C(6)H(9)O(6)		
ref:33		10 Sep 92		
	-432390.	-526470.	-2.300	
	16.8495	33.3603	-7.3624	-4.1580
	121.7569	37.2376	-0.0300	0.
Tb+4		Tb(+4)		
Tb+4		Tb(1)+(4)		
ref:21		10.Dec.97		
	-88200.	-105900.	-104.200	
	-4.3164	-18.3144	12.9330	-2.0218
	39.9963	-3.2789	3.7484	4.
TcO4-		TcO4(-1)		
TcO4-		Tc(1)O(4)-(1)		
ref:21		30.Apr.97		
	-151100.	-173200.	47.500	
	8.6135	18.7256	-13.3543	-3.5530
	18.9500	-5.9500	.9101	-1.
Th+2		Th(+2)		
Th+2		Th(1)+(2)		
ref:36		16.Jan.98		
	-19700.	-13500.	2.500	
	0.1521	-7.4017	8.6408	-2.4729
	8.2741	-5.5604	1.0176	2.
Th+3		Th(+3)		
Th+3		Th(1)+(3)		
ref:36		16.Jan.98		
	-82200.	-85100.	-43.700	
	-2.7582	-14.5095	11.4370	-2.1791
	5.2821	-10.5307	2.2450	3.
Th+4		Th(+4)		
Th+4		Th(1)+(4)		
ref:21		11.Sep.97		
	-168500.	-183800.	-101.000	
	-4.2886	-18.2500	12.9154	-2.0244
	40.2222	-3.0752	3.7093	4.
Tl+		Tl(+)		
Tl+		Tl(1)+(1)		
ref:21		12.Dec.97		
	-7740.	1280.	30.000	
	4.2884	2.6903	4.6910	-2.8901
	4.6024	-3.8900	0.0974	1.
TlOH,aq		TlOH		
TlOH(aq)		Tl(1)O(1)H(1)+(0)		
ref:21		11.Dec.97		
	-65550.	-66660.	32.500	
	5.1759	4.8579	3.8372	-2.9797
	-10.5781	-8.7585	-0.0300	0.
TlCl,aq		TlCl(0)		
TlCl(aq)		Tl(1)Cl(1)+(0)		
ref:22		5.Jan.98		
	-39815.	-37819.	48.690	
	8.1182	12.0440	1.0092	-3.2768
	-6.2744	-7.2370	-0.0380	0.
TlF,aq		TlF(0)		
TlF(aq)		Tl(1)F(1)+(0)		
ref:22		6.Jan.98		

	-75216.	-77118.	33.180	
	5.4909	5.6287	3.5307	-3.0116
	-3.7726	-6.3675	-0.0380	0.
Tl(Ae),aq		TlCH3COO		
Tl(Ae),aq		Tl(1)C(2)H(3)O(2)		
ref:33		17 Aug 92		
	-95860.	-113350.	55.200	
	11.2955	19.8003	-2.0347	-3.5974
	32.4138	6.1843	-0.0300	0.
Tl(Ae)+2		Tl(CH3COO)+2		
Tl(Ae)+2		Tl(1)C(4)H(6)O(4)-(1)		
ref:33		20 Sept 92		
	-183600.	-230620.	70.300	
	19.3516	39.4721	-9.7684	-4.4107
	76.8519	19.7269	0.5643	-1.
Tl+3		Tl(+3)		
Tl+3		Tl(1)+(3)		
ref:21		12.Dec.97		
	51300.	47000.	-46.000	
	-1.1607	-10.6078	9.9017	-2.3404
	27.0098	-3.0752	2.2752	3.
TlOH+2		TlOH(+2)		
TlOH+2		Tl(1)O(1)H(1)+(2)		
ref:21		11.Dec.97		
	-4500.	-18800.	-23.700	
	2.3952	-1.9341	6.5118	-2.6989
	13.2638	-5.1122	1.4192	2.
TlO+		TlO(+)		
TlO+		Tl(1)O(1)+(1)		
ref:21		11.Dec.97		
	-3200.	-5150.	17.700	
	2.9144	-0.6673	6.0161	-2.7513
	-38.6952	-19.5342	0.2832	1.
TlO2-		TlO2(-)		
TlO2-		Tl(1)O(2)-(1)		
ref:21		11.Dec.97		
	-41600.	-52500.	43.500	
	5.1854	4.8830	3.8238	-2.9808
	-57.6278	-28.3137	0.9699	-1.
TlCl+2		TlCl(+2)		
TlCl+2		Tl(1)Cl(1)+(2)		
ref:22		5.Jan.98		
	9375.	671.	-18.520	
	-0.2904	-8.4877	9.0790	-2.4281
	13.4748	-4.7774	1.3375	2.
Tm+2		Tm(+2)		
Tm+2		Tm(1)+(2)		
ref:21		10.Jul.97		
	-107600.	-105600.	-6.700	
	-0.0464	-7.8863	8.8296	-2.4529
	10.6286	-5.1937	1.1587	2.
Tm+3		Tm(+3)		
Tm+3		Tm(1)+(3)		
ref:2		13.Jul.87		
	-159900.	-168500.	-58.100	
	-3.2967	-15.8312	11.9724	-2.1245
	8.4826	-10.0215	2.4333	3.
TmOH+2		TmOH+2		
TmOH+2		Tm(1)O(1)H(1)+(2)		
ref:25		1.Dec.95		
	-206100.	-219000.	-14.7	
	2.5389	-1.5805	6.3673	-2.7136
	5.9225	-7.2104	1.2776	2.
TmO+		TmO+		
TmO+		Tm(1)O(1)+(1)		
ref:25		1.Dec.95		
	-194900.	-211600.	0.5	
	2.6339	-1.3489	6.2767	-2.7231
	-24.7834	-15.5417	0.5464	1.
TmO2-		TmO2-		
TmO2-		Tm(1)O(2)-(1)		
ref:25		1.Dec.95		

	-228700.	-241400.	23.9	
	4.6152	3.4877	4.3792	-2.9231
	-31.2152	-20.0842	1.2669	-1.
TmC1+2		TmC1+2		
TmC1+2		Tm(1)C1(1)+(2)		
ref:25		1.Dec.95		
	-191600.	-205300.	-33.0	
	-0.6062	-9.2588	9.3810	-2.3961
	21.6825	-2.6303	1.5579	2.
TmC12+		TmC12+		
TmC12+		Tm(1)C1(2)+(1)		
ref:25		1.Dec.95		
	-222600.	-244600.	-18.7	
	2.4761	-1.7342	6.4282	-2.7072
	15.5119	-2.4550	0.8334	1.
TmC13, aq		TmC13(aq)		
TmC13(aq)		Tm(1)C1(3)		
ref:25		1.Dec.95		
	-253400.	-287000.	-15.3	
	5.8915	6.6023	3.1580	-3.0518
	-12.6987	-9.4955	-0.0300	0.
TmC14-		TmC14-		
TmC14-		Tm(1)C1(4)-(1)		
ref:25		1.Dec.95		
	-284200.	-333100.	-24.2	
	10.7064	18.3610	-1.4676	-3.5379
	-35.0588	-23.7519	1.9951	-1.
TmF+2		TmF+2		
TmF+2		Tm(1)F(1)+(2)		
ref:25		1.Dec.95		
	-233800.	-243000.	-20.3	
	-3.2867	-15.8018	11.9486	-2.1257
	22.3012	-1.7948	1.3642	2.
TmF2+		TmF2+		
TmF2+		Tm(1)F(2)+(1)		
ref:25		1.Dec.95		
	-306000.	-325800.	-16.0	
	-3.0673	-15.2659	11.7377	-2.1478
	20.8477	-0.4598	0.7895	1.
TmF3, aq		TmF3(aq)		
TmF3(aq)		Tm(1)F(3)		
ref:25		1.Dec.95		
	-376900.	-412000.	-27.5	
	-2.8834	-14.8179	11.5640	-2.1663
	-2.6893	-6.0166	-0.0300	0.
TmF4-		TmF4-		
TmF4-		Tm(1)F(4)-(1)		
ref:25		1.Dec.95		
	-447200.	-503600.	-59.0	
	-1.5097	-11.4594	10.2349	-2.3052
	-15.0561	-18.4650	2.5152	-1.
TmHCO3+2		TmHCO3+2		
TmHCO3+2		Tm(1)H(1)C(1)O(3)+(2)		
ref:25		1.Dec.95		
	-302600.	-332200.	-22.4	
	0.2724	-7.1116	8.5345	-2.4849
	43.0537	5.3015	1.4006	2.
TmCO3+		TmCO3+		
TmCO3+		Tm(1)C(1)O(3)+(1)		
ref:25		1.Dec.95		
	-297300.	-312700.	-51.6	
	-1.0611	-10.3695	9.8181	-2.3502
	-17.4772	-15.5009	1.3268	1.
TmH2PO4+2		TmH2PO4+2		
TmH2PO4+2		Tm(1)H(2)P(1)O(4)+(2)		
ref:25		1.Dec.95		
	-433100.	-482200.	-39.1	
	1.2716	-4.6688	7.5680	-2.5859
	46.8875	5.8534	1.6444	2.
TmNO3+2		TmNO3+2		
TmNO3+2		Tm(1)N(1)O(3)+(2)		
ref:25		1.Dec.95		

	-186700.	-226000.	-49.3	
	1.0128	-5.3011	7.8182	-2.5598
	38.1891	2.2996	1.8100	2.
TmSO4+		TmSO4+		
TmSO4+		Tm(1)S(1)O(4)+(1)		
ref:25		1.Dec.95		
	-342700.	-381120.	-21.2	
	-1.3743	-4.4228	7.4814	-2.5961
	-18.7119	-14.0000	0.8683	1.
Tm(Ae)+2		TmCH3COO+2		
Tm(Ae)+2		Tm(1)C(2)H(3)O(2)+(2)		
ref:33		11 Sept 92		
	-251770.	-288500.	-38.300	
	2.5268	-1.6112	6.3814	-2.7123
	61.1431	10.8439	1.6332	2.
Tm(Ae)2+		Tm(CH3COO)2+		
Tm(Ae)2+		Tm(1)C(4)H(6)O(4)+(1)		
ref:33		11 Sept 92		
	-342620.	-408490.	-21.900	
	9.0621	14.3448	0.1131	-3.3719
	108.5166	29.7206	0.8803	1.
Tm(Ae)3, aq		Tm(CH3COO)3		
Tm(Ae)3, aq		Tm(1)C(6)H(9)O(6)		
ref:33		11 Sept 92		
	-432750.	-529900.	-12.700	
	16.3195	32.0657	-6.8520	-4.1045
	148.7181	46.6086	-0.0300	0.
Tm+4		Tm(+4)		
Tm+4		Tm(1)+(4)		
ref:21		10.Dec.97		
	-30100.	-48300.	-105.600	
	-4.3349	-18.3596	12.9511	-2.0199
	40.0047	-3.3604	3.7747	4.
U+2		U(+2)		
U+2		U(1)+(2)		
ref:36		16.Jan.98		
	-48800.	-42900.	0.600	
	0.1087	-7.5126	8.6950	-2.4683
	8.8177	-5.4789	1.0512	2.
U+3		U(+3)		
U+3		U(1)+(3)		
ref:21		11.Sep.97		
	-113550.	-116900.	-46.100	
	-2.8438	-14.7220	11.5280	-2.1703
	5.7945	-10.4492	2.2752	3.
UOH+2		UOH(+2)		
UOH+2		U(1)O(1)H(1)+(2)		
ref:37		20.Jan.98		
	-161800.	-167700.	1.200	
	2.8136	-0.9087	6.1004	-2.7413
	-6.8963	-10.8973	1.0376	2.
UO+		UO(+)		
UO+		U(1)O(1)+(1)		
ref:37		20.Jan.98		
	-152900.	-153900.	17.500	
	2.9158	-0.6648	6.0168	-2.7514
	-38.5403	-19.4935	0.2873	1.
UO2-		UO2(-)		
UO2-		U(1)O(2)-(1)		
ref:37		20.Jan.98		
	-183700.	-193600.	43.300	
	5.1728	4.8473	3.8491	-2.9793
	-57.3622	-28.2322	0.9733	-1.
U(Ae)+2		UCH3COO+2		
U(Ae)+2		U(1)C(2)H(3)O(2)+(2)		
ref:33		17 Aug 92		
	-205480.	-235460.	-21.300	
	4.8668	4.1031	4.1308	-2.9485
	103.6659	26.4274	1.3823	2.
U(Ae)2+		U(CH3COO)2+		
U(Ae)2+		U(1)C(4)H(6)O(4)+(1)		
ref:33		17 Aug 92		

	-296950.	-354230.	1.300	
	11.6455	20.6554	-2.3718	-3.6328
	192.7896	60.1258	0.5324	1.
U(Ae)3, aq		U(CH3COO)3		
U(Ae)3, aq		U(1)C(6)H(9)O(6)		
ref:33		17 Aug 92		
	-387360.	-473790.	17.700	
	19.3290	39.4128	-9.7374	-4.4082
	296.5739	97.9996	-0.0300	0.
U(But)+2		UCH3(CH2)2CO2+2		
U(But)+2		U(1)C(4)H(7)O(2)+(2)		
ref:27		14.Aug.93		
	-202157.	-247566.	-10.070	
	7.1485	9.6738	1.9467	-3.1788
	95.4995	24.1323	1.2126	2.
U(But)2+		U(CH3CH2CH2CO2)2+		
U(But)2+		U(1)C(8)H(14)O(4)+(1)		
ref:27		14.Aug.93		
	-289537.	-376461.	27.782	
	18.1945	36.6443	-8.6523	-4.2938
	194.4029	61.9728	0.1308	1.
U(For)+2		UCHO2+2		
U(For)+2		U(1)C(1)H(1)O(2)+(2)		
ref:27		26.Aug.93		
	-201409.	-221316.	-20.170	
	1.1596	-4.9432	7.6782	-2.5745
	24.5615	-1.0091	1.3642	2.
U(For)2+		U(CHO2)2+		
U(For)2+		U(1)C(2)H(2)O(4)+(1)		
ref:27		26.Aug.93		
	-288791.	-325806.	3.911	
	5.5468	5.7644	3.4796	-3.0172
	25.0070	1.9367	0.4925	1.
U(Pent)+2		UCH3(CH2)3CO2+2		
U(Pent)+2		U(1)C(5)H(9)O(2)+(2)		
ref:27		31.Aug.93		
	-200147.	-254046.	-3.570	
	9.3061	14.9410	-0.1209	-3.3966
	131.9035	37.1000	1.1144	2.
U(Prop)+2		UCH3CH2CO2+2		
U(Prop)+2		U(1)C(3)H(5)O(2)+(2)		
ref:27		15.Sep.93		
	-204317.	-242266.	-15.370	
	5.0757	4.6150	3.9293	-2.9697
	81.4601	18.9905	1.2944	2.
U(Prop)2+		U(CH3CH2CO2)2+		
U(Prop)2+		U(1)C(6)H(10)O(4)+(1)		
ref:27		15.Sep.93		
	-293529.	-366108.	15.255	
	13.8181	25.9582	-4.4517	-3.8520
	160.8356	49.6947	0.3215	1.
U+4		U(+4)		
U+4		U(1)+(4)		
ref:21		11.Sep.97		
	-126650.	-141300.	-99.600	
	-4.2836	-18.2319	12.8955	-2.0252
	40.2193	-2.9938	3.6835	4.
UOH+3		UOH(+3)		
UOH+3		U(1)O(1)H(3)+(3)		
ref:37		18.Jan.98		
	-182600.	-198400.	-47.800	
	2.1611	-2.5064	6.7390	-2.6753
	37.5484	0.4894	2.3058	3.
UO+2		UO(+2)		
UO+2		U(1)O(1)+(2)		
ref:37		20.Jan.98		
	-180600.	-192100.	-33.400	
	2.2403	-2.3083	6.6513	-2.6835
	7.3127	-7.6585	1.5684	2.
UO2, aq		UO2(O)		
UO2(aq)		U(1)O(2)+(0)		
ref:37		20.Jan.98		

	-233800.	-259700.	-26.000	
	2.4664	-1.7587	6.4404	-2.7062
	17.0840	0.8561	-0.0300	0.
UO2+		UO2(+1)		
UO2+		U(1)O(2)+(1)		
ref:21		11.Sep.97		
	-229690.	-245000.	-6.000	
	3.3767	.4614	5.5725	-2.7980
	-.9002	-7.5363	.6388	1.
UO2OH,aq		UO2OH(0)		
UO2OH(aq)		U(1)O(3)H(1)+(0)		
ref:37		20.Jan.98		
	-261600.	-295900.	-13.900	
	3.7390	1.3467	5.2242	-2.8346
	20.4246	2.0172	-0.0300	0.
UO3-		UO3(-1)		
UO3-		U(1)O(3)-(1)		
ref:37		20.Jan.98		
	-236600.	-273600.	-19.500	
	4.1388	2.3272	4.8294	-2.8751
	27.8024	-1.6697	1.9223	-1.
UO2+2		UO2(+2)		
UO2+2		U(1)O(2)+(2)		
ref:21		11.Sep.97		
	-227680.	-243550.	-23.500	
	3.0256	-4.1084	15.3326	-2.6091
	21.0700	.4300	1.4099	2.
UO2OH+		UO2OH(+)		
UO2OH+		U(1)O(3)H(1)+(1)		
ref:37		20.Jan.98		
	-277250.	-301500.	4.100	
	4.7640	3.8529	4.2318	-2.9382
	13.2240	-2.1586	0.4925	1.
UO3,aq		UO3(0)		
UO3(aq)		U(1)O(3)+(0)		
ref:37		20.Jan.98		
	-270300.	-299600.	-12.900	
	3.7801	1.4512	5.1736	-2.8389
	5.4213	-3.1975	-0.0300	0.
UO4-2		UO4(-2)		
UO4-2		U(1)O(4)-(2)		
ref:37		20.Jan.98		
	-296000.	-346200.	-27.100	
	5.0964	4.6606	3.9223	-2.9716
	13.5130	-12.0788	3.6219	-2.
UO2(Ae)+		UO2CH3COO+		
UO2(Ae)+		U(1)C(2)H(3)O(4)+(1)		
ref:33		17 Aug 92		
	-320100.	-363520.	-1.700	
	9.5990	15.6588	-0.4099	-3.4262
	83.5297	22.0118	0.5756	1.
UO2(Ae)2,aq		UO2(CH3COO)2		
UO2(Ae)2,aq		U(1)C(4)H(6)O(6)		
ref:33		17 Aug 92		
	-411840.	-484700.	13.700	
	17.0325	33.8044	-7.5309	-4.1764
	161.5228	51.0592	-0.0300	0.
UO2(Ae)3-		UO2(CH3COO)3-		
UO2(Ae)3-		U(1)C(6)H(9)O(8)-(1)		
ref:33		17 Aug 92		
	-502400.	-607960.	18.300	
	26.0101	55.7294	-16.1571	-5.0828
	275.3211	86.1855	1.3526	-1.
V+2		V(+2)		
V+2		V(1)+(2)		
ref:21		5.Jan.98		
	-52000.	-54000.	-31.000	
	-1.6188	-11.7289	10.3473	-2.2940
	13.8458	-5.2548	1.5269	2.
VOH+		VOH(+)		
VOH+		V(1)O(1)H(1)+(1)		
ref:21		11.Dec.97		

	-99800.	-114100.	-16.400	
	-1.5223	-11.4950	10.2591	-2.3037
	29.7729	2.6079	0.8002	1.
V+3		V(+3)		
V+3		V(1)+(3)		
ref:21		5.Jan.98		
	-57900.	-62000.	-55.000	
	-1.7304	-11.9981	10.4465	-2.2829
	23.4604	-4.7456	2.4115	3.
VOH+2		VOH(+2)		
VOH+2		V(1)O(1)H(1)+(2)		
ref:21		11.Dec.97		
	-111500.	-119400.	-10.600	
	-0.1213	-8.0707	8.9058	-2.4453
	2.7011	-8.1474	1.2206	2.
VO+		VO(+)		
VO+		V(1)O(1)+(1)		
ref:21		11.Dec.97		
	-106100.	-109400.	4.900	
	-3.1084	-15.3641	11.7711	-2.1437
	-28.3270	-16.5602	0.4799	1.
VOOH+		VOOH(+)		
VOOH+		V(1)O(2)H(1)+(1)		
ref:21		11.Dec.97		
	-155700.	-177600.	-17.700	
	-1.5696	-11.6082	10.2984	-2.2990
	30.8543	2.9134	0.8222	1.
VO+2		VO(+2)		
VO+2		V(1)O(1)+(2)		
ref:21		5.Jan.98		
	-106700.	-116300.	-32.000	
	-1.2971	-10.9431	10.0377	-2.3265
	38.8261	3.3616	1.5475	2.
VO2+		VO2(+)		
VO2+		V(1)O(2)+(1)		
ref:21		5.Jan.98		
	-140300.	-155300.	-10.100	
	3.2606	0.1790	5.6817	-2.7863
	4.5892	-5.8252	.7003	1.
VO4-3		VO4(-3)		
VO4-3		V(1)O(4)-(3)		
ref:21		5.Jan.98		
	-214900.	-270600.	-35.200	
	1.2579	-4.7015	7.5797	-2.5845
	77.6555	4.7060	5.3424	-3.
WO4-2		WO4(-2)		
WO4-2		W(1)O(4)-(2)		
ref:21		15.Sep.97		
	-218500.	-236500.	9.700	
	7.2074	8.7933	4.4891	-3.1424
	8.3308	-12.0991	3.0657	-2.
Xe,aq		Xe(0)		
Xe(aq)		Xe(1)+(0)		
ref:3		13.Jul.87		
	3225.	-4510.	14.620	
	7.5196	10.5793	1.5919	-3.2163
	42.1036	10.1652	-2.2214	0.
Y+3		Y(+3)		
Y+3		Y(1)+(3)		
ref:21		5.Jan.98		
	-163800.	-170900.	-60.000	
	-2.0463	-12.7741	10.7611	-2.2508
	21.5366	-5.6622	2.4890	3.
YOH+2		YOH(+2)		
YOH+2		Y(1)O(1)H(1)+(2)		
ref:21		12.Dec.97		
	-210000.	-220900.	-17.200	
	2.4985	-1.6806	6.4099	-2.7094
	8.0136	-6.6196	1.3201	2.
YO+		YO(+)		
YO+		Y(1)O(1)+(1)		
ref:21		12.Dec.97		

	-198100.	-204600.	-2.200	
	2.5778	-1.4890	6.3393	-2.7173
	-22.6289	-14.9102	0.5831	1.
YO2-		YO2(-1)		
YO2-		Y(1)O(2)-(1)		
ref:21		12.Dec.97		
	-227400.	-243600.	20.800	
	4.5219	3.2579	4.4729	-2.9136
	-27.0256	-18.7805	1.3145	-1.
Y(Ae)+2		YCH3COO+2		
Y(Ae)+2		Y(1)C(2)H(3)O(2)+(2)		
ref:33		10 Sep 92		
	-255670.	-291130.	-41.000	
	2.8774	-0.7576	6.0516	-2.7476
	59.7114	10.2023	1.6782	2.
Y(Ae)2+		YCH3COO)2+		
Y(Ae)2+		Y(1)C(4)H(6)O(4)+(1)		
ref:33		10 Sep 92		
	-346520.	-411420.	-25.600	
	9.4572	15.3127	-0.2736	-3.4119
	105.4982	28.4687	0.9437	1.
Y(Ae)3,aq		YCH3COO)3		
Y(Ae)3,aq		Y(1)C(6)H(9)O(6)		
ref:33		10 Sep 92		
	-436650.	-533170.	-17.500	
	16.7359	33.0838	-7.2550	-4.1466
	142.6300	44.4926	-0.0300	0.
Yb+2		Yb(+2)		
Yb+2		Yb(1)+(2)		
ref:21		10.Dec.97		
	-128500.	-126800.	-11.200	
	-0.1323	-8.0980	8.9173	-2.4441
	11.7413	-5.0308	1.2285	2.
Yb+3		Yb(+3)		
Yb+3		Yb(1)+(3)		
ref:2		13.Jul.87		
	-153000.	-160300.	-56.900	
	-3.4983	-16.3233	12.1658	-2.1042
	7.3533	-10.4493	2.4443	3.
YbOH+2		YbOH+2		
YbOH+2		Yb(1)O(1)H(1)+(2)		
ref:25		1.Dec.95		
	-199300.	-210700.	-13.1	
	2.5716	-1.5040	6.3438	-2.7167
	4.6390	-7.5770	1.2528	2.
YbO+		YbO+		
YbO+		Yb(1)O(1)+(1)		
ref:25		1.Dec.95		
	-188200.	-203400.	2.2	
	2.6656	-1.2698	6.2422	-2.7264
	-26.1516	-15.9287	0.5188	1.
YbO2-		YbO2-		
YbO2-		Yb(1)O(2)-(1)		
ref:25		1.Dec.95		
	-221800.	-232900.	25.8	
	4.6743	3.6324	4.3212	-2.9291
	-33.8142	-20.8990	1.2392	-1.
YbCl+2		YbCl+2		
YbCl+2		Yb(1)Cl(1)+(2)		
ref:25		1.Dec.95		
	-184600.	-196900.	-31.5	
	-0.8419	-9.8331	9.6042	-2.3724
	18.7222	-3.5928	1.5372	2.
YbCl2+		YbCl2+		
YbCl2+		Yb(1)Cl(2)+(1)		
ref:25		1.Dec.95		
	-215400.	-236000.	-16.9	
	2.2137	-2.3759	6.6832	-2.6807
	9.9034	-4.3329	0.8111	1.
YbCl3,aq		YbCl3(aq)		
YbCl3(aq)		Yb(1)Cl(3)		
ref:25		1.Dec.95		

	-246100.	-278100.	-12.9	
	5.6076	5.9109	3.4260	-3.0233
	-12.8306	-12.6696	-0.0300	0.
YbCl4-		YbCl4-		
YbCl4-		Yb(1)Cl(4)-(1)		
ref:25		1.Dec.95		
	-276900.	-323800.	-21.0	
	10.3734	17.5486	-1.1496	-3.5044
	-49.4701	-28.6028	1.9457	-1.
YbF+2		YbF+2		
YbF+2		Yb(1)F(1)+(2)		
ref:25		1.Dec.95		
	-226900.	-234900.	-19.5	
	-3.5184	-16.3679	12.1716	-2.1023
	19.4495	-2.7572	1.3552	2.
YbF2+		YbF2+		
YbF2+		Yb(1)F(2)+(1)		
ref:25		1.Dec.95		
	-299100.	-317700.	-15.2	
	-3.3257	-15.8931	11.9766	-2.1219
	15.3476	-2.3377	0.7790	1.
YbF3, aq		YbF3(aq)		
YbF3(aq)		Yb(1)F(3)		
ref:25		1.Dec.95		
	-370100.	-403900.	-26.5	
	-3.1673	-15.5093	11.8317	-2.1377
	-11.8214	-9.1906	-0.0300	0.
YbF4-		YbF4-		
YbF4-		Yb(1)F(4)-(1)		
ref:25		1.Dec.95		
	-440400.	-495300.	-57.3	
	-1.8339	-12.2511	10.5461	-2.2724
	-29.2268	-23.3159	2.4920	-1.
YbHCO3+2		YbHCO3+2		
YbHCO3+2		Yb(1)H(1)C(1)O(3)+(2)		
ref:25		1.Dec.95		
	-295792.	-323900.	-20.7	
	0.0344	-7.6923	8.7622	-2.4609
	40.0319	4.3390	1.3732	2.
YbCO3+		YbCO3+		
YbCO3+		Yb(1)C(1)O(3)+(1)		
ref:25		1.Dec.95		
	-290500.	-305400.	-50.2	
	-1.0611	-10.3695	9.8181	-2.3502
	-17.4772	-15.5009	1.3268	1.
YbH2PO4+2		YbH2PO4+2		
YbH2PO4+2		Yb(1)H(2)P(1)O(4)+(2)		
ref:25		1.Dec.95		
	-426300.	-473900.	-37.31	
	1.0354	-5.2480	7.8010	-2.5619
	43.9143	4.8910	1.6222	2.
YbNO3+2		YbNO3+2		
YbNO3+2		Yb(1)N(1)O(3)+(2)		
ref:25		1.Dec.95		
	-179800.	-217600.	-47.0	
	0.7716	-5.8922	8.0541	-2.5353
	35.0774	1.3371	1.7728	2.
YbSO4+		YbSO4+		
YbSO4+		Yb(1)S(1)O(4)+(1)		
ref:25		1.Dec.95		
	-335800.	-37200.	-20.1	
	-1.2745	-4.6665	7.5772	-2.5860
	-20.0516	-14.8899	0.8565	1.
Yb(Ae)+2		YbCH3COO+2		
Yb(Ae)+2		Yb(1)C(2)H(3)O(2)+(2)		
ref:33		17 Aug 92		
	-244760.	-280040.	-36.600	
	2.2906	-2.1905	6.6154	-2.6883
	58.1718	9.8814	1.6113	2.
Yb(Ae)2+		Yb(CH3COO)2+		
Yb(Ae)2+		Yb(1)C(4)H(6)O(4)+(1)		
ref:33		17 Aug 92		

	-335520.	-399750.	-19.600	
	8.7953	13.6959	0.3628	-3.3451
	102.7868	27.8427	0.8449	1.
Yb(Ae)3, aq		Yb(CH3COO)3		
Yb(Ae)3, aq		Yb(1)C(6)H(9)O(6)		
ref:33		10 Sep 92		
	-425590.	-520890.	-9.700	
	16.0356	31.3743	-6.5843	-4.0759
	139.5859	43.4346	-0.0300	0.
Yb(But)+2		YbCH3(CH2)2CO2+2		
Yb(But)+2		Yb(1)C(4)H(7)O(2)+(2)		
ref:27		26 Aug 93		
	-241293.	-291999.	-25.387	
	6.3100	7.6274	2.7480	-3.0942
	98.2368	24.3614	1.4382	2.
Yb(But)2+		Yb(CH3CH2CH2CO2)2+		
Yb(But)2+		Yb(1)C(8)H(14)O(4)+(1)		
ref:27		26 Aug 93		
	-328536.	-422417.	6.898	
	17.2830	34.4210	-7.7836	-4.2019
	198.6263	62.4199	0.4496	1.
Yb(For)+2		YbCHO2+2		
Yb(For)+2		Yb(1)C(1)H(1)O(2)+(2)		
ref:27		26 Aug 93		
	-240545.	-265749.	-35.487	
	0.3247	-6.9803	8.4752	-2.4903
	27.3976	-0.7799	1.6005	2.
Yb(For)2+		Yb(CHO2)2+		
Yb(For)2+		Yb(1)C(2)H(2)O(4)+(1)		
ref:27		26 Aug 93		
	-327027.	-370998.	-16.973	
	4.6352	3.5352	4.3629	-2.9250
	29.2285	2.3838	0.8111	1.
Yb(Pent)+2		YbCH3(CH2)3CO2+2		
Yb(Pent)+2		Yb(1)C(5)H(9)O(2)+(2)		
ref:27		7 Sep 93		
	-239283.	-298479.	-18.887	
	8.4697	12.9016	0.6745	-3.3123
	134.7001	37.3291	1.3464	2.
Yb(Pent)2+		Yb(CH3CH2CH2CH2CO2)2+		
Yb(Pent)2+		Yb(1)C(10)H(18)O(4)+(1)		
ref:27		7 Sep 93		
	-324503.	-434659.	22.260	
	21.8358	45.5373	-12.1516	-4.6614
	285.5660	93.3859	0.2160	1.
Yb(Prop)+2		YbCH3CH2CO2+2		
Yb(Prop)+2		Yb(1)C(3)H(5)O(2)+(2)		
ref:27		15 Sep 93		
	-243276.	-286522.	-30.687	
	4.2395	2.5689	4.7426	-2.8851
	84.2616	19.2197	1.5269	2.
Yb(Prop)2+		Yb(CH3CH2CO2)2+		
Yb(Prop)2+		Yb(1)C(6)H(10)O(4)+(1)		
ref:27		15 Sep 93		
	-332543.	-412078.	-5.628	
	12.9061	23.7300	-3.5735	-3.7599
	165.0447	50.1419	0.6388	1.
Yb+4		Yb(+4)		
Yb+4		Yb(1)+(4)		
ref:21		10 Dec 97		
	3800.	-13600.	-106.600	
	-4.3441	-18.3854	12.9683	-2.0188
	39.9511	-3.4215	3.7880	4.
Zn+2		Zn(+2)		
Zn+2		Zn(1)+(2)		
ref:21		15 Sep 97		
	-35200.	-36660.	-26.200	
	-1.0676	-8.9290	6.1282	-2.4098
	18.7400	-5.3700	1.4574	2.
ZnOH+		ZnOH(+1)		
ZnOH+		Zn(1)O(1)H(1)+(1)		
ref:21		16 Sep 97		

	-81190.	-86990.	15.000	
	1.1499	-4.9677	7.6896	-2.5735
	15.0306	-9.9975	.3260	1.
ZnO, aq		ZnO(0)		
ZnO(aq)		Zn(1)O(1)+(0)		
ref:21		16 Sep 97		
	-67420.	-78290.	-2.000	
	-1.2418	-10.8072	9.9819	-2.3321
	.0295	-5.0715	-0.0300	0.
ZnO2-2		ZnO2(-2)		
ZnO2-2		Zn(1)O(2)-(2)		
ref:21		17 Sep 97		
	-93290.	-132100.	-40.000	
	-5.5559	-9.1347	9.3301	-2.4013
	32.5837	-6.0900	3.6216	-2.
ZnCl+		ZnCl(+)		
ZnCl+		Zn(1)Cl(1)+(1)		
ref:22		17 Sep 97		
	-66850.	-66240.	23.000	
	1.6583	-3.7293	7.2088	-2.6248
	19.6947	1.0191	.2025	1.
ZnCl2, aq		ZnCl2(0)		
ZnCl2(aq)		Zn(1)Cl(1)+(0)		
ref:22		17 Sep 97		
	-98300.	-109080.	27.030	
	5.1486	4.7929	3.8592	-2.9771
	26.1528	4.0338	-0.0380	0.
ZnCl3-		ZnCl3(-)		
ZnCl3-		Zn(1)Cl(3)-(1)		
ref:22		17 Sep 97		
	-129310.	-151060.	25.000	
	9.5636	15.5732	-3.7779	-3.4227
	42.2912	5.5147	1.2513	-1.
ZnCl4-2		ZnCl4(-2)		
ZnCl4-2		Zn(1)Cl(4)-(2)		
ref:5		JAN. 91		
	-161890.	-195200.	36.000	
	14.6628	28.0213	-5.2636	-3.9374
	56.1061	5.7856	2.6662	-2.
ZnF+		ZnF(+)		
ZnF+		Zn(1)F(1)+(1)		
ref:22		6 Jan. 98		
	-104109.	-116131.	-21.810	
	-0.7395	-9.5843	9.5101	-2.3827
	28.5991	1.9435	0.8803	1.
Zn(Ae)+		Zn(CH3COOH)(+)		
Zn(Ae)+		Zn(1)C(2)H(3)O(2)+(1)		
ref:22		17 Sep 97		
	-125660.	-155120.	9.400	
	4.8484	4.0600	4.1473	-2.9468
	60.4626	14.5244	.4100	1.
Zn(Ae)2, aq		Zn(CH3COOH)2(0)		
Zn(Ae)2(aq)		Zn(1)C(4)H(6)O(4)+(0)		
ref:22		17 Sep 97		
	-216450.	-271500.	22.470	
	11.7443	20.8978	-2.4707	-3.6429
	119.1022	36.3407	-0.0380	0.
Zn(Ae)3-		Zn(CH3COO)3(-)		
Zn(Ae)3-		Zn(1)C(6)H(9)O(6)-(1)		
ref:22		17 Sep 97		
	-305740.	-394090.	25.000	
	20.0332	41.1373	-10.4257	-4.4796
	203.1827	61.4365	1.2513	-1.
Zn(Ala)+		Zn(C3H6NO2)+		
Zn(Ala)+		Zn(1)C(3)H(6)N(1)O(2)+(1)		
ref:27		14 Aug 93		
	-116613.	-161048.	17.000	
	7.8514	11.3891	1.2746	-3.2497
	71.8106	18.8350	0.2956	1.
Zn(Ala)2, aq		Zn(C3H6NO3)2		
Zn(Ala)2, aq		Zn(1)C(6)H(12)N(2)O(6)		
ref:27		14 Aug 93		

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-197112. -283389. 60.000
18.1164 36.4558 -8.5823 -4.2860
149.2015 46.7767 -0.0300 0.
Zn(But)+ ZnCH3(CH2)2CO2+
Zn(But)+ Zn(1)C(4)H(7)O(2)+(1)
ref:27 26.Aug.93
-121815. -166539. 5.790
9.0597 14.3386 0.1167 -3.3717
102.3810 28.9294 0.4615 1.
Zn(But)2,aq Zn(CH3CH2CH2CO2)2
Zn(But)2,aq Zn(1)C(8)H(14)O(4)
ref:27 26.Aug.93
-207667. -296560. 34.741
20.5524 42.4038 -10.9205 -4.5319
218.5528 70.8814 -0.0300 0.
Zn(For)+ ZnCHO2+
Zn(For)+ Zn(1)C(1)H(1)O(2)+(1)
ref:27 26.Aug.93
-121477. -140698. -4.310
3.0712 -0.2795 5.8532 -2.7673
31.4547 3.7879 0.6143 1.
Zn(For)2,aq Zn(CHO2)2
Zn(For)2,aq Zn(1)C(2)H(2)O(4)
ref:27 26.Aug.93
-206908. -245726. 11.424
7.7824 11.2204 1.3417 -3.2428
45.8240 10.8453 -0.0300 0.
Zn(Gly)+ Zn(C2H4NO2)+
Zn(Gly)+ Zn(1)C(2)H(4)N(1)O(2)+(1)
ref:27 27.Aug.93
-117818. -151609. 18.000
5.3677 5.3226 3.6628 -2.9989
45.1185 9.6102 0.2792 1.
Zn(Gly)2,aq Zn(C2H4NO2)2
Zn(Gly)2,aq Zn(1)C(2)H(8)N(2)O(4)
ref:27 27.Aug.93
-199139. -267408. 55.000
12.8771 23.6625 -3.5537 -3.7571
85.8243 24.7484 -0.0300 0.
Zn(Glyc)+ ZnCH3OCO2+
Zn(Glyc)+ Zn(1)C(2)H(3)O(3)+(1)
ref:27 30.Aug.93
-159617. -194550. 0.190
4.9216 4.2387 4.0776 -2.9541
61.8303 14.5628 0.5464 1.
Zn(Glyc)2,aq Zn(CH3OCO2)2
Zn(Glyc)2,aq Zn(1)C(4)H(6)O(6)
ref:27 30.Aug.93
-283311. -353139. 21.813
11.7430 20.8930 -2.4643 -3.6426
119.8505 36.5751 -0.0300 0.
Zn(Lac)+ ZnCH3CH2OCO2+
Zn(Lac)+ Zn(1)C(3)H(5)O(3)+(1)
ref:27 30.Aug.93
-160731. -200064. 18.000
7.0755 9.4969 2.0126 -3.1715
81.3400 22.1998 0.2792 1.
Zn(Lac)2,aq Zn(CH3CH2OCO2)2
Zn(Lac)2,aq Zn(1)C(6)H(10)O(6)
ref:27 30.Aug.93
-285376. -364728. 55.000
16.4877 32.4799 -7.0229 -4.1216
172.3186 54.8116 -0.0300 0.
Zn(Pent)+ ZnCH3(CH2)3CO2+
Zn(Pent)+ Zn(1)C(5)H(9)O(2)+(1)
ref:27 7.Sep.93
-119683. -172896. 12.290
11.2174 19.6055 -1.9494 -3.5894
138.7888 41.8971 0.3636 1.
Zn(Pent)2,aq Zn(CH3CH2CH2CH2CO2)2
Zn(Pent)2,aq Zn(1)C(10)H(18)O(4)
ref:27 7.Sep.93

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-203415. -308690. 49.747
25.1842 53.7125 -15.3636 -4.9994
307.6444 101.8474 -0.0300 0.
Zn(Prop)+ ZnCH3CH2CO2+
Zn(Prop)+ Zn(1)C(3)H(5)O(2)+(1)
ref:27 15.Sep.93
-123675. -160939. 0.490
6.9879 9.2834 2.0962 -3.1627
88.3708 23.7877 0.5464 1.
Zn(Prop)2,aq Zn(CH3CH2CO2)2
Zn(Prop)2,aq Zn(1)C(6)H(10)O(4)
ref:27 15.Sep.93
-211455. -285915. 22.506
16.1116 31.5558 -6.6473 -4.0834
183.2279 58.6033 -0.0300 0.
Zr+4 Zr(+4)
Zr+4 Zr(1)+(4)
ref:21 11.Sep.97
-133270. -150330. -110.300
-4.3670 -18.4377 12.9811 -2.0167
39.7423 -3.6660 3.8417 4.
ZrOH+3 ZrOH(+3)
ZrOH+3 Zr(1)O(1)H(1)+(3)
ref:21 1.May.97
-190400. -212500. -71.600
9.279 -5.5093 7.9015 -2.5511
56.7436 6.0097 2.6654 3.
ZrO+2 ZrO(+2)
ZrO+2 Zr(1)O(1)+(2)
ref:21 1.May.97
-187600. -204300. -53.400
1.9014 -3.1364 6.9769 -2.6492
23.3712 -3.0141 1.8611 2.
ZrO2,aq ZrO2
ZrO2(aq) Zr(1)O(2)+(0)
ref:21 1.May.97
-233400. -263800. -43.700
1.8643 -3.2264 7.0119 -2.6455
38.4166 8.2708 -0.0300 0.
143 minerals that do not undergo phase transition (nmin1)
33 minerals that undergo one phase transition (nmin2)
17 minerals that undergo two phase transitions (nmin3)
2 minerals that undergo three phase transitions (nmin4)
16 gases (ngas)
1147 aqueous species (naqs)
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1358 total
258 comment lines precede top of min1 block

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Appendix B: Listing of the PHREEQE.JNC Database

(see Tables 5.3.1_1 - 5.3.1_3 for formatting and unit conventions)

ELEMENTS							
Ca	4	40.078	Ca+2				
Mg	5	24.305	Mg+2				
Na	6	22.9898	Na+				
K	7	39.0983	K+				
Fe	8	55.847	Fe+2				
Mn	9	54.9380	Mn+2				
Al	10	26.9815	Al+3				
Ba	11	137.327	Ba+2				
Sr	12	87.62	Sr+2				
Si	13	60.0843	SiO2				
Cl	14	35.4527	Cl-				
C	15	44.0098	CO2				
S	16	96.064	SO4-2				
N	17	62.0049	NO3-				
B	18	10.811	B				
P	19	94.9714	PO4-3				
F	20	18.9984	F-				
Li	21	6.941	Li+				
Br	22	79.904	Br-				
I	23	126.9045	I-				
Bi	25	208.9804	Bi+3				
Cs	27	132.9054	Cs+				
Ni	29	58.6934	Ni+2				
Pb	32	207.2	Pb+2				
Sb	35	153.756	SbO2-				
Se	37	142.96	SeO4-2				
Sn	39	118.710	Sn+2				
Th	41	232.0381	Th+4				
U	42	238.0289	U+4				
Zr	43	91.224	Zr+4				

SPECIES							
1							
H+	101 1.0	0.0	9.0	9.0		-1.0	
0.0	0.0						
1 1.0							
2							
E-	100 -1.0	0.0	0.0			0.0	
0.0	0.0						
2 1.0							
3							
H2O	100 0.0	0.0	0.0			0.0	
0.0	0.0						
3 1.0							
4							
Ca+2	101 2.0	0.0	6.0	5.0	0.165	0.0	
0.0	0.0						
4 1.0							
5							
Mg+2	101 2.0	0.0	8.0	5.5	0.20	0.0	
0.0	0.0						
5 1.0							
6							
Na+	101 1.0	0.0	4.0	4.0	0.075	0.0	
0.0	0.0						
6 1.0							
7							
K+	101 1.0	0.0	3.0	3.5	0.015	0.0	
0.0	0.0						
7 1.0							
8							

Fe+2	100 2.0	2.0	6.0			0.0
0.0	0.0					
8 1.0						
9						
Mn+2	100 2.0	2.0	6.0			0.0
0.0	0.0					
9 1.0						
10						
Al+3	100 3.0	0.0	9.0			0.0
0.0	0.0					
10 1.0						
11						
Ba+2	100 2.0	0.0	5.0			0.0
0.0	0.0					
11 1.0						
12						
Sr+2	101 2.0	0.0	5.0	5.26	0.121	0.0
0.0	0.0					
12 1.0						
13						
H4SiO4	100 0.0	0.0	0.0			0.0
0.0	0.0					
13 1.0						
14						
Cl-	101 -1.0	0.0	3.0	3.5	0.015	0.0
0.0	0.0					
14 1.0						
15						
CO3-2	101 -2.0	4.0	4.5	5.4	0.0	2.0
0.0	0.0					
15 1.0						
16						
SO4-2	101 -2.0	6.0	4.0	5.0	-0.04	0.0
0.0	0.0					
16 1.0						
17						
NO3-	100 -1.0	5.0	3.0			0.0
0.0	0.0					
17 1.0						
18						
H3BO3	100 0.0	0.0	0.0			0.0
0.0	0.0					
18 1.0						
19						
PO4-3	100 -3.0	0.0	4.0			2.0
0.0	0.0					
19 1.0						
20						
F-	100 -1.0	0.0	3.5			0.0
0.0	0.0					
20 1.0						
21						
Li+	100 1.0	0.0	6.0			0.0
0.0	0.0					
21 1.0						
22						
Br-	100 -1.0	0.0	3.0			0.0
0.0	0.0					
22 1.0						
23						
I-	100 -1.0	-1.0	3.0			0.0
0.0	0.0					

23	1.0								
25									
Bi+3	100	3.0	0.0	0.0				0.0	
0.0	0.0								
25	1.0								
27									
Cs+	100	1.0	0.0	2.5				0.0	
0.0	0.0								
27	1.0								
29									
Ni+2	100	2.0	0.0	6.0				0.0	
0.0	0.0								
29	1.0								
32									
Pb+2	100	2.0	0.0	4.5				0.0	
0.0	0.0								
32	1.0								
33									
Pd+2	100	2.0	0.0	0.0				0.0	
0.0	0.0								
33	1.0								
35									
SbO2-	100	-1.0	0.0	0.0				0.0	
0.0	0.0								
35	1.0								
37									
SeO4-2	100	-2.0	6.0	4.0				0.0	
0.0	0.0								
37	1.0								
39									
Sn+2	100	2.0	2.0	6.0				0.0	
0.0	0.0								
39	1.0								
41									
Th+4	100	4.0	4.0	11.0				0.0	
0.0	0.0								
41	1.0								
42									
U+4	100	4.0	4.0	0.0				0.0	
0.0	0.0								
42	1.0								
43									
Zr+4	100	4.0	0.0	0.0				0.0	
0.0	0.0								
43	1.0								
100									
OH-	210	-1.000	.000	3.500	.000	.000	1.000		
-13.995	13.340	-814.968	-.135133969	41130.89	296.229				
-2639175.4									
3	1.000	1	-1.000						
101									
O2 AQ	310	.000	4.000	.000	.000	.000	.000		
-86.003	133.734	508.521	.078341650	-53472.79	-184.907				
1689617.4									
3	2.000	1	-4.000	2	-4.000				
102									
H2 AQ	210	.000	-2.000	.000	.000	.000	.000		
-3.105	-1.000	238.732	.041950712	-15527.16	-86.009				
938463.7									
1	2.000	2	2.000						
103									
HCO3-	211	-1.000	4.000	4.500	5.400	.000	1.000		

10.329	-3.513	988.487	.164396835	-52292.14	-358.627				
3165965.0									
15	1.000	1	1.000						
104									
H2CO3	310	.000	4.000	.000	.000	.000	.000		
16.673	-5.832	1961.123	.327901516	-105318.32	-712.077				
6490384.7									
15	1.000	1	2.000	3	-1.000				
105									
HSO4-	210	-1.000	6.000	.000	.000	.000	.000		
1.979	4.900	1031.228	.167740348	-57682.75	-373.661				
3449687.8									
16	1.000	1	1.000						
106									
HS-	410	-1.000	-2.000	3.500	.000	.000	.000	1.000	
33.690	-59.717	877.665	.157143140	-42510.55	-316.250				
3048568.3									
16	1.000	1	9.000	2	8.000	3	-4.000		
107									
H2S	410	.000	-2.000	.000	.000	.000	.000	.000	
40.677	-64.868	1796.831	.312081279	-91103.02	-651.361				
6054543.5									
16	1.000	1	10.000	2	8.000	3	-4.000		
108									
H2S2O3	410	.000	4.000	.000	.000	.000	.000	-2.000	
40.739	-57.184	3064.125	.522179413	-160696.36	-1112.631				
10048214.3									
16	2.000	1	12.000	2	8.000	3	-5.000		
109									
HS2O3-	410	-1.000	4.000	.000	.000	.000	.000	-1.000	
40.153	-60.684	2403.048	.410787847	-125499.26	-871.141				
8101435.5									
16	2.000	1	11.000	2	8.000	3	-5.000		
110									
H2S2O4	410	.000	6.000	.000	.000	.000	.000	-2.000	
13.407	-13.767	2885.921	.488097780	-159685.80	-1046.904				
9604362.5									
16	2.000	1	10.000	2	6.000	3	-4.000		
111									
HS2O4-	410	-1.000	6.000	.000	.000	.000	.000	-1.000	
13.041	-17.567	2586.186	.437111025	-142733.53	-938.290				
8622370.3									
16	2.000	1	9.000	2	6.000	3	-4.000		
112									
HSO3-	410	-1.000	4.000	4.000	.000	.000	.000	-1.000	
3.583	-.587	1003.072	.167352354	-56521.75	-362.744				
3358039.7									
16	1.000	1	3.000	2	2.000	3	-1.000		
113									
HSO5-	410	-1.000	8.000	.000	.000	.000	.000	.000	
-60.288	100.337	1643.797	.263756937	-109486.70	-596.329				
5340127.2									
16	1.000	3	1.000	2	-2.000	1	-1.000		
114									
H2BO3-	210	-1.000	.000	.000	.000	.000	.000	1.000	
-9.249	3.907	-749.560	-.118300018	40652.28	269.835				
-2529481.4									
18	1.000	1	-1.000						
115									
HPO4-2	210	-2.000	.000	4.000	.000	.000	.000	1.000	
12.322	-3.515	1118.483	.185856870	-59072.19	-405.509				
3552188.0									

19	1.000	1	1.000						
116									
H2PO4-	210	-1.000	.000	4.500	.000	.000	.000		
19.527	-4.520	1835.861	.305075586	-97715.09	-665.229				
5912355.3									
19	1.000	1	2.000						
117									
HF AQ	210	.000	.000	.000	.000	.000	-1.000		
3.168	3.315	502.317	.085310800	-27798.71	-181.851				
1656258.1									
1	1.000	20	1.000						
118									
HF2-	210	-1.000	.000	.000	.000	.000	-2.000		
2.551	4.960	621.371	.105034434	-33039.06	-226.098				
1790449.1									
1	1.000	20	2.000						
119									
H3SiO4-	210	-1.000	.000	.000	.000	.000	1.000		
-9.585	4.220	2517.724	.359699913	-143036.87	-909.042				
8405688.6									
13	1.000	1	-1.000						
120									
H2SiO4-2	210	-2.0	0.0	0.0		2.0			
-23.0	17.6	-294.0184	-0.072650	11204.49	108.18466				
-1119669.0									
13	1.0	1	-2.0						
121									
HCl	210	.000	.000	.000	.000	.000	-1.000		
-7.710	9.648	214.218	.054464831	-10517.85	-82.749				
788406.1									
1	1.000	14	1.000						
122									
HNO3	210	.000	5.000	.000	.000	.000	.000		
-1.303	4.019	671.846	.112252371	-36724.24	-245.315				
2095373.3									
1	1.000	17	1.000						
123									
H3PO4	210	.000	.000	.000	.000	.000	.000		
21.697	-2.620	2316.714	.386099144	-123813.98	-840.013				
7440311.8									
19	1.000	1	3.000						
124									
HIO3	410	.000	5.000	.000	.000	.000	-1.000		
-110.518	168.050	576.193	.089398004	-61456.94	-212.617				
1677277.5									
23	1.000	3	3.000	1	-5.000	2	-6.000		
125									
HSbO2	210	.000	.000	.000	.000	.000	.000		
10.995	-6.600	205.995	.036144163	-9032.29	-73.725				
617370.9									
35	1.000	1	1.000						
127									
HSe-	410	-1.000	-2.000	.000	.000	.000	.000		
81.181	-126.267	677.195	.128654474	-16648.95	-244.559				
2365666.7									
37	1.000	1	9.000	2	8.000	3	-4.000		
128									
HSeO3-	410	-1.000	4.000	.000	.000	.000	.000		
36.304	-48.097	958.119	.162509198	-43513.48	-347.650				
3192628.3									
37	1.000	1	3.000	2	2.000	3	-1.000		
129									

H2SeO3	410	.000	4.000	.000	.000	.000	-1.000		
38.877	-46.407	1466.667	.248458421	-70758.45	-532.712				
4766215.5									
37	1.000	1	4.000	2	2.000	3	-1.000		
130									
HSeO4-	210	-1.000	6.000	.000	.000	.000	-1.000		
1.906	4.200	1109.489	.183362040	-60980.15	-403.400				
3596961.1									
37	1.000	1	1.000						
300									
AlOH+2	310	2.000	.000	.000	.000	.000	1.000		
-4.957	11.902	518.723	.087307679	-30187.23	-188.639				
1627784.5									
10	1.000	3	1.000	1	-1.000				
301									
AlOH2+	310	1.000	.000	.000	.000	.000	2.000		
-10.594	23.490	635.697	.106955890	-38498.51	-230.700				
1937490.5									
10	1.000	3	2.000	1	-2.000				
302									
AlOH3	310	.000	.000	.000	.000	.000	3.000		
-16.433	34.585	1026.675	.168585830	-61549.69	-372.496				
3091927.9									
10	1.000	3	3.000	1	-3.000				
303									
AlOH4-	310	-1.000	.000	.000	.000	.000	4.000		
-22.883	43.236	254.968	.037464280	-21866.23	-89.941				
610777.9									
325									
BaOH+	310	1.000	.000	.000	.000	.000	1.000		
-13.500	20.917	-149.785	-.023556399	5059.36	54.248				
-701773.9									
11	1.000	3	1.000	1	-1.000				
326									
BaCl+	210	1.000	.000	.000	.000	.000	.000		
-4.498	3.110	551.818	.096853934	-30344.23	-201.966				
1807602.8									
11	1.000	14	1.000						
327									
BaF+	210	1.000	.000	.000	.000	.000	.000		
-1.183	2.140	773.520	.131270194	-41996.15	-282.807				
2471505.7									
11	1.000	20	1.000						
328									
BaCO3	210	.000	4.000	.000	.000	.000	2.000		
2.645	4.035	1202.115	.204575449	-63300.94	-439.639				
3529649.3									
11	1.000	15	1.000						
350									
BiOH+2	310	2.000	.000	.000	.000	.000	1.000		
-1.105	4.157	-15.546	-.004256649	1633.32	5.617				
-326041.0									
25	1.000	3	1.000	1	-1.000				
351									
BiOH2+	310	1.000	.000	.000	.000	.000	2.000		
-3.304	18.657	-701.450	-.113390384	38819.53	256.574				
-2944397.0									
25	1.000	3	2.000	1	-2.000				
352									
BiOH3	310	.000	.000	.000	.000	.000	3.000		
-8.206	30.974	-1395.429	-.228640541	76549.91	509.901				
-5606920.9									

25	1.000	3	3.000	1	-3.000			
353								
BiOH4-	310	-1.000	.000	.000	.000	.000	.000	4.000
-21.107	45.674	-1262.701	-.209757822	64579.44	462.127			
-4975299.5								
25	1.000	3	4.000	1	-4.000			
400								
CaOH+	310	1.000	.000	.000	.000	.000	.000	1.000
-12.833	18.517	48.057	.008051439	-5894.27	-17.763			
38295.7								
4	1.000	3	1.000	1	-1.000			
401								
CaCl+	210	1.000	.000	.000	.000	.000	.000	.000
-.292	1.126	756.098	.130939275	-40791.05	-277.128			
2410655.5								
4	1.000	14	1.000					
402								
CaCl2	210	.000	.000	.000	.000	.000	.000	.000
-.644	-1.394	1589.350	.272060292	-84643.89	-583.386			
5008097.6								
4	1.000	14	2.000					
403								
CaF+	210	1.000	.000	.000	.000	.000	.000	.000
.682	1.350	1067.673	.178933274	-57839.51	-389.929			
3423074.0								
4	1.000	20	1.000					
404								
CaHCO3+	311	1.000	4.000	.000	5.400	.000	.000	1.000
11.376	-3.165	2048.601	.339363192	-110541.90	-744.834			
6702119.6								
4	1.000	15	1.000	1	1.000			
405								
CaCO3	210	.000	4.000	.000	.000	.000	.000	2.000
3.327	3.795	1266.919	.215696046	-66449.82	-463.382			
3696159.4								
4	1.000	15	1.000					
406								
CaSO4	210	.000	6.000	.000	.000	.000	.000	.000
2.111	1.300	1286.832	.216006198	-67690.42	-470.850			
3822460.9								
4	1.000	16	1.000					
410								
CaH3SiO4	310	1.000	.000	.000	.000	.000	.000	1.000
-8.575	4.783	3799.621	.572165180	-211848.92	-1377.281			
12422706.4								
4	1.000	13	1.000	1	-1.000			
425								
CH4 AQ	410	.000	-4.000	.000	.000	.000	.000	.000
38.194	-64.575	1957.295	.335797111	-100765.85	-709.737			
6662566.5								
15	1.000	1	10.000	2	8.000	3	-3.000	
450								
CsOH	310	.000	.000	.000	.000	.000	.000	1.000
-15.685	17.687	-269.095	-.049506083	11122.83	97.848			
-1000453.9								
27	1.000	3	1.000	1	-1.000			
451								
CsCl	210	.000	.000	.000	.000	.000	.000	.000
-.139	.653	456.846	.075746038	-24335.19	-166.885			
1333328.4								
27	1.000	14	1.000					
452								

CsBr	210	.000	.000	.000	.000	.000	.000	.000
.022	2.220	383.709	.063348834	-20432.13	-140.117			
1125972.7								
27	1.000	22	1.000					
453								
CsI	210	.000	.000	.000	.000	.000	.000	.000
.982	-2.550	224.949	.038641780	-11614.28	-82.453			
665976.8								
27	1.000	23	1.000					
500								
FeOH+	310	1.000	2.000	.000	.000	.000	.000	1.000
-9.315	12.287	458.740	.072604481	-27012.31	-167.456			
1356234.5								
8	1.000	3	1.000	1	-1.000			
501								
FeOH2	310	.000	2.000	.000	.000	.000	.000	2.000
-20.405	27.417	180.811	.028793639	-14872.42	-66.543			
421087.1								
8	1.000	3	2.000	1	-2.000			
502								
FeOH3-	310	-1.000	2.000	.000	.000	.000	.000	3.000
-29.207	32.984	428.934	.059826568	-31339.83	-156.140			
1377230.6								
8	1.000	3	3.000	1	-3.000			
503								
FeCl+	210	1.000	2.000	.000	.000	.000	.000	.000
-.161	.723	876.802	.147143826	-47742.47	-320.423			
2858795.1								
8	1.000	14	1.000					
504								
FeCl2	210	.000	2.000	.000	.000	.000	.000	.000
-8.172	23.426	1624.758	.273351854	-91791.49	-591.719			
5121162.5								
8	1.000	14	2.000					
505								
FeF+	210	1.000	2.000	.000	.000	.000	.000	.000
1.427	.748	1006.203	.166186287	-54475.98	-366.935			
3230866.7								
8	1.000	20	1.000					
514								
Fe+3	210	3.000	3.000	9.000	.000	.000	.000	.000
-13.011	10.200	-101.736	-.017163759	3781.73	34.213			
-311058.6								
8	1.000	2	-1.000					
515								
FeOH+2	410	2.000	3.000	.000	.000	.000	.000	1.000
-15.216	20.367	113.510	.015220631	-9193.40	-42.482			
239182.9								
8	1.000	3	1.000	2	-1.000	1	-1.000	
516								
FeOH2+	410	1.000	3.000	.000	.000	.000	.000	2.000
-18.661	29.367	-469.648	-.078431304	21664.30	170.453			
-1783755.2								
8	1.000	3	2.000	1	-2.000	2	-1.000	
517								
FeOH3	410	.000	3.000	.000	.000	.000	.000	3.000
-25.029	38.384	-642.841	-.108881515	29996.50	234.158			
-2644016.6								
8	1.000	3	3.000	1	-3.000	2	-1.000	
518								
FeOH4-	410	-1.000	3.000	.000	.000	.000	.000	4.000
-34.631	47.984	-619.528	-.109884387	24467.08	226.853			

-2287898.8									
8	1.000	3	4.000	1	-4.000	2	-1.000		
521									
FeCl+2	310		2.000		3.000		.000	.000	.000
-11.531	11.163				686.031		.115696727	-38860.60	-253.292
2225664.2									
8	1.000	14	1.000	2	-1.000				
524									
FeF+2	310		2.000		3.000		.000	.000	.000
-7.011	15.059				848.998		.139466396	-48129.57	-309.804
2704569.6									
8	1.000	20	1.000	2	-1.000				
550									
IO3-	410		-1.000		5.000		4.500	.000	.000
-111.324	165.650				-80.344		-.019674636	-25739.47	26.554
-399135.9									
23	1.000	3	3.000	1	-6.000	2	-6.000		
551									
IO4-	410		-1.000		7.000		3.500	.000	.000
-165.044	250.667				285.173		.036433312	-61406.97	-106.000
637400.5									
23	1.000	3	4.000	1	-8.000	2	-8.000		
575									
KOH	310		.000		.000		.000	.000	1.000
-14.895	16.483				349.824		.050960980	-21962.98	-128.042
940603.2									
7	1.000	3	1.000	1	-1.000				
576									
KCl	210		.000		.000		.000	.000	.000
-1.751	2.803				941.395		.149125852	-52074.36	-342.358
3039483.1									
7	1.000	14	1.000						
577									
KBr	210		.000		.000		.000	.000	.000
-1.737	2.990				242.347		.042919392	-13652.55	-88.844
777675.4									
7	1.000	22	1.000						
578									
KI	210		.000		.000		.000	.000	.000
-1.598	2.041				231.496		.042873438	-12537.22	-85.636
717633.6									
7	1.000	23	1.000						
579									
KHSO4	310		.000		6.000		.000	.000	-1.000
-179.909	277.670				1488.638		.241725974	-132606.48	-545.199
4729756.6									
7	1.000	16	1.000	1	1.000				
580									
KSO4-	210		-1.000		6.000		.000	.000	.000
.880	.690				1102.996		.179334687	-59776.88	-401.927
3506858.5									
7	1.000	16	1.000						
582									
KAlO4	410		.000		.000		.000	.000	4.000
-24.224	50.603				1193.112		.186224965	-72937.50	-430.881
3374636.9									
7	1.000	10	1.000	3	4.000	1	-4.000		
600									
LiOH	310		.000		.000		.000	.000	1.000
-13.649	13.369				-20.761		-.006191102	-1294.13	6.302
-203984.6									
21	1.000	3	1.000	1	-1.000				

601									
LiCl	210		.000		.000		.000	.000	.000
-1.511	.805				526.414		.090058339	-28586.59	-193.139
1690109.2									
21	1.000	14	1.000						
625									
MgOH+	310		1.000		.000		.000	.000	1.000
-11.682	26.729				636.776		.098556534	-38168.05	-231.461
2036313.2									
5	1.000	3	1.000	1	-1.000				
626									
MgCl+	210		1.000		.000		.000	.000	.000
-.135	.367				1088.098		.180272411	-58882.26	-397.534
3483003.2									
5	1.000	14	1.000						
627									
MgF+	210		1.000		.000		.000	.000	.000
1.352	.567				1284.027		.208518452	-70287.10	-467.329
4202395.6									
5	1.000	20	1.000						
628									
MgHCO3+	310		1.000		4.000		.000	.000	1.000
11.365	-2.998				2477.682		.404496646	-133861.26	-900.524
8030882.7									
5	1.000	1	1.000	15	1.000				
629									
MgCO3	210		.000		4.000		.000	.000	2.000
2.979	2.182				1363.900		.226965746	-72097.13	-498.117
4069666.6									
5	1.000	15	1.000						
635									
MgH3SiO4	310		1.000		.000		.000	.000	1.000
-8.325	3.759				4051.337		.608609927	-225649.90	-1468.304
13239398.7									
5	1.000	13	1.000	1	-1.000				
650									
MnOH+	310		1.000		2.000		.000	.000	1.000
-10.620	14.417				287.434		.044885174	-18036.78	-104.992
787107.2									
9	1.000	3	1.000	1	-1.000				
651									
MnOH2	310		.000		2.000		.000	.000	2.000
-22.201	22.117				-3.606		-.000866611	-4831.90	.247
-243738.1									
9	1.000	3	2.000	1	-2.000				
652									
MnOH3-	310		-1.000		2.000		.000	.000	3.000
-34.800	39.634				10.159		-.007005930	-9359.18	-4.313
-71771.5									
9	1.000	3	3.000	1	-3.000				
653									
MnOH4-2	310		-2.000		2.000		.000	.000	4.000
-48.287	56.234				-1067.057		-.186190332	45507.93	388.739
-3578904.8									
9	1.000	3	4.000	1	-4.000				
654									
MnCl+	210		1.000		2.000		.000	.000	.000
-.139	4.553				965.727		.162105090	-52863.14	-352.225
3081470.6									
9	1.000	14	1.000						
657									
MnF+	210		1.000		2.000		.000	.000	.000

3770819.7	.880	.596	1194.770	.197973966	-64321.91	-436.301	
9	1.000	20	1.000				
660							
MnSO4	210	.000	8.000	.000	.000	.000	
	1.913	3.550	1452.985	.241233096	-76839.40	-531.000	
4324641.8							
9	1.000	16	1.000				
662							
Mn+3	210	3.000	3.000	.000	.000	.000	
	-25.509	22.200	-350.610	-.057782802	14257.71	123.895	
-1072167.2							
9	1.000	2	-1.000				
663							
MnO4-2	410	-2.000	6.000	.000	.000	.000	
	-118.345	169.567	-942.744	-.173063096	18728.17	343.932	
-3365013.4							
9	1.000	3	4.000	1 -8.000	2 -4.000		
675							
NaOH	310	.000	.000	.000	.000	.000	1.000
	-14.795	12.823	416.532	.065022580	-25700.92	-153.295	
1316230.2							
6	1.000	3	1.000	1 -1.000			
676							
NaCl	210	.000	.000	.000	.000	.000	.000
	-.777	1.206	435.280	.072276061	-24040.10	-158.876	
1435884.1							
6	1.000	14	1.000				
677							
NaF	210	.000	.000	.000	.000	.000	.000
	-.998	1.723	886.198	.145078346	-47509.85	-323.890	
2697174.5							
6	1.000	20	1.000				
678							
NaBr	210	.000	.000	.000	.000	.000	.000
	-1.357	1.643	632.427	.104303732	-34380.69	-231.113	
1982896.3							
6	1.000	22	1.000				
679							
NaI	210	.000	.000	.000	.000	.000	.000
	-1.540	1.932	570.484	.095763909	-30846.25	-209.002	
1781773.4							
6	1.000	23	1.000				
682							
NaSO4-	210	-1.000	6.000	.000	.000	.000	.000
	.923	-.647	1469.473	.235914956	-79536.64	-535.343	
4671368.2							
6	1.000	16	1.000				
684							
NaAlOH4	410	.000	.000	.000	.000	.000	4.000
	-23.627	45.489	1448.202	.230900017	-86305.07	-525.223	
4304910.5							
6	1.000	10	1.000	3 4.000	1 -4.000		
685							
NaB(OH)4	410	.000	.000	.000	.000	.000	1.000
	.278	-.147	191.667	.036737069	-10360.30	-70.778	
670417.5							
6	1.000	18	1.000	3 1.000	1 -1.000		
686							
NaH3SiO4	310	.000	.000	.000	.000	.000	1.000
	-7.754	1.631	3757.083	.561173967	-208056.83	-1362.488	
12183994.2							

6	1.000	13	1.000	1 -1.000			
725							
NiOH+	310	1.000	.000	.000	.000	.000	1.000
	-10.803	13.567	621.981	.097869700	-37136.99	-226.372	
2021226.0							
29	1.000	3	1.000	1 -1.000			
726							
NiOH2	310	.000	.000	.000	.000	.000	2.000
	-20.706	24.987	595.585	.096080522	-39053.07	-215.168	
1641594.0							
29	1.000	3	2.000	1 -2.000			
727							
NiOH3-	310	-1.000	.000	.000	.000	.000	3.000
	-31.003	30.774	657.860	.096381627	-44762.59	-239.867	
2317425.3							
29	1.000	3	3.000	1 -3.000			
728							
NiOH4-2	310	-2.000	.000	.000	.000	.000	4.000
	-44.036	47.734	-444.237	-.085913899	11439.71	162.328	
-1264209.3							
29	1.000	3	4.000	1 -4.000			
729							
NiCl+	210	1.000	.000	.000	.000	.000	.000
	-.996	1.433	860.735	.145492560	-46393.90	-315.232	
2712784.2							
29	1.000	14	1.000				
730							
NiF+	210	1.000	.000	.000	.000	.000	.000
	1.120	.735	938.690	.154645514	-50955.26	-342.182	
3016570.8							
29	1.000	20	1.000				
775							
NH3 AQ	410	.000	-3.000	.000	.000	.000	1.000
	109.906	-174.961	362.506	.075072550	10293.78	-131.497	
1410912.7							
17	1.000	1	9.000	2 8.000	3 -3.000		
776							
NH4+	410	1.000	-3.000	2.500	.000	.000	.000
	119.147	-187.371	549.591	.103123837	3037.58	-199.394	
1956469.2							
17	1.000	1	10.000	2 8.000	3 -3.000		
778							
N2 AQ	410	.000	.000	.000	.000	.000	.000
	207.269	-313.538	1494.655	.267559947	-24081.60	-544.108	
5331034.3							
17	2.000	1	12.000	2 10.000	3 -6.000		
779							
NO2-	410	-1.000	3.000	3.000	.000	.000	.000
	27.767	-43.888	-81.292	-.009696531	12033.72	29.920	
-217519.0							
17	1.000	1	2.000	2 2.000	3 -1.000		
825							
PbOH+	310	1.000	.000	.000	.000	.000	1.000
	-6.192	-.803	-61.676	-.012242871	3789.23	20.182	
-312443.6							
32	1.000	3	1.000	1 -1.000			
826							
PbOH2	310	.000	.000	.000	.000	.000	2.000
	-16.894	23.397	-338.453	-.056356175	14784.26	122.803	
-1341649.3							
32	1.000	3	2.000	1 -2.000			
827							

PbOH3-	310	-1.000	.000	.000	.000	.000	3.000
-27.917	32.214	-752.526	-.130113036	35175.17	273.305		
-2742511.3							
32	1.000	3	3.000	1	-3.000		
830							
PbCl+	210	1.000	.000	.000	.000	.000	.000
1.437	1.083	783.284	.132501234	-42110.09	-286.110		
2475238.9							
32	1.000	14	1.000				
831							
PbCl2	210	.000	.000	.000	.000	.000	.000
2.003	1.946	1297.314	.220894673	-69252.17	-474.537		
4028009.6							
32	1.000	14	2.000				
832							
PbCl3-	210	-1.000	.000	.000	.000	.000	.000
1.688	1.879	1648.091	.279928337	-87854.47	-603.200		
5100824.5							
32	1.000	14	3.000				
833							
PbCl4-2	210	-2.000	.000	.000	.000	.000	.000
1.430	.303	1647.021	.280297498	-87799.03	-603.213		
5149501.2							
32	1.000	14	4.000				
834							
PbF+	210	1.000	.000	.000	.000	.000	.000
2.060	-.661	996.558	.165226865	-53308.67	-363.815		
3135515.6							
32	1.000	20	1.000				
835							
PbF2	210	.000	.000	.000	.000	.000	.000
3.420	-2.703	1599.953	.265609161	-85547.60	-584.360		
5081486.3							
32	1.000	20	2.000				
836							
PbHS2	510	.000	-2.000	.000	.000	.000	2.000
82.089	-134.805	3557.832	.611454222	-178989.38	-1289.641		
11859835.6							
32	1.000	16	2.000	1	18.000	2	16.000
837							
PbHS3-	510	-1.000	-2.000	.000	.000	.000	3.000
117.078	-196.525	4841.737	.832664974	-244229.09	-1753.096		
16369938.4							
32	1.000	16	3.000	1	27.000	2	24.000
850							
PdOH+	310	1.000	.000	.000	.000	.000	1.000
-1.091	-.763	415.150	.064184337	-22633.86	-151.298		
1325776.4							
33	1.000	3	1.000	1	-1.000		
851							
PdOH2	310	.000	.000	.000	.000	.000	2.000
-2.190	1.537	256.363	.041558345	-12866.44	-94.705		
582292.2							
33	1.000	3	2.000	1	-2.000		
852							
PdCl+	210	1.000	.000	.000	.000	.000	.000
6.099	-7.647	959.383	.161106997	-50353.56	-350.611		
3123209.3							
33	1.000	14	1.000				
853							
PdCl2	210	.000	.000	.000	.000	.000	.000
10.733	-15.814	1819.284	.305982103	-95522.32	-665.417		

5967755.9							
33	1.000	14	2.000				
854							
PdCl3-	210	-1.000	.000	.000	.000	.000	.000
13.094	-24.281	2572.586	.427939382	-135633.62	-940.824		
8519672.8							
33	1.000	14	3.000				
855							
PdCl4-2	210	-2.000	.000	.000	.000	.000	.000
15.162	-36.348	2795.782	.465087629	-146547.22	-1024.014		
9430857.8							
33	1.000	14	4.000				
901							
S2-2	410	-2.000	-2.000	.000	.000	.000	.000
57.645	-104.534	1451.874	.264883400	-69953.19	-523.794		
5112593.0							
16	2.000	1	16.000	2	14.000	3	-8.000
902							
S3-2	410	-2.000	-2.000	.000	.000	.000	.000
94.457	-161.401	2907.217	.517957838	-142258.09	-1051.762		
9997170.3							
16	3.000	1	24.000	2	20.000	3	-12.000
903							
S4-2	410	-2.000	-2.000	.000	.000	.000	.000
131.049	-217.968	4346.529	.767958213	-213906.87	-1573.717		
14850442.5							
16	4.000	1	32.000	2	26.000	3	-16.000
904							
S5-2	410	-2.000	-2.000	.000	.000	.000	.000
167.422	-274.235	5580.331	.983261211	-275274.95	-2020.000		
19162497.9							
16	5.000	1	40.000	2	32.000	3	-20.000
905							
S203-2	410	-2.000	4.000	4.000	.000	.000	.000
38.467	-61.784	1536.132	.266874951	-77744.80	-555.799		
5227983.1							
16	2.000	1	10.000	2	8.000	3	-5.000
906							
S204-2	410	-2.000	6.000	5.000	.000	.000	.000
10.549	-18.567	1611.209	.275984254	-89406.41	-583.562		
5414621.5							
16	2.000	1	8.000	2	6.000	3	-4.000
907							
S205-2	410	-2.000	8.000	.000	.000	.000	.000
2.348	-2.150	1819.335	.306422449	-101699.51	-660.365		
5936649.9							
16	2.000	1	6.000	2	4.000	3	-3.000
908							
S206-2	410	-2.000	10.000	4.000	.000	.000	.000
-8.418	17.766	1701.044	.282382169	-98348.05	-616.986		
5591123.9							
16	2.000	1	4.000	2	2.000	3	-2.000
909							
S208-2	210	-2.000	14.000	4.000	.000	.000	.000
-65.501	113.400	1950.999	.315193634	-128651.93	-707.231		
6313368.3							
16	2.000	2	-2.000				
910							
S306-2	410	-2.000	10.000	.000	.000	.000	.000
25.902	-36.701	3334.767	.565013805	-179919.23	-1211.070		
10919710.6							
16	3.000	1	12.000	2	8.000	3	-6.000

911								
S406-2	410	-2.000	10.000	.000	.000	.000	.000	
76.128	-106.168	4830.090	.823862961	-251911.60	-1751.763			
15995568.7								
16	4.000	1	20.000	2	14.000	3	-10.000	
912								
S506-2	410	-2.000	10.000	.000	.000	.000	.000	
97.474	-150.435	5922.101	1.016944428	-307733.36	-2149.155			
19757256.1								
16	5.000	1	28.000	2	20.000	3	-14.000	
913								
S03-2	410	-2.000	4.000	4.500	.000	.000	.000	
-3.623	-2.817	-76.012	-.011796480	1877.60	28.569			
-96236.1								
16	1.000	1	2.000	2	2.000	3	-1.000	
914								
S02	410	.000	4.000	.000	.000	.000	.000	
5.443	3.572	1826.363	.308250395	-101820.26	-662.569			
6059447.6								
16	1.000	1	4.000	2	2.000	3	-2.000	
925								
SeO3-2	410	-2.000	4.000	.000	.000	.000	1.000	
29.018	-46.817	-49.881	-.003537584	11119.27	17.864			
-137209.0								
37	1.000	1	2.000	2	2.000	3	-1.000	
975								
SnOH+	310	1.000	.000	.000	.000	.000	1.000	
-3.414	6.617	144.023	.020579352	-8438.16	-52.041			
311262.2								
39	1.000	3	1.000	1	-1.000			
976								
SnOH2	310	.000	.000	.000	.000	.000	2.000	
-7.079	10.317	-150.092	-.025521383	7295.27	54.228			
-713824.4								
39	1.000	3	2.000	1	-2.000			
977								
SnOH3-	310	-1.000	.000	.000	.000	.000	3.000	
-16.599	16.634	-369.411	-.068739688	17082.04	134.129			
-1411683.5								
39	1.000	3	3.000	1	-3.000			
1000								
SrOH+	311	1.000	.000	5.000	5.000	.000	1.000	
-13.302	19.987	99.133	.014069720	-8839.39	-35.913			
167232.5								
12	1.000	3	1.000	1	-1.000			
1001								
SrCl+	210	1.000	.000	.000	.000	.000	.000	
-2.248	1.813	940.954	.158962656	-50551.22	-344.352			
2936080.2								
12	1.000	14	1.000					
1002								
SrF+	210	1.000	.000	.000	.000	.000	.000	
.139	1.150	1046.500	.173031837	-56581.59	-382.094			
3315152.4								
12	1.000	20	1.000					
1004								
SrCO3	210	.000	4.000	.000	.000	.000	2.000	
2.865	4.435	1395.434	.233818757	-73753.32	-509.674			
4110630.5								
12	1.000	15	1.000					
1025								
SiF6-2	410	-2.000	.000	.000	.000	.000	-4.000	

26.275	-16.959	6090.487	.959294846	-328137.72	-2210.280			
19516155.2								
13	1.000	1	4.000	20	6.000	3	-4.000	
1075								
Th+2	210	2.000	2.000	.000	.000	.000	.000	
-109.071	170.300	226.627	.040869947	-47591.17	-77.397			
288991.6								
41	1.000	2	2.000					
1076								
Th+3	210	3.000	3.000	.000	.000	.000	.000	
-63.258	98.700	-252.869	-.038257331	-5884.17	94.546			
-1172965.8								
41	1.000	2	1.000					
1100								
U+2	210	2.000	2.000	.000	.000	.000	.000	
-57.065	98.400	145.999	.026809334	-27966.70	-47.805			
91846.8								
42	1.000	2	2.000					
1101								
U+3	210	3.000	3.000	.000	.000	.000	.000	
-9.602	24.400	-209.792	-.031242733	8072.53	78.494			
-1048770.8								
42	1.000	2	1.000					
1102								
UOH+2	410	2.000	3.000	.000	.000	.000	1.000	
-15.787	41.917	26.700	.006081412	-6457.05	-6.425			
-599625.4								
42	1.000	3	1.000	2	1.000	1	-1.000	
1103								
UOH2+	410	1.000	3.000	.000	.000	.000	2.000	
-22.311	55.717	-462.825	-.072635741	18599.63	172.365			
-2375169.5								
42	1.000	3	2.000	2	1.000	1	-2.000	
1104								
UOH3	410	.000	3.000	.000	.000	.000	3.000	
-30.805	72.834	-1180.441	-.192069636	56299.60	434.517			
-5076801.9								
42	1.000	3	3.000	2	1.000	1	-3.000	
1105								
UOH4-	410	-1.000	3.000	.000	.000	.000	4.000	
-41.287	84.334	-1077.185	-.177428469	46716.58	397.355			
-4544090.4								
42	1.000	3	4.000	2	1.000	1	-4.000	
1106								
UOH+3	310	3.000	4.000	.000	.000	.000	1.000	
-.541	11.217	379.997	.059338534	-21825.06	-136.117			
1047519.1								
42	1.000	3	1.000	1	-1.000			
1107								
UOH2+2	310	2.000	4.000	.000	.000	.000	2.000	
-2.007	17.517	-6.700	-.001285081	-305.81	4.530			
-453944.4								
42	1.000	3	2.000	1	-2.000			
1108								
UOH3+	310	1.000	4.000	.000	.000	.000	3.000	
-5.003	23.134	586.501	.093017064	-32834.64	-211.604			
1288275.0								
42	1.000	3	3.000	1	-3.000			
1109								
UOH4	310	.000	4.000	.000	.000	.000	4.000	
-4.563	18.234	653.911	.106053771	-36035.34	-237.333			
1603178.3								

42	1.000	3	4.000	1	-4.000				
1110									
UOH5-	310	-1.000	4.000	.000	.000	.000	.000	5.000	
-16.576	25.050	1001.671	.151691166	-60056.11	-362.685				
3146386.8									
42	1.000	3	5.000	1	-5.000				
1111									
UO2+	410	1.000	5.000	.000	.000	.000	.000	.000	
-7.576	32.934	227.164	.035074896	-13864.38	-80.552				
55431.6									
42	1.000	3	2.000	2	-1.000	1	-4.000		
1112									
UO2OH	410	.000	5.000	.000	.000	.000	.000	1.000	
-25.738	50.350	794.224	.124922917	-49435.17	-288.371				
1969210.6									
42	1.000	3	3.000	2	-1.000	1	-5.000		
1113									
UO2OH2-	410	-1.000	5.000	.000	.000	.000	.000	2.000	
-44.063	72.650	398.946	.056001524	-34140.25	-143.222				
817387.5									
42	1.000	3	4.000	2	-1.000	1	-6.000		
1114									
UO2+2	410	2.000	6.000	.000	.000	.000	.000	.000	
-9.049	34.384	606.976	.092741480	-35609.11	-218.681				
1499467.2									
42	1.000	3	2.000	2	-2.000	1	-4.000		
1115									
UO2OH+	410	1.000	6.000	.000	.000	.000	.000	1.000	
-14.267	44.750	502.746	.074921403	-31206.93	-180.266				
1011014.0									
42	1.000	3	3.000	2	-2.000	1	-5.000		
1116									
UO2OH2	410	.000	6.000	.000	.000	.000	.000	2.000	
-19.361	46.650	448.930	.068968459	-28298.55	-162.731				
775796.2									
42	1.000	3	4.000	2	-2.000	1	-6.000		
1117									
UO2OH3-	410	-1.000	6.000	.000	.000	.000	.000	3.000	
-28.295	51.667	626.678	.088689559	-41071.67	-226.717				
1541177.4									
42	1.000	3	5.000	2	-2.000	1	-7.000		
1118									
UO2OH4-2	410	-2.000	6.000	.000	.000	.000	.000	4.000	
-42.075	68.367	-245.020	-.055519263	3375.35	90.568				
-1415944.2									
42	1.000	3	6.000	2	-2.000	1	-8.000		
1150									
ZrOH+3	310	3.000	.000	.000	.000	.000	.000	1.000	
.324	6.147	656.392	.103728549	-37020.84	-237.085				
2117953.6									
43	1.000	3	1.000	1	-1.000				
1151									
ZrOH2+2	310	2.000	.000	.000	.000	.000	.000	2.000	
-1.728	14.347	264.855	.042403809	-15352.91	-94.571				
558087.9									
43	1.000	3	2.000	1	-2.000				
1152									
ZrOH3+	310	1.000	.000	.000	.000	.000	.000	3.000	
-5.091	21.664	1410.055	.225286719	-79446.00	-511.653				
4462372.5									
43	1.000	3	3.000	1	-3.000				
1153									

ZrOH4	310	.000	.000	.000	.000	.000	.000	4.000	
-9.709	23.164	1191.054	.192686380	-67348.64	-433.022				
3481201.9									
43	1.000	3	4.000	1	-4.000				
1154									
ZrOH5-	310	-1.000	.000	.000	.000	.000	.000	5.000	
-16.004	21.980	934.554	.141916041	-56184.39	-337.646				
2760831.4									
43	1.000	3	5.000	1	-5.000				
LOOK MIN									
Akerman	5	.000	45.319	-68.970	1				
1 -6.0	3	-1.0	4 2.0	5 1.0	13 2.0				
-7479.300	-1.127733552	427046.53	2707.535	-24097829.2					
Alaband	3	.000	.045	-5.750	1				
9 1.0	106 1.0	1 -1.0							
-990.802	-1.164530798	53531.70	361.330	-2998374.8					
Albite	4	.000	-20.119	30.843	1				
6 1.0	303 1.0	13 3.0	3 -8.0						
-10348.331	-1.555244968	567564.23	3746.228	-33913353.5					
Alb-high	4	.000	-18.800	28.213	1				
6 1.0	303 1.0	13 3.0	3 -8.0						
-10436.474	-1.569454292	572661.25	3778.291	-34156627.7					
Alunite	5	12.000	-.348	-55.414	1				
10 3.0	3 6.0	1 -6.0	7 1.0	16 2.0					
-5507.427	-.916819492	300187.74	2003.800	-16416616.0					
Analcime	5	.000	6.948	-20.421	1				
1 -4.0	3 -1.0	6 1.0	10 1.0	13 2.0					
-7446.063	-1.119406059	415245.43	2693.508	-24080007.2					
Andalus	4	.000	15.944	-56.221	1				
1 -6.0	3 1.0	10 2.0	13 1.0						
-4851.444	-.747867671	274825.20	1753.752	-15196130.6					
Anglesit	2	6.000	-7.853	2.690	1				
32 1.0	16 1.0								
-1757.967	-.290646263	94363.84	640.089	-5652288.5					
Anhydrit	2	6.000	-4.306	-4.440	1				
4 1.0	16 1.0								
-1732.6926	-0.29047750	93359.2040	631.641790	-5430711.9					
Annite	5	6.000	29.469	-62.132	1				
1 -10.0	7 1.00	8 3.00	10 1.00	13 3.0					
-11924.617	-1.804025576	668489.65	4316.082	-38226809.4					
Anorth	4	.000	-19.189	14.044	1				
4 1.0	303 2.0	13 2.0	3 -8.0						
-7859.229	-1.202238613	429124.33	2847.113	-25405419.4					
Anthophy	4	.000	66.796	-115.554	1				
1 -14.0	3 -8.0	5 7.0	13 8.0						
-29526.408	-4.431294078	1655352.19	10683.925	-95503064.9					
Antig	4	.000	477.194	-804.099	1				
1 -96.0	5 48.0	13 34.0	3 11.0						
-135487.825	-20.36302000	7617232.16	49043.019	-432537669.9					
Aragonit	2	4.000	-8.336	-2.654	1				
15 1.0	4 1.0								
-1875.816	-.316393683	99680.75	683.927	-5764723.1					
Artinite	4	4.000	9.327	-27.660	1				
1 -2.0	5 2.0	15 1.0	3 5.0						
-2478.304	-.399899523	136957.56	901.577	-7412746.9					
Barite	2	6.000	-9.971	6.200	1				
11 1.0	16 1.0								
-1846.064	-.310722479	97011.28	674.192	-5768118.7					
Na-Beid	5	.000	5.647	-37.253	1				
3 -2.68	1 -7.32	6 0.33	10 2.33	13 3.67					
-13722.572	-2.073105312	764018.99	4963.790	-44339021.5					

K-Beid	5	.000	5.309	-36.055	1		
3 -2.68	1	-7.32	7 0.33	10 2.33	13 3.67		
-13794.274-2.086296955			767361.09	4990.575	-44533824.3		
Ca-Beid	5	.000	5.591	-38.820	1		
3 -2.68	1	-7.32	4 0.165	10 2.33	13 3.67		
-13676.962-2.067368826			761750.04	4947.070	-44196302.6		
Mg-Beid	5	.000	5.554	-39.550	1		
3 -2.68	1	-7.32	5 0.165	10 2.33	13 3.67		
-13675.311-2.065894557			761987.31	4945.951	-44210057.9		
Berndtit	5	-2.000	-92.228	146.334	1		
3 -8.0	39 1.0		16 2.0	1 16.0	2 14.0		
-3338.034 -.577782924			162045.35	1212.010	-11064594.2		
Boehmite	3	.000	7.564	-27.075	1		
1 -3.0	10 1.0		3 2.0				
-885.683 -.143391305			51203.55	320.101	-2471843.0		
Brucite	3	.000	16.298	-26.611	1		
1 -2.0	5 1.0		3 2.0				
-452.893 -.06887724			29129.29	163.991	-1223060.0		
Bunsenit	3	.000	12.472	-23.917	1		
1 -2.0	29 1.0		3 1.0				
-666.525 -.106871087			39719.19	242.020	-1886279.7		
CaAlpyrx	5	.000	35.976	-86.411	1		
1 -8.0	3 2.0		4 1.0	10 2.0	13 1.00		
-5136.742 -.793286373			296156.66	1857.336	-15994942.7		
Calcite	2	4.000	-8.480	-2.633	1		
15 1.0	4 1.0		85519.84	582.500	-5010794.3		
-1599.510 -.270987115			-3.163	.066	1		
Cassiter	4	2.000	39 1.0	3 2.0			
1 -4.0	2 -2.0		32909.93	243.842	-1880950.7		
-664.775 -.103903817			7.457	-17.786	1		
Celadon	6	.000	5 1.0	7 1.0	10 1.0		
1 -6.0	3 -4.0						
13 4.0			795968.52	5185.065	-46494406.6		
-14328.054-2.150756077			-5.677	-1.770	1		
Celestite	2	6.000					
12 1.0	16 1.0		98130.55	662.153	-5787958.5		
-1818.393 -.301055879			-13.538	6.835	1		
Cerussit	2	4.000	97339.27	673.154	-5828973.4		
15 1.0	32 1.0		-3.728	7.507	1		
-1848.113 -.308591128			172688.44	1129.990	-10347838.8		
Chalcedy	2	.000	32.842	-87.051	1		
13 1.0	3 -2.0		10 2.0	13 1.0	3 5.0		
-3123.761 -.465693023			332564.88	2112.780	-17912014.6		
Chm-7A	5	4.000	-87.115	43.063	1		
1 -10.0	8 2.0		303 2.0	13 3.0	100 8.0		
-5840.044 -.902158261			1077279.97	7250.236	-63592949.0		
Chlorite	5	.000	31.125	-52.112	1		
3 -10.0	5 5.0		13 2.0	5 3.0			
-19991.001-3.133247435			455736.100	2931.70270	-25791661.0		
Chrysotl	4	.000	70.612	-150.126	1		
3 1.0	1 -6.0		10 2.0	13 3.0	3 6.0		
-8098.4599-1.21792980			798424.08	5097.978	-44108901.7		
Cnc-7A	5	.000	67.239	-146.359	1		
1 -16.0	5 5.0		10 2.0	13 3.0	3 6.0		
-14087.645-2.140140582			801753.94	5124.825	-44353020.9		
Cnc-14A	5	.000	43.257	-109.404	1		
1 -16.0	5 5.0		10 3.0	13 3.0	3 1.0		
-14162.235-2.152335846			714177.92	4573.839	-40177450.4		
Czo	5	.000	52.303	-149.668	1		
1 -13.0	4 2.0						
-12642.665-1.929004296							
Cordier	5	.000					

1 -16.0	3 -2.00	5 2.00	10 4.00	13 5.0
-20252.193-3.073237500		1142309.32	7322.248	-64800136.6
Corundum	3	.000	18.312	-61.812
1 -6.0	10 2.0		3 3.0	
-1807.728 -.294292276		106064.49	652.820	-5095360.1
a-Crs	2	.000	-3.449	6.980
13 1.0	3 -2.0			
-3109.751 -.463608807		172039.14	1124.875	-10304943.3
b-Crs	2	.000	-3.005	5.900
13 1.0	3 -2.0			
-3223.421 -.482555307		177782.61	1166.693	-10569625.9
Dph-7A	5	10.000	55.655	-127.234
1 -16.0	8 5.0		10 2.0	13 3.0
-13593.925-2.072201691		767343.59	4920.092	-42732843.3
Dph-14A	5	10.000	50.365	-120.147
1 -16.0	8 5.0		10 2.0	13 3.0
-13619.951-2.076289129		767594.51	4929.156	-42849750.2
Diaspore	3	.000	7.160	-26.391
1 -3.0	10 1.0		3 2.0	
-980.072 -.159069041		55862.29	354.813	-2726112.4
Diopside	5	.000	20.964	-31.973
1 -4.0	3 -2.0		4 1.0	5 1.0
-7122.056-1.070227146		400510.90	2577.775	-23089799.4
Dolomite	3	8.000	-18.144	-7.306
4 1.0	5 1.0		15 2.0	
-3657.993 -.612488930		195522.75	1331.265	-11329561.5
Ord-Dol	3	8.000	-18.144	-7.306
4 1.0	5 1.0		15 2.0	
-3677.022 -.615681181		196492.78	1338.260	-11381374.3
Dis-Dol	3	8.000	-16.600	-10.233
4 1.0	5 1.0		15 2.0	
-3580.249 -.599514810		192169.62	1302.489	-11117943.9
Enstatit	4	.000	11.327	-19.773
1 -2.0	3 -1.0		5 1.0	13 1.00
-3818.110 -.574363172		214630.23	1381.763	-12291481.2
Epidote	6	3.000	32.930	-92.362
1 -13.0	4 2.0		514 1.0	10 2.00
3 1.0				13 3.0
-12602.342-1.922159744		707991.75	4560.056	-39985393.7
Ord-ep	6	3.000	32.930	-92.338
1 -13.0	4 2.0		514 1.0	10 2.00
3 1.0				13 3.0
-12691.663-1.935137259		712918.44	4592.331	-40269334.2
Fayalite	3	4.000	19.111	-36.390
1 -4.0	8 2.0		13 1.0	
-4152.095 -.629499494		235235.02	1502.800	-13215977.6
K-fs	5	.000	-4.488	-5.486
1 -4.0	3 -4.0		7 1.0	10 1.0
-10567.643-1.592335431		585341.68	3825.157	-34348101.0
Feo	3	2.000	13.532	-25.347
1 -2.0	8 1.0		3 1.0	
-364.484 -.056961006		24454.74	131.220	-1041733.9
Fluorite	2	.000	-10.037	2.900
4 1.0	20 2.0			
-1722.337 -.288926229		90264.70	628.331	-5250972.6
Forsteri	3	0.000	19.111	-36.390
1 -4.0	5 2.0		13 1.0	
-4152.095 -.629499494		235235.02	1502.800	-13215977.6
Galena	5	-2.000	-48.544	79.587
3 -4.0	32 1.0		16 1.0	1 8.0
-1759.170 -.302652609		85853.89	638.116	-5873232.6
Gehlenit	5	.000	56.300	-117.096

1 -10.0	3 3.0	4 2.0	10 2.0	13 1.0
-5452.296	-.844960385	318834.33	1972.405	-16839986.2
Gibbsite 3	.000	7.756	-24.566	1
10 1.0	3 3.0	1 -3.0		
-1035.632	-.167985591	58000.08	376.073	-2811524.2
Graphite 4	.000	-32.159	43.565	1
3 -3.0	15 1.0	1 6.0	2 4.0	
-1279.976	-.223805226	63446.79	464.655	-4268148.1
Grn 4	6.000	22.670	-39.510	1
1 -6.0	8 3.0	13 2.0	3 1.0	
-7780.486	-1.173213917	436236.99	2816.853	-24918143.6
Halite 2	.000	1.585	.894	1
6 1.0	14 1.0			
-661.998	-.110080533	35472.55	242.853	-2088650.8
Halloysite 4	.000	8.266	-38.201	1
1 -6.0	3 1.0	13 2.0	10 2.0	
-8053.662	-1.224137760	448994.88	2913.681	-25667467.9
Hd 5	2.000	19.606	-29.758	1
1 -4.0	3 -2.0	4 1.0	8 1.0	
-7126.053	-1.073940699	399968.50	2580.012	-23087063.4
Hematite 3	6.000	.109	-30.930	1
514 2.0	3 3.0	1 -6.0		
-1436.588	-.230339374	80055.00	517.950	-3979539.1
Huntite 3	16.000	-31.014	-26.841	1
4 1.0	5 3.0	15 4.0		
-7439.772	-1.239652579	400664.70	2705.258	-23064297.0
Hydr-mgs 4	16.000	-10.461	-55.186	1
1 -2.0	5 5.0	15 4.0	3 6.0	
-8331.097	-1.368707717	453917.27	3028.087	-25470794.2
Illite 6	.000	-43.606	58.384	1
1 1.2	7 0.6	5 0.25	303 2.3	
-12794.603	-1.939187958	694687.37	4634.697	-41698245.6
Fe(c) 2	.000	16.031	-22.050	1
8 1.0	2 2.0			
-475.935	-.077889507	31501.05	171.825	-1389856.8
Jadeite 5	.000	8.389	-20.182	1
1 -4.0	3 -2.0	6 1.0	10 1.0	
-7574.843	-1.143280623	421692.17	2741.870	-24431462.4
Kaolinit 4	.000	6.810	-36.275	1
1 -6.0	3 1.0	13 2.0	10 2.0	
-8286.561	-1.262845831	460406.93	2999.259	-26294093.0
Kyanite 4	.000	15.674	-55.190	1
1 -6.0	3 1.0	10 2.0	13 1.0	
-4743.347	-.729179993	269292.11	1713.973	-14924987.2
Lmt 4	.000	13.667	-44.132	1
1 -8.0	4 1.0	10 2.0	13 4.0	
-14714.624	-2.217489144	820687.64	5323.100	-47542555.1
Lawsonite 5	.000	22.213	-58.509	1
1 -8.0	4 1.0	10 2.0	13 2.0	
-8332.911	-1.268250104	468146.76	3015.253	-26487073.6
Lime 3	.000	18.581	-32.987	1
1 -1.0	4 1.0	100 1.0		
-1178.684	-.195360728	69659.25	429.388	-3610840.6
Magnesit 2	4.000	-8.035	-7.122	1
5 1.0	15 1.0			
-2074.278	-.344113358	111341.02	754.505	-6362306.1
Magnetit 4	8.000	10.472	-51.767	1
1 -8.0	514 2.0	8 1.0	3 4.0	
-2027.628	-.324745692	115078.19	732.416	-5632950.5
Mng 3	2.000	4.369	-15.807	1
1 -1.0	9 1.0	100 1.0		

-1135.067	-.185282737	63919.91	412.142	-3513780.6
Margarit 5	.000	41.066	-124.805	1
1 -14.0	4 1.0	10 4.0	13 2.0	
-10063.791	-1.551460700	570971.57	3639.291	-31362453.8
Merwinit 4	.000	68.514	-102.788	1
1 -8.0	4 3.0	5 1.0	13 2.00	
-7887.939	-1.195224658	455128.87	2857.088	-25190378.6
Max-mc 5	.000	-.448	-5.486	1
1 -4.0	3 -4.0	7 1.0	10 1.0	
-10609.943	-1.597968871	587973.57	3840.169	-34525302.1
Mnn 4	6.000	13.981	-25.151	1
1 -6.0	3 -4.0	8 3.0	13 4.0	
-14021.559	-2.102326013	782168.43	5073.166	-45715584.5
Mtc 4	.000	29.585	-46.776	1
1 -4.0	4 1.0	5 1.0	13 1.0	
-4134.594	-.626378004	236613.20	1497.400	-13146403.8
Na-Mont 6	.000	2.484	-22.284	1
3 -4.0	1 -6.0	6 0.33	5 0.33	
13 4.0				
-14342.806	-2.158849015	797303.81	5188.598	-46585806.2
K-Mont 6	.000	2.142	-21.082	1
3 -4.0	1 -6.0	7 0.33	5 0.33	
13 4.0				
-14334.449	-2.158810234	796739.44	5185.874	-46592829.1
Ca-Mont 6	.000	2.495	-23.942	1
3 -4.0	1 -6.0	4 0.165	5 0.33	
13 4.0				
-14345.021	-2.160786176	797538.22	5189.374	-46577188.5
Mg-Mont 5	.000	2.388	-24.529	1
3 -4.0	1 -6.0	5 0.495	10 1.67	
-14294.860	-2.151283400	795281.01	5170.446	-46461250.8
Muscovit 4	.000	14.050	-57.895	1
10 3.0	13 3.0	7 1.0	1 -10.0	
-12152.686	-1.847242051	678653.92	4396.344	-38863206.8
Nephelin 4	.000	13.801	-32.282	1
1 -4.0	6 1.0	10 1.0	13 1.0	
-4370.329	-.664264106	245846.09	1581.632	-13870692.1
Nesqueh 2	4.000	-5.333	-5.126	1
5 1.0	15 1.0	3 3.0		
49851.470	7.653242489	-2688423.62	-18120.452	152568176.4
Ni(c) 2	-2.000	7.990	-12.900	1
29 1.0	2 2.0			
-421.178	-.068444742	26903.17	151.035	-1278704.3
Na-Mont 6	6.000	-11.526	-7.550	1
3 -2.68	1 -7.32	6 0.33	514 2.00	
13 3.67				
-13454.443	-2.025559977	743668.89	4866.612	-43517908.0
K-Mont 6	6.000	-11.865	-6.358	1
3 -2.68	1 -7.32	7 0.33	514 2.00	
13 3.67				
-13467.076	-2.029747569	743836.79	4871.933	-43534605.6
Ca-Mont 6	6.000	-11.582	-9.120	1
3 -2.68	1 -7.32	4 0.165	514 2.00	
13 3.67				
-13507.124	-2.036717794	746289.43	4886.171	-43627645.0
Mg-Mont 6	6.000	-11.620	-9.847	1
3 -2.68	1 -7.32	5 0.165	514 2.00	
13 3.67				
-13573.212	-2.045849257	750111.05	4909.747	-43839568.1
Paragoni 4	.000	17.522	-65.739	1
1 -10.0	6 1.0	10 3.0	13 3.0	
-12224.958	-1.856130652	683757.75	4421.870	-39030905.0

Pargasit	6	.000	101.994	-210.371	1			
1 -22.0		6 1.0	4 2.0	5 4.0	10 3.0			
13 6.0								
-24800.265	-3.757500513	1403655.86	8973.054	-78987142.2				
Pd(c)	2	-2.000	-30.933	42.080	1			
33 1.0	2 2.0							
-391.307	-.064011947	13062.76	140.850	-1144809.2				
PdO	3	.000	.064	-5.837	1			
1 -2.0	33 1.0	3 1.0						
-557.284	-.088194330	29888.63	202.315	-1530740.2				
Periclas	3	.000	7.340	-22.544	1			
1 -1.0	5 1.0	100 1.0						
-1297.046	-.209430861	73987.99	470.249	-3993898.9				
Phlogopi	5	.000	37.267	-73.975	1			
1 -10.0	7 1.0	5 3.0	10 1.0	13 3.0				
-12259.827	-1.850841351	688872.78	4436.987	-39166474.0				
Prehnite	4	.000	32.931	-74.539	1			
1 -10.0	4 2.0	10 2.0	13 3.0					
-11829.504	-1.796157920	665091.71	4281.148	-37885439.1				
Pyrite	4	.000	-16.231	11.250	1			
1 -2.0	2 -2.0	8 1.0	106 2.0					
-1548.858	-.259099104	77697.68	566.855	-4743914.2				
Pyrophyll	4	.000	.440	-24.416	1			
1 -6.0	3 -4.0	10 2.0	13 4.0					
-14331.760	-2.159302294	796991.87	5183.719	-46571317.3				
Pyrrhot	3	.000	-3.719	-1.900	1			
8 1.0	106 1.0	1 -1.0						
-1025.621	-.172010922	54305.77	374.152	-3090835.2				
Quartz	2	.000	-3.999	7.875	1			
13 1.0	3 -2.0							
-3119.839	-.464968292	172433.11	1128.522	-10340602.0				
Rhodochr	2	6.000	-10.082	-1.764	1			
9 1.0	15 1.0							
-1697.910	-.284421270	90619.33	617.635	-5298772.0				
Hi-sa	5	.000	.751	-8.136	1			
1 -4.0	3 -4.0	7 1.0	10 1.0	13 3.0				
-10603.862	-1.596705488	588225.89	3837.627	-34508836.6				
Sepiolit	4	.000	15.222	-18.801	1			
5 2.0	13 3.0	1 -4.0	3 -0.5					
-10839.905	-1.622706355	603023.20	3924.346	-35037234.6				
Siderite	2	6.000	-10.521	-4.262	1			
8 1.0	15 1.0							
-1965.635	-.328101570	104973.89	714.917	-6059054.5				
Silliman	4	.000	16.308	-56.988	1			
1 -6.0	3 1.0	10 2.0	13 1.0					
-4987.000	-.769270055	282118.80	1803.157	-15588136.2				
Sn(c)	2	-2.000	4.816	-2.100	1			
39 1.0	2 2.0							
-467.216	-.075720505	27571.64	169.223	-1475424.2				
SnO	3	.000	1.318	-2.077	1			
1 -2.0	39 1.0	3 1.0						
-332.250	-.050717303	18657.56	120.450	-1060842.3				
SnS	5	-2.000	-49.312	79.231	1			
3 -4.0	39 1.0	16 1.0	1 8.0	2 8.0				
-1937.432	-.332181853	95083.09	703.165	-6372481.6				
Strontit	2	4.000	-10.643	1.545	1			
12 1.0	15 1.0							
-1992.808	-.333033249	105442.71	726.315	-6170664.1				
Sylvite	2	.000	.846	4.167	1			
7 1.0	14 1.0							
-590.444	-.100416932	31258.87	217.463	-1930064.5				
Talc	4	.000	-62.832	44.491	1			

3 -10.0	5 3.0	13 4.0	100 6.0					
-19416.602	-2.992242873	1056811.85	7035.287	-62848878.0				
Tremolit	5	.000	61.237	-97.131	1			
1 -14.0	3 -8.0	4 2.0	5 5.0	13 8.0				
-28763.041	-4.316868131	1610956.14	10409.407	-93268346.2				
Uraninit	3	4.000	-4.827	-18.614	1			
1 -4.0	42 1.0	3 2.0						
-522.960	-.084066499	28482.59	186.188	-1159366.4				
Wairakit	5	.000	18.076	-56.830	1			
1 -8.0	3 -2.0	4 1.0	10 2.0	13 4.0				
-14790.873	-2.233044351	827113.64	5349.918	-47775054.1				
Witherit	2	4.000	-13.325	7.615	1			
11 1.0	15 1.0							
-1596.724	-.269593183	83916.36	581.785	-5090742.0				
Wollast	4	.000	13.761	-18.302	1			
1 -2.0	3 -1.0	4 1.0	13 1.00					
-3496.501	-.526508819	197363.48	1266.189	-11362479.4				
Zoisite	5	.000	43.302	-109.494	1			
1 -13.0	4 2.0	10 3.0	13 3.0	3 1.0				
-12711.565	-1.940677793	717652.86	4599.215	-40357316.6				
PCO2	1	4.000	-1.469	-4.849	1			
104 1.0								
601.628	.098938252	-32950.42	-220.638	2122221.1				
O2 gas	1	4.000	-2.898	-2.900	1			
101 1.0								
340.129	.057594960	-19638.01	-125.175	1369242.5				
H2 gas	1	-2.000	-3.105	-1.000	1			
102 1.0								
323.711	.054983622	-18216.82	-119.338	1172092.5				
N2 gas	1	.000	-3.186	-2.495	1			
778 1.0								
662.955	.111292748	-36151.15	-243.796	2238739.5				
H2S gas	1	-2.000	-.988	-4.070	1			
107 1.0								
395.925	.068107112	-21036.57	-146.269	1357485.8				
CH4 gas	1	-4.000	-2.850	-3.130	1			
425 1.0								
644.123	.107080467	-35748.80	-236.314	2288661.1				
NH3 gas	1	-3.000	1.797	-8.419	1			
775 1.0								
-303.189	-.045603086	17285.59	108.962	-801252.8				

END

**Appendix C: Calculated and Experimental Equilibrium
 Constants of Mineral and Gas Solubility
 Reactions as Functions of Pressure and
 Temperature**

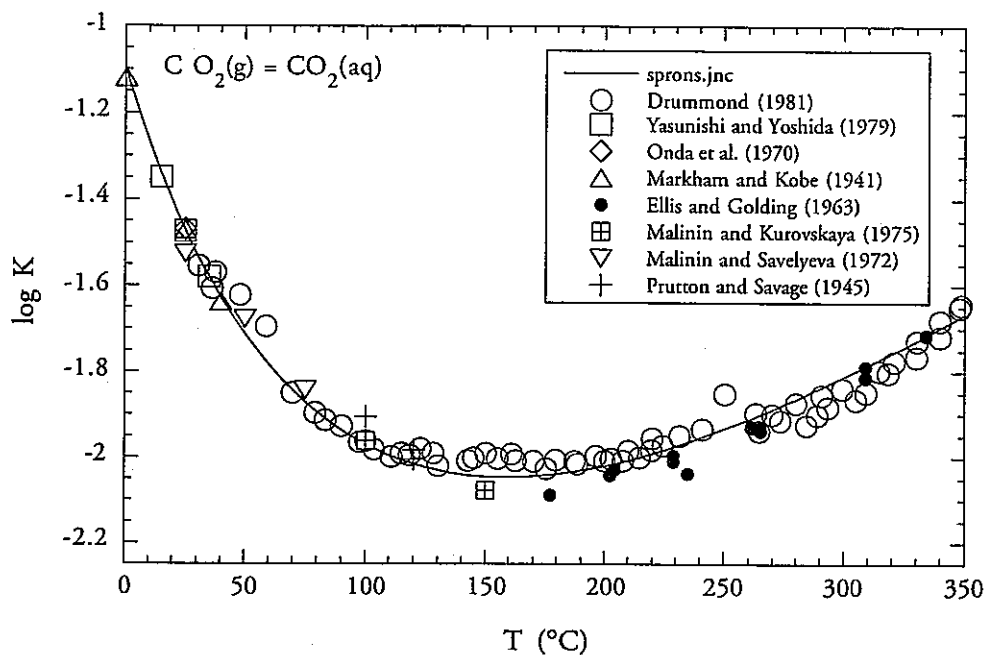


Fig. C.1: Logarithm of the equilibrium constant for the reaction $\text{CO}_2(\text{g}) \rightleftharpoons \text{CO}_2(\text{aq})$ between 0 and 350°C at P_{sat} .

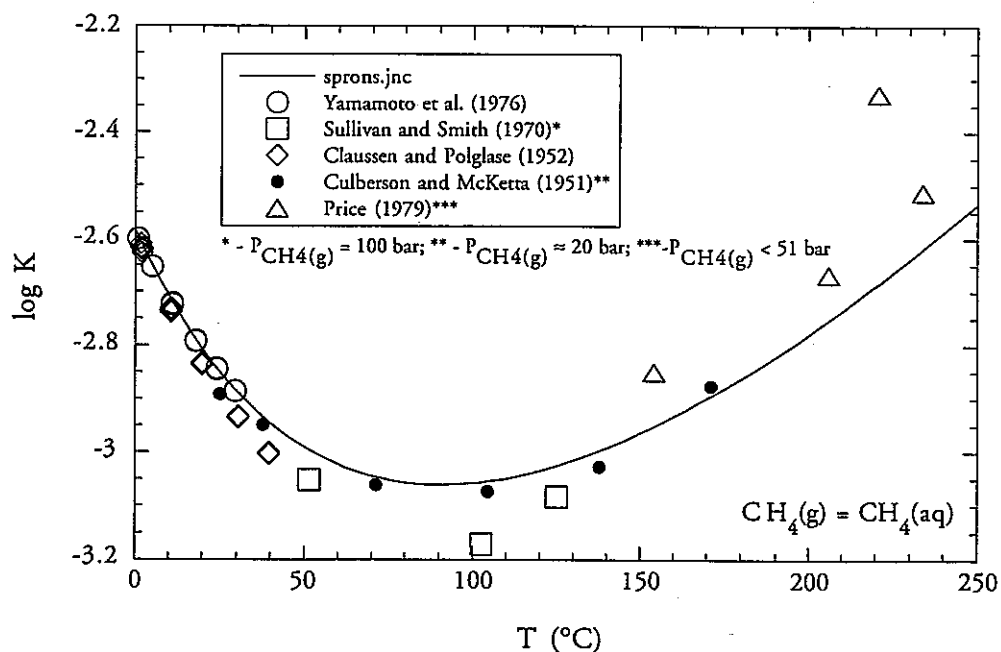


Fig. C.2: Logarithm of the equilibrium constant for the reaction $\text{CH}_4(\text{g}) \rightleftharpoons \text{CH}_4(\text{aq})$ between 0 and 250°C, and pressures from P_{sat} to 100 bars.

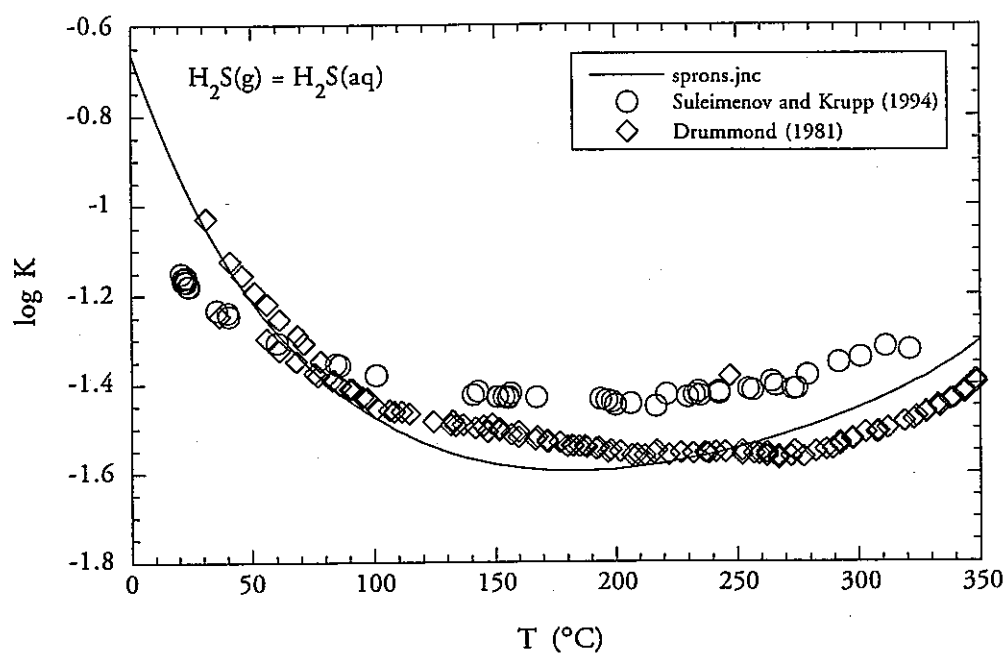


Fig. C.3: logarithm of the equilibrium constant for the reaction $\text{H}_2\text{S}(\text{g}) \rightleftharpoons \text{H}_2\text{S}(\text{aq})$ between 0 and 350°C at P_{sat} .

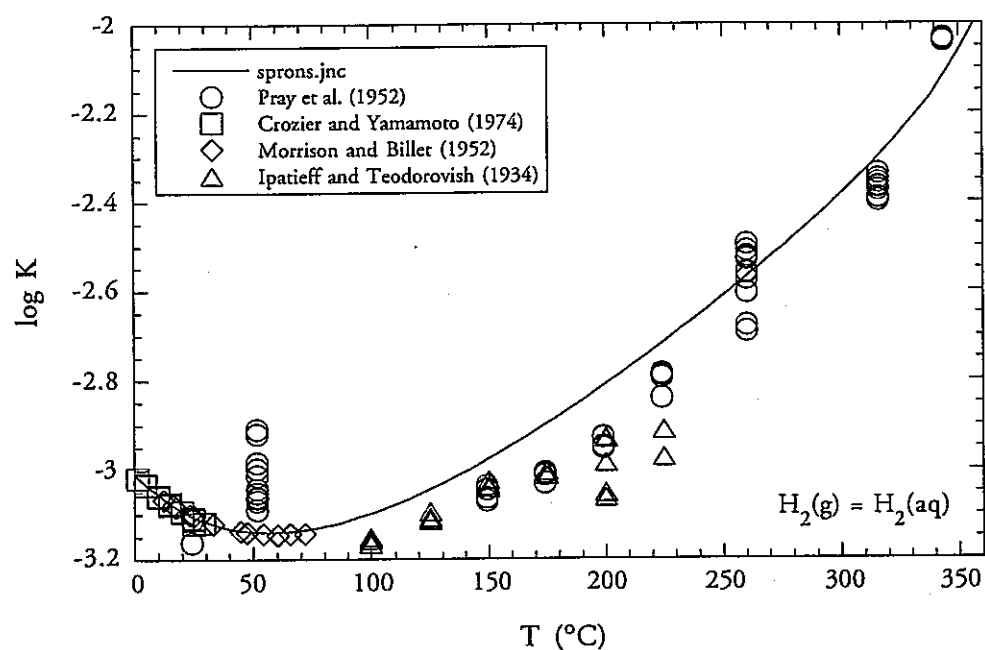


Fig. C.4: Logarithm of the equilibrium constant for the reaction $\text{H}_2(\text{g}) \rightleftharpoons \text{H}_2(\text{aq})$ between 0 and 350°C at P_{sat} .

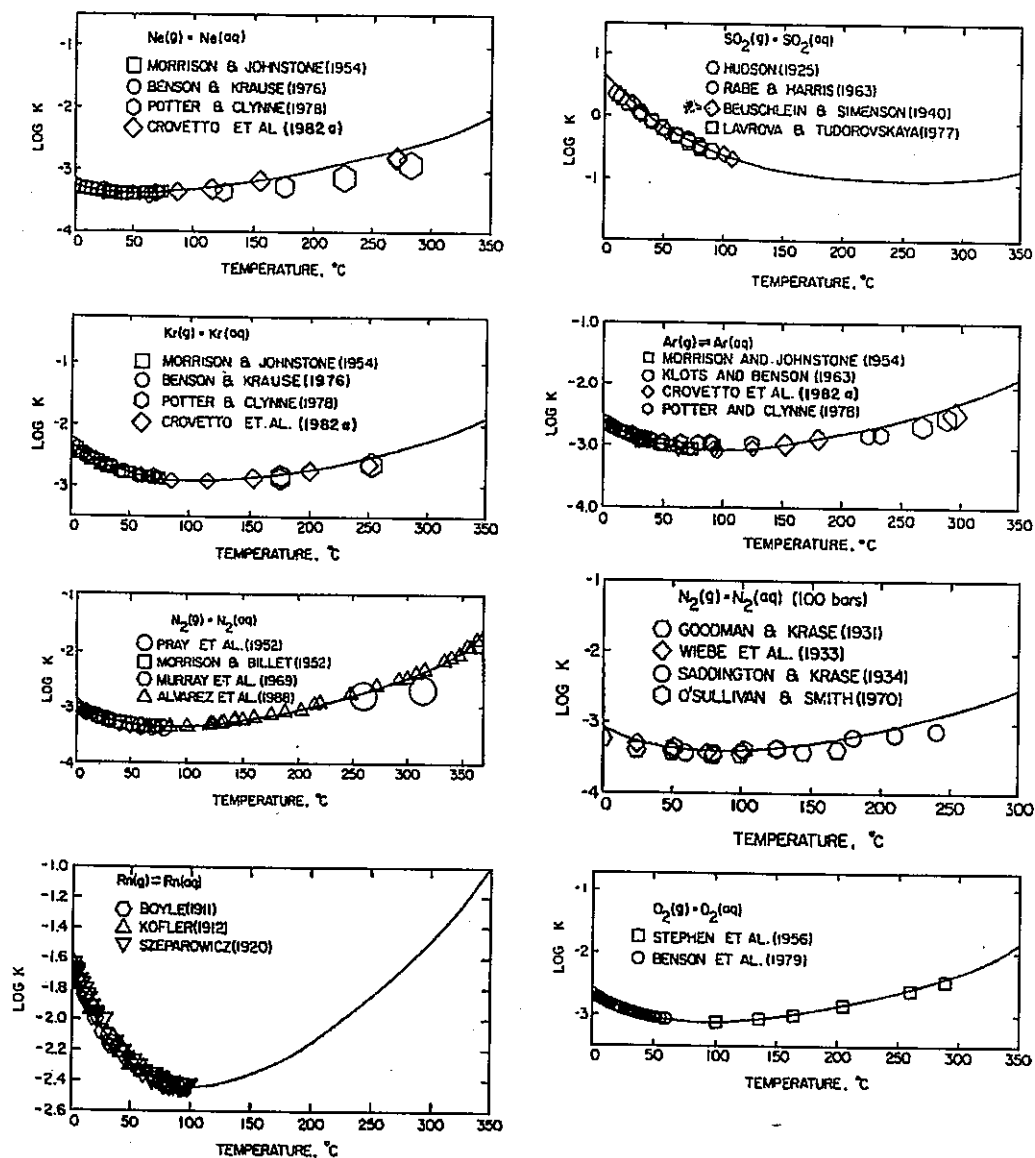


Fig. C.5: Logarithms of the equilibrium constants of gas solubility reactions as functions of temperature at P_{sat} [and at 100 bars for the reaction $N_2(g) \rightleftharpoons N_2(aq)$]. The figures were generated by Shock *et al.* (1989) using thermodynamic data for gas and aqueous species that are included in SPRONS.JNC.

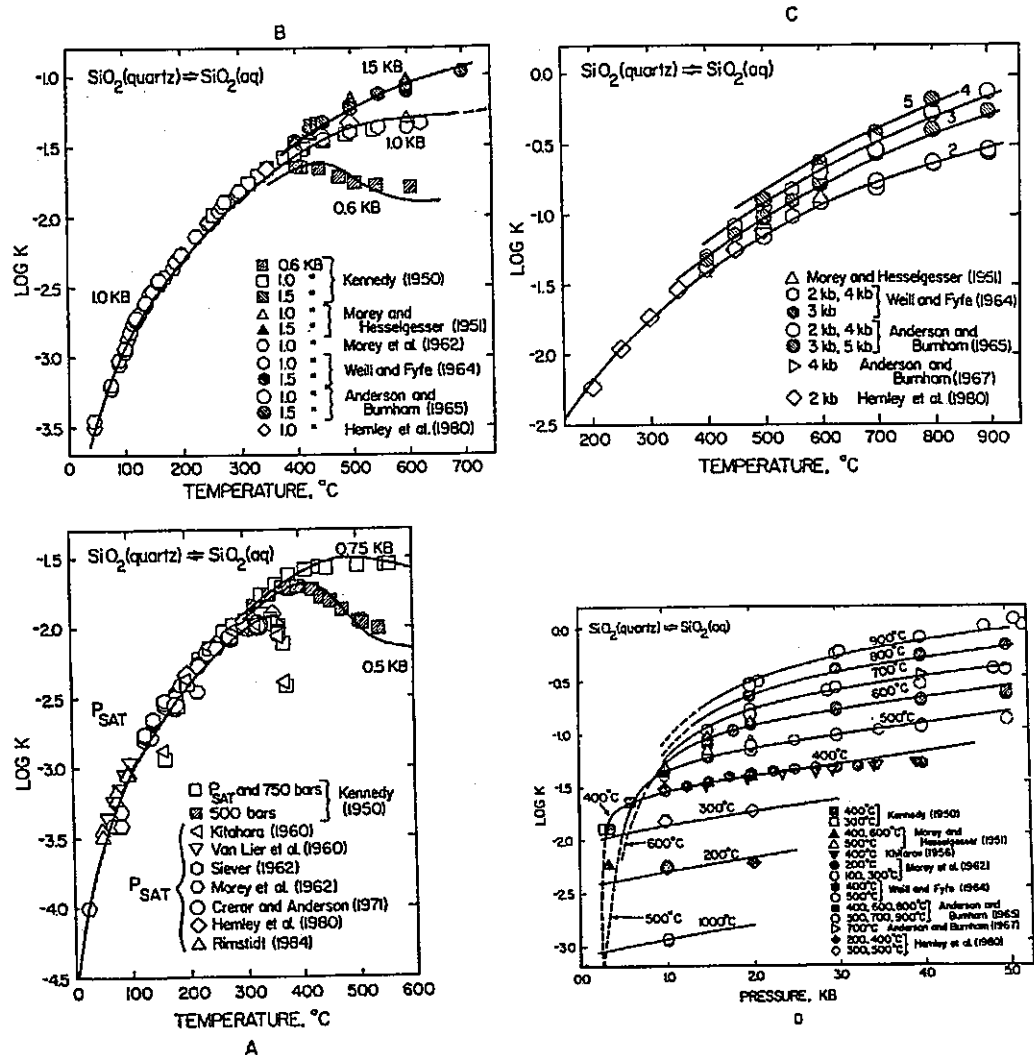


Fig. C.6: Logarithms of the equilibrium constant for the reaction $\text{SiO}_2(\text{quartz}) \rightleftharpoons \text{SiO}_2(\text{aq})$ as a function of temperature at P_{sat} and constant pressure (diagrams A, B and C), and pressure at constant temperature (diagram D) [from Shock *et al.*, 1989].

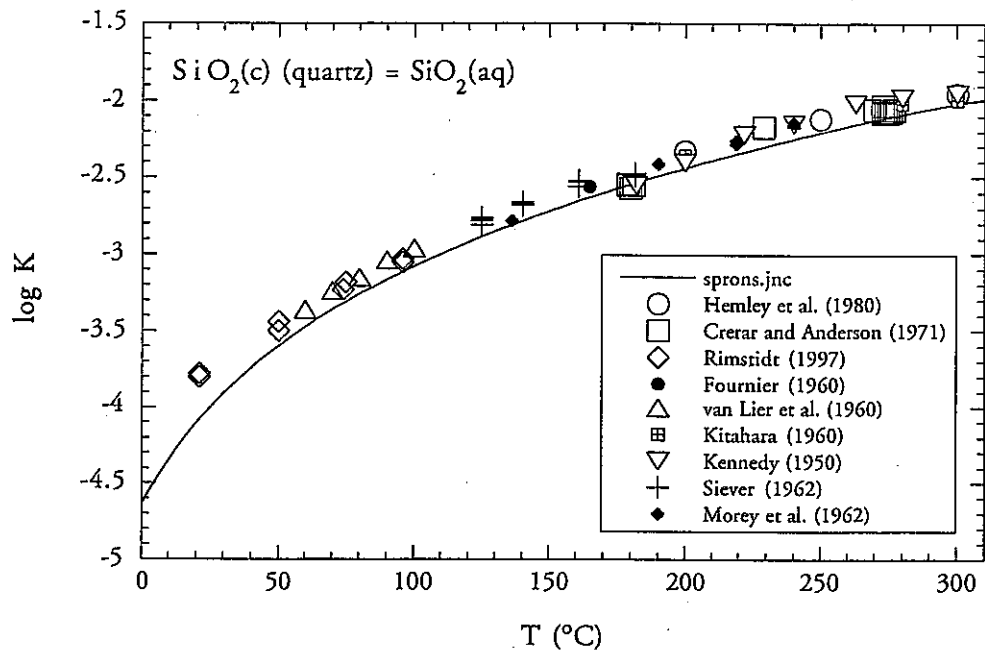


Fig. C.7: Calculated and experimental silica molalities ($\approx \log K$) in solutions equilibrated with quartz between 0 and 300°C at P_{sat} . Data from “unequilibrated” experiments by Fournier (1960) and Morey *et al.* (1962) at temperatures < 100°C not shown (see Fig. C.8).

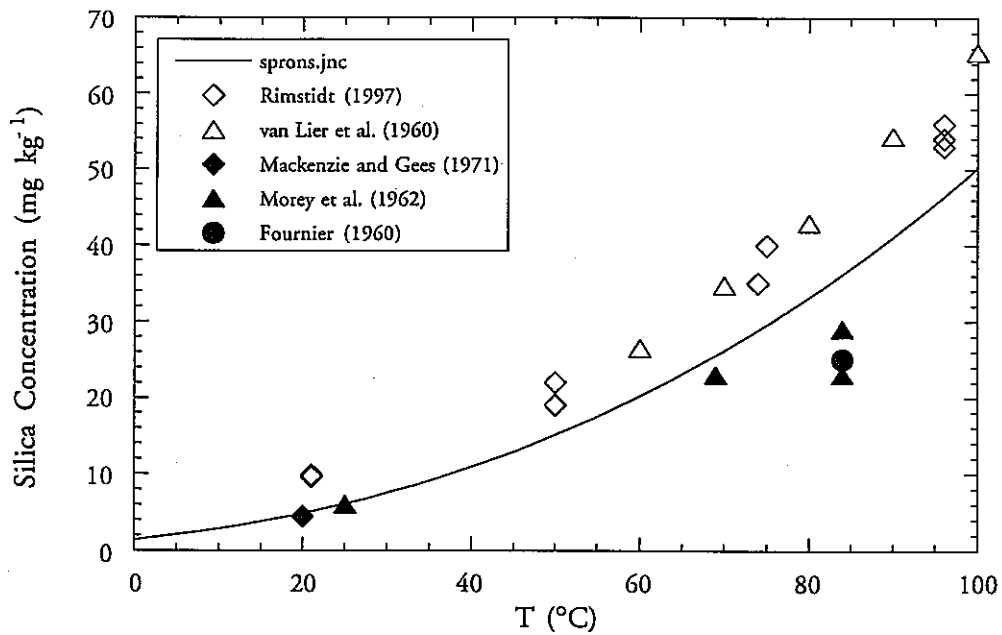


Fig. C.8: Calculated and experimental silica concentrations in solutions coexisting with quartz between 0 and 100°C at 1 bar. [open symbols denote equilibrated experiments; filled symbols refer to unequilibrated experiments (see text)].

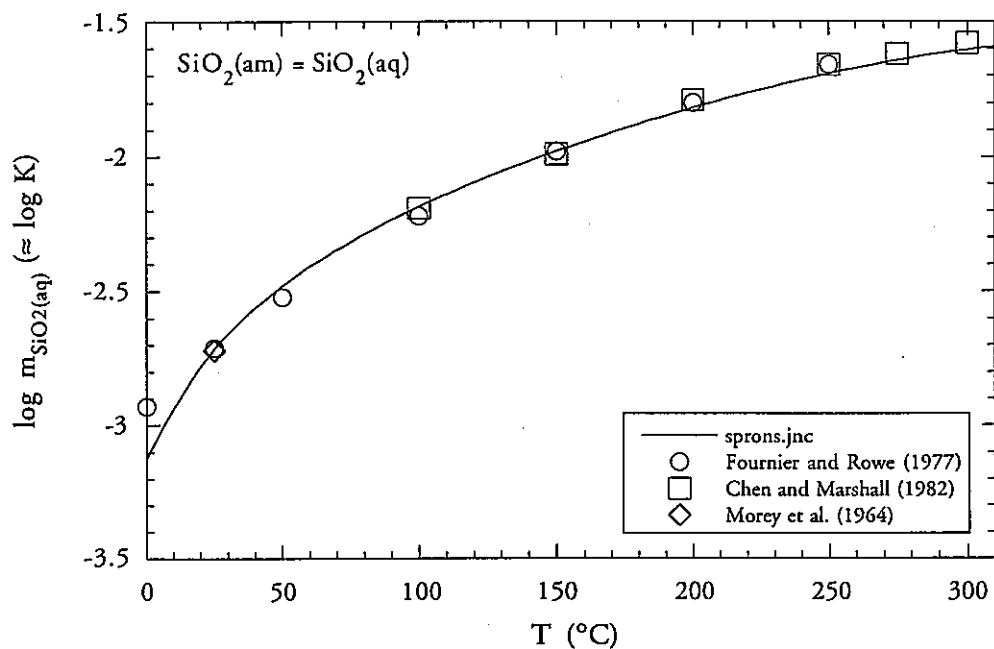


Fig. C.9: Calculated and experimental silica molalities ($\approx \log K$) in solutions equilibrated with amorphous silica between 0 and 300°C at P_{sat} .

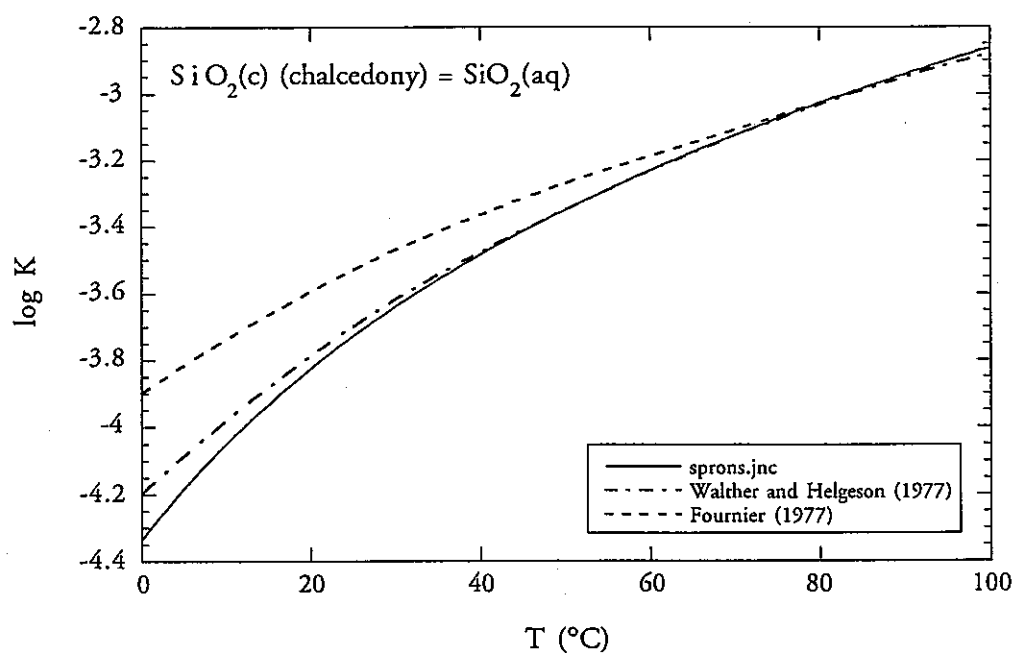


Fig. C.10: Solubility of chalcedony between 0 and 100°C at 1 bar calculated using temperature functions from Walther and Helgeson (1977) and Fournier (1977), which are based on experimental solubilities between 92 and 259°C determined by Fournier and Rowe (1966) and Fournier (1973).

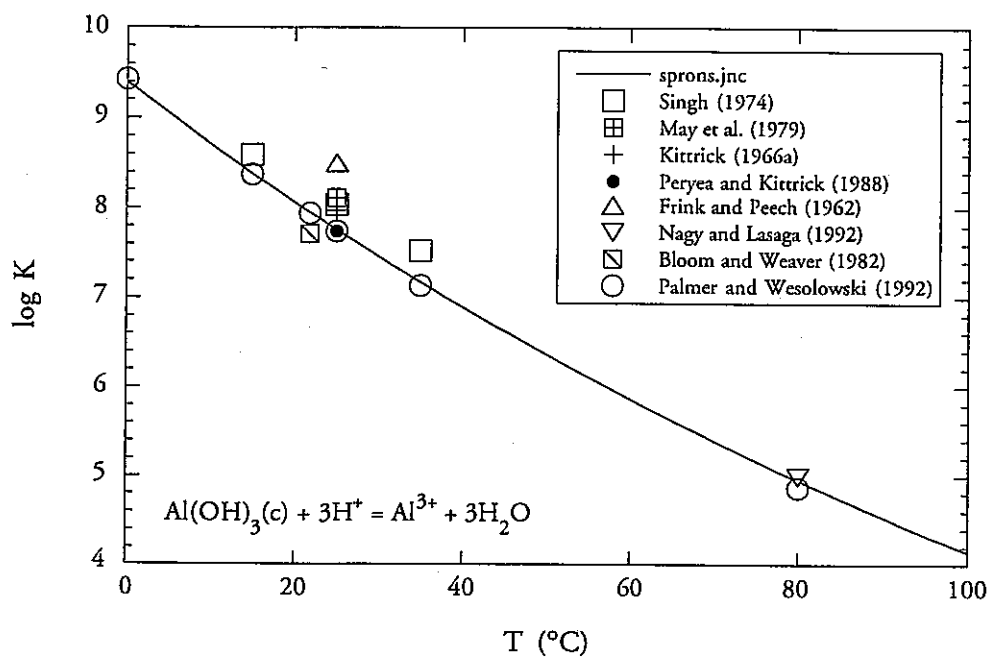


Fig. C.11: Logarithm of the equilibrium constant for the reaction gibbsite + $3\text{H}^+ \rightleftharpoons \text{Al}^{3+} + 3\text{H}_2\text{O}(\text{l})$ between 0 and 100°C at 1 bar.

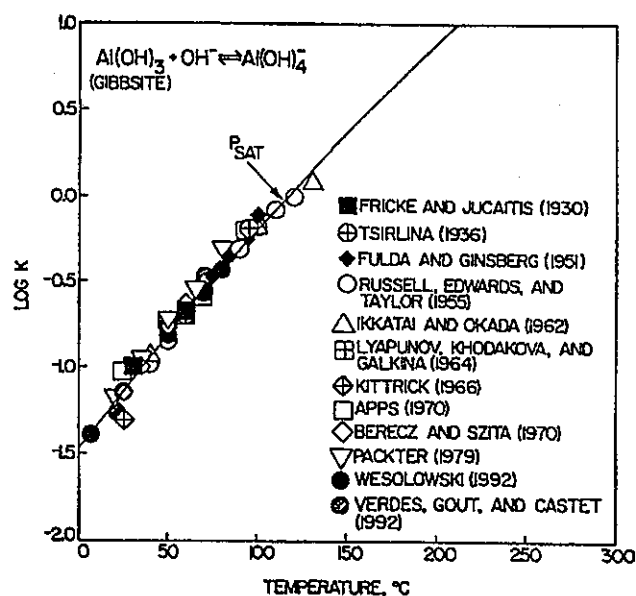


Fig. C.12: Logarithm of the equilibrium constant for the reaction gibbsite + $\text{OH}^- \rightleftharpoons \text{Al}(\text{OH})_4^-$ between 0 and 300°C at P_{sat} (from Pokrovskii and Helgeson, 1995).

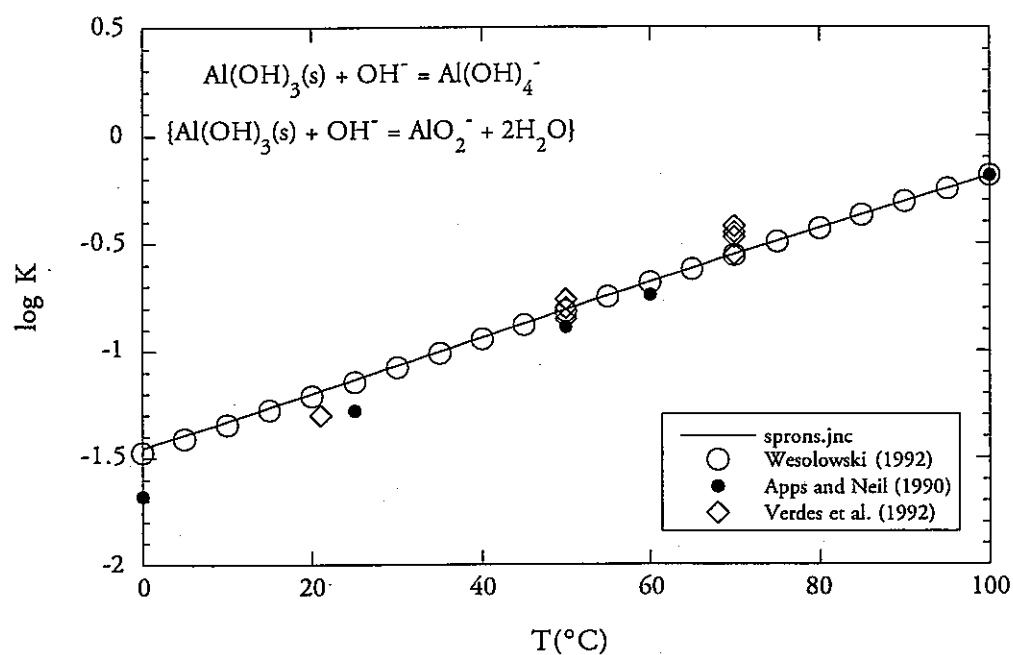


Fig. C.13: Logarithm of the equilibrium constant for the reaction gibbsite + $\text{OH}^- \rightleftharpoons \text{Al(OH)}_4^-$ between 0 and 100°C at 1 bar.

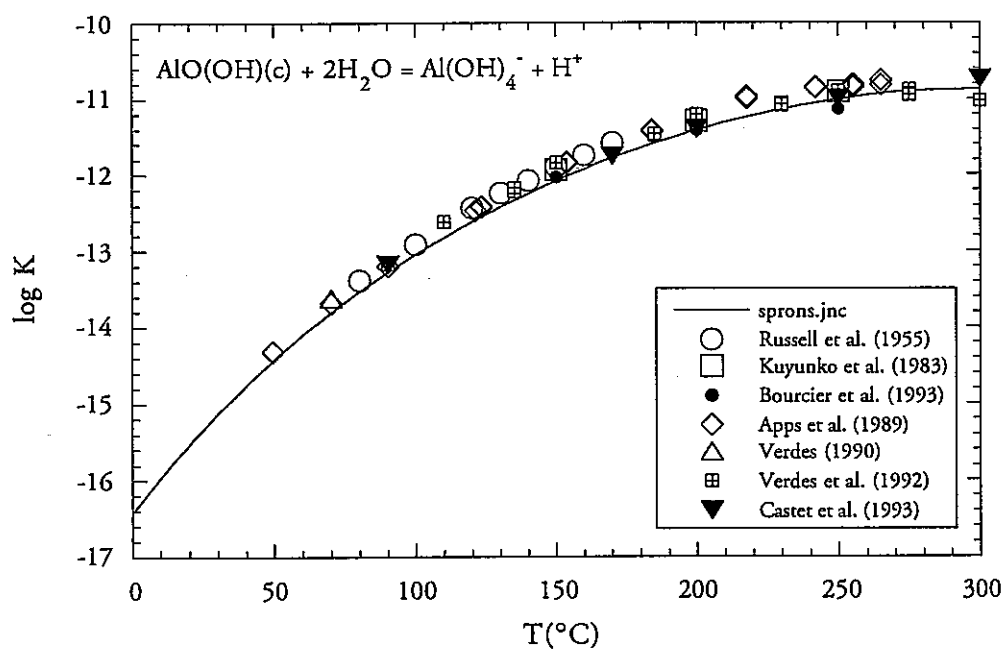


Fig. C.14: Logarithm of the equilibrium constant for the reaction boehmite + $2\text{H}_2\text{O}(l) \rightleftharpoons \text{Al(OH)}_4^-$ between 0 and 300°C at P_{sat} .

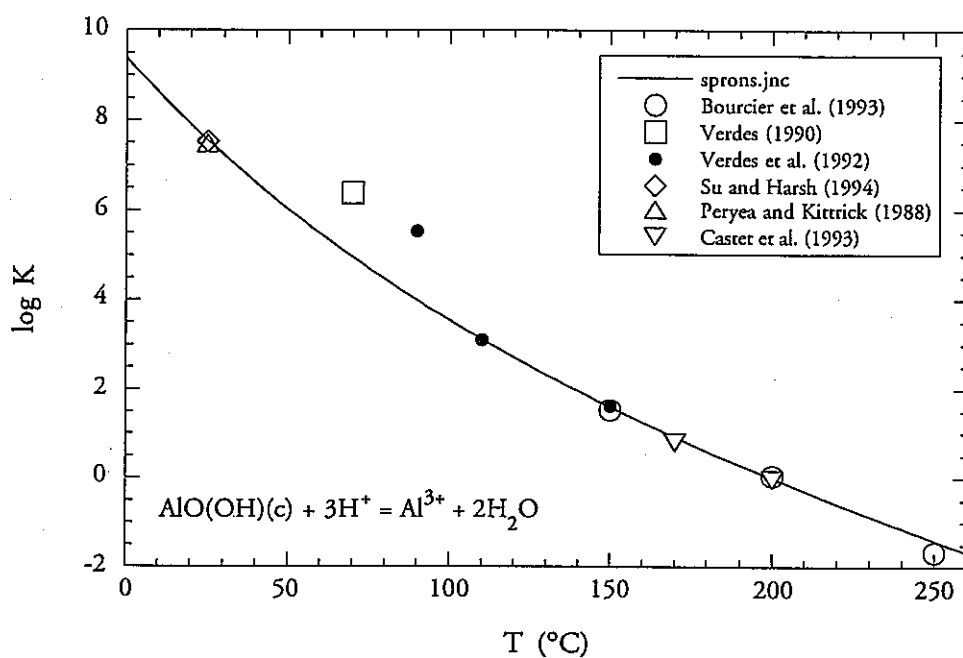


Fig. C.15: Logarithm of the equilibrium constant for the reaction boehmite + 3H⁺ ⇌ Al³⁺ + 2H₂O(l) between 0 and 250°C at *P_{sat}*.

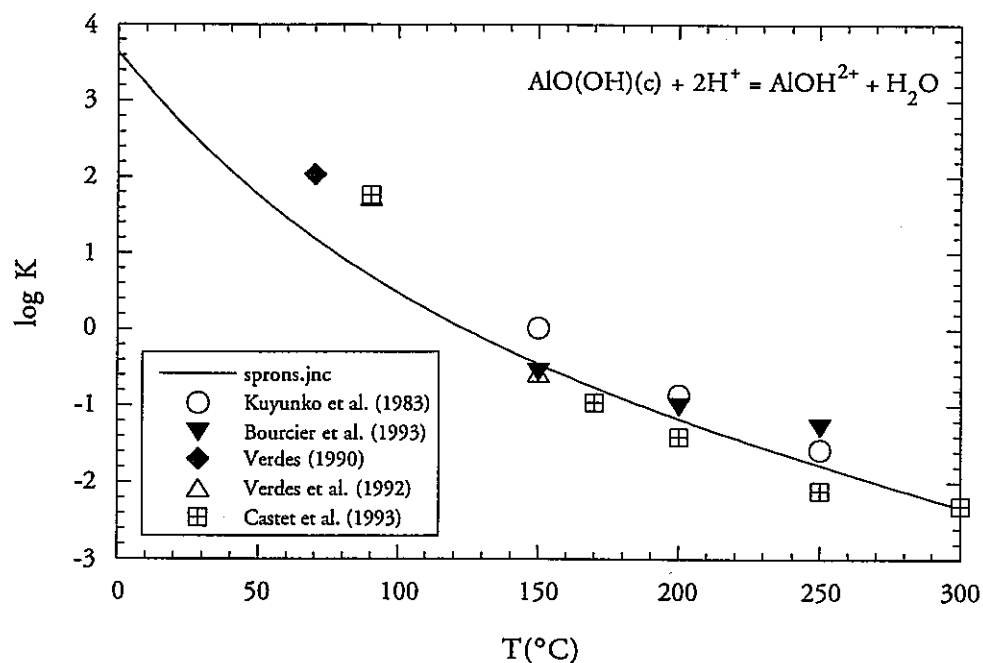


Fig. C.16: Logarithm of the equilibrium constant for the reaction boehmite + 2H⁺ ⇌ AlOH²⁺ + H₂O(l) between 0 and 300°C at *P_{sat}*.

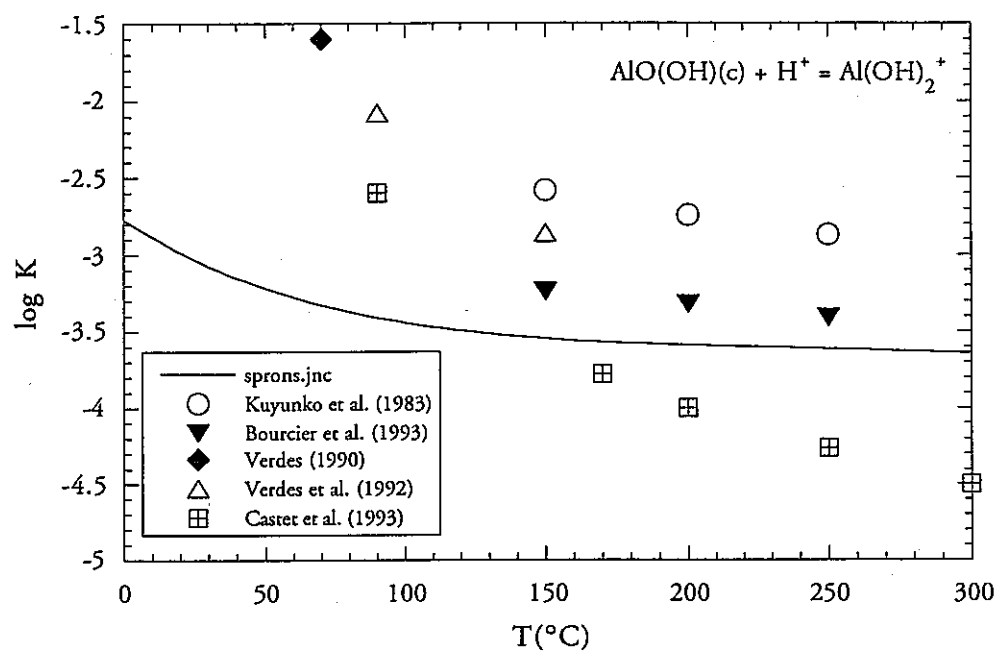


Fig. C.17: Logarithm of the equilibrium constant for the reaction boehmite + $\text{H}^+ \rightleftharpoons \text{Al(OH)}_2^+$ between 0 and 300°C at P_{sat} .

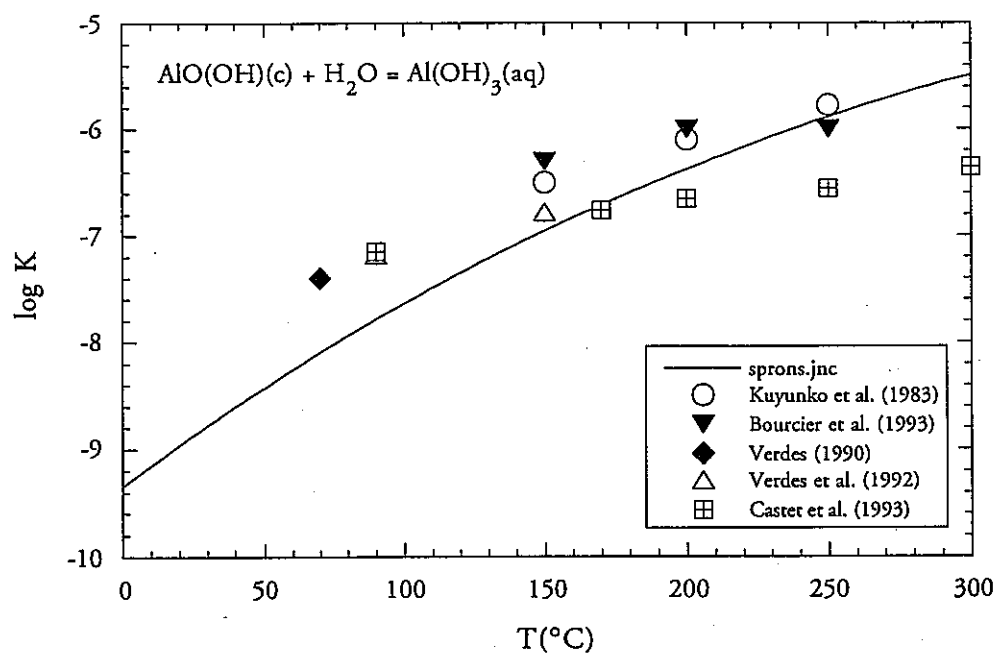


Fig. C.18: Logarithm of the equilibrium constant for the reaction boehmite + $\text{H}_2\text{O}(l) \rightleftharpoons \text{Al(OH)}_3(\text{aq})$ between 0 and 300°C at P_{sat} .

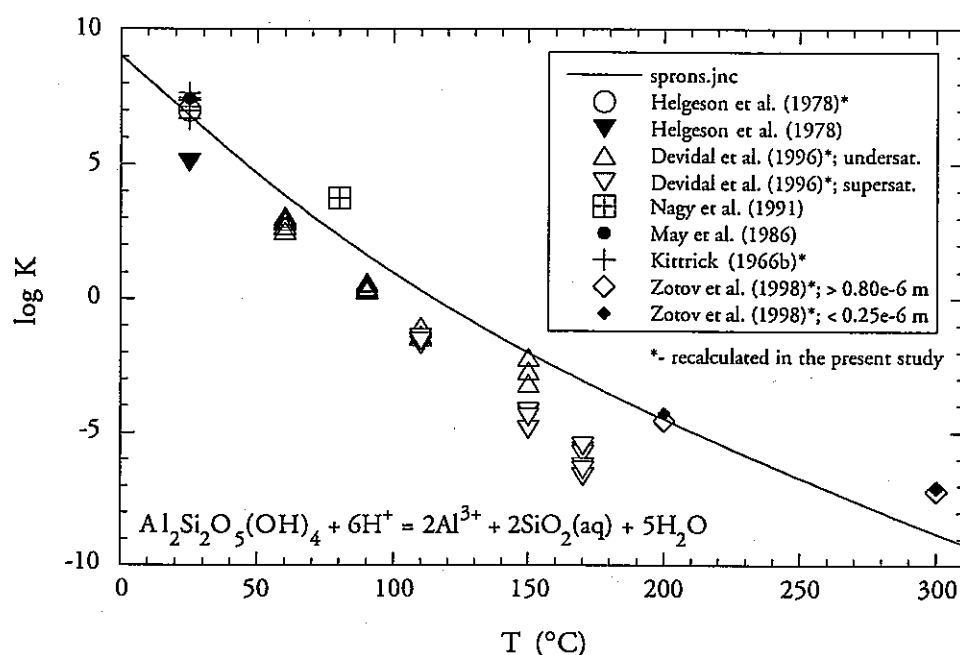


Fig. C.19: Logarithm of the equilibrium constant for the reaction kaolinite + $6\text{H}^+ = 2\text{Al}^{3+} + 2\text{SiO}_2(\text{aq}) + 5\text{H}_2\text{O}$ between 0 and 300°C at P_{sat} .

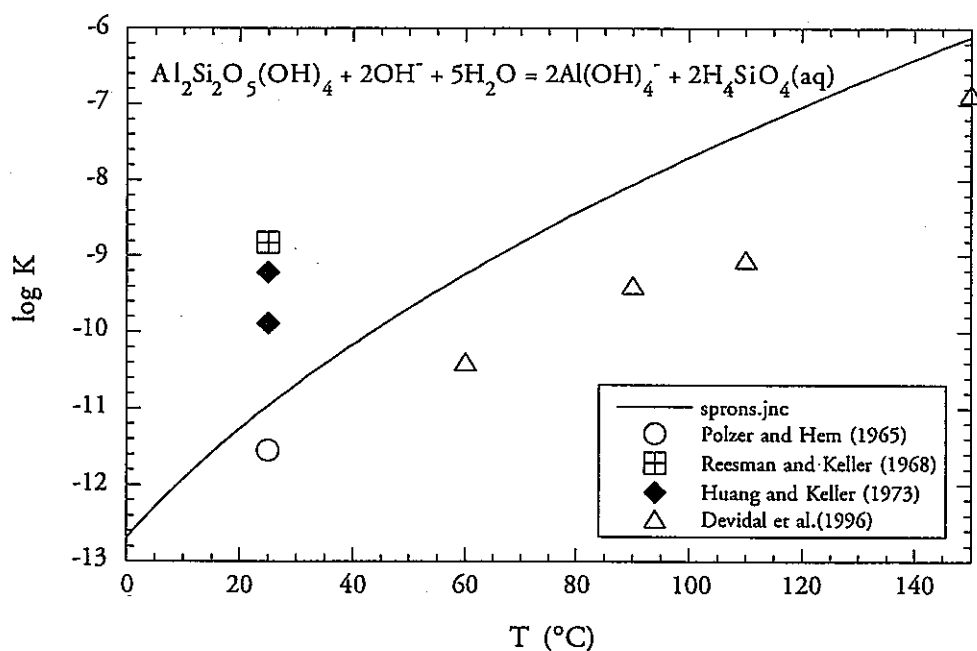


Fig. C.20: Logarithm of the equilibrium constant for the reaction kaolinite + $2\text{OH}^- + 5\text{H}_2\text{O} = 2\text{Al}(\text{OH})_4^- + 2\text{H}_4\text{SiO}_4(\text{aq})$ between 0 and 150°C at P_{sat} . All experimental equilibrium constants evaluated by Devidal *et al.* (1996) based on their experimental results and results reported by Polzer and Hem (1965), Reesman and Keller (1968) and Huang and Keller (1973).

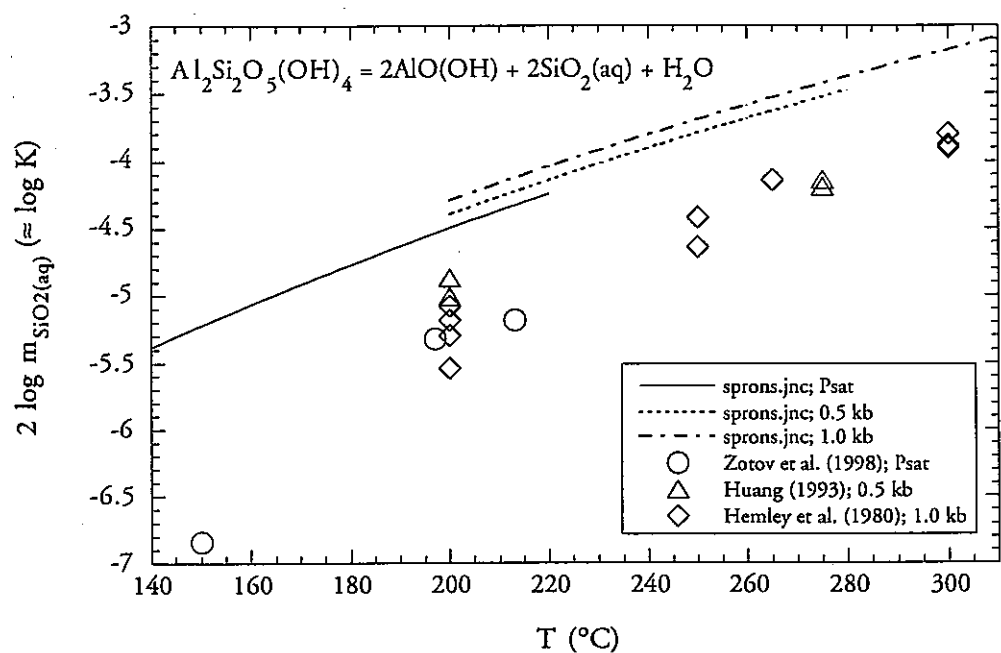


Fig. C.21: Logarithm of the equilibrium constant [$\approx 2 \log m_{\text{SiO}_2(\text{aq})}$] for the reaction kaolinite $\rightleftharpoons$ 2 boehmite + 2 $\text{SiO}_2(\text{aq})$ + H_2O between 140 and 300°C and pressures from P_{sat} to 1 kb.

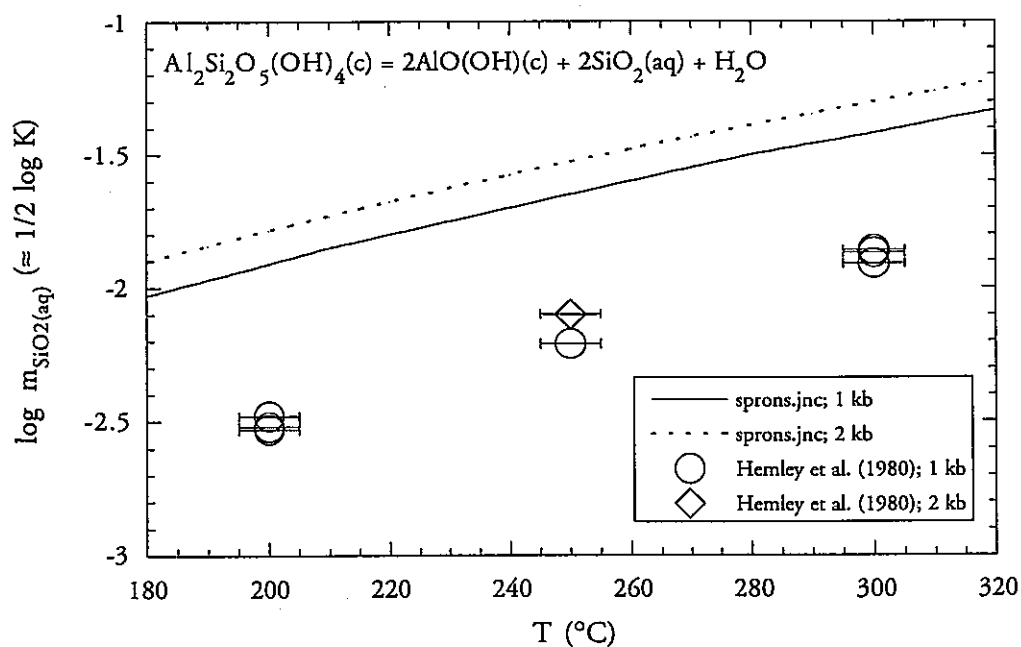


Fig. C.22: Plot of $1/2 \log K$ [$\approx \log m_{\text{SiO}_2(\text{aq})}$] for the reaction kaolinite $\rightleftharpoons$ 2 diaspor + 2 $\text{SiO}_2(\text{aq})$ + H_2O between 180 and 320°C and pressures of 1 and 2 kb.

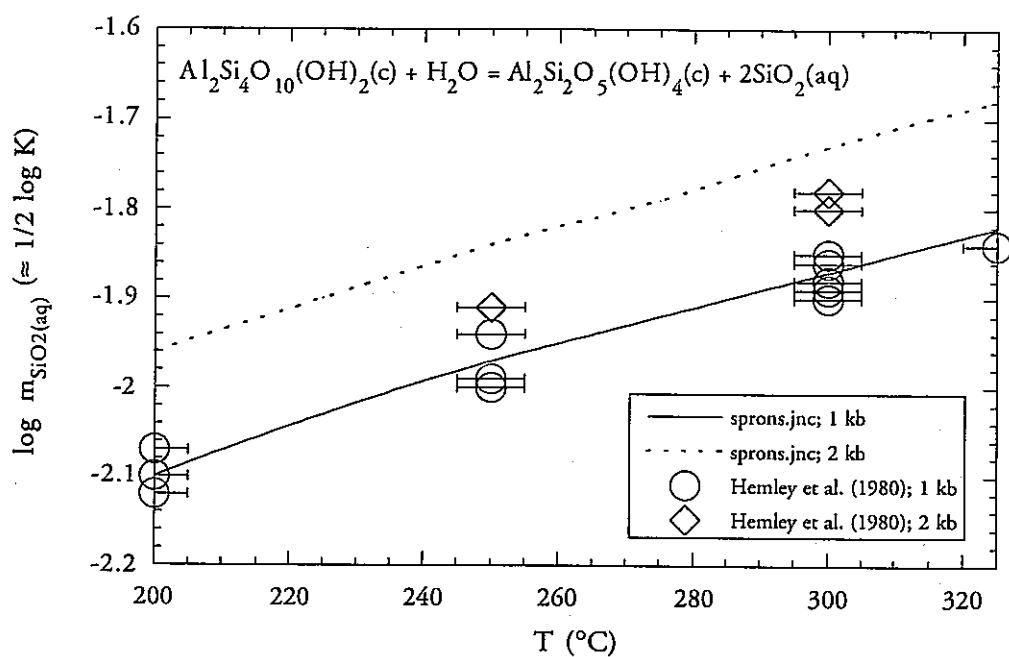


Fig. C.23: Plot of $1/2 \log K$ [$\approx \log m_{\text{SiO}_2(\text{aq})}$] for the reaction pyrophyllite + $\text{H}_2\text{O} \rightleftharpoons$ kaolinite + $2 \text{SiO}_2(\text{aq})$ between 200 and 320°C and pressures of 1 and 2 kb.

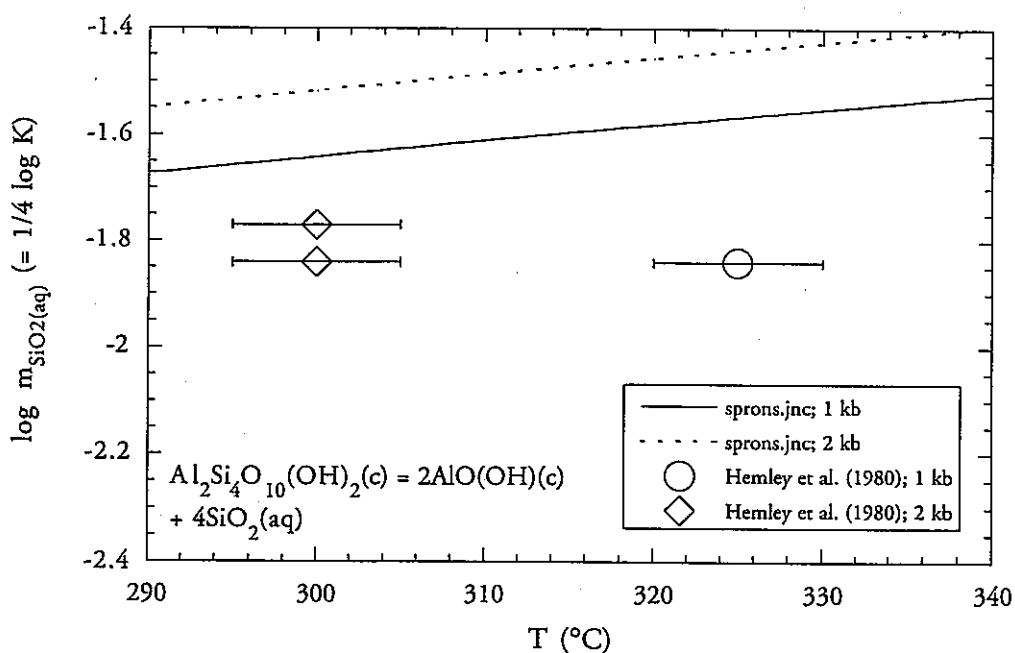


Fig. C.24: Plot of $1/4 \log K$ [$\approx \log m_{\text{SiO}_2(\text{aq})}$] for the reaction pyrophyllite $\rightleftharpoons$ 2 diaspor + $4 \text{SiO}_2(\text{aq})$ between 290 and 340°C and pressures of 1 and 2 kb.

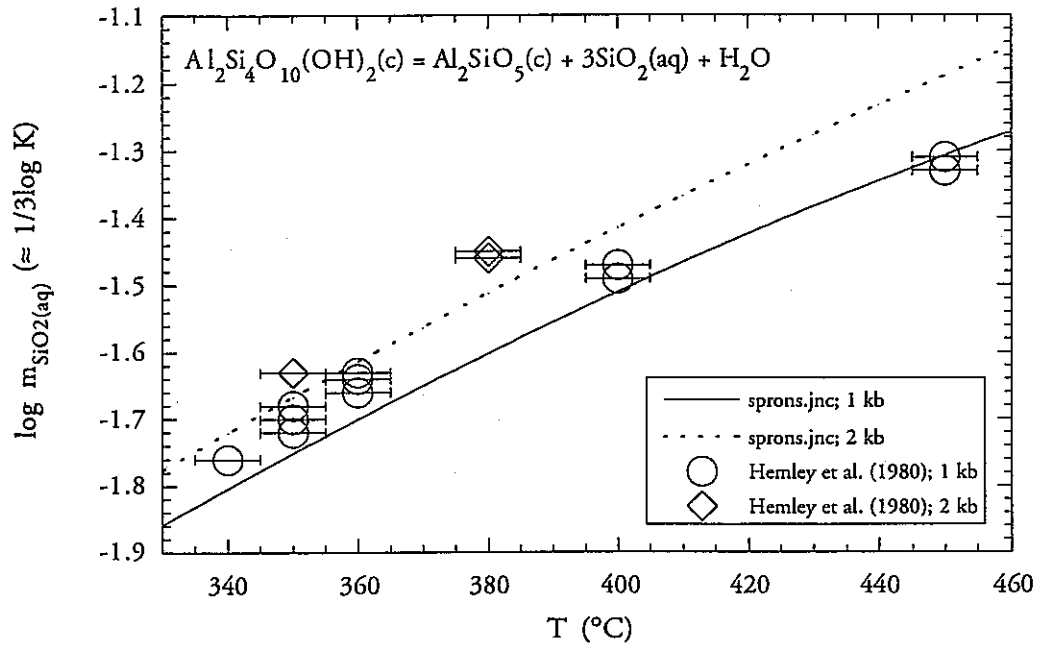


Fig. C.25: Plot of $1/3 \log K [\approx \log m_{\text{SiO}_2(\text{aq})}]$ for the reaction pyrophyllite $\rightleftharpoons$ andalusite + 3 $\text{SiO}_2(\text{aq}) + \text{H}_2\text{O}$ between 340 and 460°C and pressures of 1 and 2 kb.

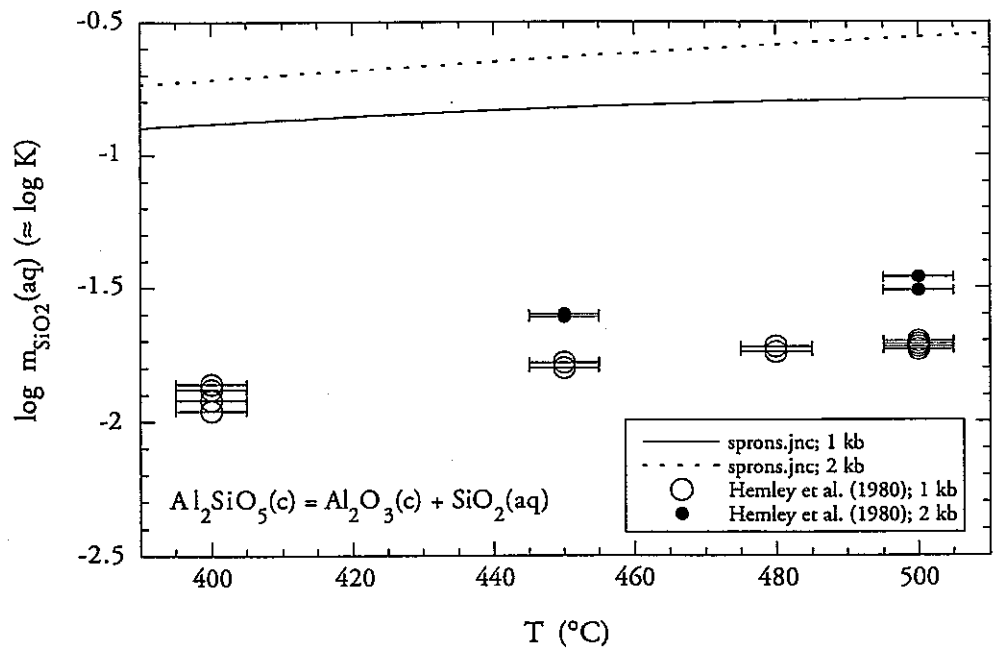


Fig. C.26: Logarithm of the equilibrium constant $[\approx \log m_{\text{SiO}_2(\text{aq})}]$ for the reaction andalusite $\rightleftharpoons$ corundum + $\text{SiO}_2(\text{aq})$ between 390 and 510°C and pressures of 1 and 2 kb.

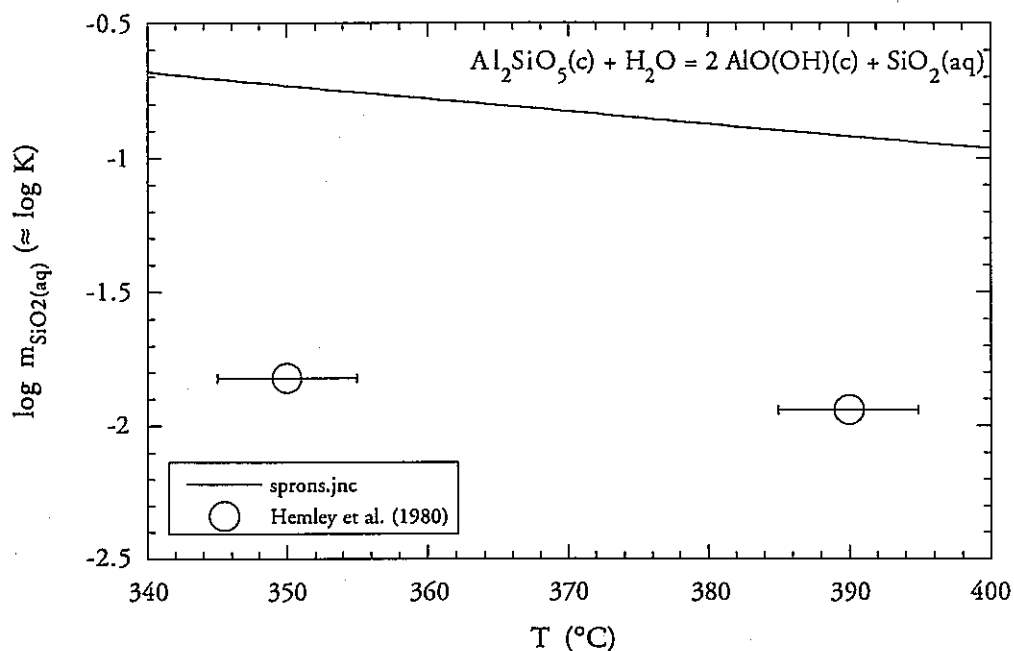


Fig. C.27: Logarithm of the equilibrium constant [$\approx \log m_{\text{SiO}_2(\text{aq})}$] for the reaction andalusite + $\text{H}_2\text{O} \rightleftharpoons$ diaspore + $\text{SiO}_2(\text{aq})$ between 340 and 400°C at 1 kb.

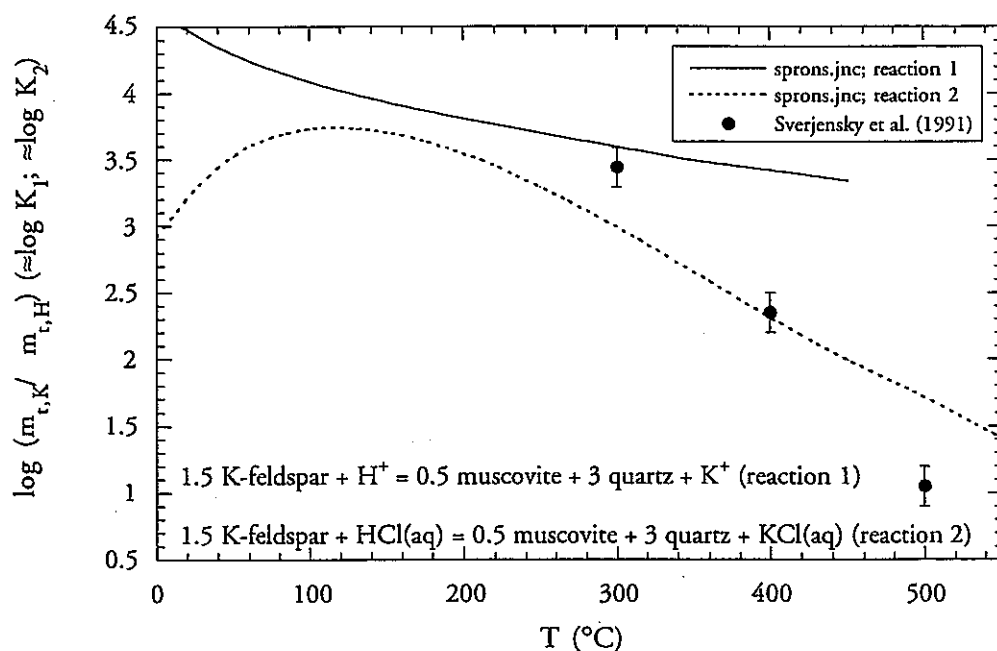


Fig. C.28: Logarithm of the equilibrium constant [$\approx \log (m_{\text{t,K}} / m_{\text{t,H}})$] for reactions in the system $\text{K}_2\text{O}-\text{Al}_2\text{O}_3-\text{SiO}_2-\text{HCl}$ between 0 and 500°C at 0.5 kb.

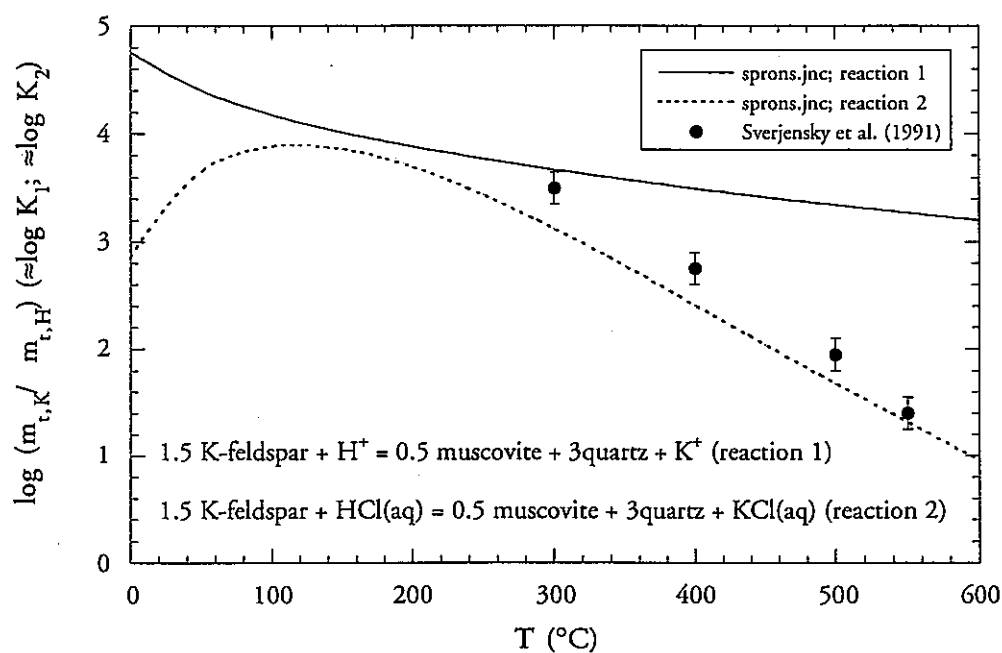


Fig. C.29: Logarithm of the equilibrium constant [$\approx \log (m_{t,K} / m_{t,H})$] for reactions in the system $\text{K}_2\text{O-Al}_2\text{O}_3\text{-SiO}_2\text{-HCl}$ between 0 and 600°C at 1 kb.

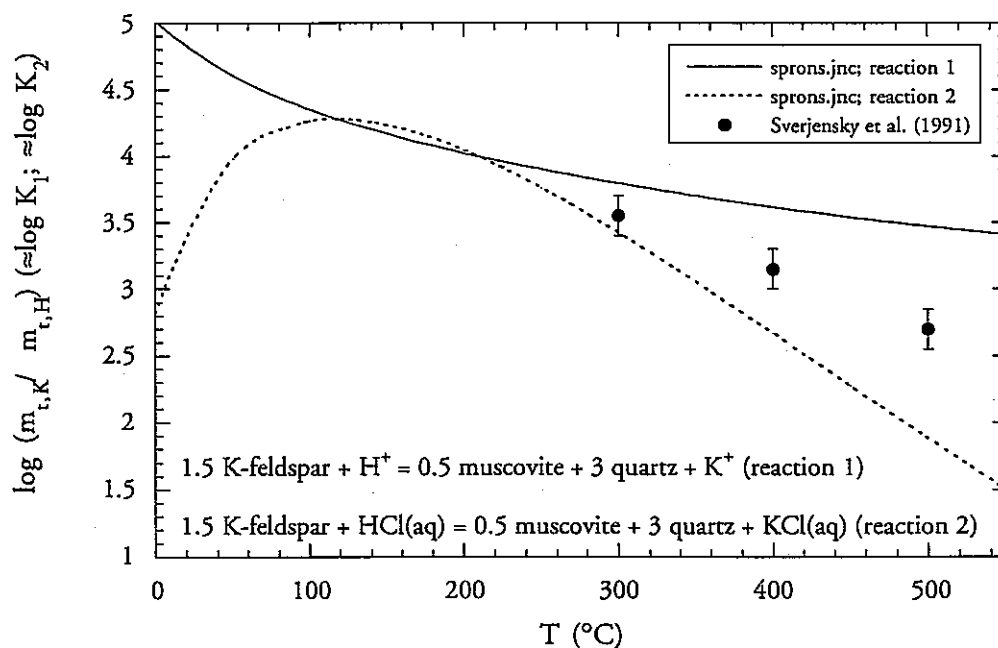


Fig. C.30: Logarithm of the equilibrium constant [$\approx \log (m_{t,K} / m_{t,H})$] for reactions in the system $\text{K}_2\text{O-Al}_2\text{O}_3\text{-SiO}_2\text{-HCl}$ between 0 and 550°C at 2 kb.

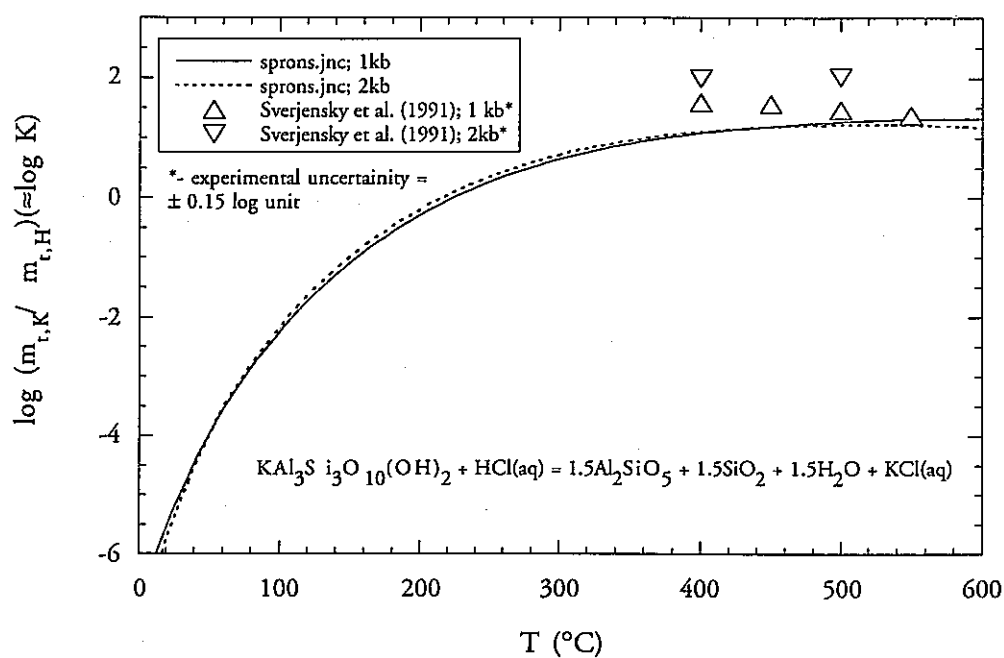


Fig. C.31: Logarithm of the equilibrium constant [$\approx \log (m_{t,K} / m_{t,H})$] for the reaction muscovite + HCl(aq) $\rightleftharpoons$ 1.5 andalusite + 1.5 quartz + 1.5 H₂O + KCl(aq) between 0 and 600°C and pressures of 1 and 2 kb.

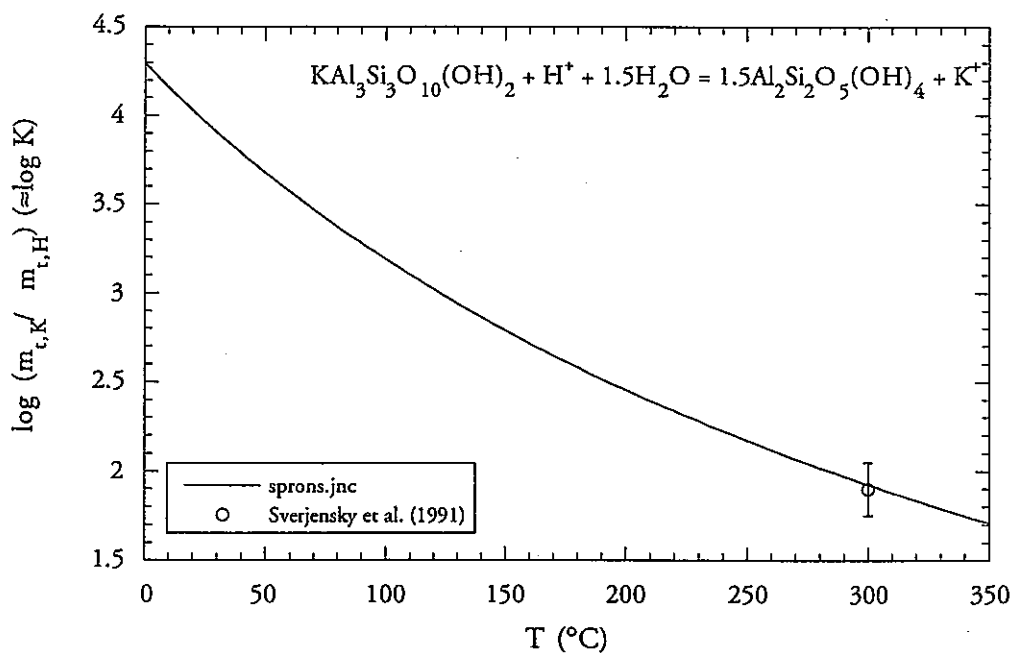


Fig. C.32: Logarithm of the equilibrium constant [$\approx \log (m_{t,K} / m_{t,H})$] for the reaction muscovite + H⁺ + 1.5 H₂O $\rightleftharpoons$ 1.5 kaolinite + K⁺ between 0 and 350°C at 1 kb.

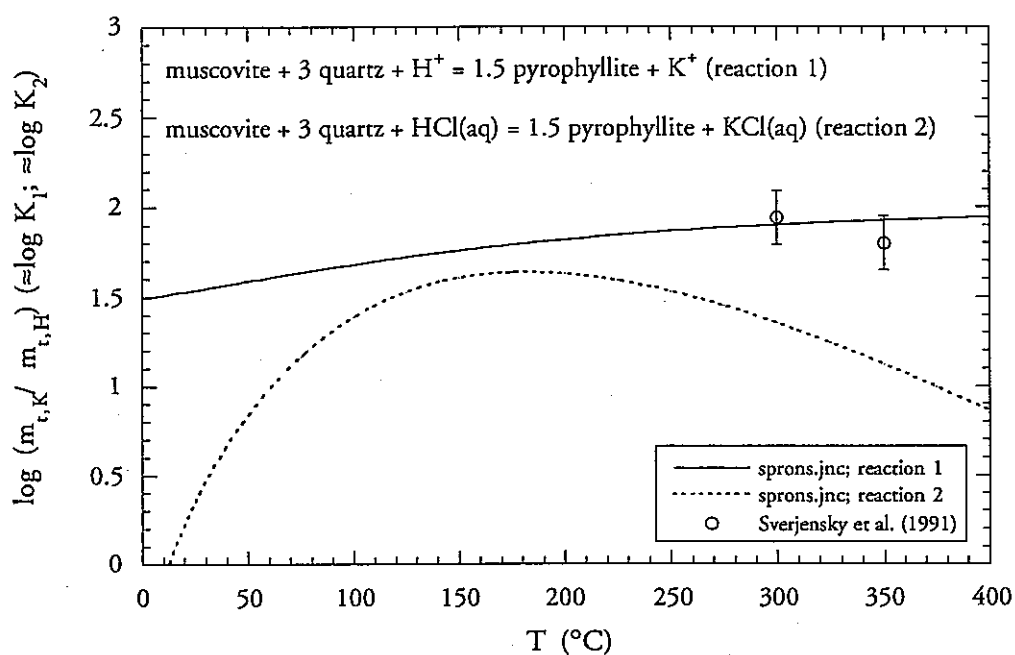


Fig. C.33: Logarithm of the equilibrium constant [$\approx \log (m_{t,K} / m_{t,H})$] for reactions in the system $K_2O-Al_2O_3-SiO_2-HCl$ between 0 and 400°C at 1 kb.

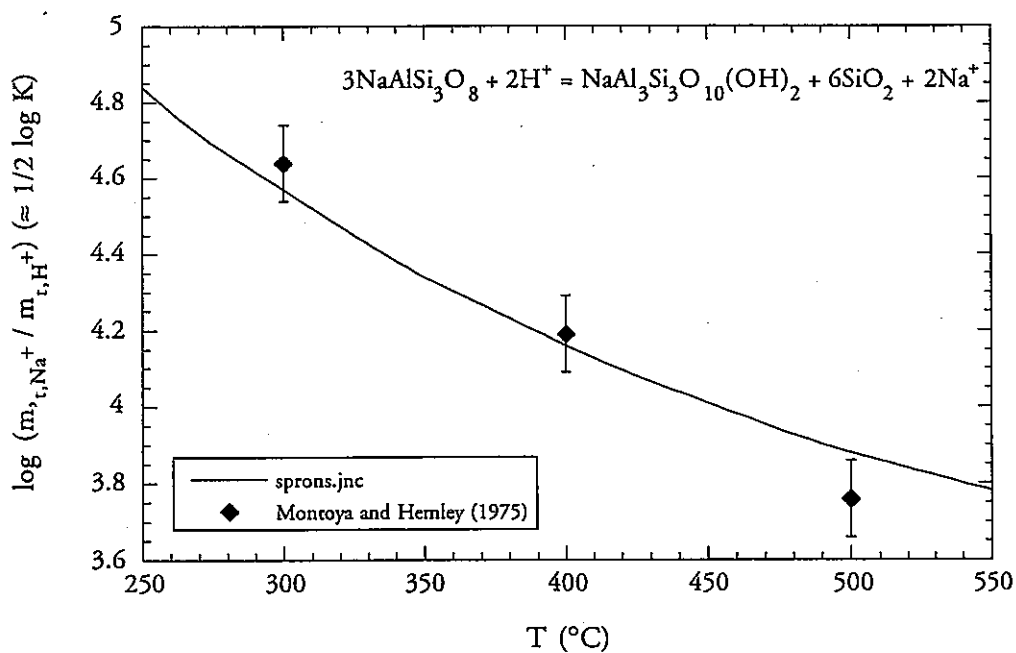


Fig. C.34: Logarithm of the equilibrium constant [$\approx 2 (\log m_{t,Na} / m_{t,H})$] for the reaction 3 albite + 2 H^+ $\rightleftharpoons$ paragonite + 6 quartz + 2 Na^+ between 250 and 550°C at 1 kb.

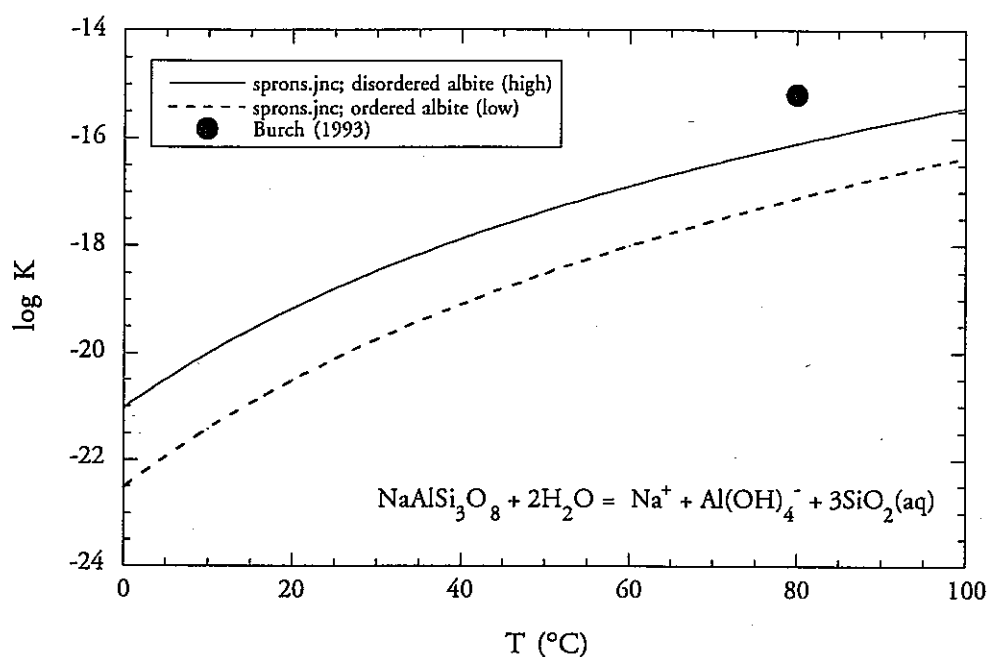


Fig. C.35: Logarithm of the equilibrium constant for the reaction albite + 2 H₂O ⇌ Al(OH)₄⁻ + 3 SiO₂(aq) + Na⁺ between 0 and 100°C at 1 bar.

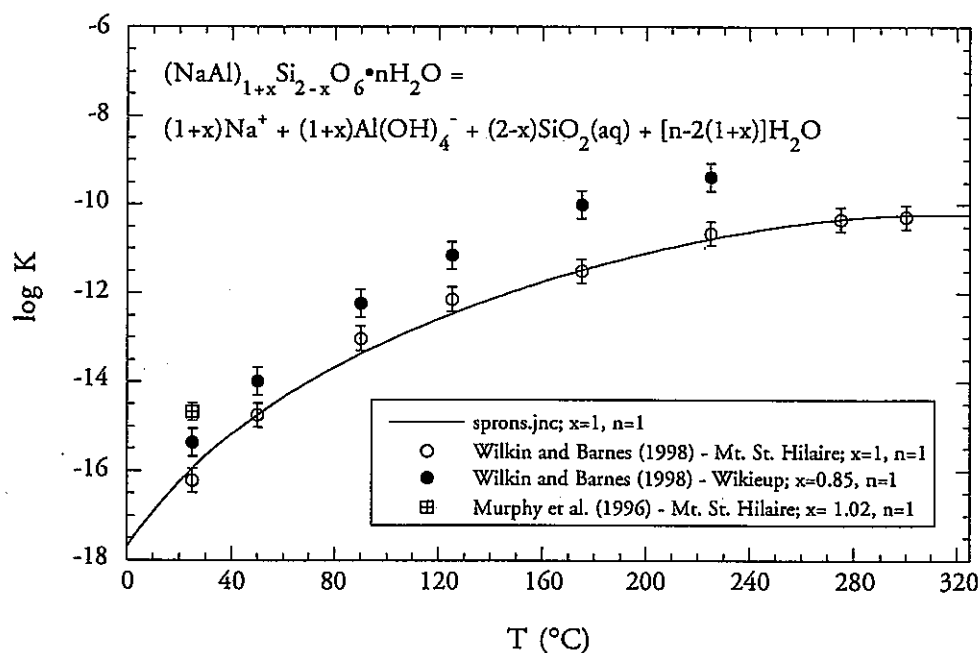


Fig. C.36: Logarithm of the equilibrium constant for the analcime hydrolysis reaction between 0 and 300°C at *P*_{sat}.

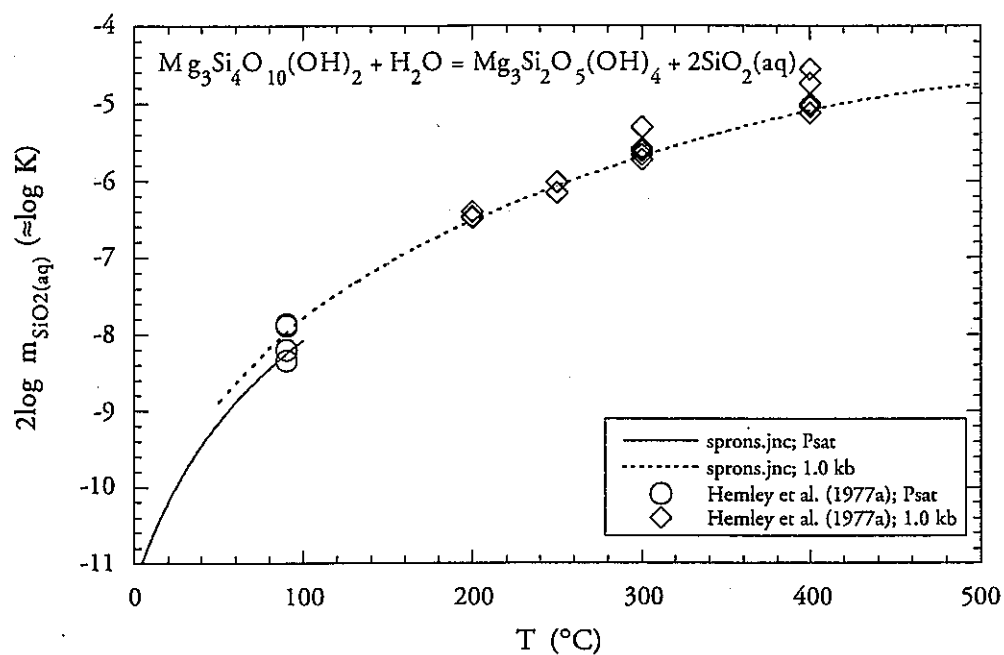


Fig. C.37: Logarithm of the equilibrium constant [$\approx 2 \log m_{\text{SiO}_2(\text{aq})}$] for the reaction talc + $\text{H}_2\text{O} \rightleftharpoons$ chrysotile + $2 \text{SiO}_2(\text{aq})$ between 0 and 500°C at P_{sat} and 1 kb.

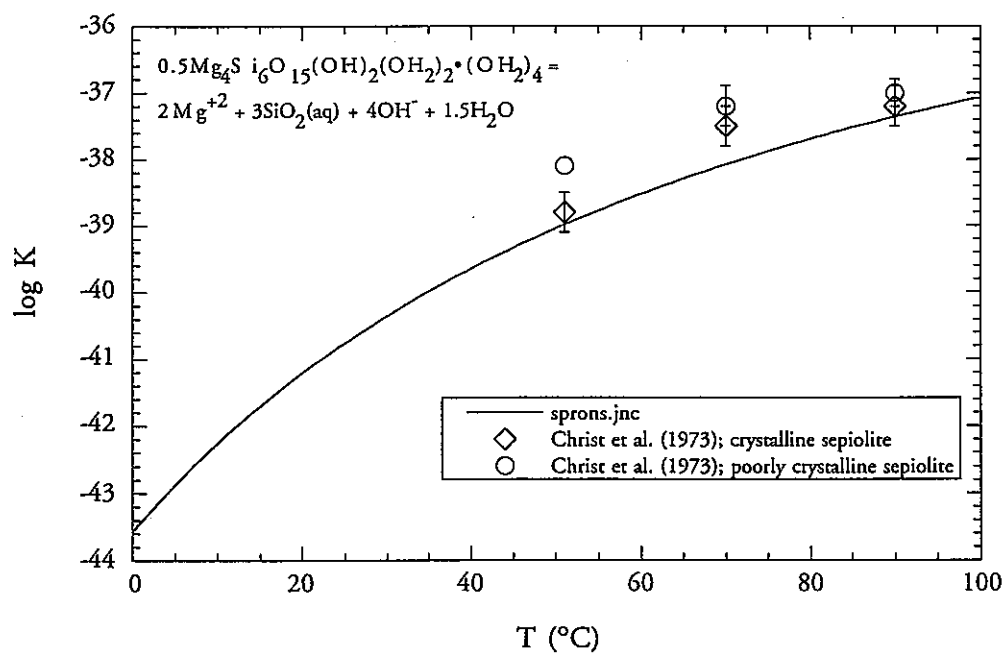


Fig. C.38: Logarithm of the equilibrium constant for the reaction sepiolite $\rightleftharpoons 2 \text{Mg}^{2+} + 3 \text{SiO}_2(\text{aq}) + 4 \text{OH}^- + 1.5 \text{H}_2\text{O}$ between 0 and 100°C at 1 bar.

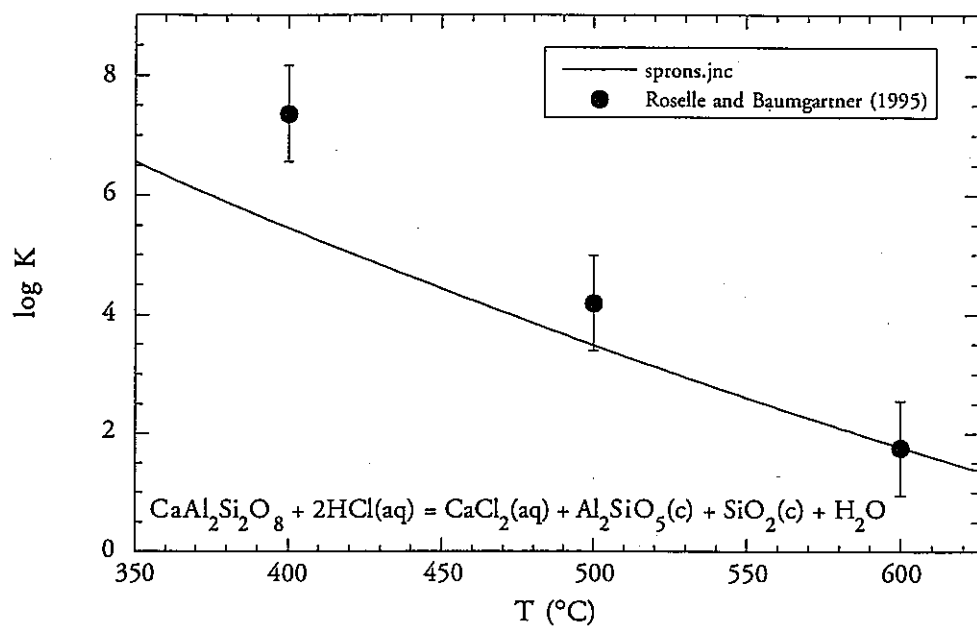


Fig. 39: Logarithm of the equilibrium constant for the reaction anorthite + 2 HCl(aq) $\rightleftharpoons$ CaCl₂(aq) + andalusite + quartz + H₂O between 350 and 625°C at 2 kb.

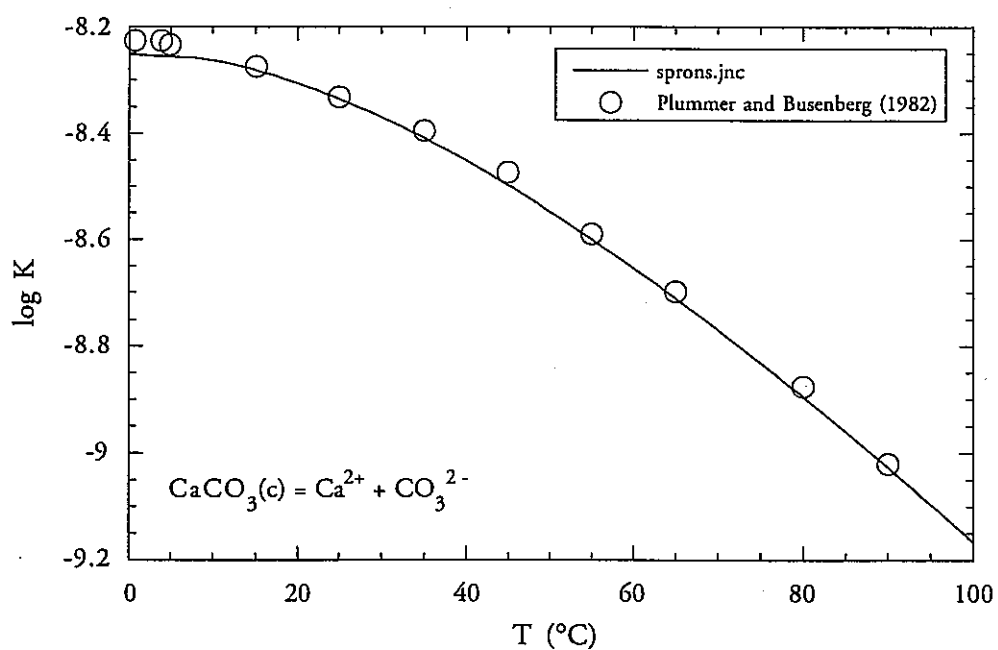


Fig C.40: Logarithm of the solubility product of aragonite between 0 and 100°C at 1 bar.

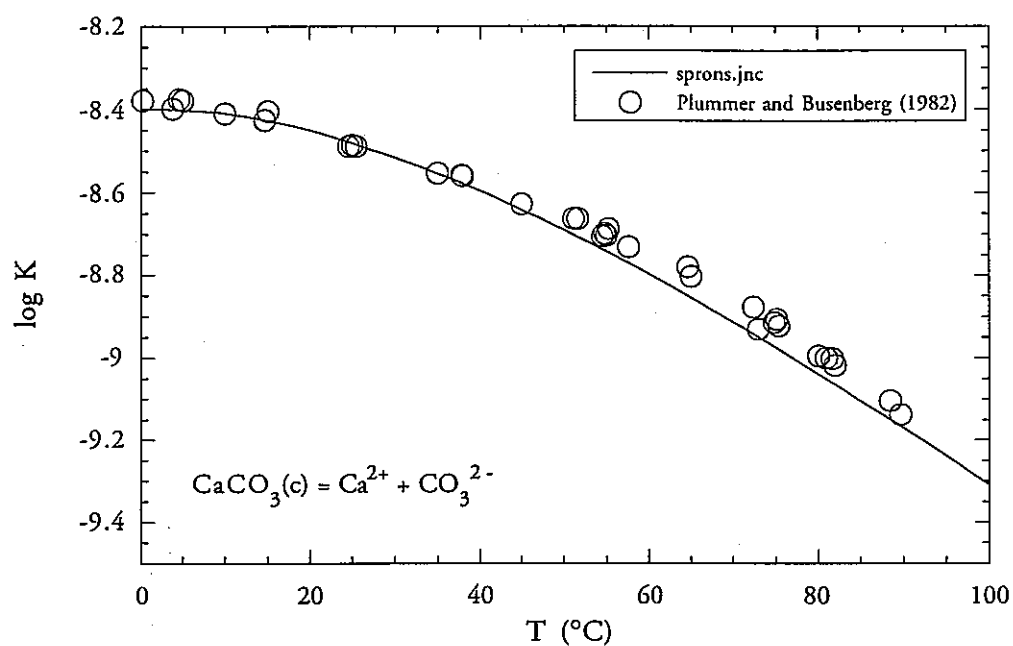


Fig C.41: Logarithm of the solubility product of calcite between 0 and 100°C at 1 bar.

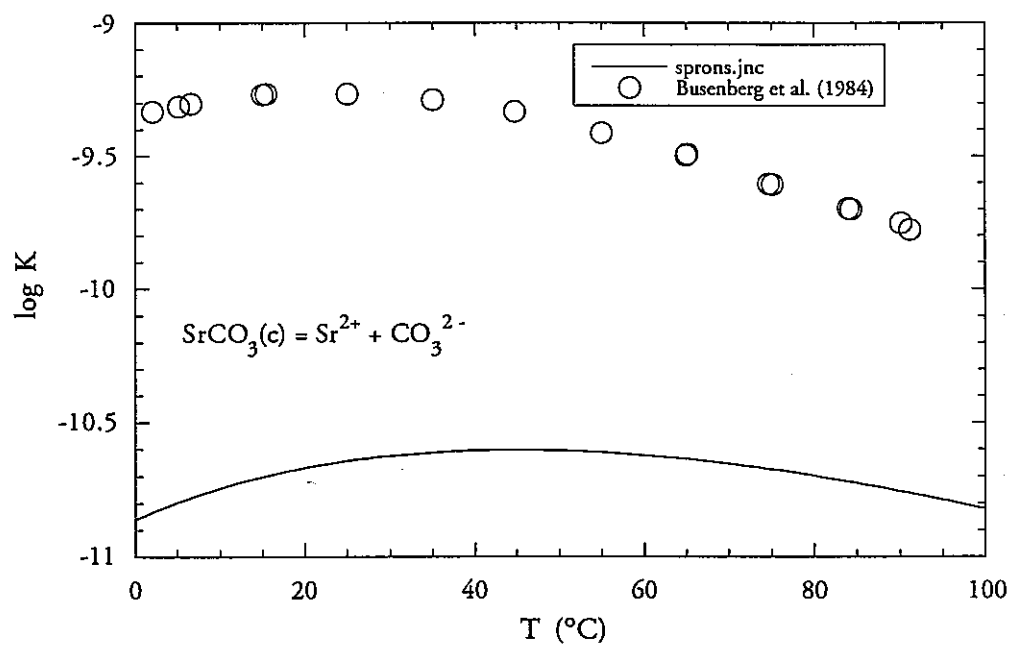


Fig C.42: Logarithm of the solubility product of strontianite between 0 and 100°C at 1 bar.

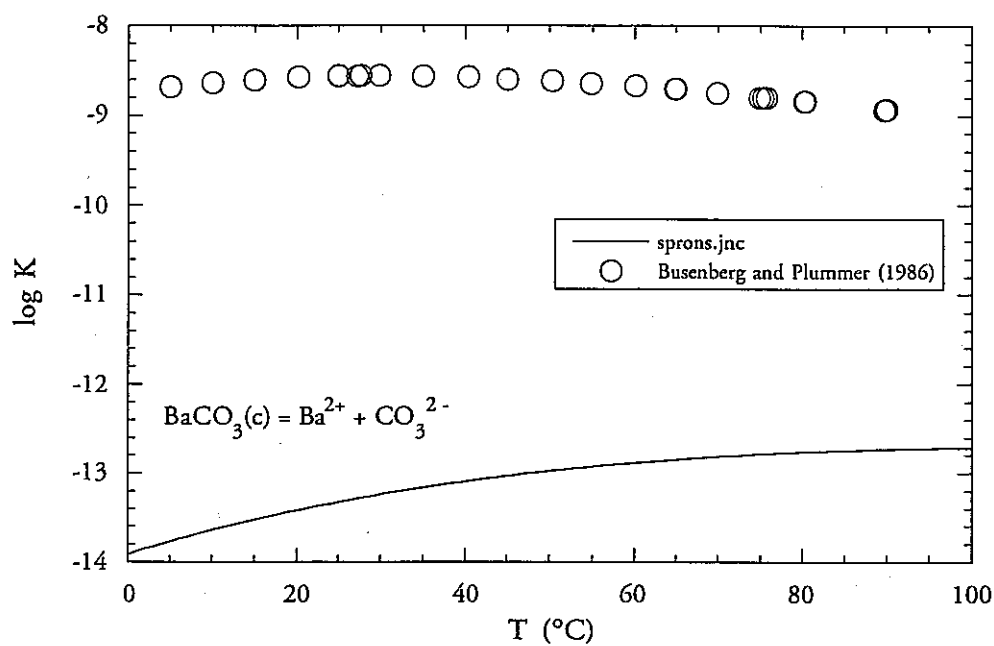


Fig C.43: Logarithm of the solubility product of witherite between 0 and 100°C at 1 bar.

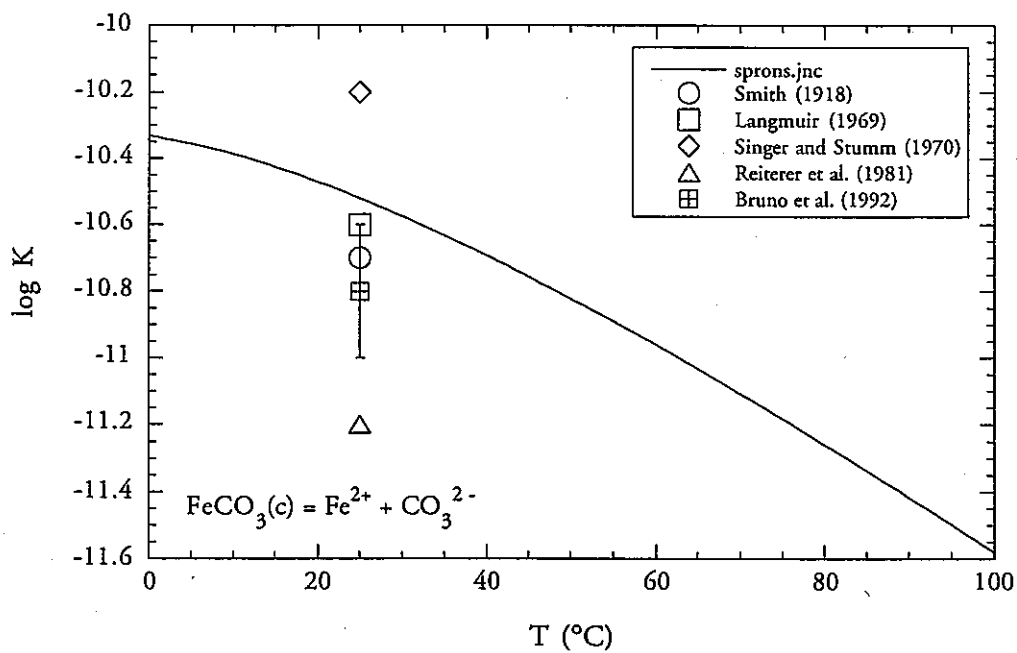


Fig C.44: Logarithm of the solubility product of siderite between 0 and 100°C at 1 bar.

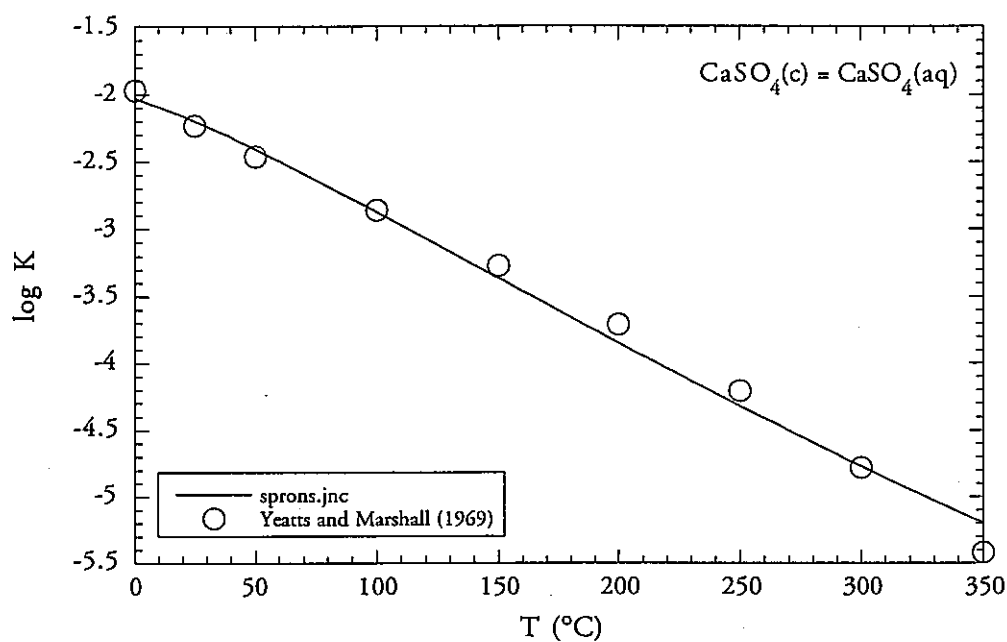


Fig C.45: Logarithm of the equilibrium constant for the reaction CaSO_4 (anhydrite) $\rightleftharpoons \text{CaSO}_4(\text{aq})$ between 0 and 350°C at P_{sat} .

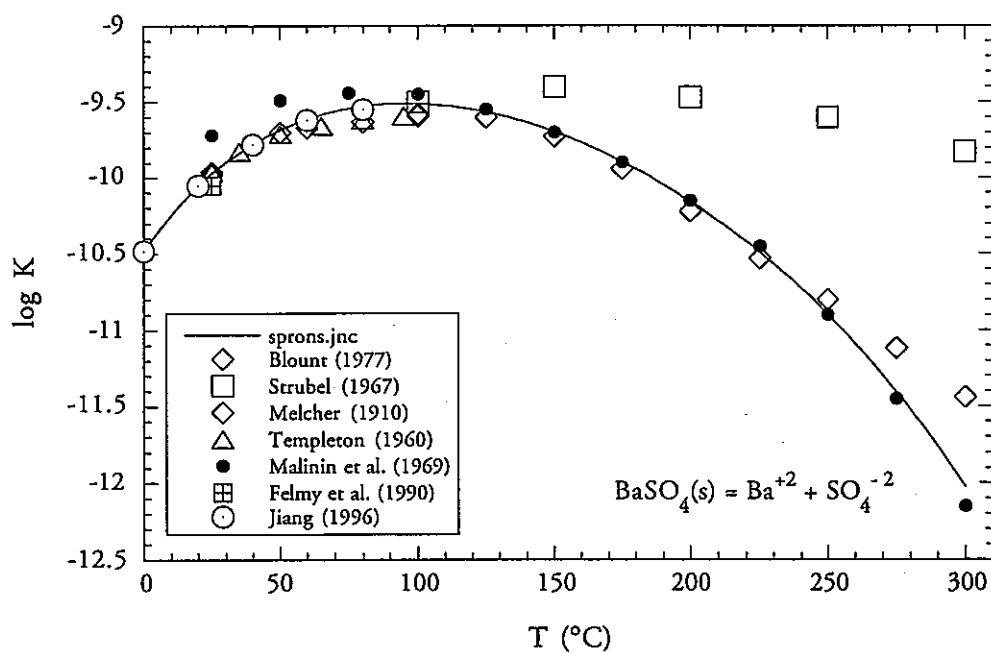


Fig C.46: Logarithm of the solubility product of barite between 0 and 300°C at P_{sat} .

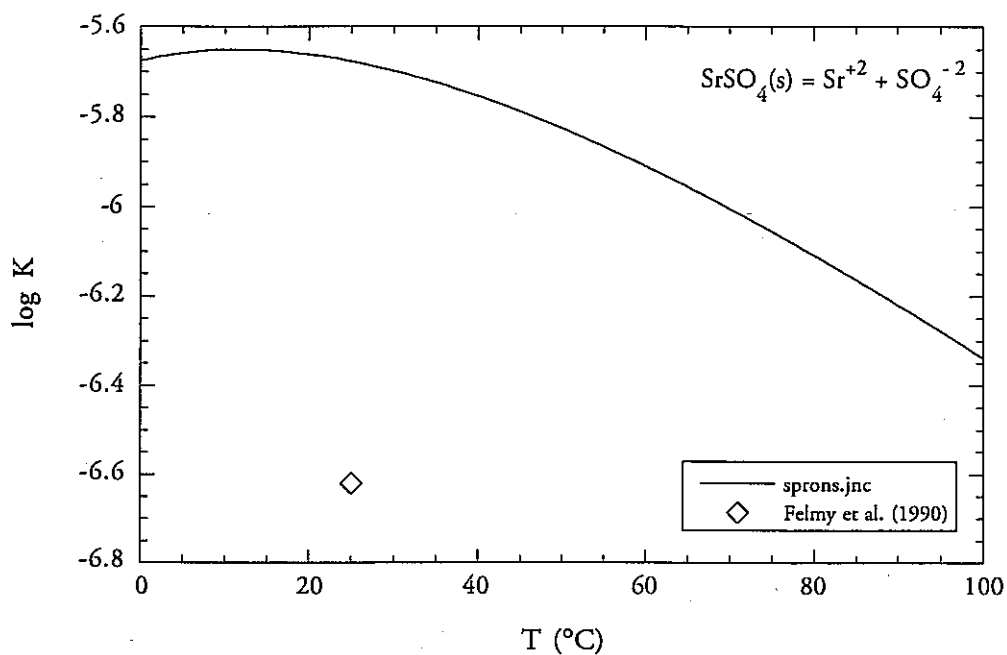


Fig C.47: Logarithm of the solubility product of celestite between 0 and 100°C at 1 bar.

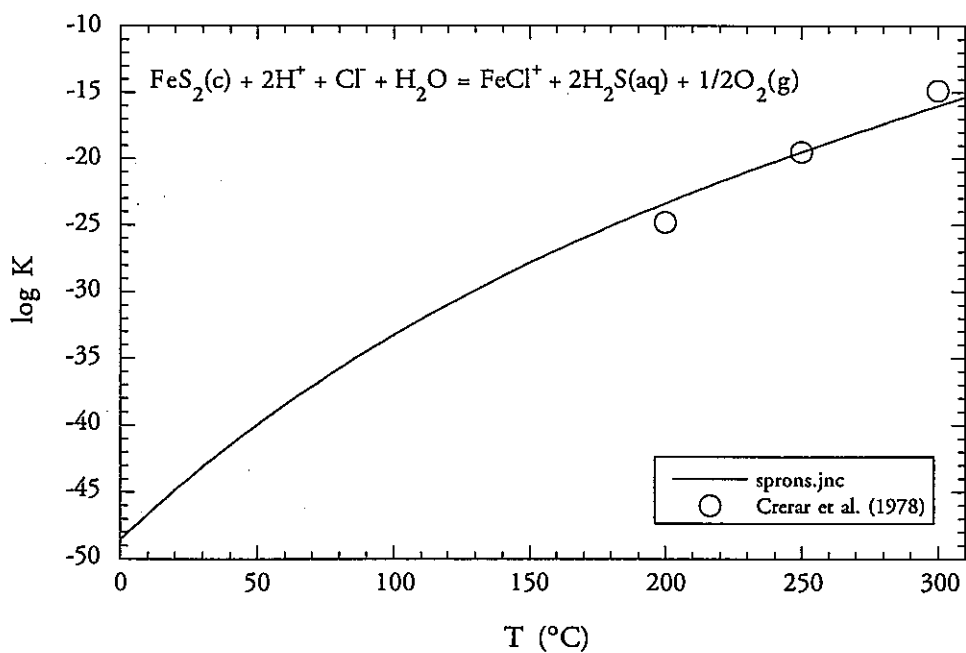


Fig C.48: Logarithm of the equilibrium constant for the reaction pyrite + 2 H⁺ + Cl⁻ + H₂O = FeCl⁺ + 2 H₂S(aq) + 1/2 O₂(g) between 0 and 300°C at P_{sat}.

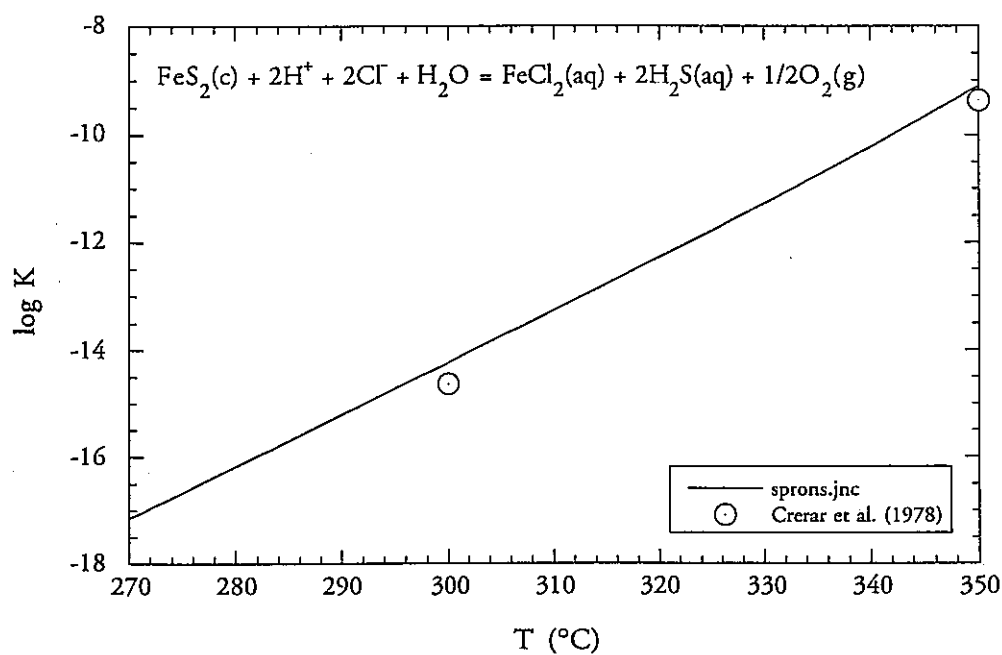


Fig C.49: Logarithm of the equilibrium constant for the reaction pyrite + 2 H⁺ + 2 Cl⁻ + H₂O = FeCl₂(aq) + 2 H₂S(aq) + 1/2 O₂(g) between 270 and 350°C at *P*_{sat}.

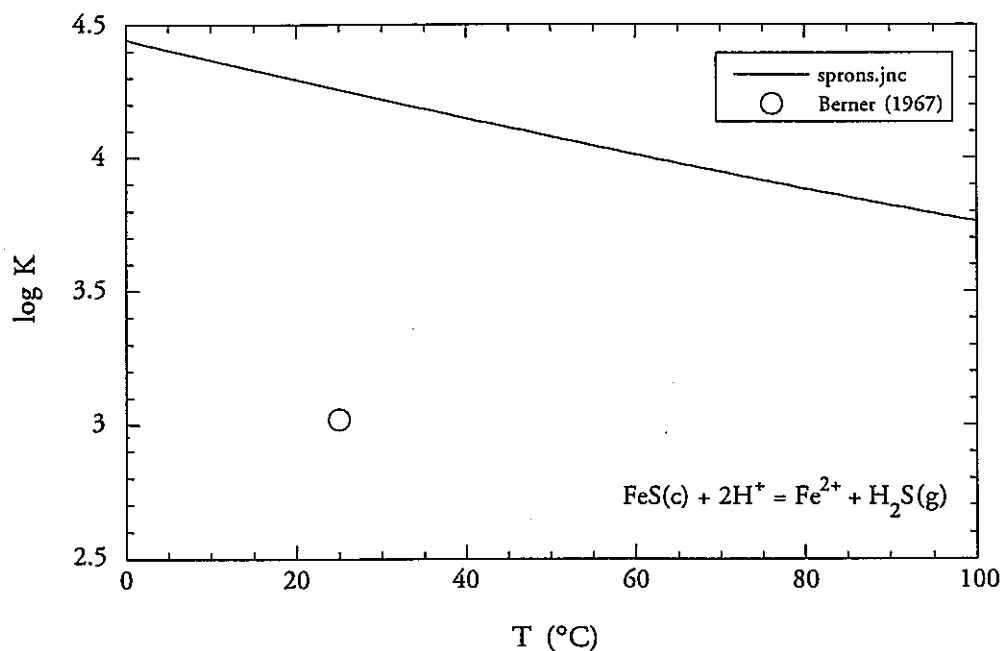


Fig C.50: Logarithm of the equilibrium constant for the reaction pyrrhotite + 2 H⁺ = Fe²⁺ + H₂S(aq) between 0 and 100°C at 1 bar [calculated in the present study using $\log K = -21.9$ at 25°C (Maronny, 1959) for the reaction $\text{H}_2\text{S}(\text{g}) = 2\text{H}^+ + \text{S}^{2-}$].

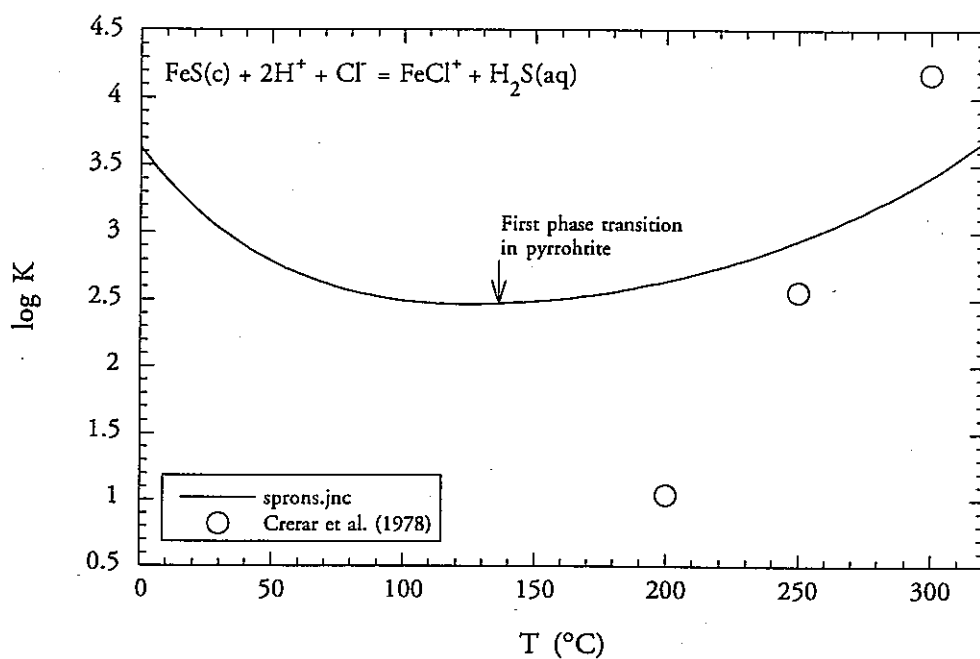


Fig C.51: Logarithm of the equilibrium constant for the reaction pyrrhotite + 2 H⁺ + Cl⁻ = FeCl⁺ + H₂S(aq) between 0 and 300°C at P_{sat} .

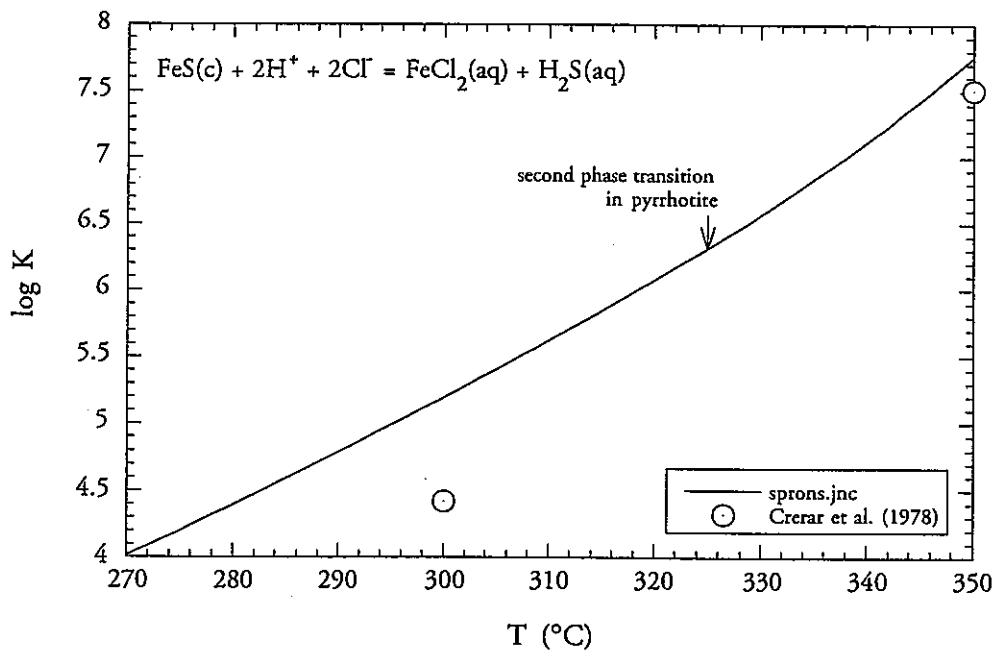


Fig C.52: Logarithm of the equilibrium constant for the reaction pyrrhotite + 2 H⁺ + 2 Cl⁻ = FeCl₂(aq) + H₂S(aq) between 270 and 350°C at P_{sat} .

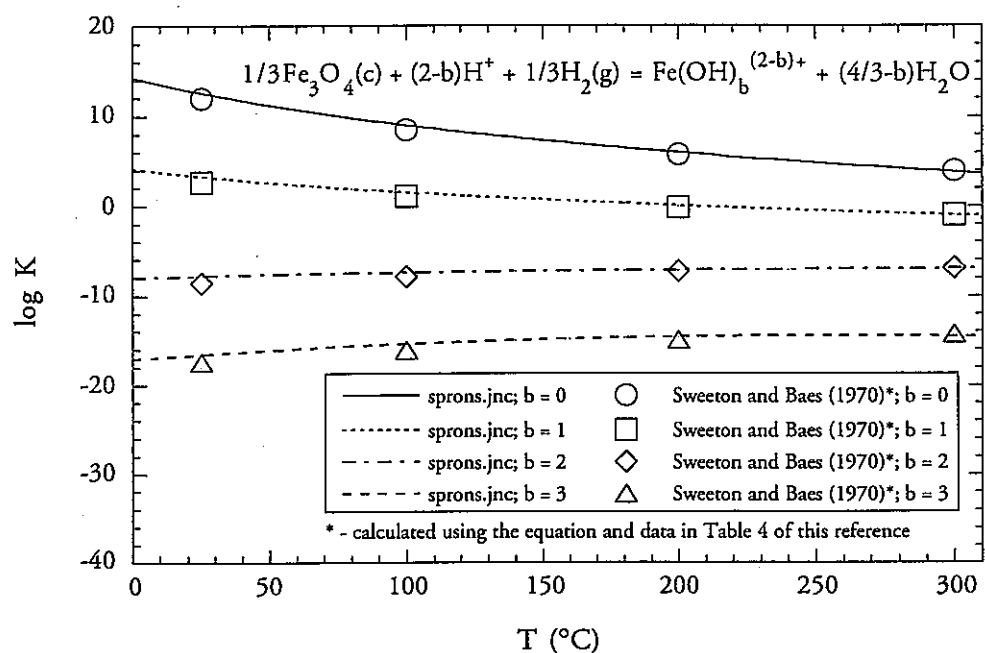


Fig C.53: Logarithm of the equilibrium constant for the reaction $\frac{1}{3}$ magnetite + $(2-b)\text{H}^+ + \frac{1}{3}\text{H}_2(\text{g}) \rightleftharpoons \text{Fe}(\text{OH})_b^{(2-b)+} + \left(\frac{4}{3}-b\right)\text{H}_2\text{O}$ ($b = 0, 1, 2, 3$) between 0 and 300°C at P_{sat} .

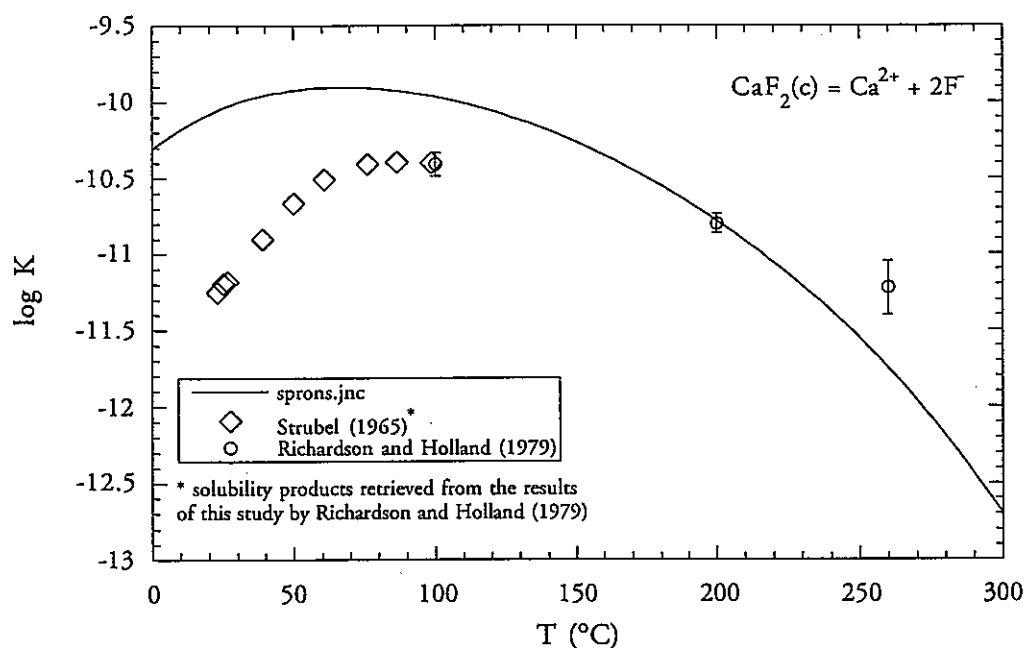


Fig C.54: Logarithm of the solubility product of fluorite between 0 and 300°C at P_{sat} .

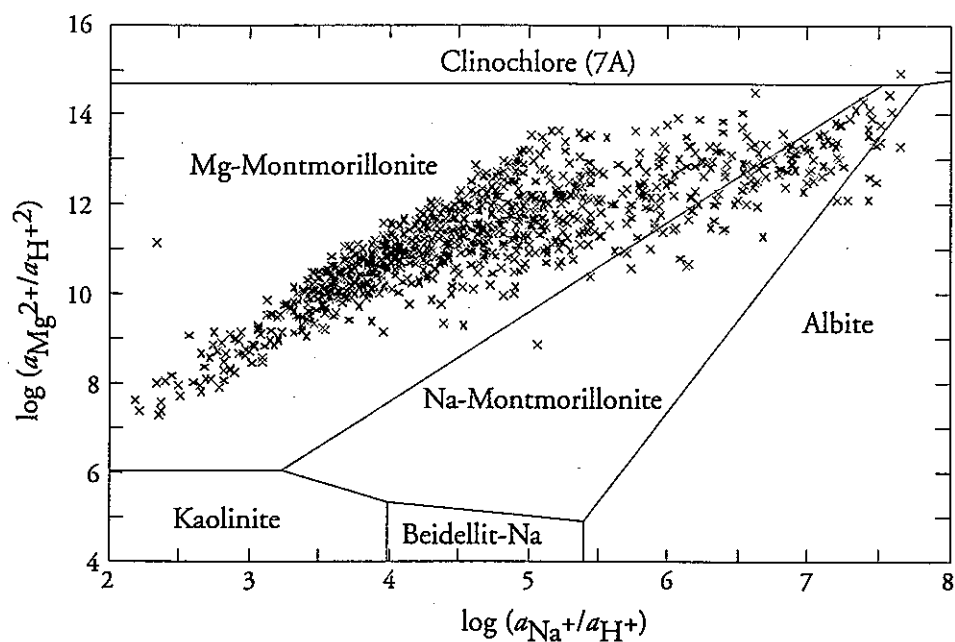


Fig. C.55: Stability relations among minerals and aqueous solutions in the system $\text{Na}_2\text{O}-\text{MgO}-\text{Al}_2\text{O}_3-\text{SiO}_2-\text{H}_2\text{O}$ at 19°C and 1 bar. Symbols represent compositions of deep groundwaters in Japan (≈ 1000 samples).

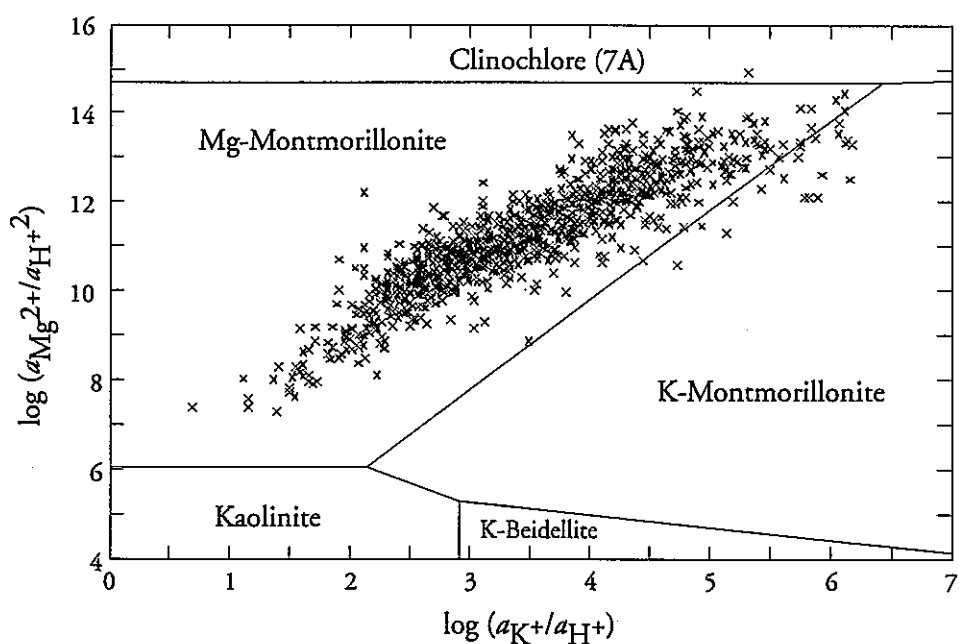


Fig. C.56: Stability relations among minerals and aqueous solutions in the system $\text{K}_2\text{O}-\text{MgO}-\text{Al}_2\text{O}_3-\text{SiO}_2-\text{H}_2\text{O}$ at 19°C and 1 bar. Symbols represent compositions of deep groundwaters in Japan (≈ 1000 samples).

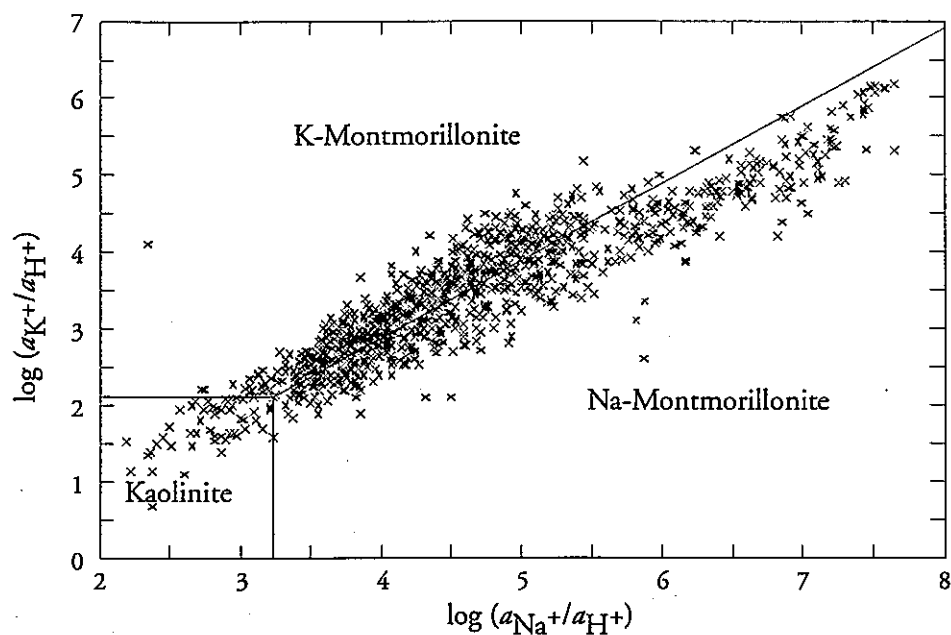


Fig. C.57: Stability relations among minerals and aqueous solutions in the system $\text{Na}_2\text{O}-\text{K}_2\text{O}-\text{MgO}-\text{Al}_2\text{O}_3-\text{SiO}_2-\text{H}_2\text{O}$ at 19°C and 1 bar. Symbols represent compositions of deep groundwaters in Japan (≈ 1000 samples).

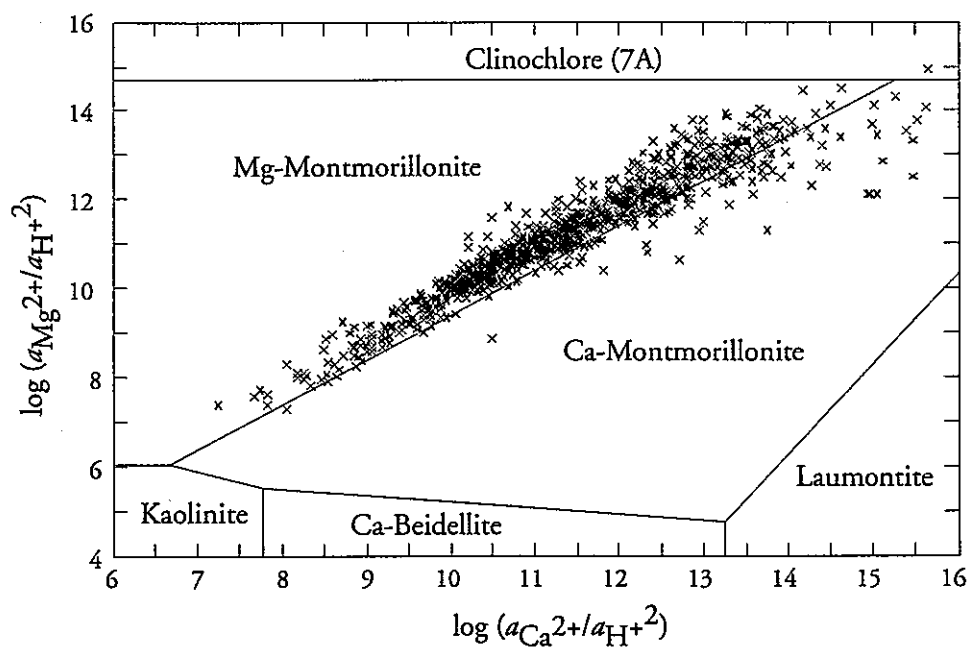


Fig. C.58: Stability relations among minerals and aqueous solutions in the system $\text{CaO}-\text{MgO}-\text{Al}_2\text{O}_3-\text{SiO}_2-\text{H}_2\text{O}$ at 19°C and 1 bar. Symbols represent compositions of deep groundwaters in Japan (≈ 1000 samples).

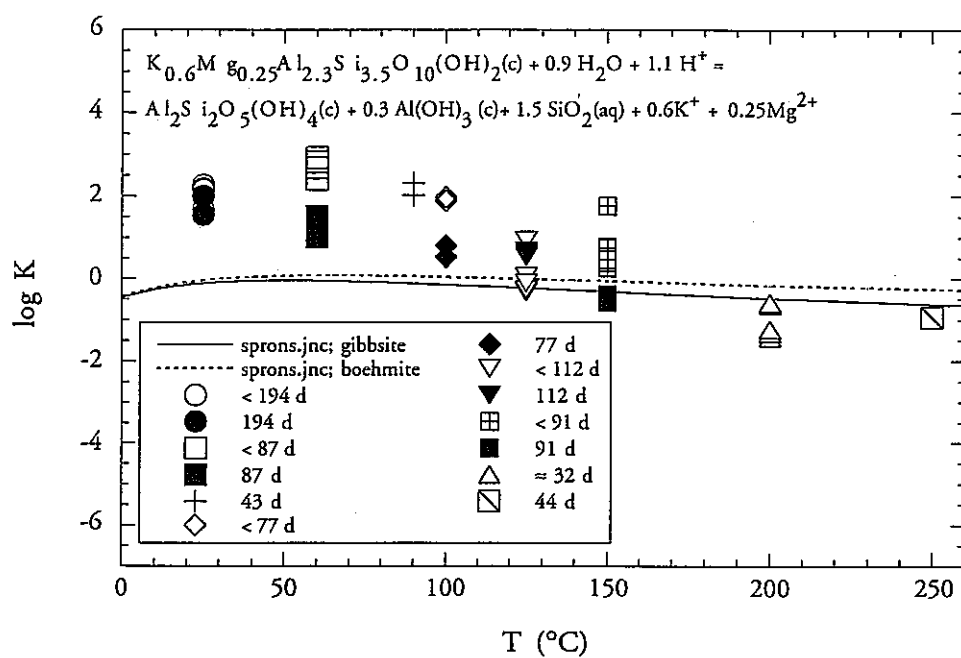


Fig C.59: Logarithm of the equilibrium constant for the reaction illite + 0.9 H₂O + 1.1 H⁺ ⇌ kaolinite + gibbsite + 1.5 SiO₂(aq) + 0.6 K⁺ + 0.25 Mg²⁺ between 0 and 250°C at *P*_{sat} [experimental data from Aja *et al.* (1991b)].

Appendix D: Calculated and Experimental Equilibrium Constants of Aqueous Reactions as Functions of Pressure and Temperature

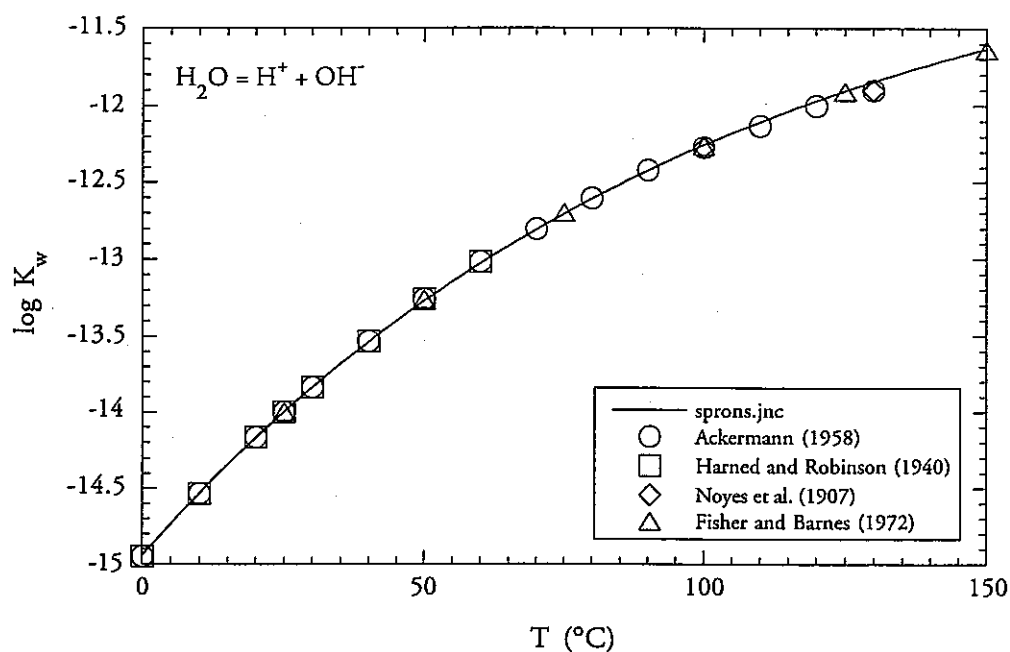


Fig. D.1: Logarithm of the equilibrium constant for the reaction $\text{H}_2\text{O} \rightleftharpoons \text{H}^+ + \text{OH}^-$ between 0 and 150°C at P_{sat} .

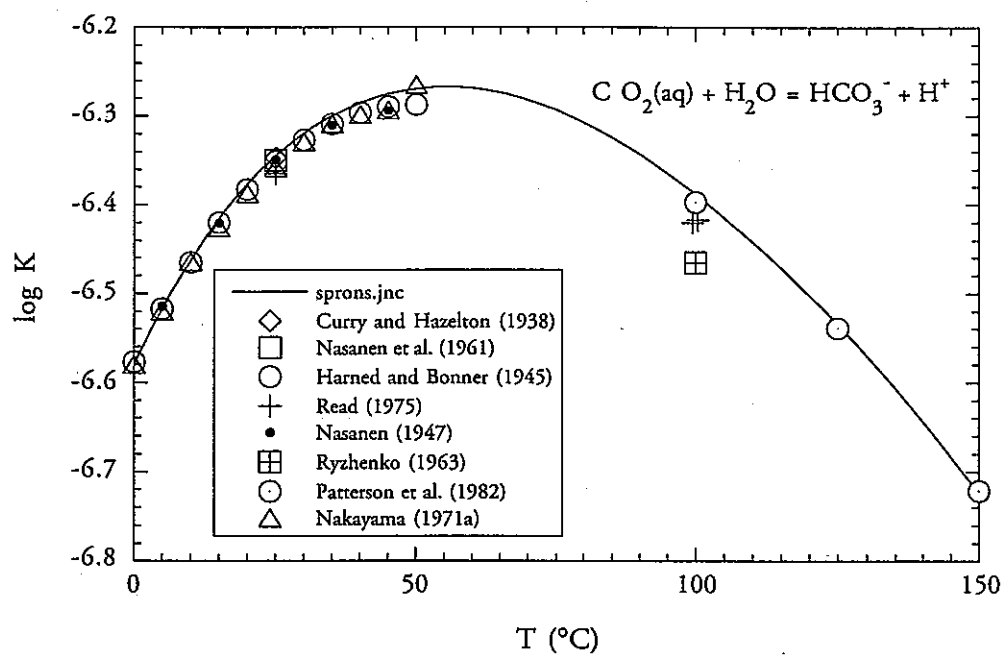


Fig. D.2: Logarithm of the equilibrium constant for the reaction $\text{CO}_2(\text{aq}) + \text{H}_2\text{O} \rightleftharpoons \text{HCO}_3^- + \text{H}^+$ between 0 and 150°C at P_{sat} .

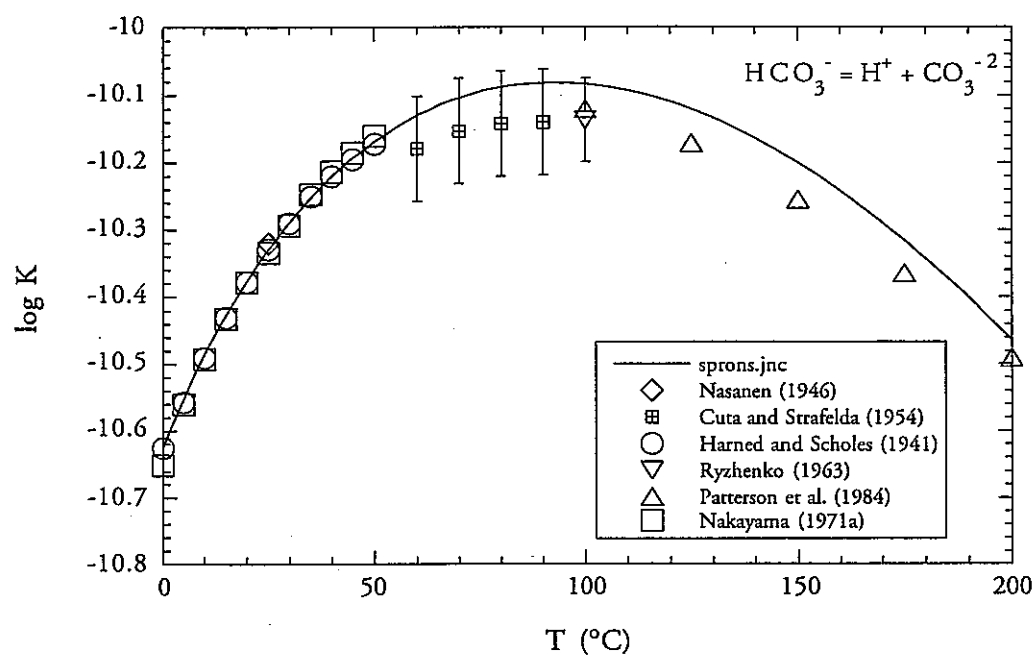


Fig. D.3: Logarithm of the equilibrium constant for the reaction $\text{HCO}_3^- \rightleftharpoons \text{H}^+ + \text{CO}_3^{2-}$ between 0 and 200°C at P_{sat} .

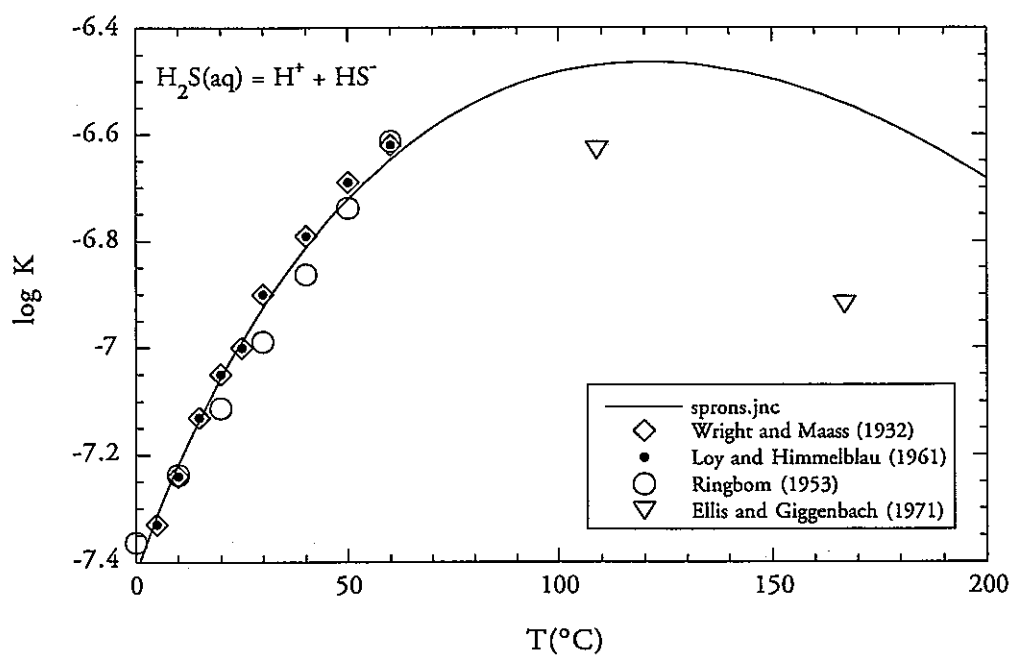


Fig. D.4: Logarithm of the equilibrium constant for the reaction $\text{H}_2\text{S}(\text{aq}) \rightleftharpoons \text{H}^+ + \text{HS}^-$ between 0 and 200°C at P_{sat} .

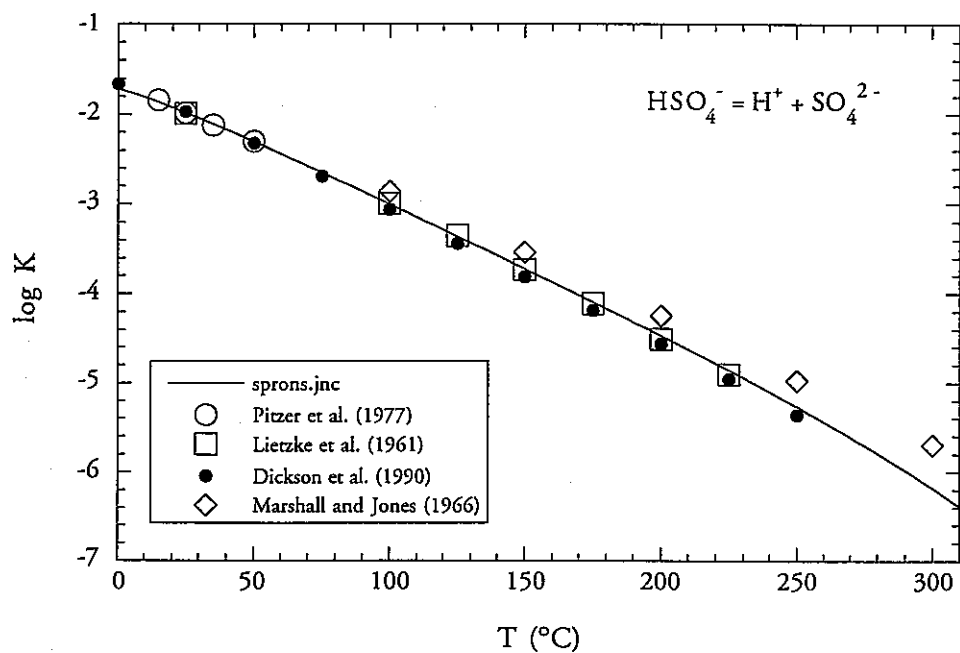


Fig. D.5: Logarithm of the equilibrium constant for the reaction $\text{HSO}_4^- \rightleftharpoons \text{H}^+ + \text{SO}_4^{2-}$ between 0 and 300 $^{\circ}\text{C}$ at P_{sat} .

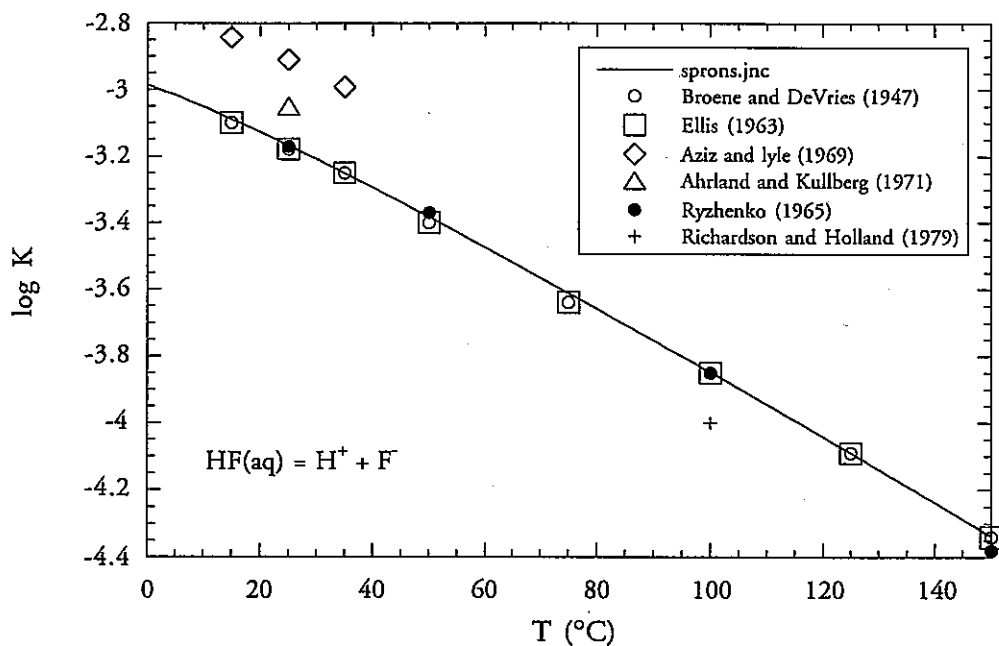


Fig. D.6: Logarithm of the equilibrium constant for the reaction $\text{HF}(\text{aq}) \rightleftharpoons \text{H}^+ + \text{F}^-$ between 0 and 150 $^{\circ}\text{C}$ at P_{sat} .

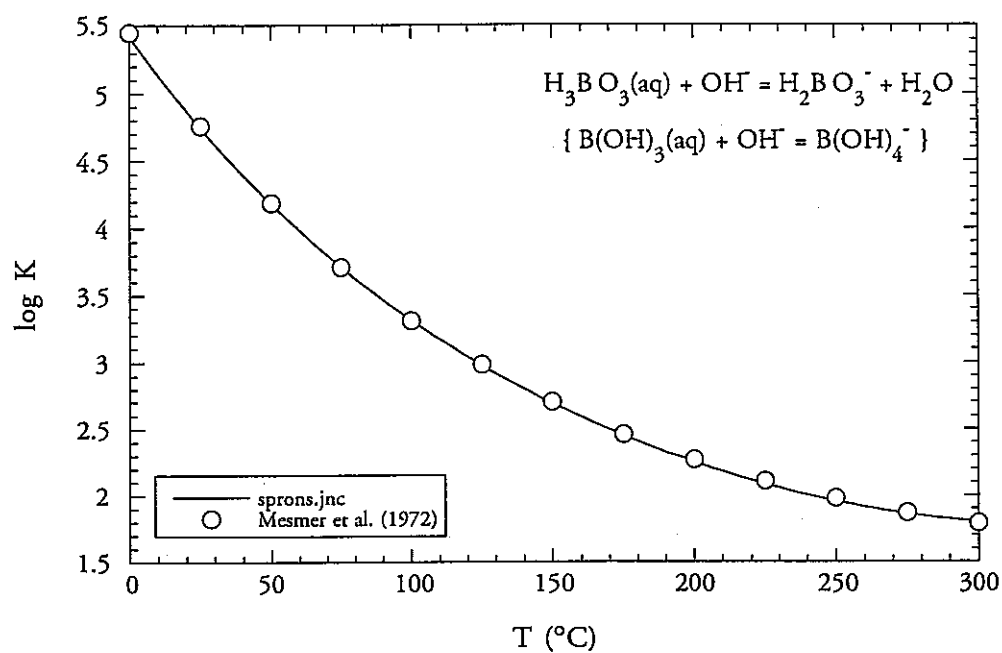


Fig. D.7: Logarithm of the equilibrium constant for the reaction $\text{H}_3\text{BO}_3(\text{aq}) + \text{OH}^- \rightleftharpoons \text{H}_2\text{BO}_3^- + \text{H}_2\text{O}$ between 0 and 300°C at P_{sat} .

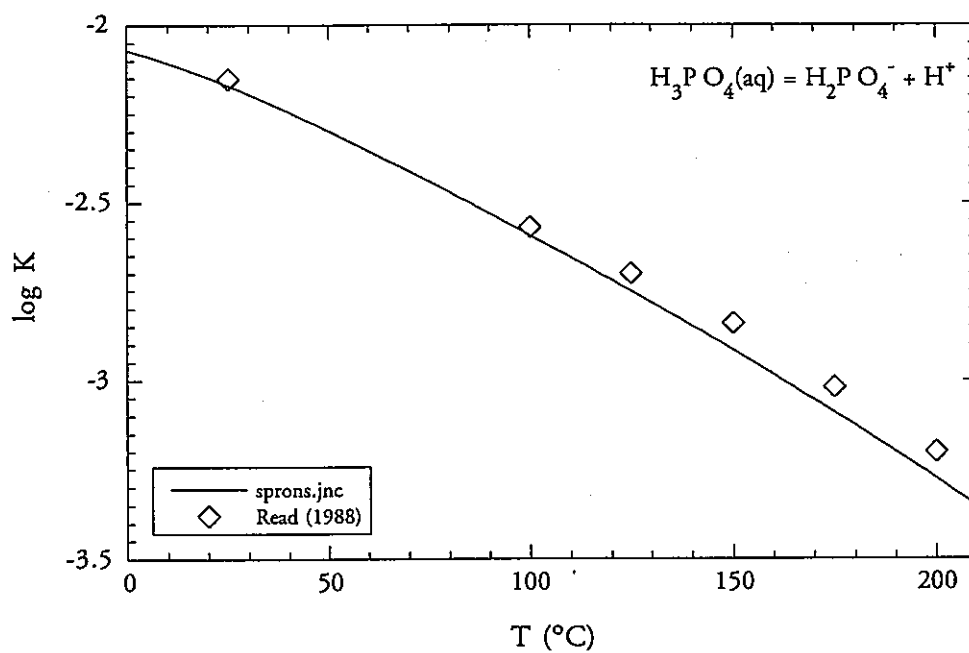


Fig. D.8: Logarithm of the equilibrium constant for the reaction $\text{H}_3\text{PO}_4(\text{aq}) \rightleftharpoons \text{H}_2\text{PO}_4^- + \text{H}^+$ between 0 and 200°C at P_{sat} .

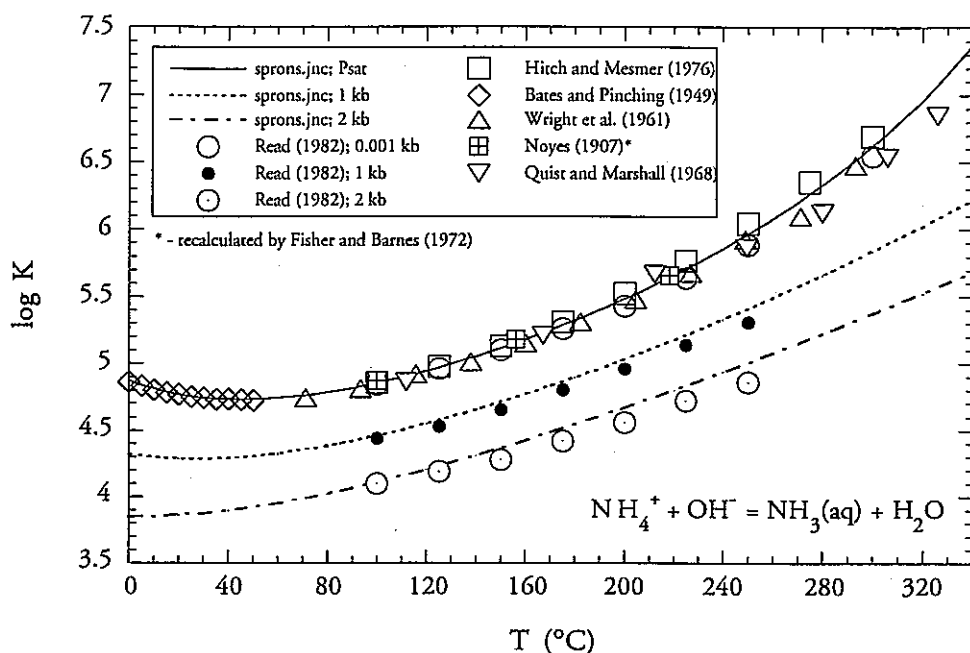


Fig. D.9: Logarithm of the equilibrium constant for the reaction $\text{NH}_4^+ + \text{OH}^- \rightleftharpoons \text{NH}_3(\text{aq}) + \text{H}_2\text{O}$ between 0 and 310°C and pressures from P_{sat} to 2 kb.

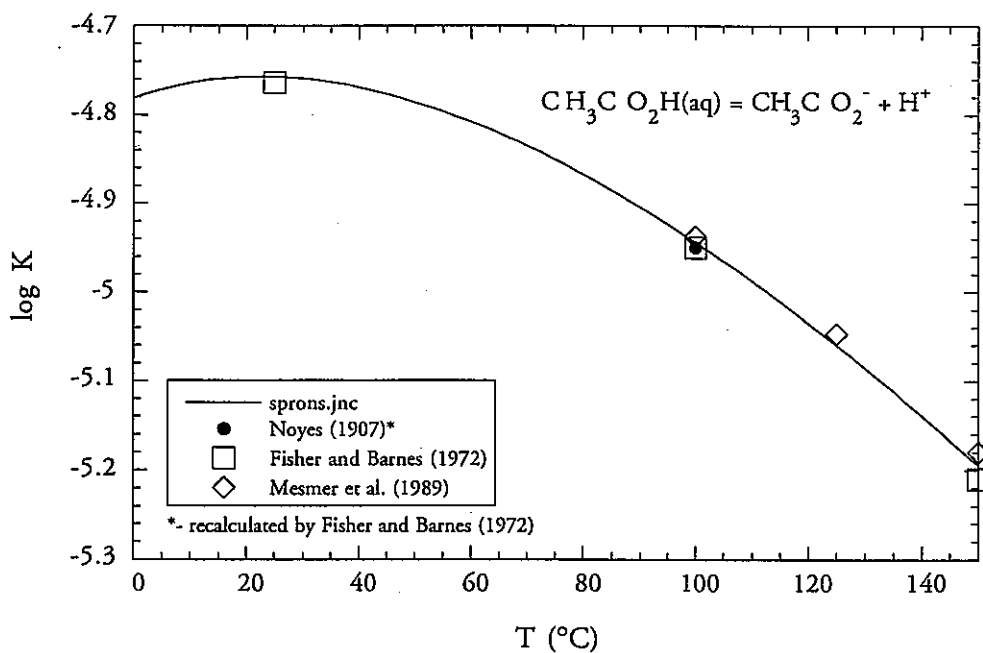


Fig. D.10: Logarithm of the equilibrium constant for the reaction $\text{CH}_3\text{COOH} \rightleftharpoons \text{CH}_3\text{COO}^- + \text{H}^+$ between 0 and 150°C at P_{sat} .

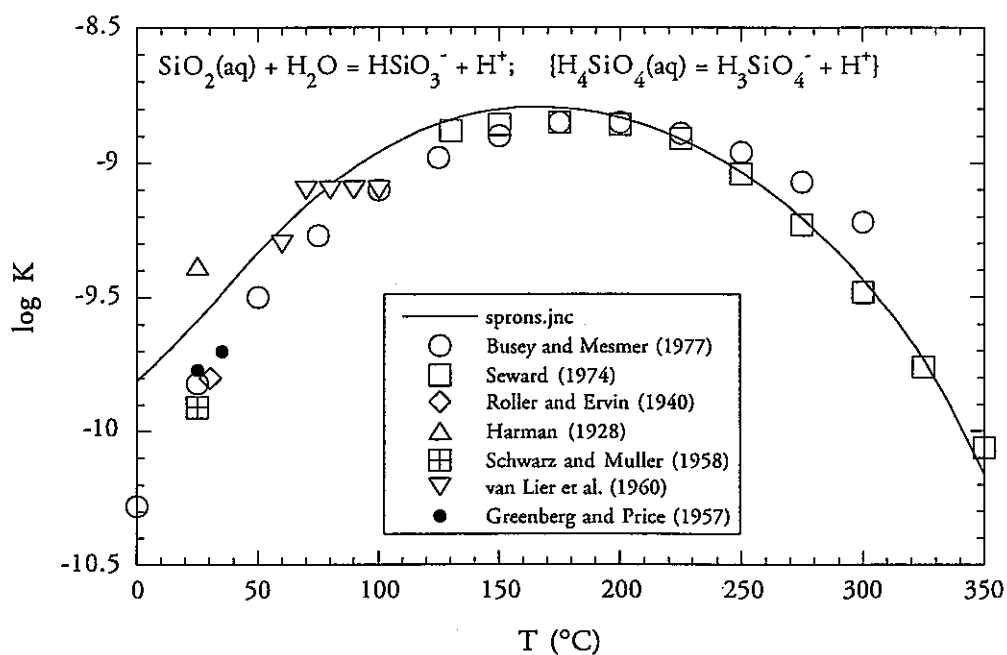


Fig. D.11: Logarithm of the equilibrium constant for the reaction $\text{H}_4\text{SiO}_4(\text{aq}) = \text{H}_3\text{SiO}_4^- + \text{H}^+$ between 0 and 350°C at P_{sat} .

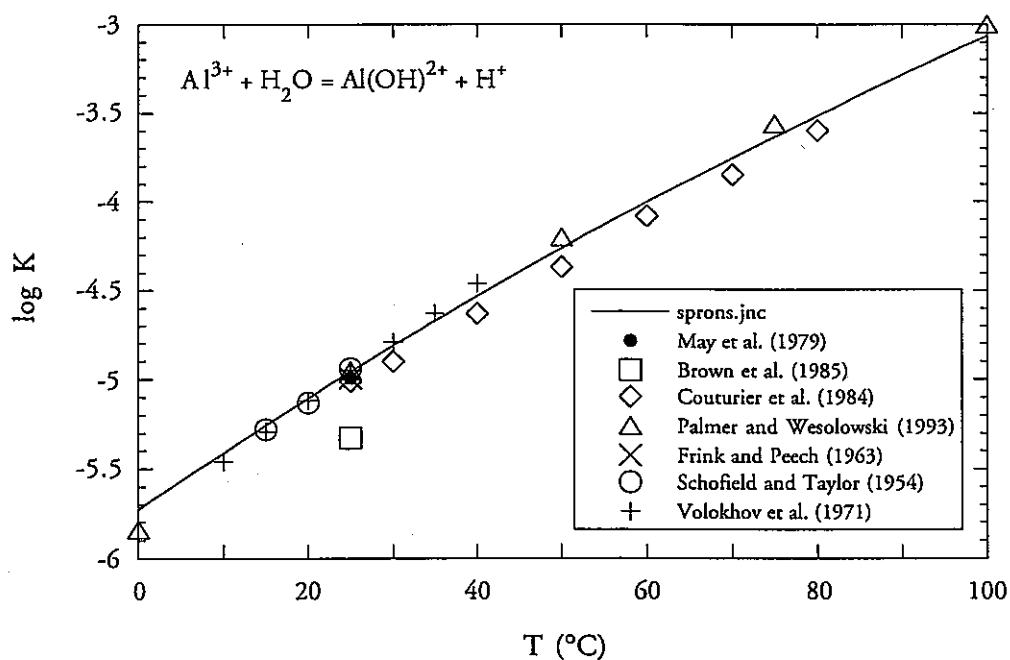


Fig. D.12: Logarithm of the equilibrium constant for the reaction $\text{Al}^{3+} + \text{H}_2\text{O} = \text{Al}(\text{OH})^{2+} + \text{H}^+$ between 0 and 100°C at 1 bar.

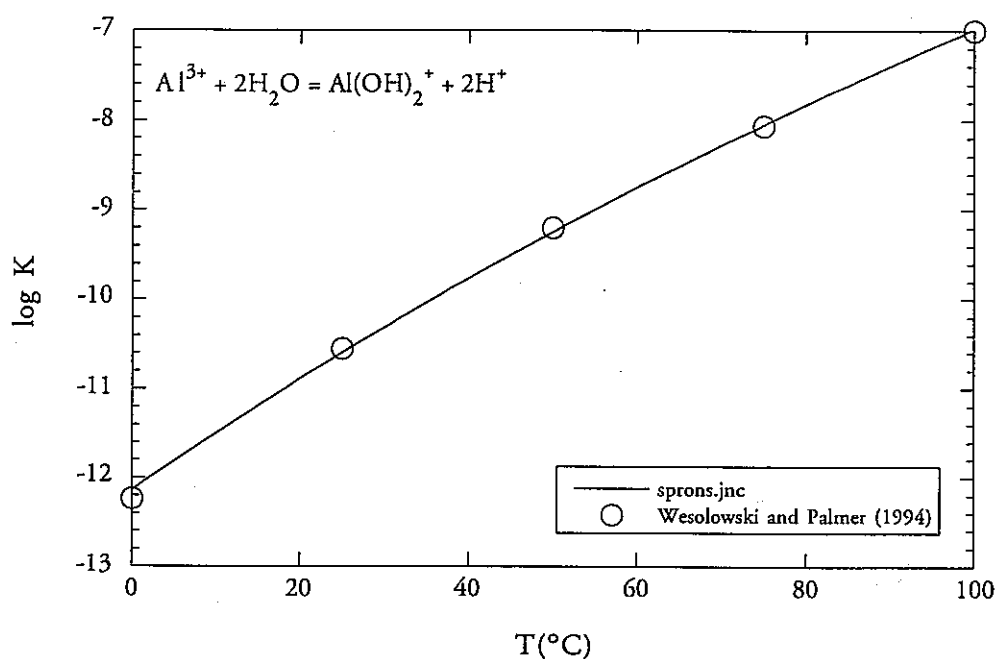


Fig. D.13: Logarithm of the equilibrium constant for the reaction $\text{Al}^{3+} + 2\text{H}_2\text{O} \rightleftharpoons \text{Al}(\text{OH})_2^+ + 2\text{H}^+$ between 0 and 100°C at 1 bar.

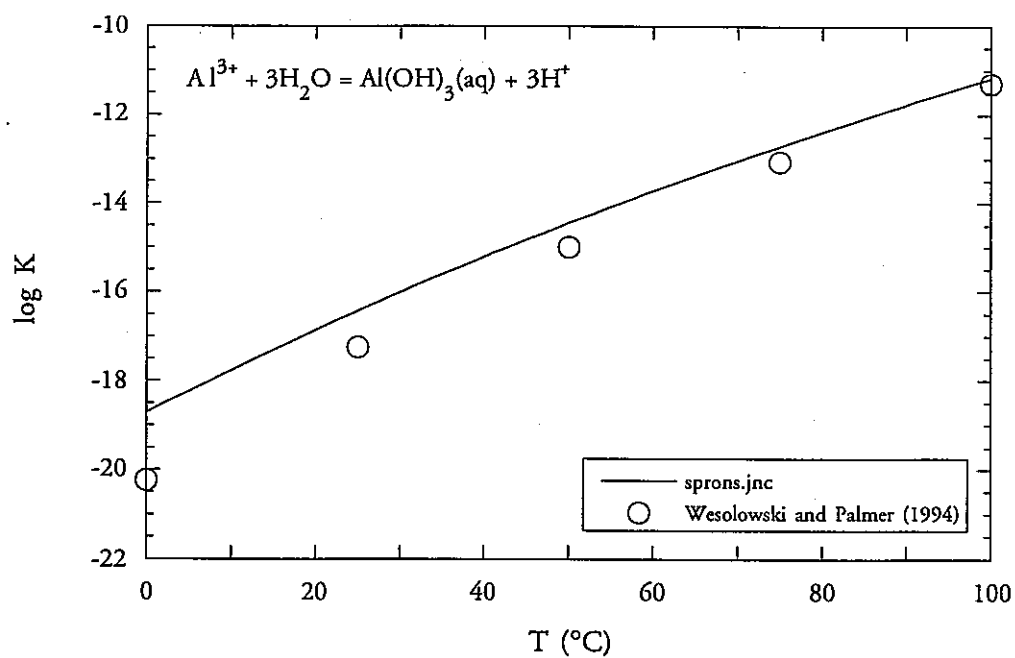


Fig. D.14: Logarithm of the equilibrium constant for the reaction $\text{Al}^{3+} + 3\text{H}_2\text{O} \rightleftharpoons \text{Al}(\text{OH})_3(\text{aq}) + 3\text{H}^+$ between 0 and 100°C at 1 bar.

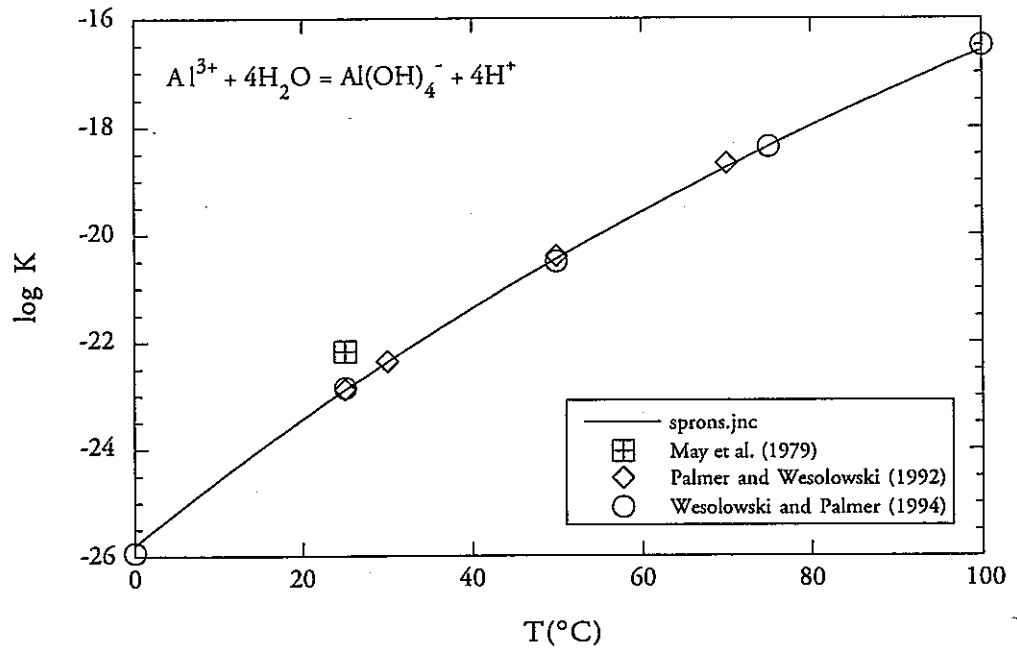


Fig. D.15: Logarithm of the equilibrium constant for the reaction $\text{Al}^{3+} + 4\text{H}_2\text{O} \rightleftharpoons \text{Al}(\text{OH})_4^- + 4\text{H}^+$ between 0 and 100°C at 1 bar.

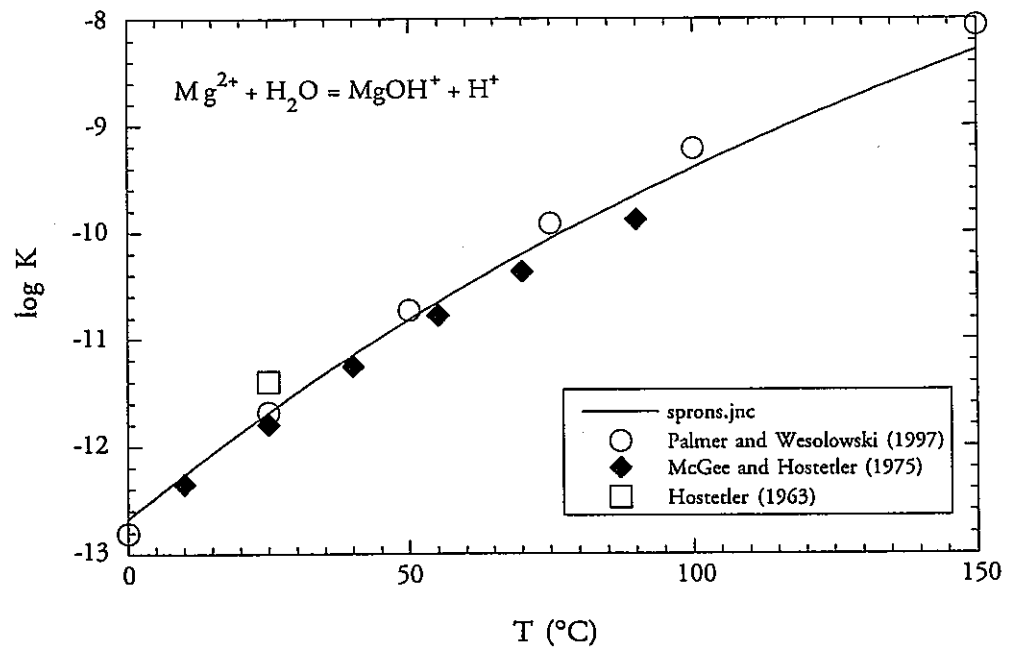


Fig. D.16: Logarithm of the equilibrium constant for the reaction $\text{Mg}^{2+} + \text{H}_2\text{O} \rightleftharpoons \text{MgOH}^+ + \text{H}^+$ between 0 and 150°C at P_{sat} .

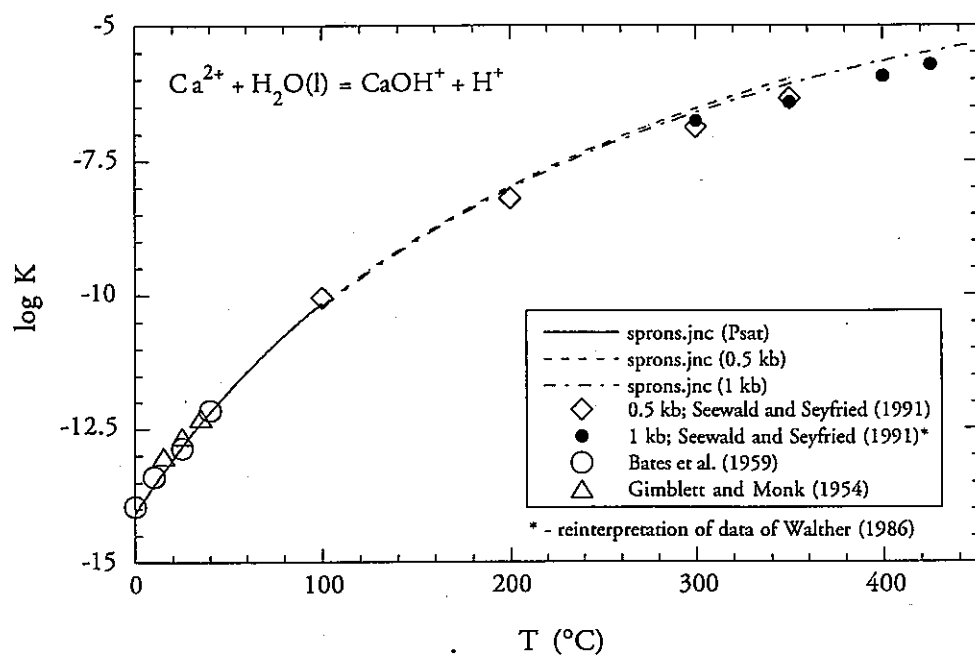


Fig. D.17: Logarithm of the equilibrium constant for the reaction $\text{Ca}^{2+} + \text{H}_2\text{O} \rightleftharpoons \text{CaOH}^+ + \text{H}^+$ between 0 and 450°C and pressures from P_{sat} to 1 kb.

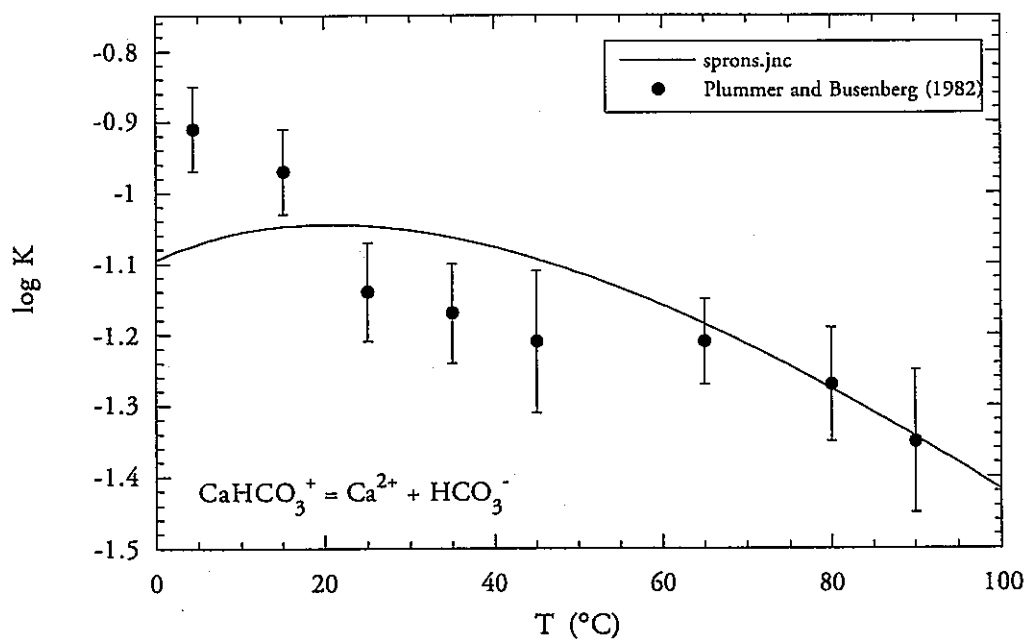


Fig. D.18: Logarithm of the equilibrium constant for the reaction $\text{CaHCO}_3^+ \rightleftharpoons \text{Ca}^{2+} + \text{HCO}_3^-$ between 0 and 100°C at 1 bar.

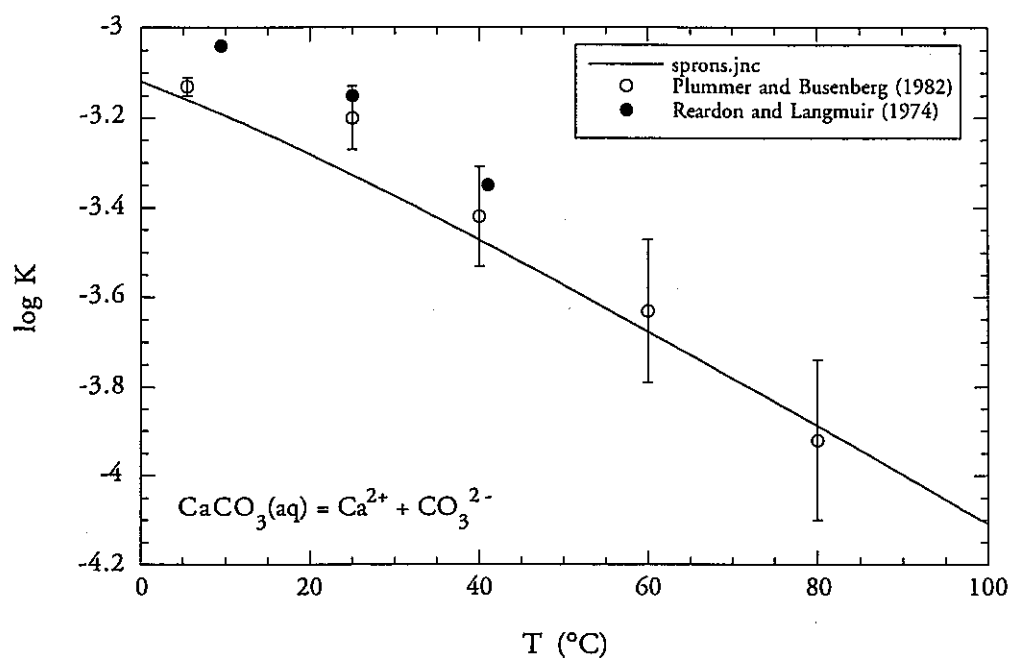


Fig. D.19: Logarithm of the equilibrium constant for the reaction $\text{CaCO}_3(\text{aq}) \rightleftharpoons \text{Ca}^{2+} + \text{CO}_3^{2-}$ between 0 and 100°C at 1 bar.

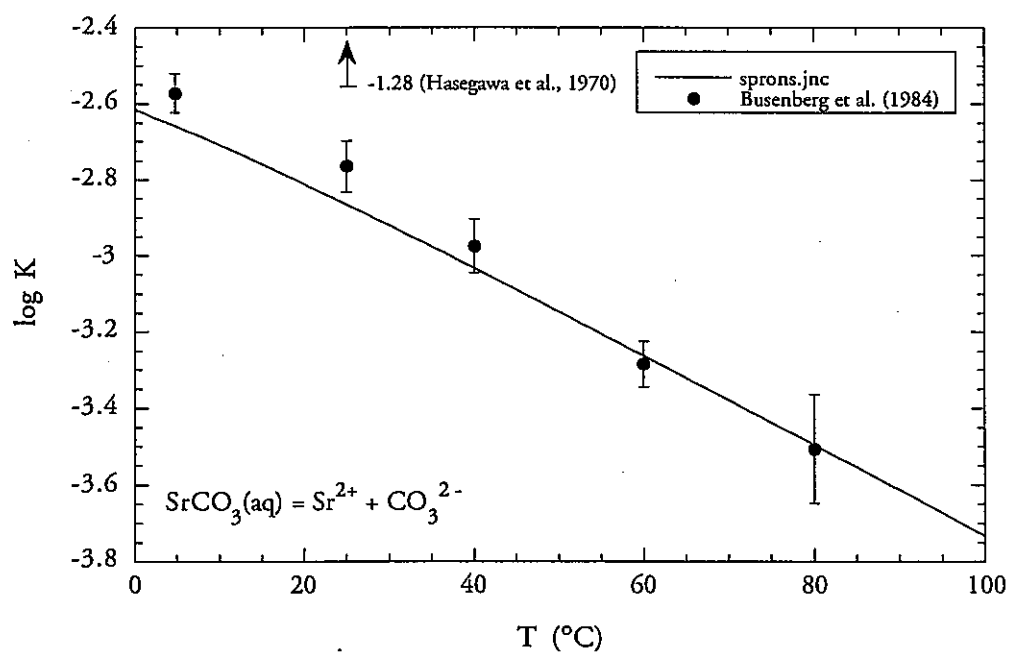


Fig. D.20: Logarithm of the equilibrium constant for the reaction $\text{SrCO}_3(\text{aq}) \rightleftharpoons \text{Sr}^{2+} + \text{CO}_3^{2-}$ between 0 and 100°C at 1 bar.

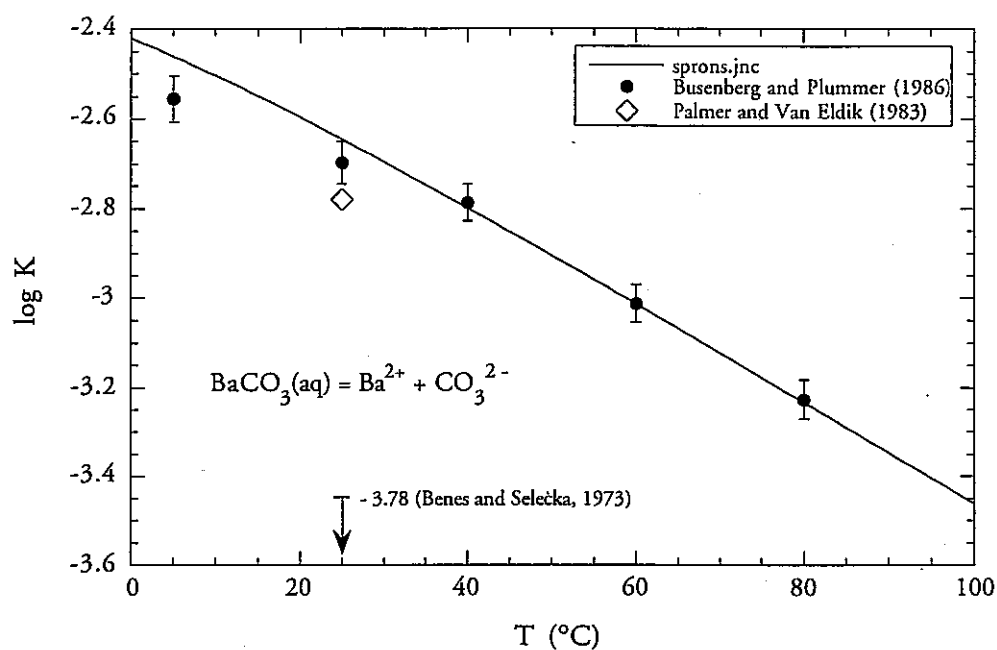


Fig. D.21: Logarithm of the equilibrium constant for the reaction $\text{BaCO}_3(\text{aq}) \rightleftharpoons \text{Ba}^{2+} + \text{CO}_3^{2-}$ between 0 and 100°C at 1 bar.

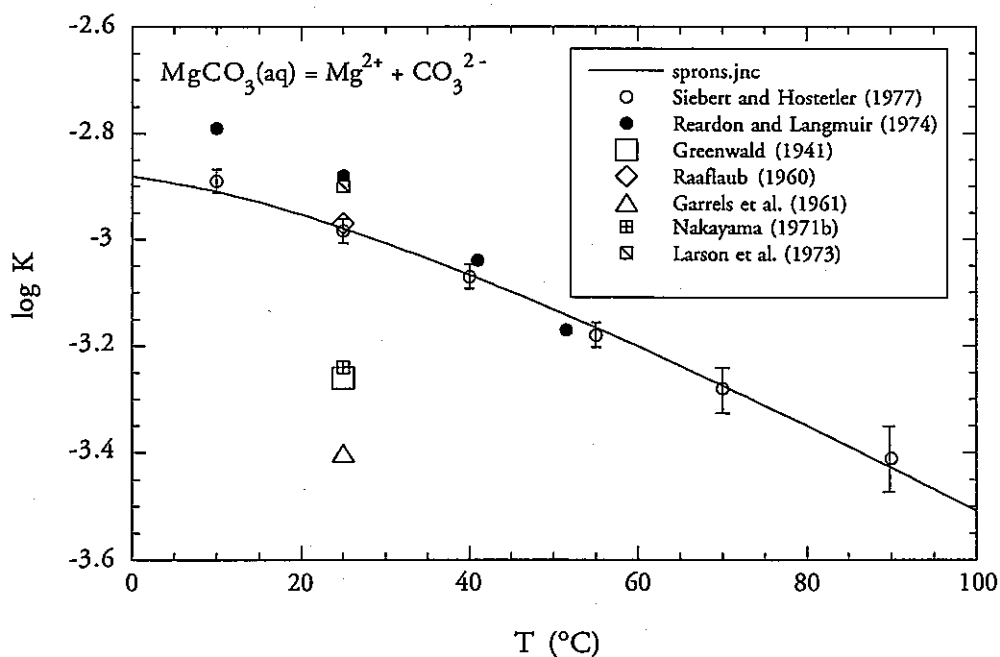


Fig. D.22: Logarithm of the equilibrium constant for the reaction $\text{MgCO}_3(\text{aq}) \rightleftharpoons \text{Mg}^{2+} + \text{CO}_3^{2-}$ between 0 and 100°C at 1 bar.

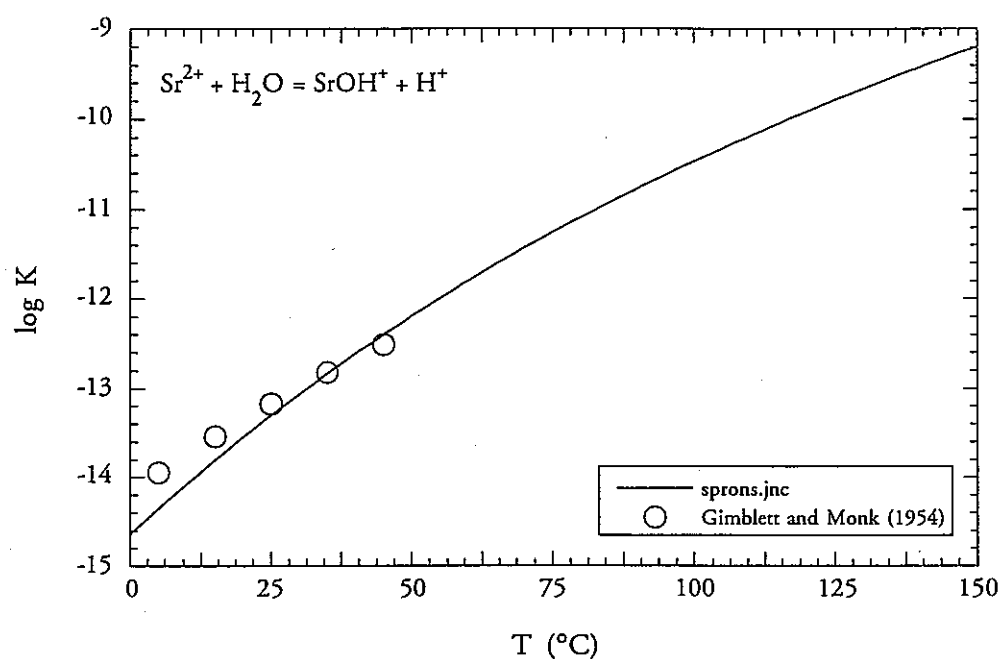


Fig. D.23: Logarithm of the equilibrium constant for the reaction $\text{Sr}^{2+} + \text{H}_2\text{O} \rightleftharpoons \text{SrOH}^+ + \text{H}^+$ between 0 and 150°C at P_{sat} .

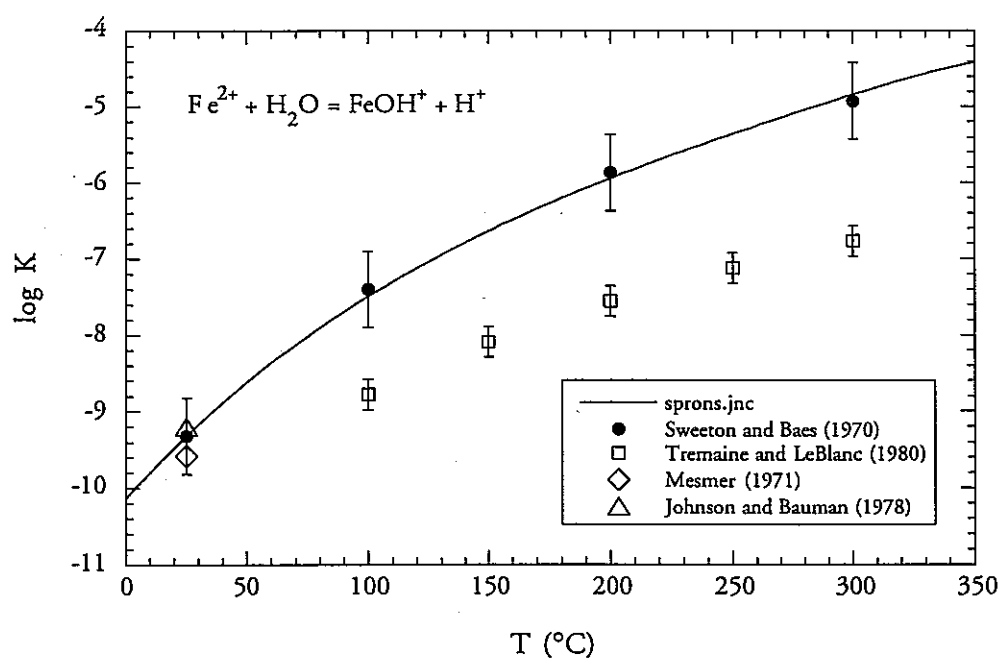


Fig. D.24: Logarithm of the equilibrium constant for the reaction $\text{Fe}^{2+} + \text{H}_2\text{O} \rightleftharpoons \text{FeOH}^+ + \text{H}^+$ between 0 and 350°C at P_{sat} .

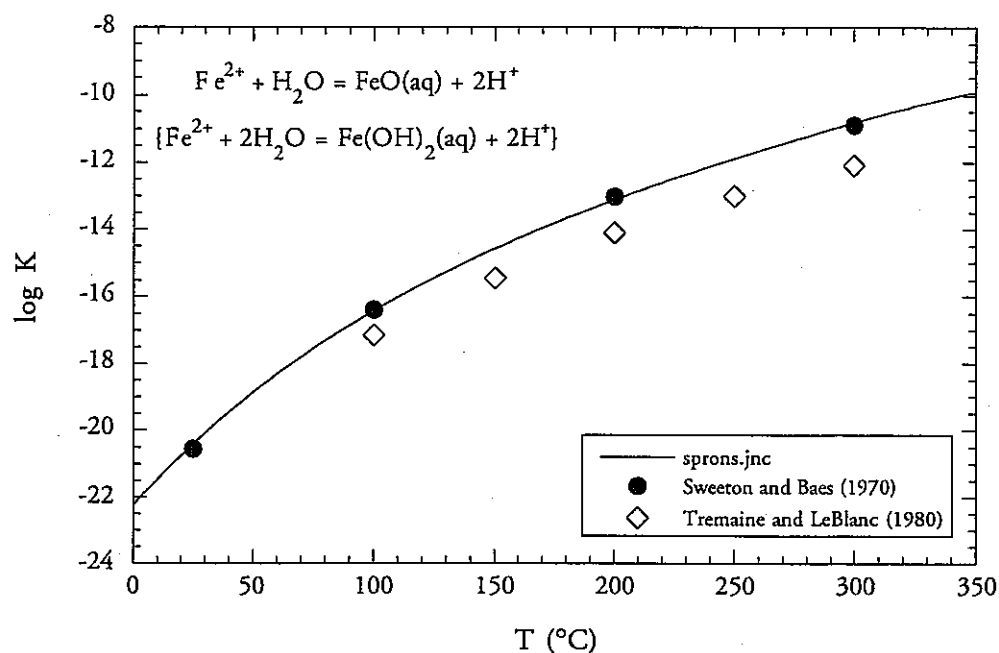


Fig. D.25: Logarithm of the equilibrium constant for the reaction $\text{Fe}^{2+} + 2\text{H}_2\text{O} = \text{Fe}(\text{OH})_2(\text{aq}) + 2\text{H}^+$ between 0 and 350 $^{\circ}\text{C}$ at P_{sat} .

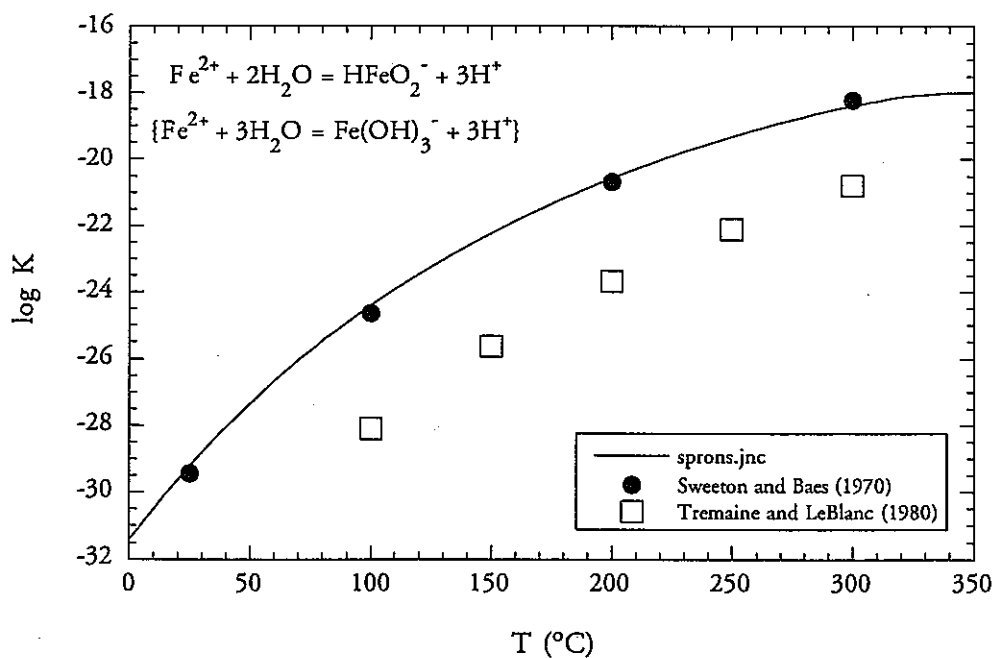


Fig. D.26: Logarithm of the equilibrium constant for the reaction $\text{Fe}^{2+} + 3\text{H}_2\text{O} = \text{Fe}(\text{OH})_3^- + 3\text{H}^+$ between 0 and 350 $^{\circ}\text{C}$ at P_{sat} .

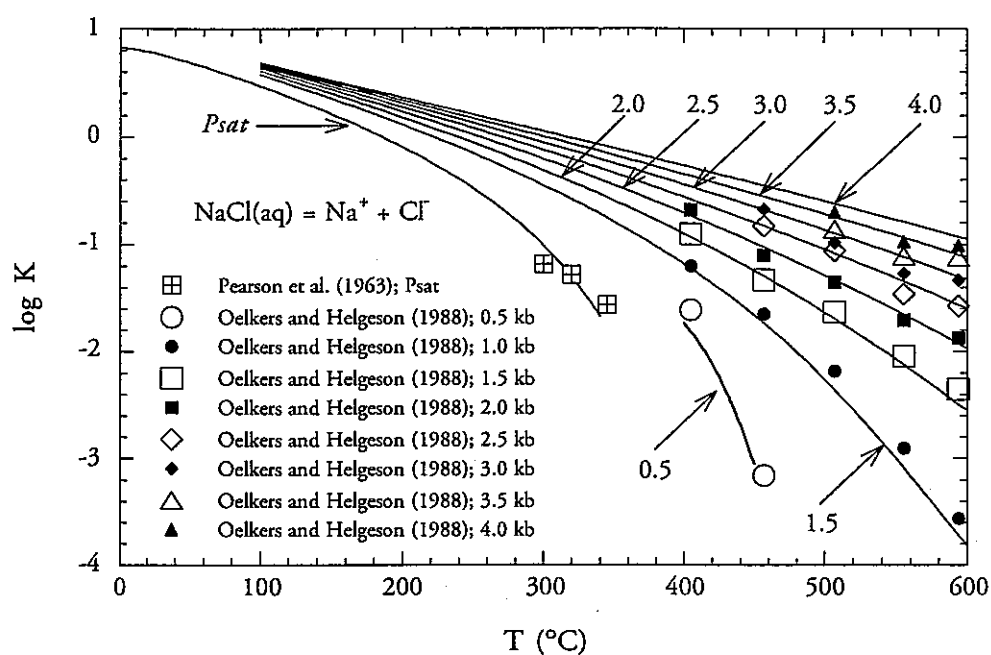


Fig. D.27: Logarithm of the equilibrium constant for the reaction $\text{NaCl(aq)} \rightleftharpoons \text{Na}^+ + \text{Cl}^-$ between 0 and 600°C and pressures from P_{sat} to 4 kb.

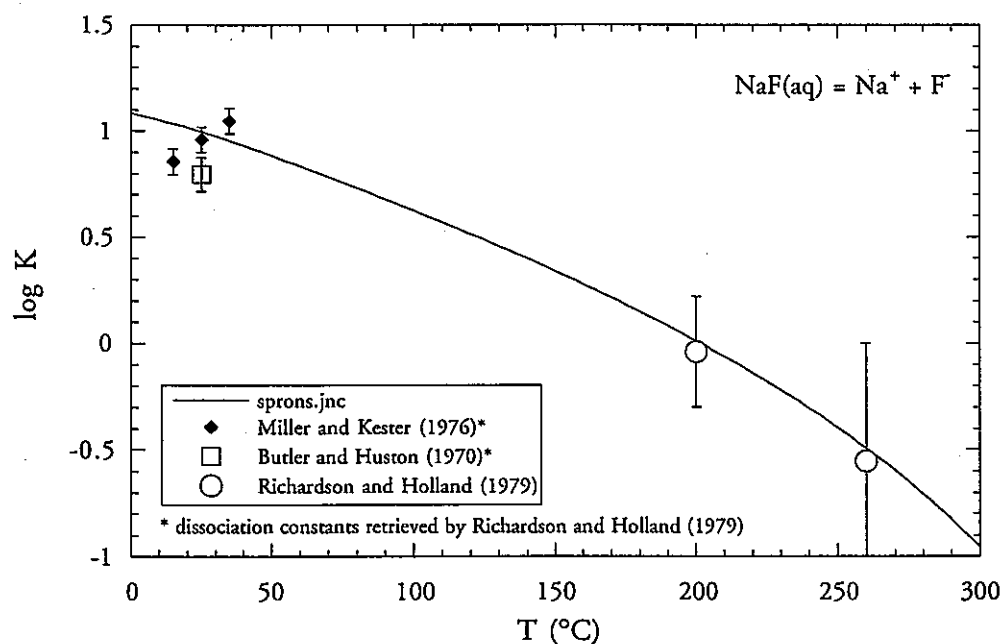


Fig. D.28: Logarithm of the equilibrium constant for the reaction $\text{NaF(aq)} \rightleftharpoons \text{Na}^+ + \text{F}^-$ between 0 and 300°C at P_{sat} .

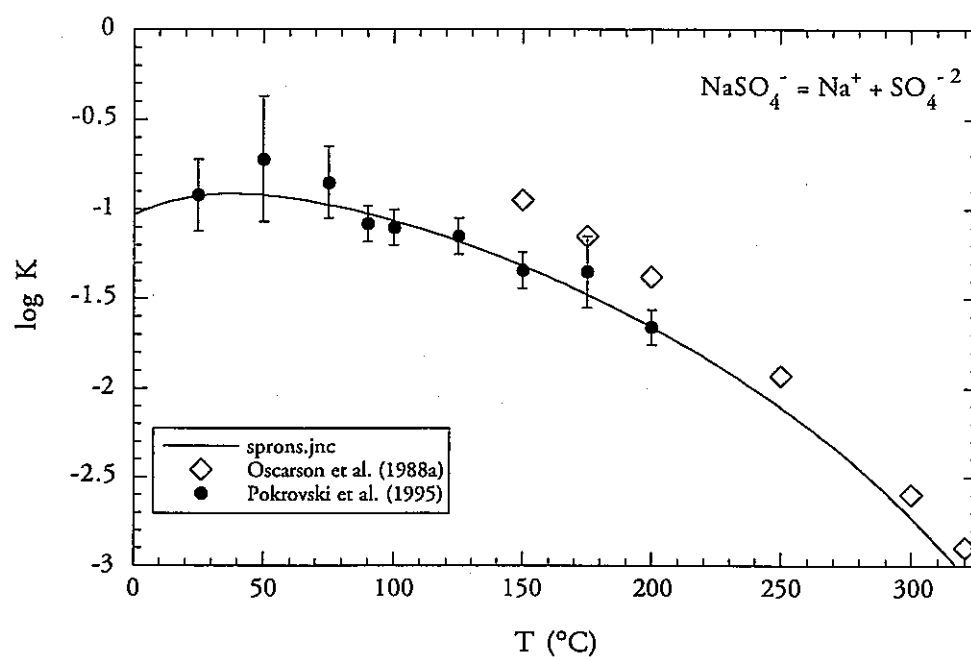


Fig. D.29: Logarithm of the equilibrium constant for the reaction $\text{NaSO}_4^- = \text{Na}^+ + \text{SO}_4^{2-}$ between 0 and 330 $^{\circ}\text{C}$ at P_{sat} .

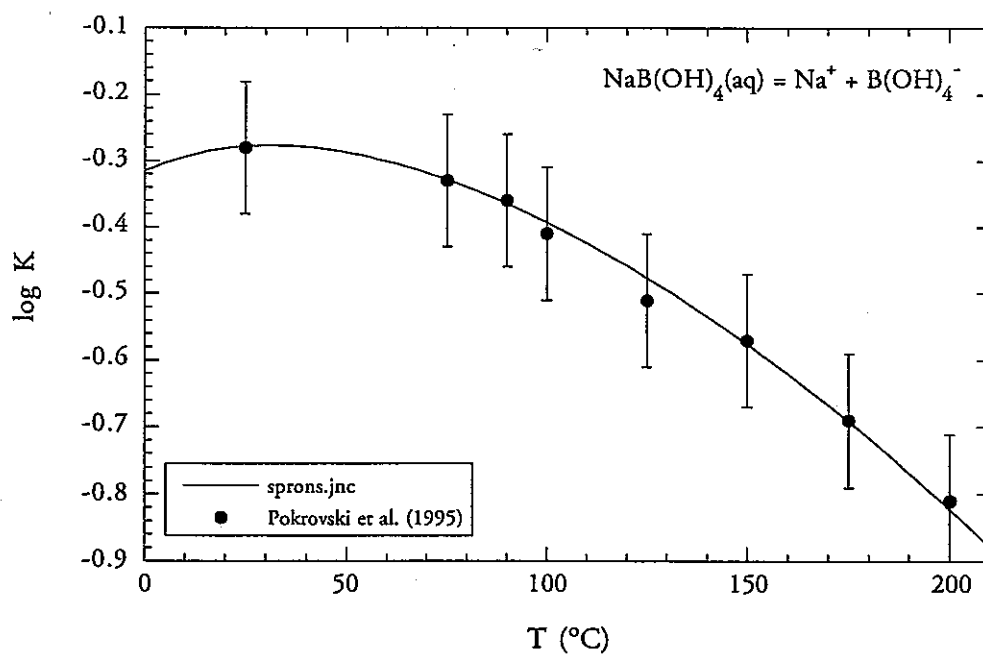


Fig. D.30: Logarithm of the equilibrium constant for the reaction $\text{NaB(OH)}_4(\text{aq}) = \text{Na}^+ + \text{B(OH)}_4^-$ between 0 and 200 $^{\circ}\text{C}$ at P_{sat} .

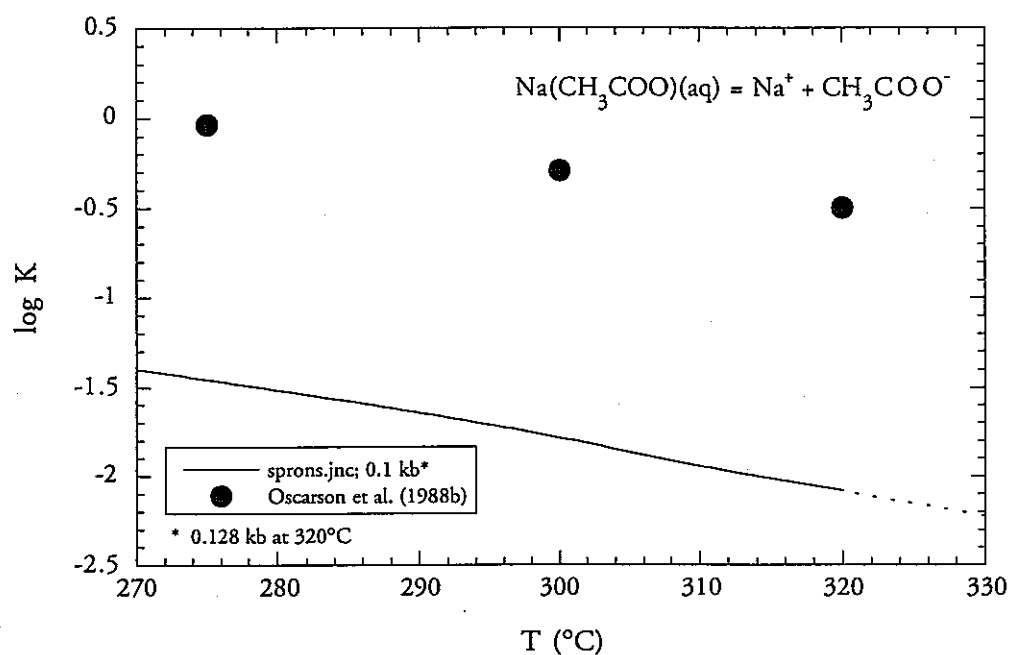


Fig. D.31: Logarithm of the equilibrium constant for the reaction $\text{Na}(\text{CH}_3\text{COO})(\text{aq}) \rightleftharpoons \text{Na}^+ + \text{CH}_3\text{COO}^-$ between 270 and 320°C at P_{sat} .

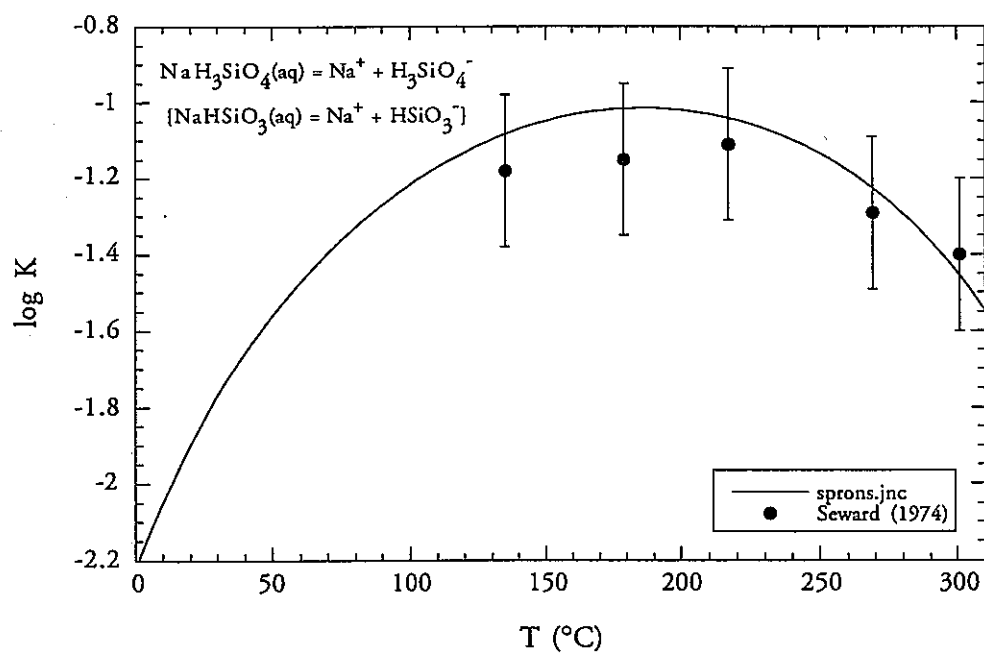


Fig. D.32: Logarithm of the equilibrium constant for the reaction $\text{NaH}_3\text{SiO}_4(\text{aq}) \rightleftharpoons \text{Na}^+ + \text{H}_3\text{SiO}_4^-$ between 0 and 300°C at P_{sat} .

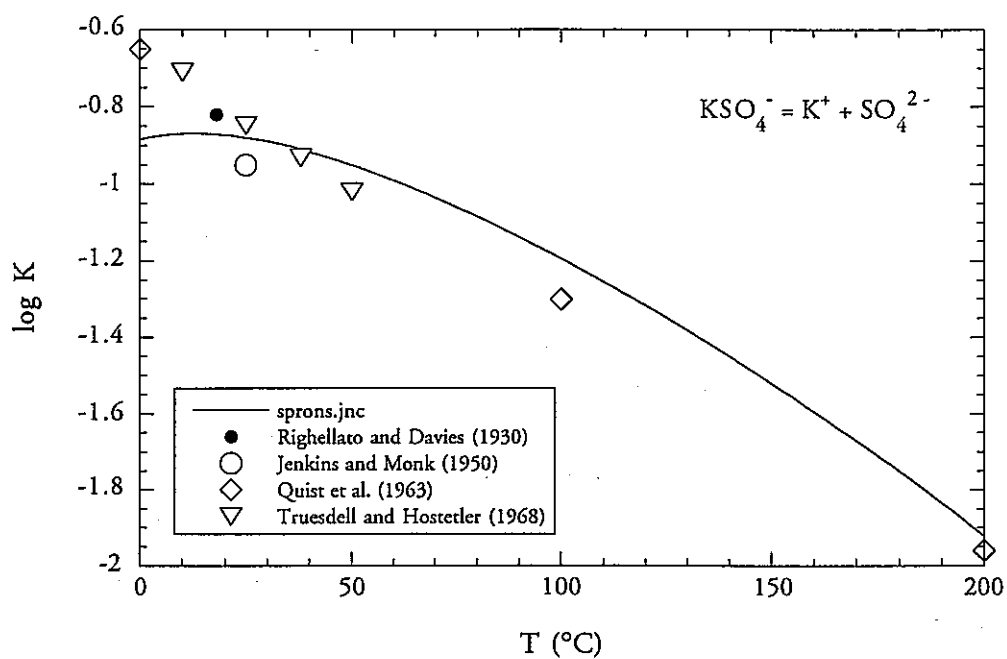


Fig. D.33: Logarithm of the equilibrium constant for the reaction $\text{KSO}_4^- = \text{K}^+ + \text{SO}_4^{2-}$ between 0 and 200 $^{\circ}\text{C}$ at P_{sat} .

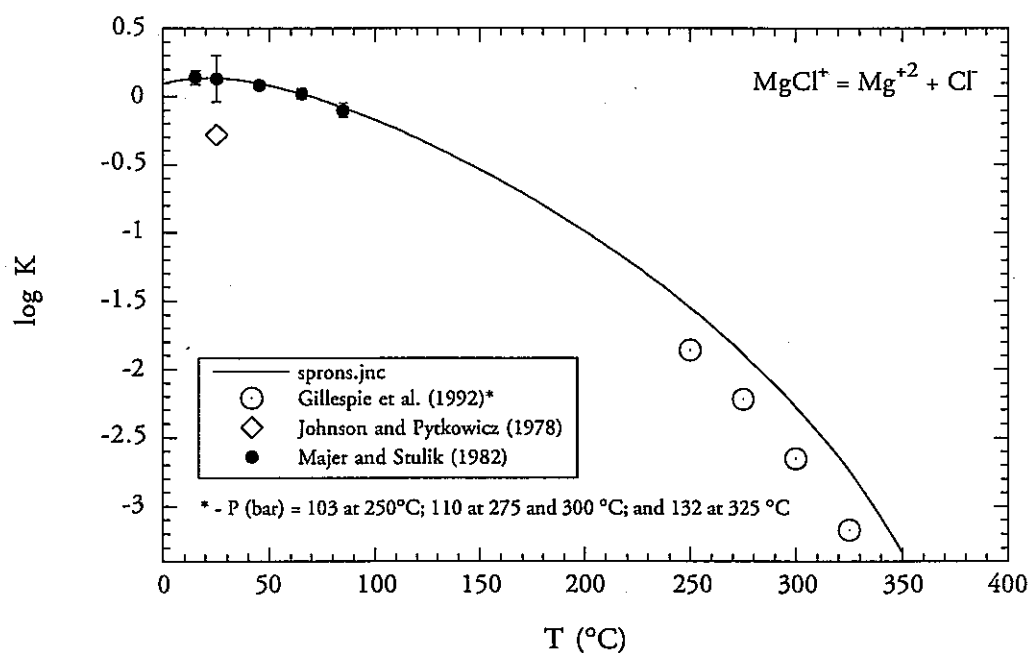


Fig. D.34: Logarithm of the equilibrium constant for the reaction $\text{MgCl}^+ = \text{Mg}^{2+} + \text{Cl}^-$ between 0 and 350 $^{\circ}\text{C}$ at P_{sat} .

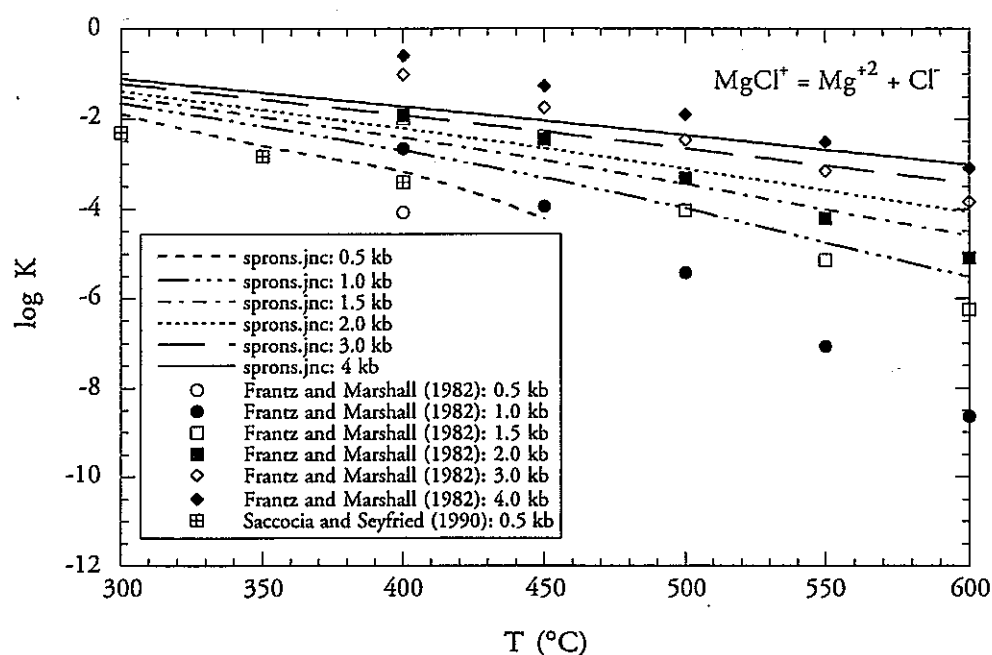


Fig. D.35: Logarithm of the equilibrium constant for the reaction $\text{MgCl}^+ \rightleftharpoons \text{Mg}^{2+} + \text{Cl}^-$ between 300 and 600°C and pressures from 0.5 to 4 kb.

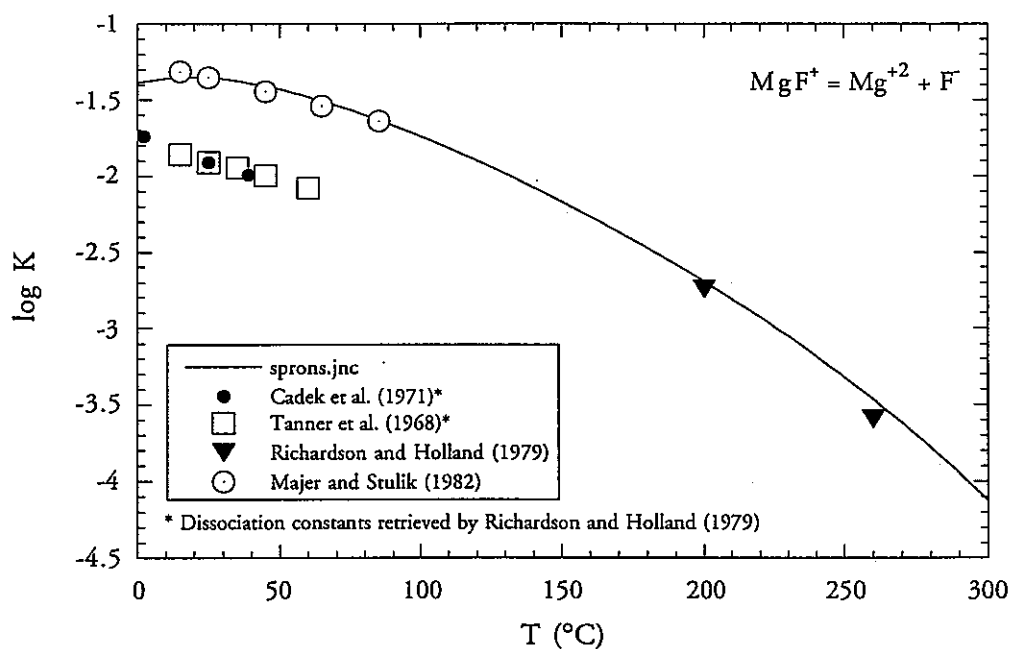


Fig. D.36: Logarithm of the equilibrium constant for the reaction $\text{MgF}^+ \rightleftharpoons \text{Mg}^{2+} + \text{F}^-$ between 0 and 300°C at P_{sat} .

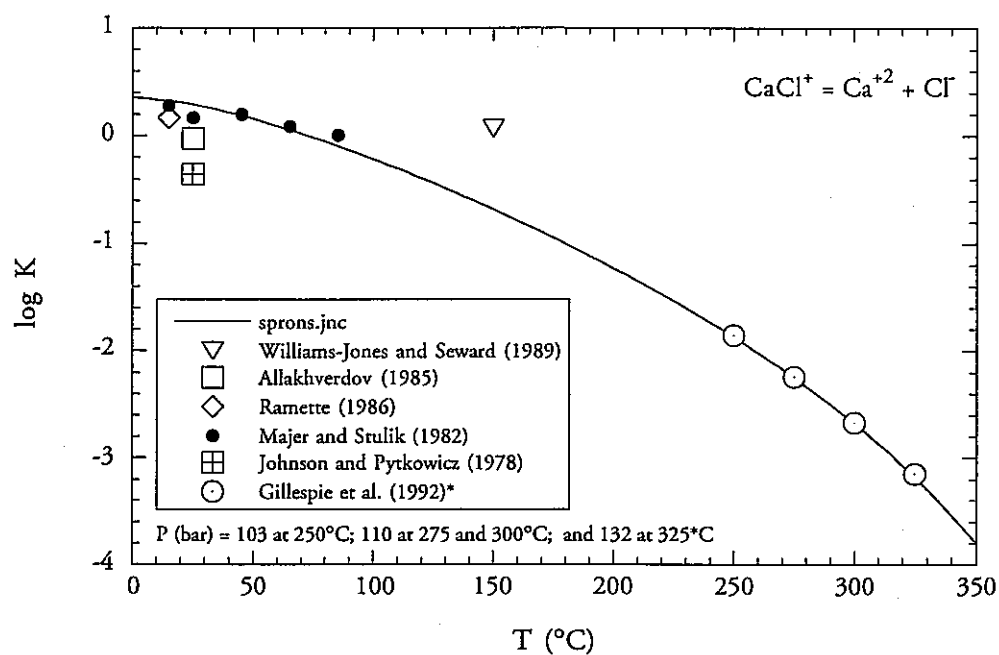


Fig. D.37: Logarithm of the equilibrium constant for the reaction $\text{CaCl}^+ \rightleftharpoons \text{Ca}^{2+} + \text{Cl}^-$ between 0 and 350 $^{\circ}\text{C}$ at P_{sat} .

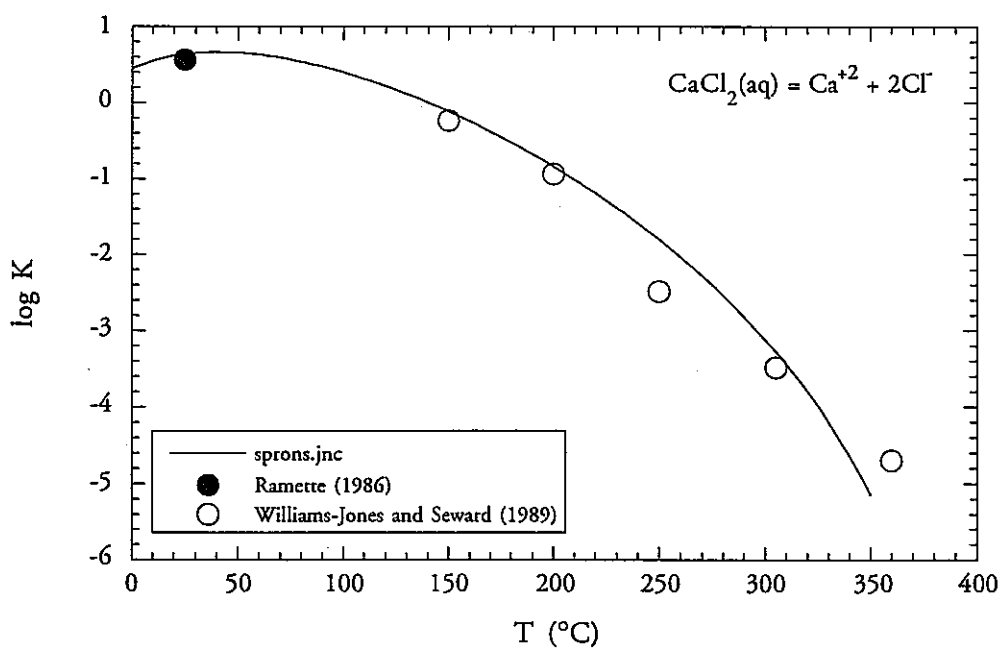


Fig. D.38: Logarithm of the equilibrium constant for the reaction $\text{CaCl}_2(\text{aq}) \rightleftharpoons \text{Ca}^{2+} + 2\text{Cl}^-$ between 0 and 350 $^{\circ}\text{C}$ at P_{sat} .

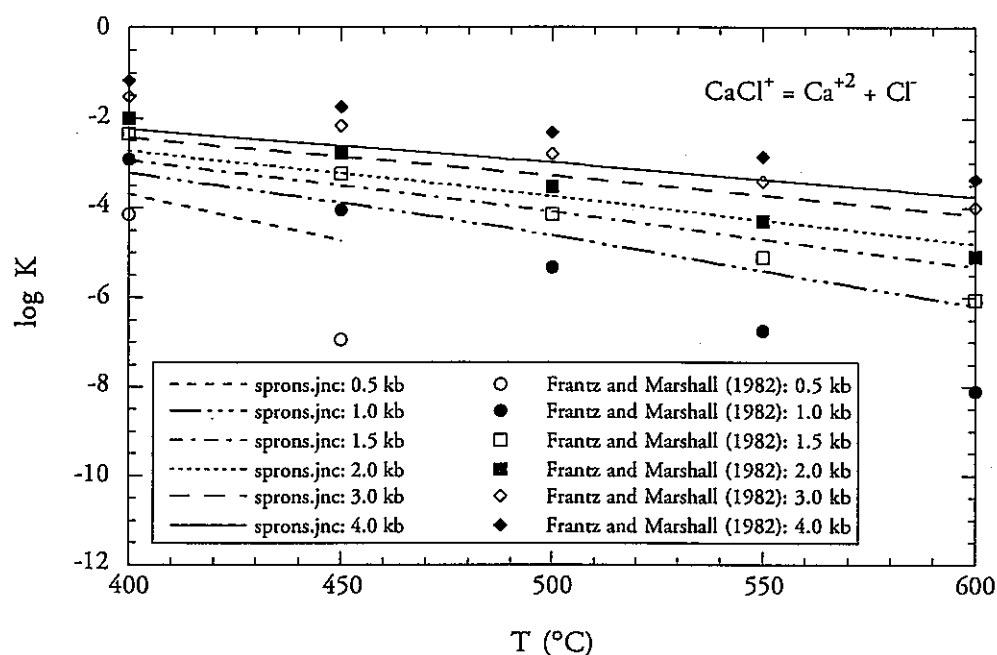


Fig. D.39: Logarithm of the equilibrium constant for the reaction $\text{CaCl}^+ \rightleftharpoons \text{Ca}^{2+} + \text{Cl}^-$ between 400 and 600°C and pressures from 0.5 to 4 kb.

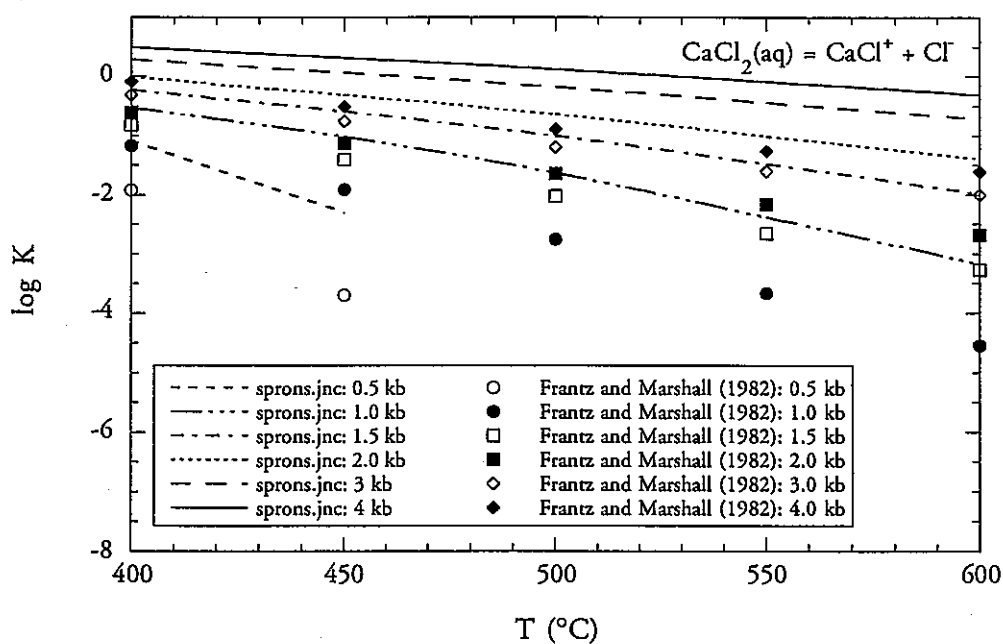


Fig. D.40: Logarithm of the equilibrium constant for the reaction $\text{CaCl}_2(\text{aq}) \rightleftharpoons \text{CaCl}^+ + \text{Cl}^-$ between 400 and 600°C and pressures from 0.5 to 4 kb.

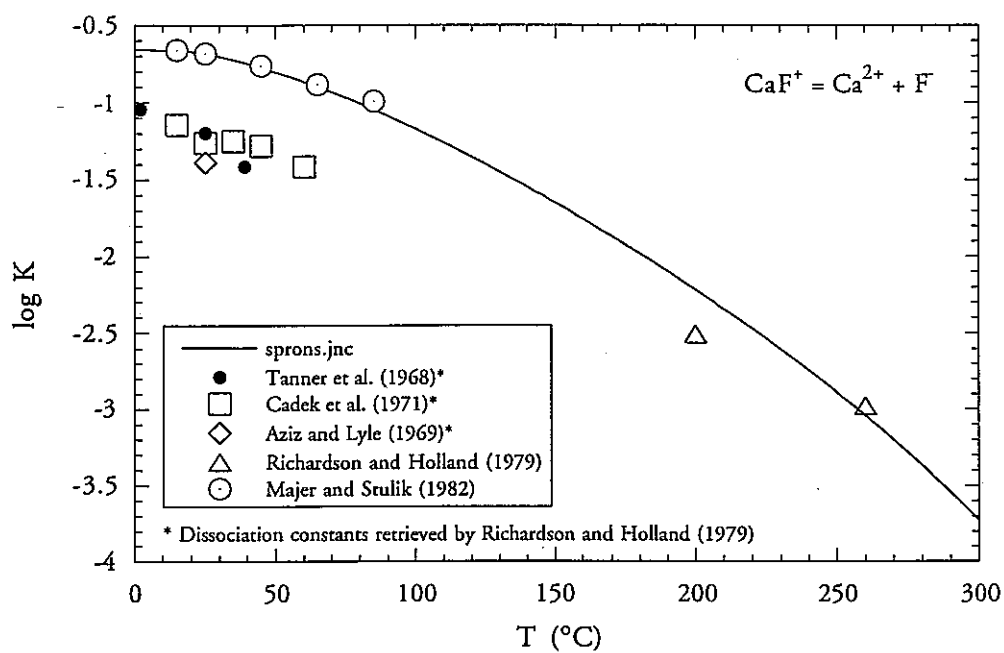


Fig. D.41: Logarithm of the equilibrium constant for the reaction $\text{CaF}^+ \rightleftharpoons \text{Ca}^{2+} + \text{F}^-$ between 0 and 300°C at P_{stat} .

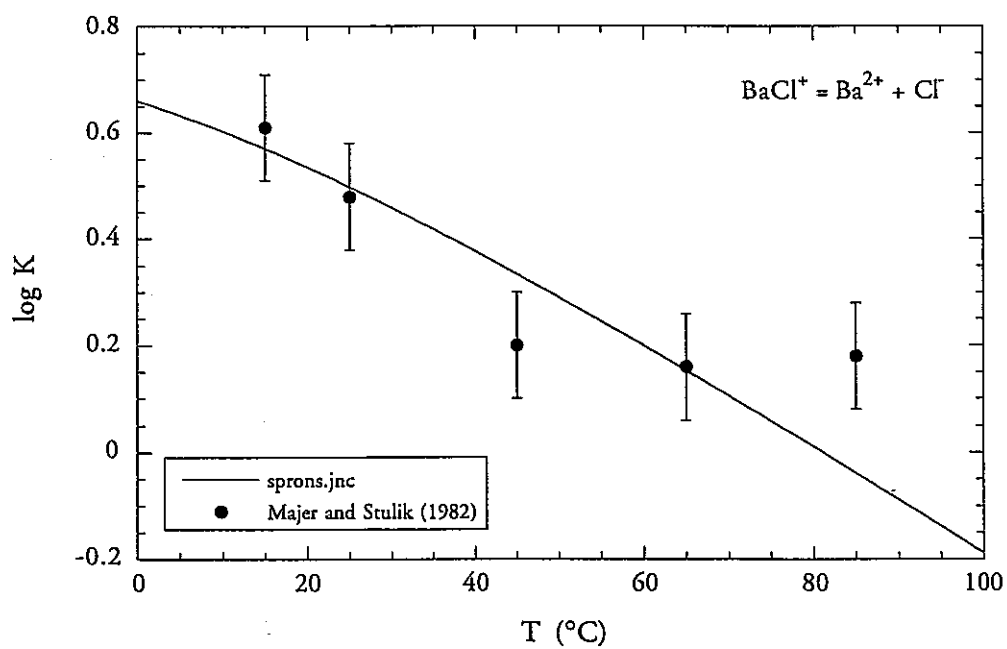


Fig. D.42: Logarithm of the equilibrium constant for the reaction $\text{BaCl}^+ \rightleftharpoons \text{Ba}^{2+} + \text{Cl}^-$ between 0 and 100°C at 1 bar.

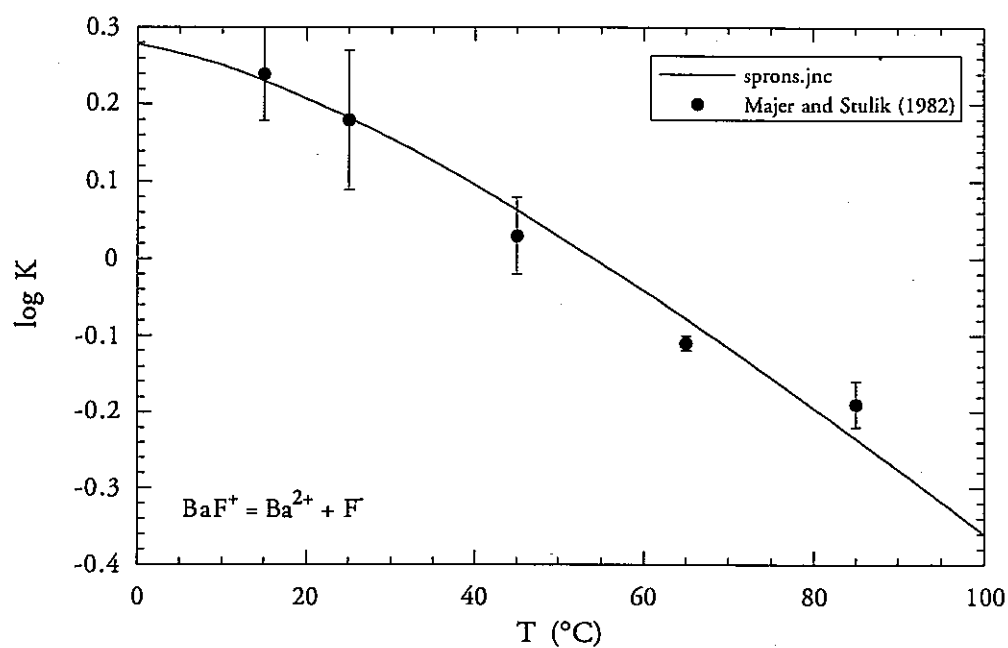


Fig. D.43: Logarithm of the equilibrium constant for the reaction $\text{BaF}^+ \rightleftharpoons \text{Ba}^{2+} + \text{F}^-$ between 0 and 100°C at 1 bar.

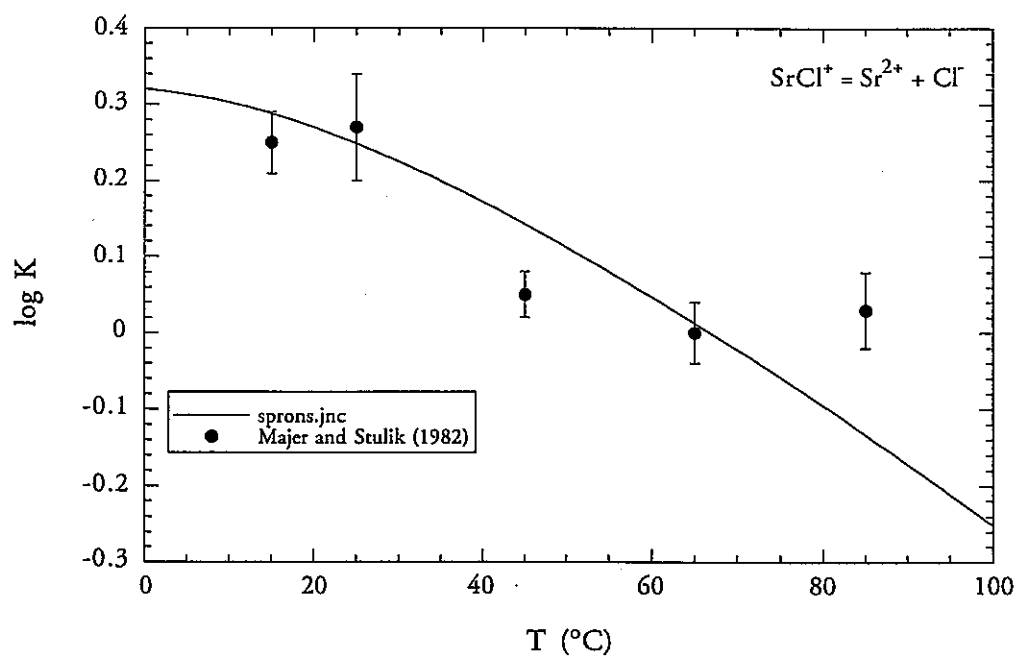


Fig. D.44: Logarithm of the equilibrium constant for the reaction $\text{SrCl}^+ \rightleftharpoons \text{Sr}^{2+} + \text{Cl}^-$ between 0 and 100°C at 1 bar.

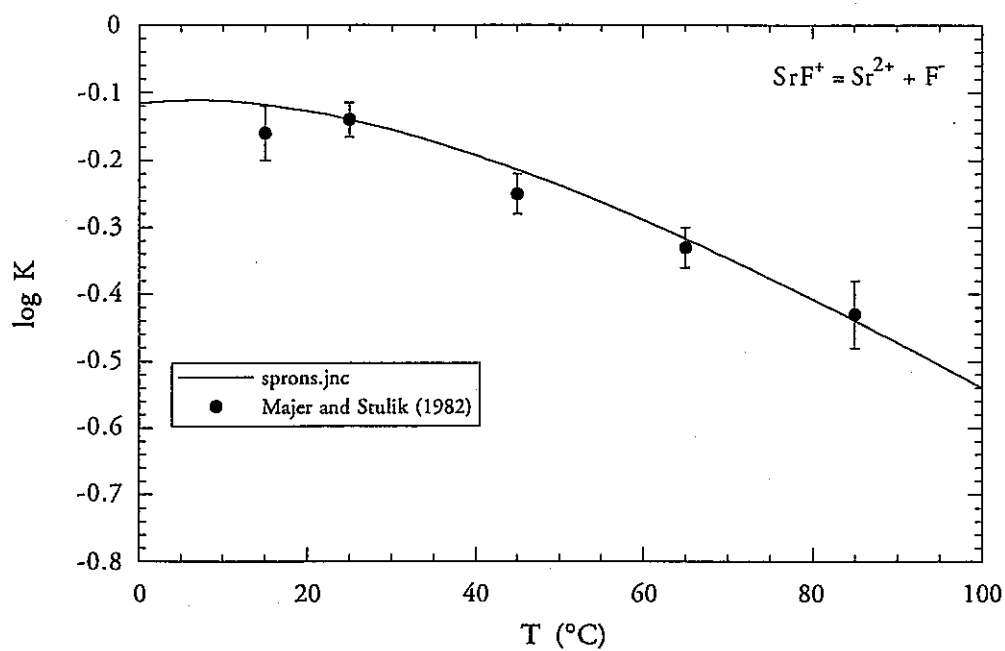


Fig. D.45: Logarithm of the equilibrium constant for the reaction $\text{SrF}^+ \rightleftharpoons \text{Sr}^{2+} + \text{F}^-$ between 0 and 100°C at 1 bar.

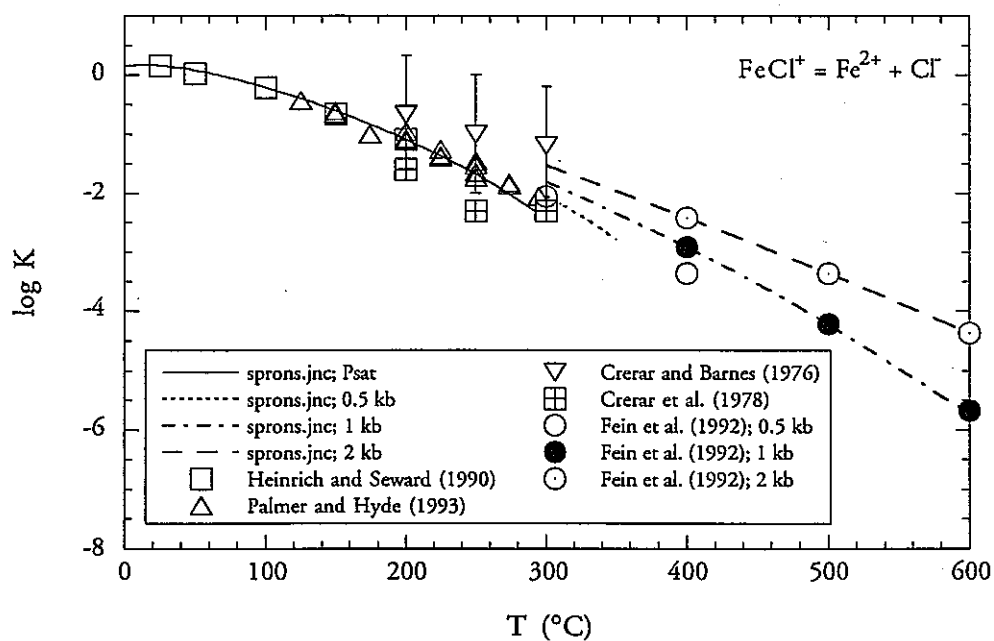


Fig. D.46: Logarithm of the equilibrium constant for the reaction $\text{FeCl}^+ \rightleftharpoons \text{Fe}^{2+} + \text{Cl}^-$ between 0 and 600°C and pressures from P_{sat} to 2 kb.

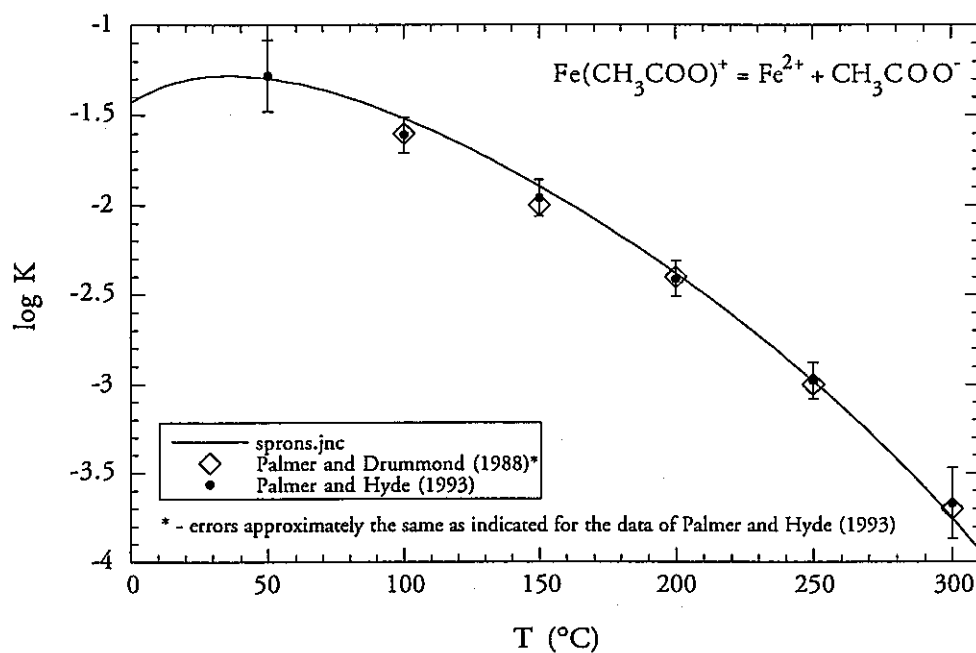


Fig. D.47: Logarithm of the equilibrium constant for the reaction $\text{Fe}(\text{CH}_3\text{COO})^+ \rightleftharpoons \text{Fe}^{2+} + \text{CH}_3\text{COO}^-$ between 0 and 300°C at P_{sat} .

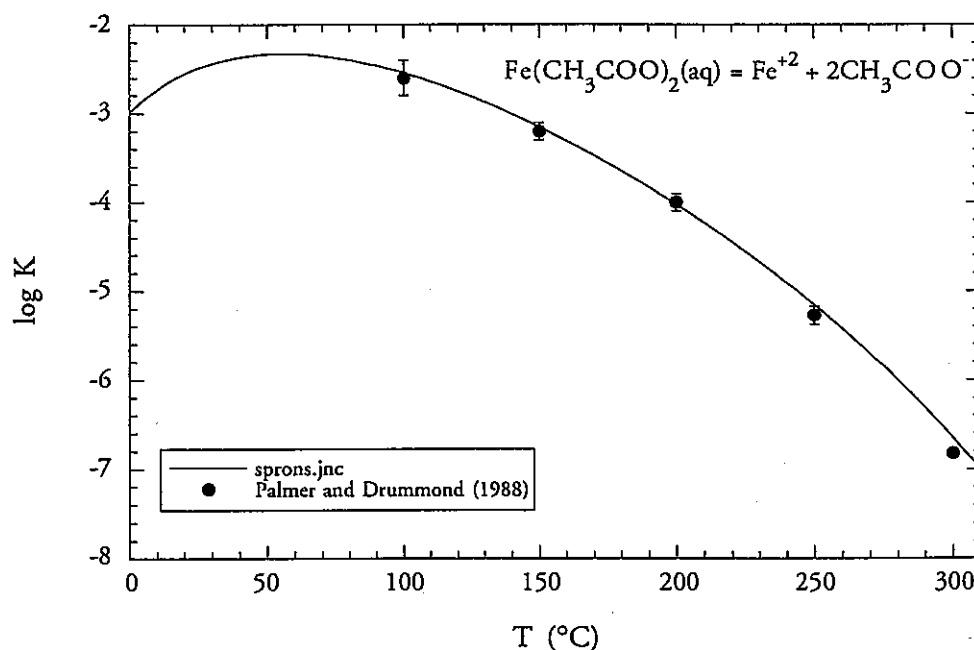


Fig. D.48: Logarithm of the equilibrium constant for the reaction $\text{Fe}(\text{CH}_3\text{COO})_2(\text{aq}) \rightleftharpoons \text{Fe}^{2+} + 2\text{CH}_3\text{COO}^-$ between 0 and 300°C at P_{sat} .

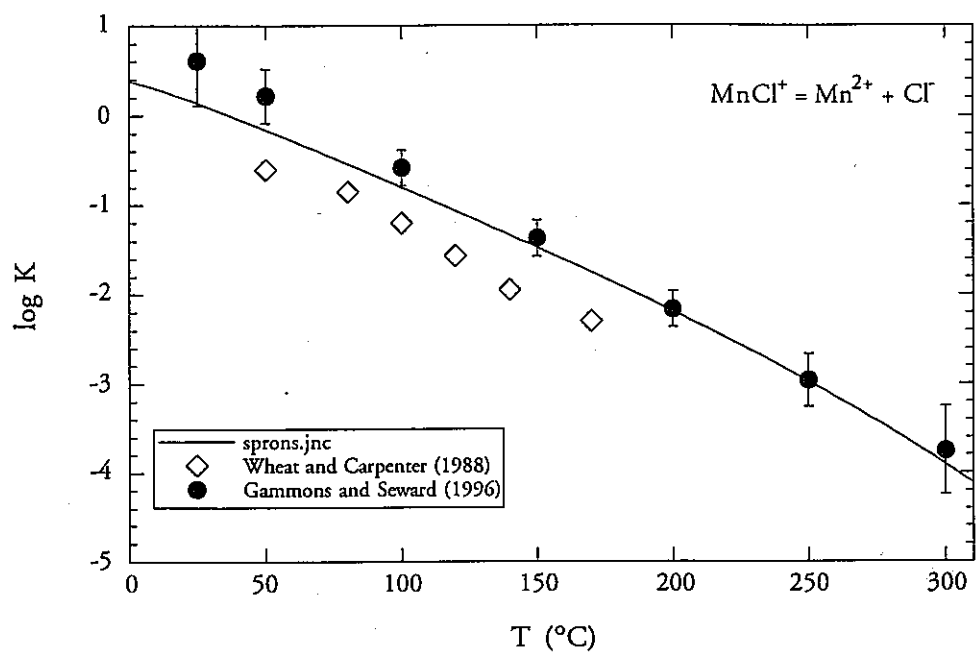


Fig. D.49: Logarithm of the equilibrium constant for the reaction $\text{MnCl}^+ \rightleftharpoons \text{Mn}^{2+} + \text{Cl}^-$ between 0 and 300 $^{\circ}\text{C}$ at P_{sat} .

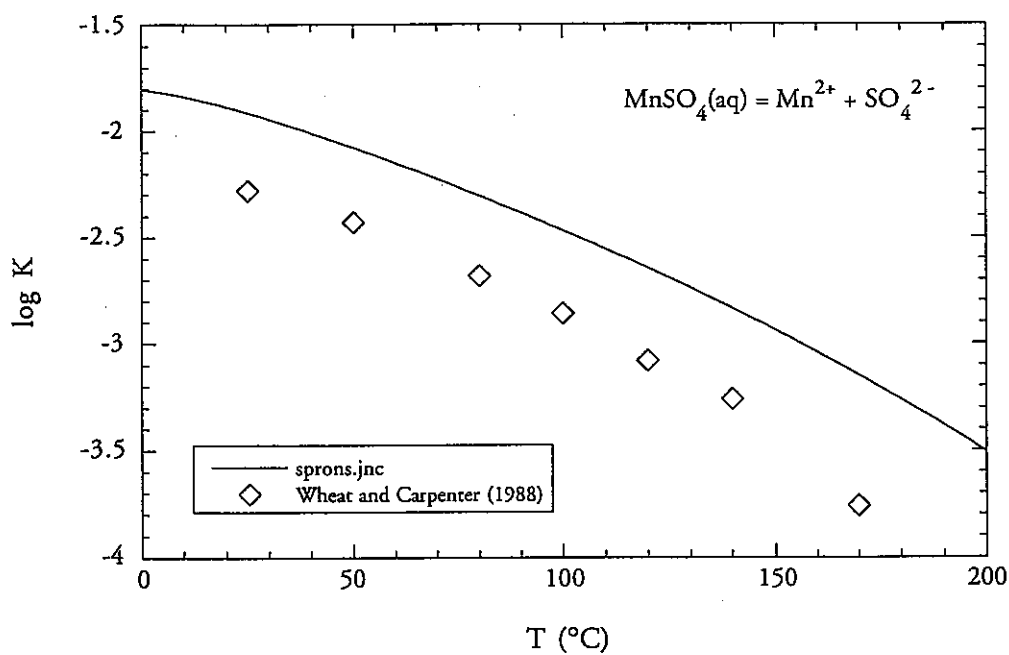


Fig. D.50: Logarithm of the equilibrium constant for the reaction $\text{MnSO}_4(\text{aq}) \rightleftharpoons \text{Mn}^{2+} + \text{SO}_4^{2-}$ between 0 and 200 $^{\circ}\text{C}$ at P_{sat} .

Appendix E: Calculation Examples:

- 1- Apparent conventional partial molal Gibbs free energy of formation of $\text{SiO}_2(aq)$ at 100°C and 1 bar**
- 2- Apparent molal Gibbs free energy of formation of quartz at 1 bar and the temperature of transition between α - and β -quartz**

Introduction

To illustrate how equations in the SUPCRT model are evaluated as functions of temperature and pressure, we calculate the apparent conventional partial molal Gibbs free energy of formation of $\text{SiO}_2(aq)$ at 100°C and 1 bar, and the apparent molal Gibbs free energy of formation of quartz at 574.85°C and 1 bar representing P - T conditions at which α -quartz is converted to β -quartz (Helgeson et al., 1978).

1 Example #1: Calculation of $\Delta\bar{G}_{\text{SiO}_2(aq), 1 \text{ bar}, 100^\circ\text{C}}^\circ$

The appropriate equation is (Section 4.1):

$$\Delta G_{P,T}^\circ \equiv \Delta G_f^\circ + (G_{P,T}^\circ - G_{P_r,T_r}^\circ), \quad (4.1_1)$$

where,

$$\begin{aligned} \bar{G}_{P,T}^\circ - \bar{G}_{P_r,T_r}^\circ = & -\bar{S}_{P_r,T_r}^\circ(T - T_r) - c_1 \left[T \ln \left(\frac{T}{T_r} \right) - T + T_r \right] + a_1(P - P_r) + a_2 \ln \left(\frac{\Psi + P}{\Psi + P_r} \right) \\ & - c_2 \left\{ \left[\left(\frac{1}{T - \Theta} \right) - \left(\frac{1}{T_r - \Theta} \right) \right] \left[\frac{\Theta - T}{\Theta} \right] - \frac{T}{\Theta^2} \ln \left[\frac{T_r(T - \Theta)}{T(T_r - \Theta)} \right] \right\} \\ & + \left(\frac{1}{T - \Theta} \right) \left[a_3(P - P_r) + a_4 \ln \left(\frac{\Psi + P}{\Psi + P_r} \right) \right] + \omega \left(\frac{1}{\epsilon} - 1 \right) - \omega_{P_r,T_r} \left(\frac{1}{\epsilon_{P_r,T_r}} - 1 \right) \\ & + \omega_{P_r,T_r} Y_{P_r,T_r}(T - T_r), \end{aligned} \quad (4.2.2_20)$$

and parameters are as defined in Sections 4.1 and 4.2.2. Evaluating the right-hand-side of the latter equation term-by-term using thermodynamic data for $\text{SiO}_2(aq)$ from Appendix A, and conventions in Table 5.1_3, gives:

first term;

$$-\bar{S}_{P_r,T_r}^\circ(T - T_r) = -18.0 \text{ (cal mol}^{-1}\text{K}^{-1}) (373.15 \text{ (K)} - 298.15 \text{ (K)}) = -1350 \text{ (cal mol}^{-1}\text{)},$$

second term;

$$\begin{aligned} -c_1 \left[T \ln \left(\frac{T}{T_r} \right) - T + T_r \right] = \\ -29.1 \text{ (cal mol}^{-1}\text{K}^{-1}) \left[373.15 \text{ (K)} \ln \left(\frac{373.15 \text{ (K)}}{298.15 \text{ (K)}} \right) - 373.15 \text{ (K)} + 298.15 \text{ (K)} \right] = \\ -254.01 \text{ (cal mol}^{-1}\text{)}, \end{aligned}$$

third term;

$$a_1(P - P_r) = 19.0(\text{cal mol}^{-1}\text{bar}^{-1})(1.013(\text{bar}) - 1(\text{bar})) = 0.25(\text{cal mol}^{-1}),$$

fourth term;

$$a_2 \ln \left[\frac{\Psi + P}{\Psi + P_r} \right] = 170(\text{cal mol}^{-1}) \ln \left[\frac{2600(\text{bar}) + 1.013(\text{bar})}{2600(\text{bar}) + 1.0(\text{bar})} \right] = 8.5 \times 10^{-4}(\text{cal mol}^{-1}),$$

(where Ψ is a pressure- and temperature-independent coefficient = 2600 bar, see Section 4.2.2.2)

fifth term;

$$\begin{aligned} -c_2 \left\{ \left[\left(\frac{1}{T - \Theta} \right) - \left(\frac{1}{T_r - \Theta} \right) \right] \left[\frac{\Theta - T}{\Theta} \right] - \frac{T}{\Theta^2} \ln \left[\frac{T_r(T - \Theta)}{T(T_r - \Theta)} \right] \right\} = \\ 51.2 \times 10^4(\text{cal K mol}^{-1}) \left\{ \left[\left(\frac{1}{373.15(\text{K}) - 228(\text{K})} \right) - \left(\frac{1}{298.15(\text{K}) - 228(\text{K})} \right) \right] \right\} \\ \left[\frac{228(\text{K}) - 373.15(\text{K})}{228(\text{K})} \right] - \frac{373.15(\text{K})}{228(\text{K})^2} \ln \left[\frac{298.15(\text{K})(373.15(\text{K}) - 228(\text{K}))}{373.15(\text{K})(298.15(\text{K}) - 228(\text{K}))} \right] = \\ 563.2(\text{cal mol}^{-1}), \end{aligned}$$

(where Θ is a pressure- and temperature-independent coefficient = 228°K, see Section 4.2.2.2)

sixth term;

$$\begin{aligned} \left(\frac{1}{T - \Theta} \right) \left[a_3(P - P_r) + a_4 \ln \left(\frac{\Psi + P}{\Psi + P_r} \right) \right] = \left(\frac{1}{373.15(\text{K}) - 228(\text{K})} \right) \\ \times \left[20(\text{cal K mol}^{-1}\text{bar}^{-1})(1.013(\text{bar}) - 1.0(\text{bar})) \right. \\ \left. - 2.7 \times 10^4(\text{cal K mol}^{-1}) \ln \left(\frac{2600(\text{bar}) + 1.013(\text{bar})}{2600(\text{bar}) + 1.0(\text{bar})} \right) \right] = \\ 9 \times 10^{-4}(\text{cal mol}^{-1}), \end{aligned}$$

seventh term;

$$\omega \left(\frac{1}{\varepsilon} - 1 \right) = 0.1291 \times 10^5(\text{cal mol}^{-1}) \left(\frac{1}{55.58} - 1 \right) = -12677.72(\text{cal mol}^{-1}),$$

[noting that $\omega = \omega_{p, T_r}$ at temperatures less than 448°K (Tanger and Helgeson, 1988), and that values of ε , the dielectric constant of $\text{H}_2\text{O}(l)$ (dimensionless) is taken from Weast (1979)]

eighth term;

$$-\omega_{P_r, T_r} \left(\frac{1}{\varepsilon_{P_r, T_r}} - 1 \right) = -0.1291 \times 10^5 \text{ (cal mol}^{-1}) \left(\frac{1}{78.36} - 1 \right) = 12745.25 \text{ (cal mol}^{-1}),$$

and,

ninth term;

$$\begin{aligned} \omega_{P_r, T_r} Y_{P_r, T_r} (T - T_r) = \\ [0.1291 \times 10^5 \text{ (cal mol}^{-1})] [-5.81 \times 10^{-5} \text{ (K}^{-1})] (373.15 \text{ (K)} - 298.15 \text{ (K)}) = \\ -56.25 \text{ (cal mol}^{-1}), \end{aligned}$$

where the solvent born function, Y_{P_r, T_r} , at the reference pressure and temperature equals $-5.81 \times 10^{-5} \text{ K}^{-1}$ (Tanger and Helgeson, 1988).

Adding the results calculated above for the nine terms on the right-hand-side of Eqn. (4.2.2_20) gives:

$$\bar{G}_{P, T}^{\circ} - \bar{G}_{P_r, T_r}^{\circ} = -1029 \text{ cal mol}^{-1}.$$

Substituting this value into Eqn. (4.1_1) with $\Delta G_{f, \text{SiO}_2(aq)}^{\circ}$ from Appendix A gives the apparent partial molal Gibbs free energy of $\text{SiO}_2(aq)$ formation at 100°C and 1 bar:

$$\begin{aligned} \Delta \bar{G}_{P, T}^{\circ} &= \Delta G_f^{\circ} + (\bar{G}_{P, T}^{\circ} - \bar{G}_{P_r, T_r}^{\circ}) = -199190 \text{ (cal mol}^{-1}) - 1029 \text{ (cal mol}^{-1}) \\ &= -200219 \text{ cal mol}^{-1} \end{aligned}$$

This agrees closely with the value calculated using SUPCRT and SPRONS.JNC ($-200229 \text{ cal mol}^{-1}$). The minor discrepancy may be due to differences in precision between the two calculation approaches, and/or to differences in the dielectric constant of $\text{H}_2\text{O}(l)$ at 25 and 100°C as calculated in SUPCRT *vs.* the values reported by Weast (1979).

2 Example #2: Calculation of $\Delta G_{\text{quartz, 1 bar, 574.85}^{\circ}\text{C}}^{\circ}$

The appropriate equation is (Section 4.1):

$$\Delta G_{P, T}^{\circ} \equiv \Delta G_f^{\circ} + (\bar{G}_{P, T}^{\circ} - \bar{G}_{P_r, T_r}^{\circ}), \quad (4.1_1)$$

where,

$$\begin{aligned}
G_{P,T}^{\circ} - G_{P_r,T_r}^{\circ} = & -S_{P_r,T_r}^{\circ}(T - T_r) + \sum_{i=1}^{1+\phi_T} a_i \left\{ T_{i+1} - T_i - T_{i+1} \ln \left(\frac{T_{i+1}}{T_i} \right) \right\} \\
& + \sum_{i=1}^{1+\phi_T} \left\{ \frac{(-c_i - b_i T_{i+1} T_i^2)(T_{i+1} - T_i)^2}{2 T_{i+1} T_i^2} \right\} + V_{P_r,T}^{\circ}(P - P_r) \\
& - \sum_{i=1}^{\phi_P} \int_{P_{i,T}}^P \Delta V_{i,i}^{\circ} dP - \sum_{i=1}^{\phi_T} \frac{\Delta H_{i,i}^{\circ}}{T_{i,i}} (T - T_{i,i});
\end{aligned}
\tag{4.2.1_6}$$

and parameters are as defined in Sections 4.1 and 4.2.1. The right-hand-side of the latter equation can be evaluated term-by-term using thermodynamic data for quartz from Appendix A, conventions specified in Table 5.1_2, and the constraints $T_i = T_r = 298.15^{\circ}\text{K}$ and $T_{i+1} = 848^{\circ}\text{K}$ (i.e., 574.85°C - the temperature of transition from α -quartz to β -quartz at 1 bar):

first term;

$$-S_{P_r,T_r}^{\circ}(T - T_r) = -9.88(\text{cal mol}^{-1}\text{K}^{-1})(848(\text{K}) - 298.15(\text{K})) = -5432.4(\text{cal mol}^{-1}),$$

second term;

$$\begin{aligned}
a_1 \left\{ T_{i+1} - T_i - T_{i+1} \ln \left(\frac{T_{i+1}}{T_i} \right) \right\} = \\
11.22(\text{cal mol}^{-1}\text{K}^{-1}) \left\{ 848(\text{K}) - 298.15(\text{K}) - 848(\text{K}) \ln \left(\frac{848(\text{K})}{298.15(\text{K})} \right) \right\} \\
= -3776.1(\text{cal mol}^{-1}),
\end{aligned}$$

[noting that although $\phi_T = 1$ in this example, the second summation implicit in this term does not apply because T_{i+1} then equals T_i (i.e., 848°K)],

third term;

$$\begin{aligned}
\left\{ \frac{(-c_i - b_i T_{i+1} T_i^2)(T_{i+1} - T_i)^2}{2 T_{i+1} T_i^2} \right\} = \\
\left\{ \frac{(2.7 \times 10^5 - 8.2 \times 10^{-3} \times 848(\text{K}) \times 298.15(\text{K})^2)(848(\text{K}) - 298.15(\text{K}))^2}{2 \times 848(\text{K}) \times 298.15(\text{K})} \right\} = \\
-698.1(\text{cal mol}^{-1}),
\end{aligned}$$

[noting that units for c_i and b_i are cal K mol^{-1} and $\text{cal mol}^{-1} \text{K}^{-2}$, respectively, and that although $\phi_T = 1$, the second summation in this term does not apply because T_{i+1} then equals T_i (i.e., 848°K)],

fourth term;

$$V_{P_r, T}^\circ (P - P_r) = 0,$$

(because the pressure at the transition temperature in this example is equal to P_r)

fifth term;

$$-\sum_{i=1}^{\phi_P} \int_{P_{t,i,T}}^P \Delta V_{t,i}^\circ dP = 0,$$

(because the pressure at the transition temperature in this example is equal to P_r), and

sixth term;

$$-\sum_{i=1}^{\phi_T} \frac{\Delta H_{t,i}^\circ}{T_{t,i}} (T - T_{t,i}) = 0$$

(because $T = T_{t,i}$).

Adding the results calculated above for the six terms on the right-hand-side of Eqn. (4.2.1_6) gives:

$$\bar{G}_{P,T}^\circ - \bar{G}_{P_r, T_r}^\circ = -9907 \text{ cal mol}^{-1}.$$

Substituting this value into Eqn. (4.1_1) with $\Delta G_{f, \text{quartz}}^\circ$ from Appendix A gives the apparent standard molal Gibbs free energy of formation quartz at 574.85°C and 1 bar:

$$\begin{aligned} \Delta \bar{G}_{P,T}^\circ &= \Delta G_f^\circ + (\bar{G}_{P,T}^\circ - \bar{G}_{P_r, T_r}^\circ) = -204646 \text{ (cal mol}^{-1}) - 9907 \text{ (cal mol}^{-1}) \\ &= -214553 \text{ cal mol}^{-1}, \end{aligned}$$

which is in exact agreement with the value calculated using SPRONS.JNC and SUPCRT.